

CONTRATO

IICA / INDRHI / CSU



INSTITUTO INTERAMERICANO
DE COOPERACION PARA LA
AGRICULTURA (IICA)

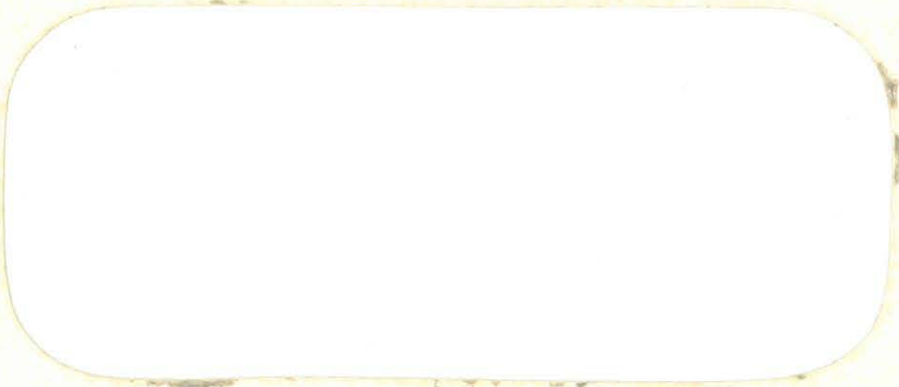


INSTITUTO NACIONAL DE RECURSOS
HIDRAULICOS (INDRHI)



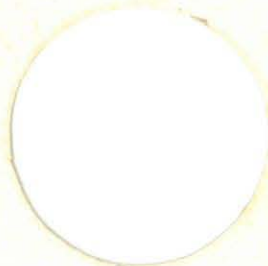
UNIVERSIDAD DEL
ESTADO DE COLORADO
(CSU)

ESTUDIOS SOBRE LA OPERACION Y SEGURIDAD DEL SISTEMA DE EMBALSES DE VALDESIA



IICA-
PM-A1/
00-86-
02-25

DOCUMENTO No.

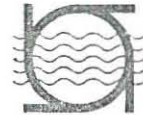




INSTITUTO INTERAMERICANO
DE COOPERACION PARA LA
AGRICULTURA (IICA)

CONTRATO
IICA / INDRHI / CSU

IICA-CIDIA



INSTITUTO NACIONAL DE RECURSOS
HIDRAULICOS (INDRHI)

Centro Interamericano de
Documentación e
Información Agrícola
? 0 1 ? 1993
IICA — CIDIA



UNIVERSIDAD DEL
ESTADO DE COLORADO
(CSU)

REPUBLICA DE VENEZUELA
INSTITUTO VENEZOLANO DE INVESTIGACIONES CIENTÍFICAS Y TECNOLÓGICAS
LIBRERÍA

SECOND TECHNICAL REPORT ^{1/}
VOLUME I
HYDROLOGICAL STUDIES

25
June
1986

^{1/} Additional comments to this report
are included in Addendum (Document
No. 32)

3V 00-0202

11CA
OM-A1/DO
86-002
no. 25

00001511



Technical Report on the Project

OPERATIONAL AND SAFETY STUDIES OF THE VALDESIA RESERVOIR

for

Contract IICA/INDRHI/CSU
(Loan 1655-DO from World Bank)

VOLUME I

HYDROLOGIC STUDIES

by

Jayantha T. B. Obeysekera
Guillermo Q. Tabios III
Fernando A. Pons
Jose D. Salas
Hseih Wen Shen

VOLUME I.
HYDROLOGIC STUDIES

TABLE OF CONTENTS

1.1	INTRODUCTION.....	I-1
1.2	METHODS OF INVESTIGATION.....	I-1
1.3	SUMMARY OF CONCLUSIONS.....	I-3
1.4	ORGANIZATION OF THE VOLUME.....	I-4
1.5	NIZAO WATERSHED.....	I-6
1.5.1	Physiography.....	I-6
1.5.2	Vegetation.....	I-6
1.5.3	Hydrometeorological Data.....	I-6
1.5.4	Streamflow Data.....	I-13
1.5.5	Quality of Data.....	I-25
	Appendix 1.5.A. Calibration of Stage-Discharge Curves.....	I-29
1.6	DESIGN STORMS.....	I-44
1.6.1	Historic Storms.....	I-44
1.6.2	Ishyetal Mapping.....	I-44
1.6.3	Development of Depth-Area-Duration (DAD) Curves.....	I-49
1.6.4	Standard Project Storm (SPS).....	I-56
1.6.5	Temporal Distribution of SPS.....	I-60
1.6.6	Probable Maximum Precipitation (PMP).....	I-60
	Appendix 1.6.A Mass Curves of Rainfall for Selected Storms.....	I-63
	Appendix 1.6.B Program Listing for Rainfall Isohyet Computations.....	I-89
	Appendix 1.6.C. Depth-Area-Duration Curves for Selected Storms..	I-121
1.7	RAINFALL RUNOFF MODELING.....	I-147
1.7.1	Selection of Model.....	I-147
1.7.2	HEC-1 Model Calibration.....	I-147
1.7.3	Design Flood Hydrographs.....	I-159
1.7.4	Reconstruction of Hydrographs for Hurricane DAVID	I-161
1.7.5	Effects of Natural Storages in the Watershed.....	I-161
	Appendix 1.7.A. Sample Output of Rainfall Weighting Using Thiessen Polygon Method.....	I-169
	Appendix 1.7.B. Results of HEC-1 Model Run Using Hurricane DAVID Data to Assess Effects of Upstream Natural Storages.....	I-191
1.8	STREAMFLOW FORECASTING MODEL.....	I-265
1.8.1	Selection of Model.....	I-265
1.8.2	Development of SACKW Model.....	I-266
1.8.3	Description of Model Components.....	I-267
1.8.3.1	Sacramento Soil Moisture Accounting Model.....	I-267
1.8.3.2	Kinematic Wave Routing Model.....	I-271
1.8.3.3	Watershed Partitioning and Timing Considerations.	I-273
1.8.4	Model Calibration.....	I-275
1.8.5	Model Testing by Forecasting.....	I-299
1.8.6	References.....	I-312
	Appendix 1.8.A. SACKW Model User's Manual.....	I-319
	1.8.A.1. Introduction.....	I-319
	1.8.A.2 Program Description.....	I-319
	1.8.A.3 Input and Output Information.....	I-323

1.8.A.4	Some Guidelines for Modle Usage and Parameter Calibration.....	I-324
1.8.A.5	Sample Model Application.....	I-327
Appendix 1.8.B.	Program Listing of SACKW Model.....	I-341
Appendix 1.8.C.	Program Input Requirements and Description.....	I-386
1.9.	STOCHASTIC GENERATION OF STREAMFLOWS AND TURBINE OPERATING HOURS.....	I-396
1.9.1	Introduction.....	I-396
1.9.2	Description of Hydrologic Data Used.....	I-396
1.9.3	Filling-in and Extension of Historical Data.....	I-405
1.9.3.1	Filling-in of Missing Data at Palo De Caja.....	I-405
1.9.3.2	Extension of Records of Paso Del Ermitano and Rancho Arriba.....	I-409
1.9.4	Stochastic Modeling of Streamflow.....	I-414
1.9.5	Streamflow Data Generation.....	I-417
1.9.5.1	Data Generation Scheme.....	I-417
1.9.5.2	Analysis of Generated Data.....	I-418
1.9.6	Modeling and Generation of Turbine Operating Hours.....	I-419
1.9.7	Final Remarks.....	I-422
1.9.8	References.....	I-424
Appendix 1.9.A.	Standardization and Normalization.....	I-425
Appendix 1.9.B.	Historical and Extended Series Statistics of Monthly and Weekly Data of Paso Del Ermitano and Rancho Arriba.....	I-431
Appendix 1.9.C.	Fourier Series Fitting of Periodic Statistical Parameters.....	I-492
Appendix 1.9.D.	Historical (Extended Series) and Generated Monthly and Weekly Statistics for Paso Del Ermitano (PASODE), Palo De Caja (PALODE) and Rancho Arriba (RANCHO) Using Models A, B, and C in the Original Domain of Flows.....	I-513
Appendix 1.9.E.	Historical (Extended Series) and Generated Monthly and Weekly Statistics of Palo De Caja, Paso Del Ermitano and Rancho Arriba for Model B in the Original, Logarithmic and Log-Wilson-Hilferty Domain of Flows.....	I-550
Appendix 1.9.F.	Historical and Generated Statistics of Monthly Turbine Operating Hours Time Series of Valdesia Reservoir.....	I-661

VOLUME I. HYDROLOGIC STUDIES

1.1 INTRODUCTION

The Valdesia reservoir system, located on the Nizao River in the Dominican Republic, was designed to provide irrigation water to the Nizao project areas and hydroelectric energy to the national electrical network system. The reservoir system consists of a main reservoir, dam and spillway, a power plant and outflow regulating works, together with an afterbay, diversion and spillway system a short distance downstream.

The study on the operational management of the Valdesia system reported in a series of volumes including this one, involved several interrelated areas. This volume reports in detail the basic hydrologic studies including rainfall-runoff modeling, flood forecasting and stochastic data generation which are essential components of the entire study. The products reported in this volume are used for developing emergency and normal operation plans for the Valdesia system.

1.2 METHODS OF INVESTIGATION

Hydrologic studies are a prerequisite to any water resources management project. The required studies necessary for this project can be categorized under four topics:

1. Design storms
2. Rainfall-runoff modeling
3. Streamflow forecasting
4. Stochastic data generation

Hypothetical design storms are required to compute hypothetical floods which are necessary for developing emergency operating procedures. Two types of hypothetical storms are considered: (a) Standard Project Storms (SPS); and (b) Probable Maximum Precipitation.

The rationale behind the use of SPS is discussed in many documents of the U.S. Army Corps of Engineers. Both types of hypothetical storms require depth-area-duration curves. These are developed from about 25 observed historic storms. Since two regimes of storms, namely hurricane and non-hurricane, are present in the Dominican Republic, two types of SPS are developed. The time distribution of SPS is also derived from the observed storms. The PMP is based on the Hurricane model of the U.S. National Weather Service (U.S. Weather Bureau, 1961) modified by the counterparts in the Dominican Republic.

An event type rainfall-runoff model suitable for conditions in the Nizao basin is necessary to compute hypothetical floods from hypothetical storms. The HEC-1 model of the U.S. Army Corps of Engineers is used for this purpose. It is calibrated by using data of historic storms and floods and a few flood data derived from Valdesia reservoir levels during storm events. The calibrated model is used to compute hypothetical floods for three different antecedent basin conditions. The model is also used to reconstruct the possible hydrographs from the precipitation that occurred during the Hurricane DAVID which struck the island on August 30, 1986.

For real-time forecasting, a modified version of the U.S. National Weather Service River Forecast Model is employed. The modification is necessary to (a) develop a version which will fit in the computing facilities at INDRHI/CDE; and (b) incorporate kinematic wave flood routing procedure in the model. The model is calibrated using several years of daily streamflow data and hourly precipitation data.

Synthetically generated data are necessary to develop and test the normal operating rules. Multivariate stochastic models of streamflow of

three gauging stations (Ermitano, Palo de Caja, Rancho Arriba) are developed from extended existing historic data. The generated data at Ermitano is employed to generate another synthetic series of number of hours of energy generation at Valdesia dam. Both the series are used for developing and testing of normal operating rules.

1.3 SUMMARY OF CONCLUSIONS

Following is a summary of conclusions based on the hydrologic studies reported in this section:

1. Significant amount of errors and inconsistencies are found in various precipitation and streamflow data collected. A careful review of data collection processing and reporting of all hydrologic data is warranted.

2. The standard project storm (for 48-hour duration) based on non-hurricane storms is 260 mm whereas the same based on hurricane precipitation is 493 mm.

3. The watershed average PMP simulated by the hurricane model of U.S. Weather Bureau is 1338 mm.

4. The nonhurricane SPS simulated a peak inflow to Valdesia reservoir of $3469 \text{ m}^3/\text{sec}$ for dry antecedent conditions whereas the same for wet antecedent conditions is $7544 \text{ m}^3/\text{s}$.

5. The hurricane SPS simulated a peak inflow to Valdesia reservoir of $10185 \text{ m}^3/\text{s}$ for dry antecedent conditions whereas the same for wet antecedent conditions is $16548 \text{ m}^3/\text{s}$.

6. The flood peak simulated by using the calibrated HEC-1 and observed precipitation during hurricane DAVID is $5332 \text{ m}^3/\text{s}$ for dry antecedent conditions and it is $10358 \text{ m}^3/\text{s}$ for wet antecedent

conditions. The calibrated SAC-KW model resulted in peak of 7074 cms for the same event.

7. The PMP supplied by counterparts produces a PMF of 20,000 cms for dry antecedent conditions. For wet antecedent conditions it increases to 23,000 cms.

8. The calibration of the developed flood forecasting model for Nizao basin was performed on a year-to-year basis using the data from 1972 to 1975. It is concluded from this exercise that the best model calibration is in year 1972. Subsequent use of the 1972-model parameters to forecast the streamflow regime during Hurricane DAVID (August, 1979) gave a highest hourly peak flow at Paso del Ermitano of 7074 m³/sec. For stochastic modeling and data generation, a trivariate, contemporaneous first-order autoregressive process with seasonal parameters has been found to adequately describe the streamflows of Rancho Arriva, Palo de Caja and Paso del Ermitano. On both monthly and weekly levels, the streamflows are concluded to be log-Pearson Type III distributed as indicated by the normalizing transformation used which is the combination of logarithmic and Wilson-Hufferty transformations. A similar model structure has also been found for modeling and generation of turbine operating hours time series at the Valdesia reservoir with bivariate dependence on Paso del Ermitano streamflows. Knowledge of the time series structure of the above said processes could be valuable to future analysis and applications such as data transposition, regional flood estimation and sampling frequency design.

1.4 ORGANIZATION OF THE VOLUME

The work involved in hydrologic studies are reported in five subsections. The general characteristics of the watershed physiography,

vegetation, etc. and the details of availability and quality of data are included in Section 1.5. The development of hypothetical storms Standard Project Storm and the Probable Maximum Precipitation is discussed in Section 1.6. The details of rainfall-runoff modeling including the calibration of the selected HEC-1 model and its application to compute hypothetical floods are included in Section 1.7. In Section 1.8, the development, calibration, testing and application of the SAC-KW model for real-time flood forecasting is discussed. Finally, Section 1.9 deals with the development and application of the stochastic models of streamflow and number of hours of energy generation.

1.5 NIZAO WATERSHED

1.5.1 Physiography

The Nizao Watershed is located in the south central part of the Dominican Republic (Figure 1.5.1). The drainage area up to the Valdesia reservoir which is located approximately 50 km away from the confluence of Nizao river and the Atlantic Ocean is about 900 sq. km. The watershed has a distinct elongated shape with a predominant orientation in the NW-SE direction (see Figure 1.5.2). Most of the headwater areas have high relief with main channel slopes reaching as much as 8 to 10 percent. The drop in elevation from the highest point to the Valdesia dam site is about 2500 meters. No significant flood plains exist in the entire Nizao watershed.

1.5.2 Vegetation

The watershed is covered primarily with forest and pasture. Less than 8 percent of the watershed is covered with agricultural lands.

1.5.3 Hydrometeorological Data

Precipitation data: A list of the precipitation data received from INDRHI and the Meteorology Agency is included in Table 1.5.1. Table 1.5.2 shows the available hourly precipitation data that was obtained in computer tape. For purposes of data analysis the nine computer files corresponding to these stations were used to create yearly files containing the hourly data. The precipitation data availability for station in and around Nizao watershed is shown in the form of a bar chart in Figure 1.5.3.

Climatological Data: Table 1.5.3 shows the stations in and around Nizao, for which climatological data is available. The evaporation data was used in the calibration of the real-time streamflow forecast model.

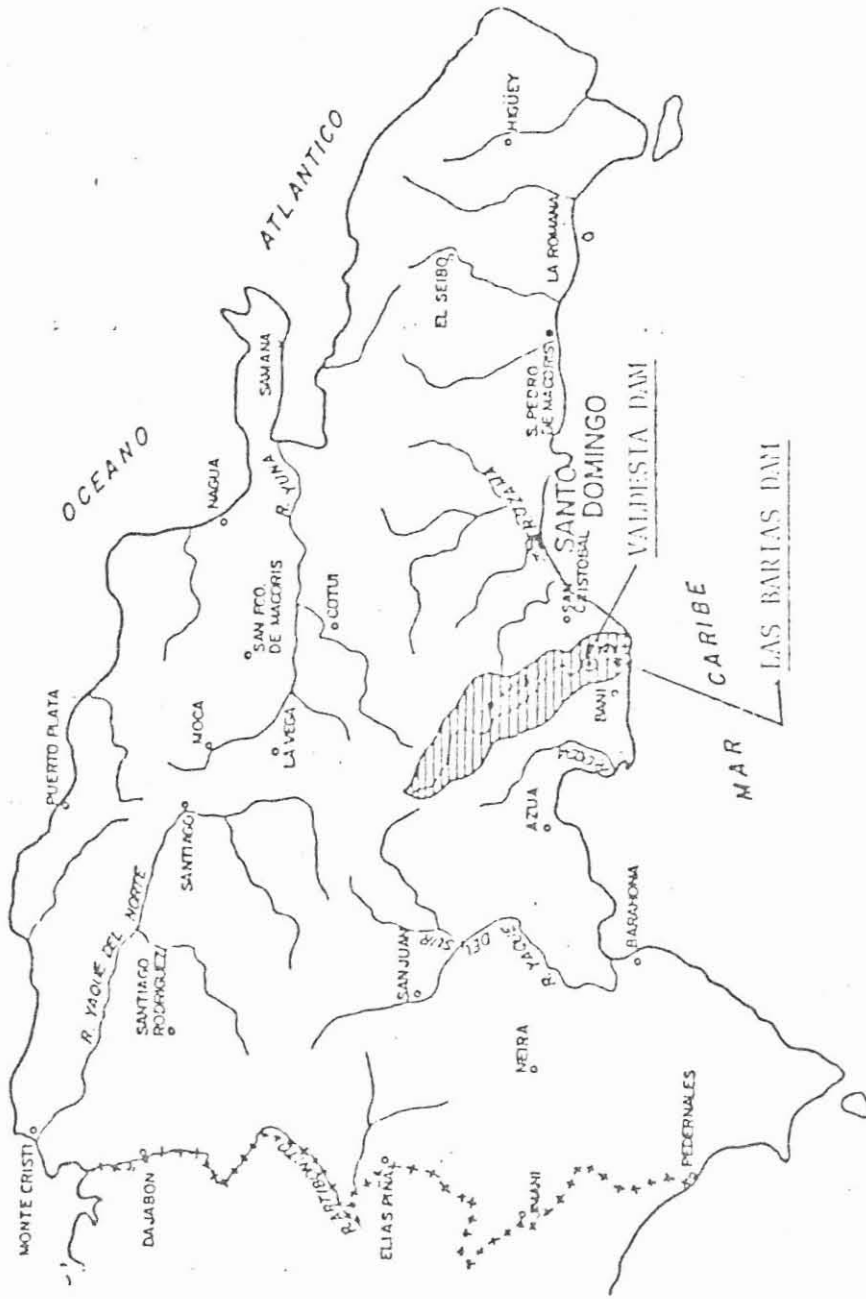


Figure 1.5.1. Location of the Nizao watershed in the Dominican Republic.

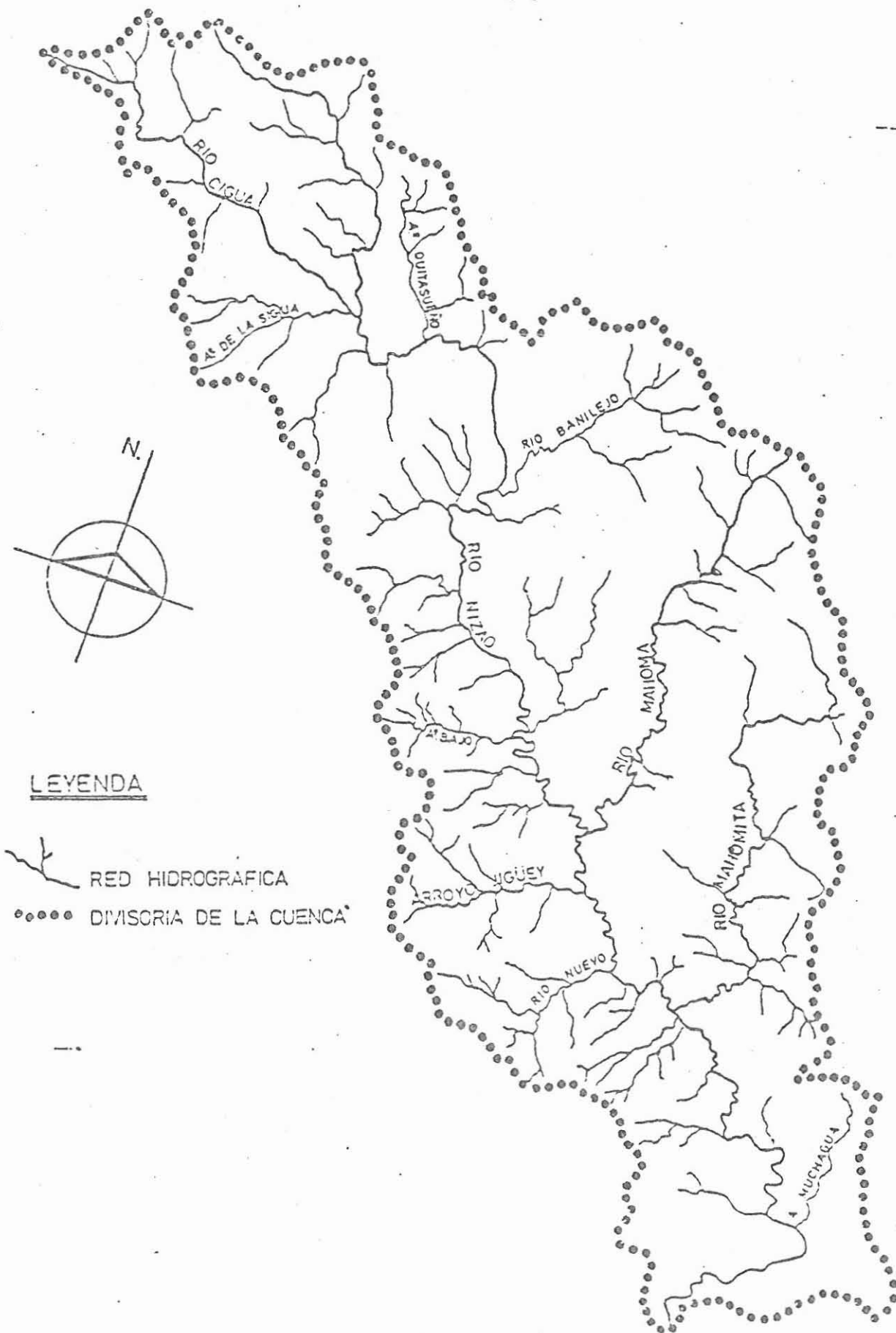


Figure 1.5.2. Nizac river basin.

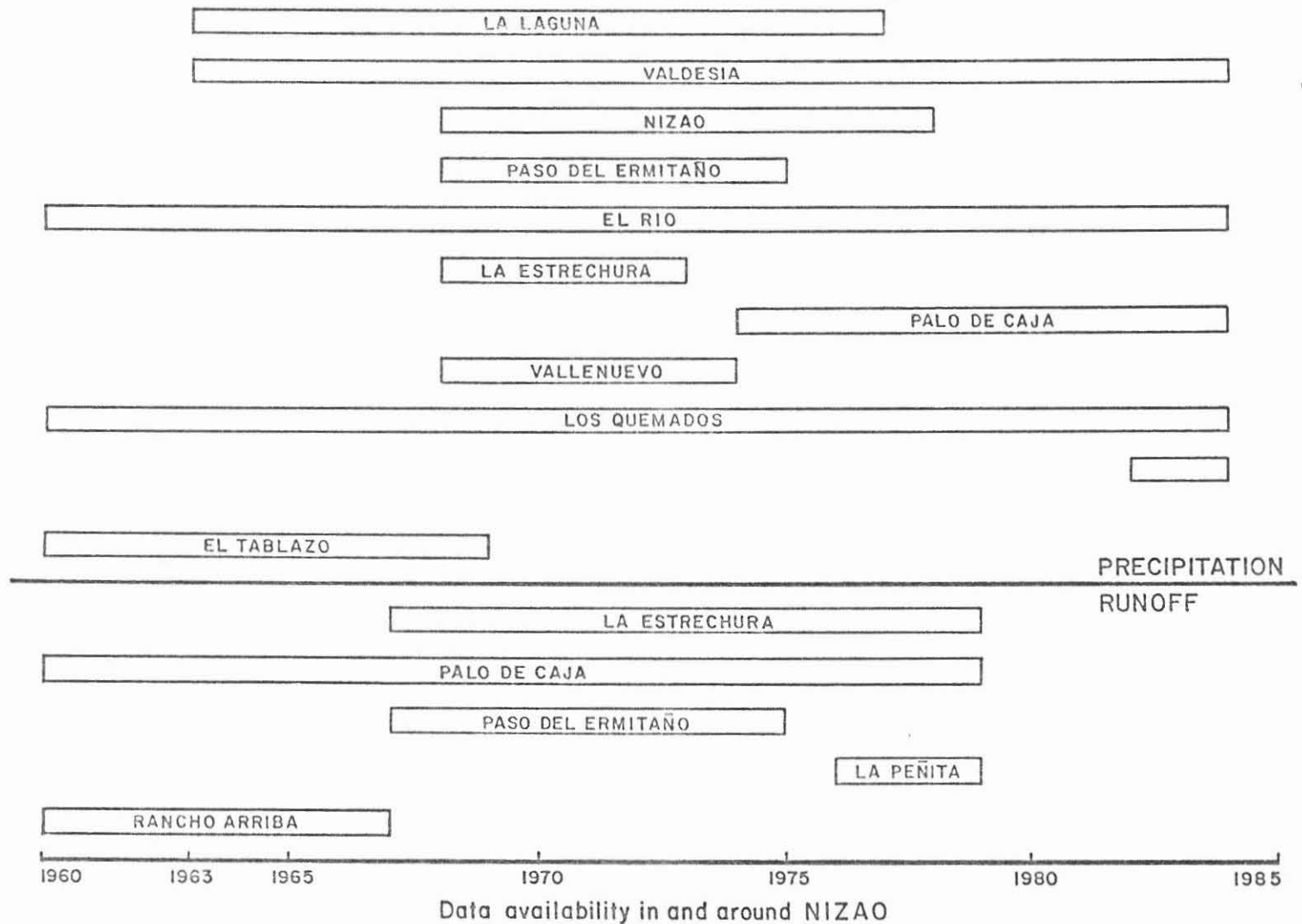


Figure 1.5.3. Precipitation and runoff data availability for stations in and around Nizao watershed.

TABLE 1.5.1 PRECIPITATION DATA

Station Name	Basin	Type	Start	End	Lat.	Lon.	Tap
La Laguna	Nizao	Day	12/62	12/77	18°32'30"	70°24'45"	Yes
Nizao	Nizao	Day	1/68	3/78	18°36'53"	70°27'07"	Yes
Paso Del Ermitano	Nizao	Day	4/68	11/75	18°26'00"	70°16'00"	Yes
Los Cacaos	Nizao	Day	8/67	7/70	18°31'40"	70°18'00"	Yes
Azua Hatillo	Ocoa	Day	8/69	3/84	18°23'40"	70°32'20"	Yes
Valdesia	Nizao	Day	2/63	7/84	18°24'30"	70°16'50"	Yes
La Estrechura	Nizao	Day	1/68	12/73	18°43'40"	70°29'00"	Yes
Presas Mana	Haina	Day	11/82	7/84	18°36'28"	70°12'55"	Yes
Presas Isa	Haina	Day	11/82	7/84	18°36'28"	70°12'32"	Yes
Quija Quieta	Nizao	Day	10/76	4/79	18°13'49"	70°27'31"	Yes
Engombre	Haina	12 hr	1/77	7/83	18°27'00"	70°00'07"	No
Palo De Caja	Nizao	Day	5/74	10/84	18°31'50"	70°24'00"	Yes
Valle Nuevo	Y. Del Sur	Day	1/68	6/74	18°49'27"	70°40'58"	Yes
Constanza	Y. Del Sur	Day	1/68	12/79	18°54'40"	70°43'00"	Yes
Guayabal	Y. Del Sur	Day	3/79	9/84			Yes
Los Quemados	Yuna	Day	1/60	10/84	18°53'30"	70°27'30"	Yes
Juma-Bonao	Yuna	Day	12/70	9/84	18°54'00"	70°23'10"	Yes
El Río (Constanza)	Y. Del Sur	Day	6/60	9/84	18°58'30"	70°37'40"	Yes
Esta Bania	Grande Del Med.	Day	9/69	10/84	18°27'20"	70°38'45"	Yes
El Tablazo	Nigua	Day	8/60	1/69	18°29'10"	70°10'50"	Yes
Rancho Arriba	Nizao	Mon	1/39	12/80	18°42'	70°27'	No
Padre Las Casas	Y. Del Sur	Mon	1/38	12/83			No
Bani	Bani	Mon	1/36	12/83	18°16'	70°20'	No
Villa Autagracia	Haina	Mon	1/38	12/83			No
Azua	Via	Mon	1/31	12/83			No
Valdesia	Nizao	Hour	2/63	5/83	18°24'30"	70°16'50"	No
La Laguna	Nizao	Hour	12/62	11/77	18°32'30"	70°24'45"	No
Nizao	Nizao	Hour	1/63	4/78	18°36'53"	70°27'07"	No
Medina	Haina	12 hr	10/79	7/84	18°32'06"	70°08'40"	No
Quija Quieta	Nizao	12 hr	4/79	4/79	18°13'49"	70°27'31"	No
Medina	Haina	Day	3/76	12/84	18°32'06"	70°08'40"	Yes
Rancho Arriba	Nizao	Day	3/39	12/84	18°42'	70°27'	No
Padre Las Casas	Y. Del Sur	Day	10/38	12/84			No
Bani	Bani	Day	1/36	12/84	18°16'	70°20'	No
Villa Altagracia	Haina	Day	8/38	12/84			No
Azua	Via	Day	1/31	12/84			No

TABLE 1.5.2 HOURLY PRECIPITATION

STATION NAME	BASIN	START	END
Valdesia	Nizao	2/63	5/83
La Laguna	Nizao	12/62	5/78
Nizao	Nizao	1/63	4/78
Engombe	Haina	5/72	6/84
Palo De Caja	Nizao	2/79	9/83
Valle Nuevo	Y. Del Sur	9/77	3/83
El Eio (Const.)	Y. Del Sur	1/77	12/84
Los Quemados	Yuna	1/65	7/84
Juma-Bonao	Yuna	7/71	5/82

TABLE 1.5.3 CLIMATOLOGICAL DATA

CLIMATOLOGICAL REPORTS: Precipitation, evaporation, temperature, humidity, wind speed, cloudiness, radiation, pressure) (printouts)

CODE	STATION	BASIN	START	END	LAT.	LON.
34001	Engombe	Haina	10/68	7/84	18°27'00"	70°00'07"
34002	Medina	Haina	10/79	7/84	18°32'06"	70°08'40"
38002	Valdesia	Nizao	10/67	7/84	18°24'30"	70°16'50"
38001	Nizao	Nizao	10/67	4/78	18°36'53"	70°27'07"
38009	Quija Quieta	Nizao	10/76	4/79	18°13'49"	70°27'31"

EVAPORATION (Tape)

CODE	STATION	BASIN	START	END	LAT.	LONG.	TYPE
34001	Engombe	Haina	/77	/84	18°27'00"	70°00'07"	Daily

TABLE 1.5.4 RUNOFF DATA

Station Name	Basin	Type	Start	End	Lat.	Long.	Tape
La Estrechura	Nizao	Day	10/67	8/79	18°43'47"	70°29'00"	Yes
Palo De Caja	Nizao	Day	10/56	8/79	18°33'17"	70°22'52"	Yes
Paso Del Ermitano	Nizao	Day	11/67	10/75	18°26'02"	70°15'43"	Yes
Rio Abajo	Nizao	Day	5/58	10/67	18°35'08"	70°25'05"	Yes
La Penita	Nizao	Day	10/76	7/79	18°27'19"	70°16'32"	Yes
Caobal	Haina	Day	9/57	7/84	18°35'08"	70°08'57"	Yes
Los Corozos	Haina	Day	6/82	7/84	18°31'23"	70°07'10"	Yes
Arroyo Limon	Ocoa	Day	3/70	11/83	18°29'37"	70°30'43"	Yes
El Recodo	Bani	Day	2/79	9/83	18°22'27"	70°20'24"	Yes
Los Quemados	Yuna	Day	4/62	8/79	18°53'31"	70°27'25"	Yes
Blanco	Yuna/Bl.	Day	11/77	6/84	18°52'56"	70°31'17"	Yes
Maimon	Yuna/Mai.	Day	1/68	6/84	18°53'47"	70°17'71"	Yes
El Tablazo	Nigua	Day	1/59	3/84	18°28'39"	70°10'15"	Yes
Rancho Arriba	Nizao	Day	5/59	10/67	18°42'58"	70°27'59"	Yes
El Cacao	Nizao	Day	1/62	11/83	18°31'41"	70°17'59"	Yes
Los Ranchitos	Ocoa	Day	1/61	12/67	18°26'58"	70°29'55"	Yes
Carrizal	Jura	Day	10/64	10/81	18°32'27"	70°49'14"	Yes
Palomino	Y. Del. S.	Day	1/78	12/83	18°48'06"	70°58'26"	Yes
Mendez	Ocoa	Day	1/56	4/61	18°28'29"	70°30'48"	Yes
La Higuana	Nizao	Day	1/56	12/61	18°22'46"	70°16'22"	Yes

1.5.4 Streamflow Data

Daily Runoff Data: Table 1.5.4 shows the daily runoff data available for gauging stations in Nizao and other surrounding watersheds. The daily streamflow data at stations La Estrechura, Palo De Caja, Paso Del Ermitano, and Rancho Arriba was used for stochastic streamflow generation and in the calibration of the real-time streamflow forecast model. The data availability at these stations is summarized in the form of a bar chart in Figure 1.5.3.

Storm Hydrograph (Stage) Data: Table 1.5.5 presents the data availability on selected storm hydrographs for stream gauging stations La Estrechura, Palo de Caja, Paso del Ermitano and La Penita. It is noted that the original raw data corresponds to stages observed during storm events and calibrated rating curves needed to be employed to convert them to actual discharges.

Rating Curves: INDRHI provided stage-discharge relations for stations La Estrechura, Palo De Caja, La Penita, and Paso del Ermitano to be used in transforming the hourly stage data to discharges. The plots of these curves are shown in Figures 1.5.4 to 1.5.13. In view of some inconsistencies present in these curves, the raw stage-discharge data were used to develop a new set of rating curves for this study. The development of these new rating curves is explained in detail in Appendix 1.5.A.

Reservoir Levels: In order to supplement the gauged storm hydrograph data, reservoir levels at Valdesia dam for certain major events were obtained from CDE. Hourly reservoir levels were obtained for following periods:

1. August 1-13, 1980

TABLE 1.5.5 HYDROGRAPH DATA OF SELECTED STORMS

Starting				Ending				Estrec	Palo de	La Peni	Ermitan	STORM ID
YR	MON	DAY	HR	YR	MON	DAY	HR					
69	Jul	18	1	69	Jul	22	24	Incom	N/A	N/A	Incomp	A
70	Jul	7	1	70	Jul	12	24	Incom	N/A	N/A	Comp	B
70	Nov	6	1	70	Nov	12	24	N/A	N/A	N/A	Comp	C
70	Dec	7	1	70	Dec	16	24	N/A	N/A	N/A	Comp	D
72	Mar	10	1	72	Mar	13	24	Comp	Comp	N/A	Comp	E
72	May	20	1	72	May	25	24	Incom	Comp	N/A	Comp	F
74	Aug	1	1	74	Aug	13	13	Incom	Comp	N/A	Comp	G
74	Sep	11	1	74	Sep	19	24	Incom	Comp	N/A	Comp	H
74	Oct	21	1	74	Oct	26	24	Incom	Comp	N/A	Incom	I
75	Sep	16	1	75	Sep	21	11	Incom	N/A	N/A	Comp	J
75	Oct	22	1	75	Oct	27	14	Incom	N/A	N/A	Comp	K
75	Nov	1	1	75	Nov	12	12	Incom	N/A	N/A	Comp	L
77	May	20	1	77	May	25	24	Incom	Comp	Comp	N/A	M
77	Nov	21	1	77	Nov	27	24	Incom	Comp	Comp	N/A	N
77	Dec	27	1	78	Jan	6	24	N/A	Comp	Comp	N/A	O
78	May	23	1	78	May	31	24	Incom	Comp	Comp	N/A	P
78	Aug	2	1	78	Aug	8	24	Incom	Comp	Comp	N/A	Q
79	Jun	29	1	79	Jul	5	24	N/A	N/A	Comp	N/A	R

Notes: Incom = Some hourly data is missing
 Comp = Complete - we have this data at CSU
 N/A = Not available - no record available

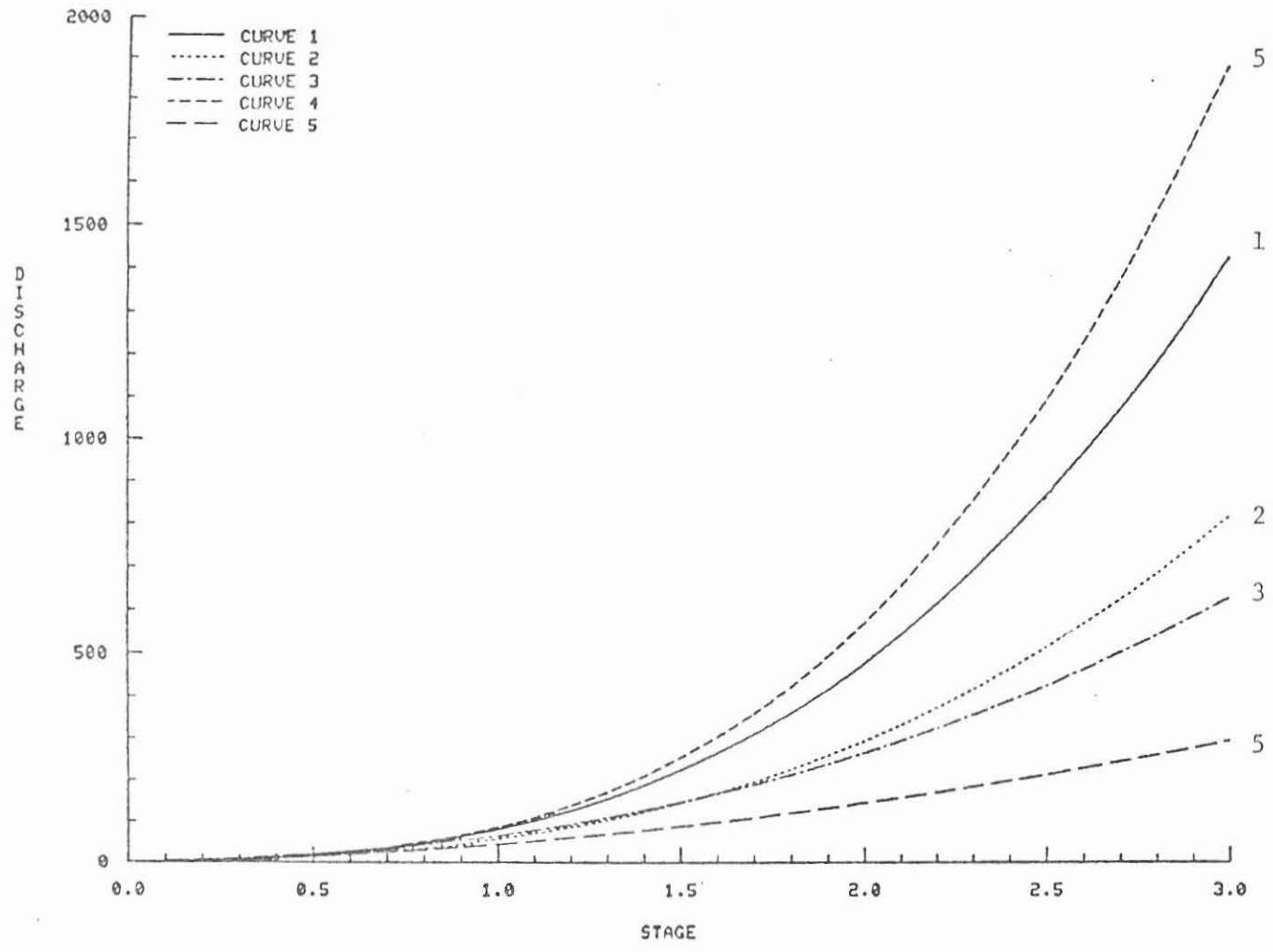


Figure 1.5.4 RATING CURVES FOR LA ESTRECHURA

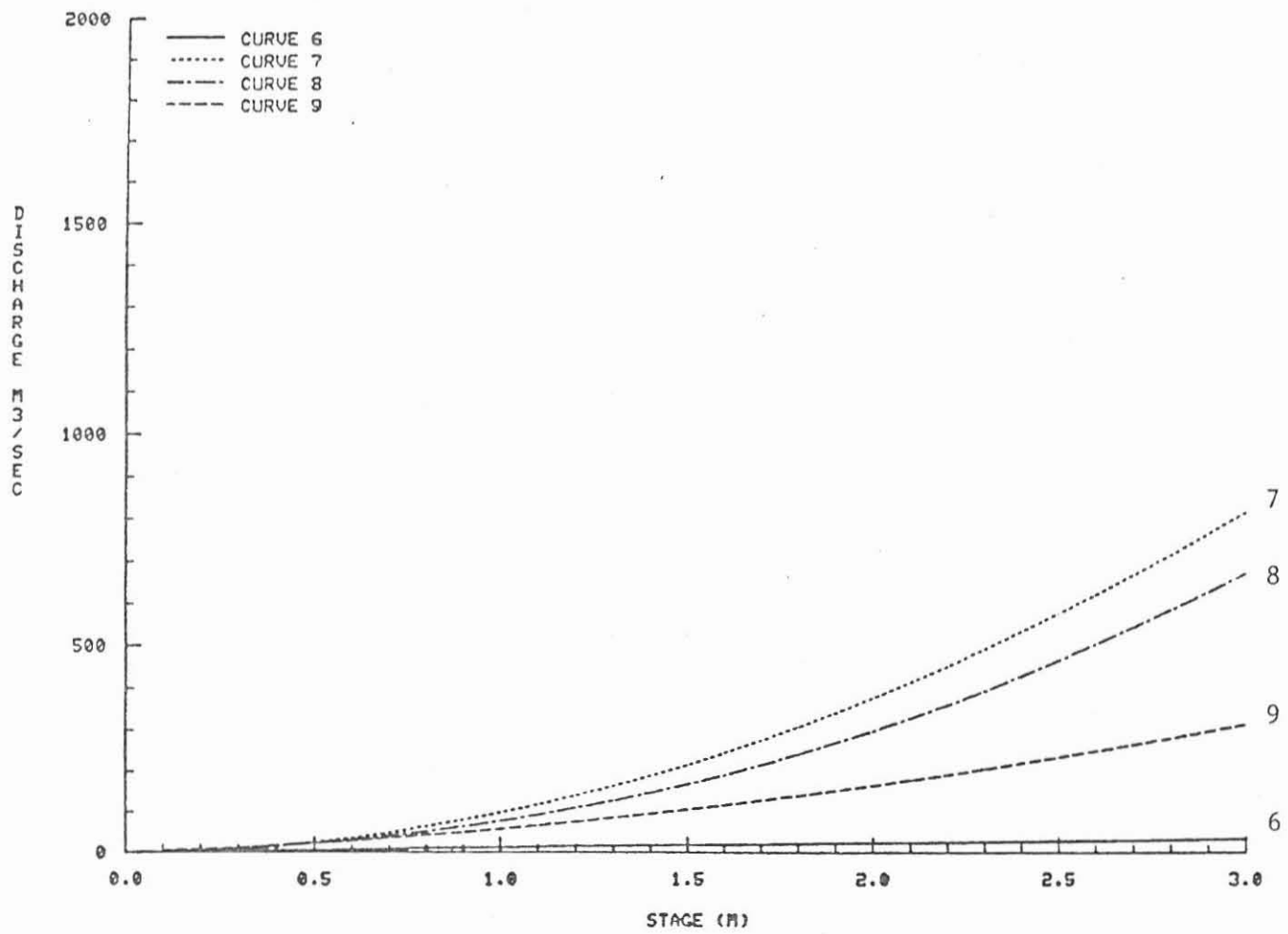


Figure 1.5.5 RATING CURVES FOR LA ESTRECHURA

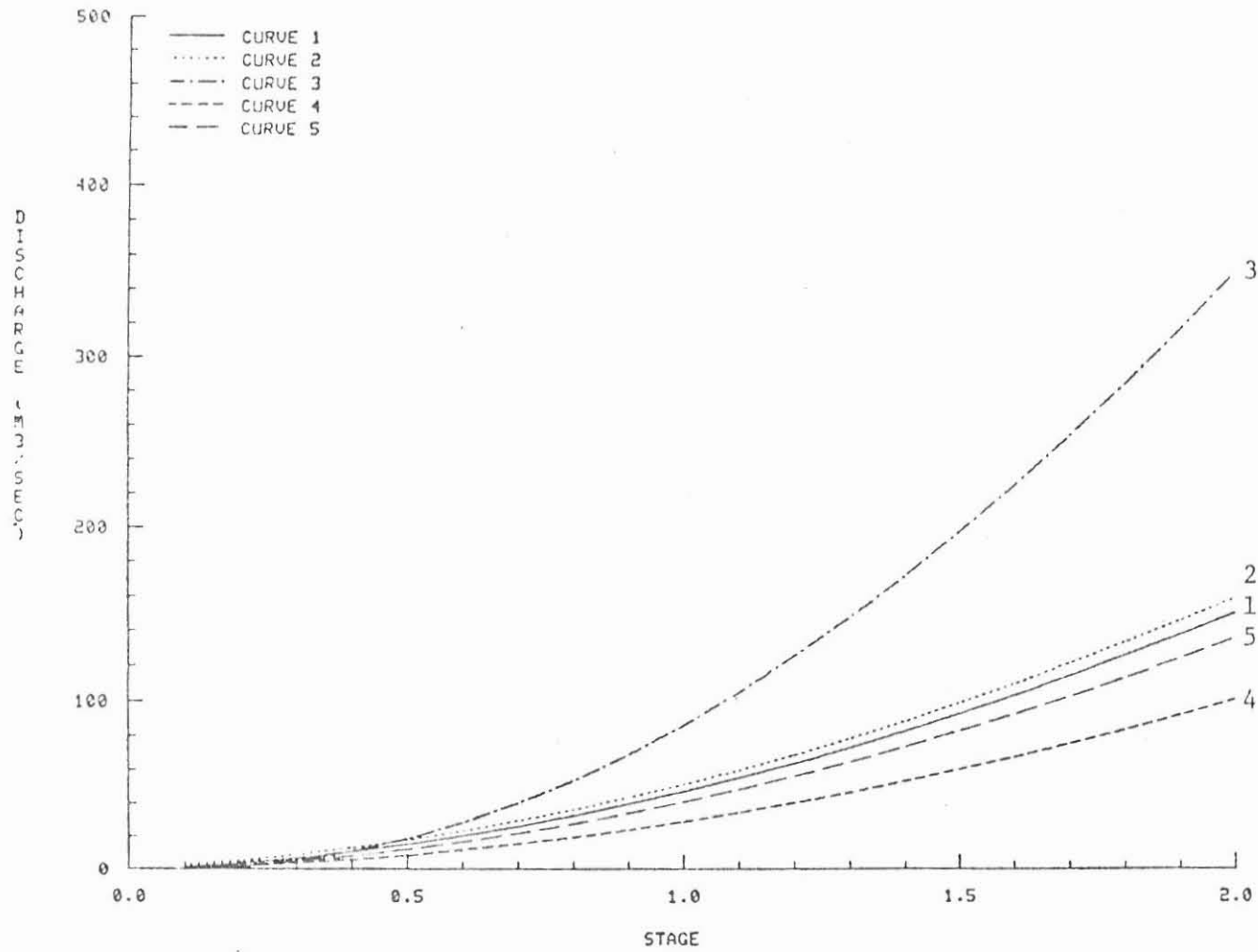


Figure 1.5.6 RATING CURVES FOR PALO DE CAJA

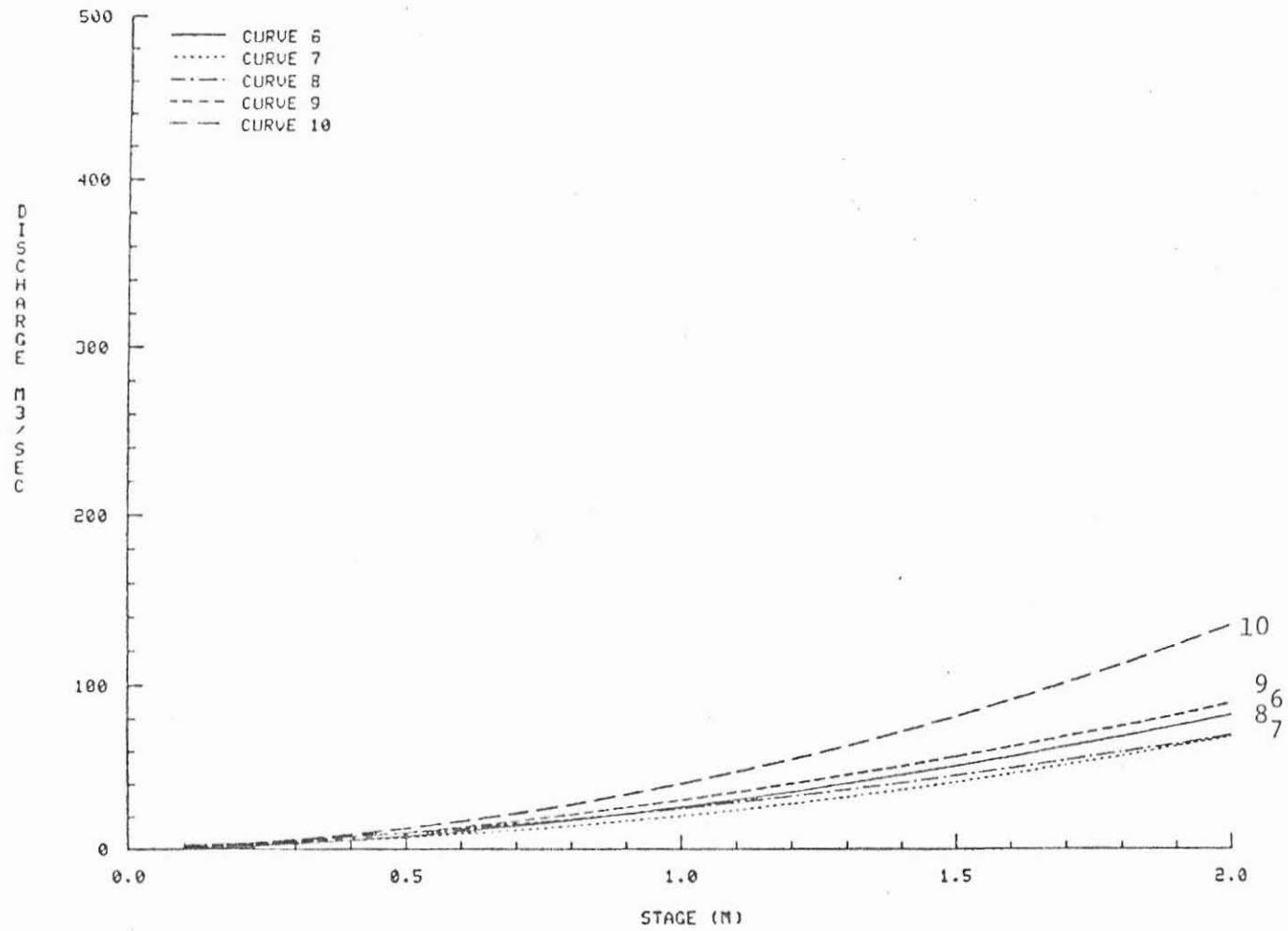


Figure 1.5.7 RATING CURVES FOR PALO DE CAJA

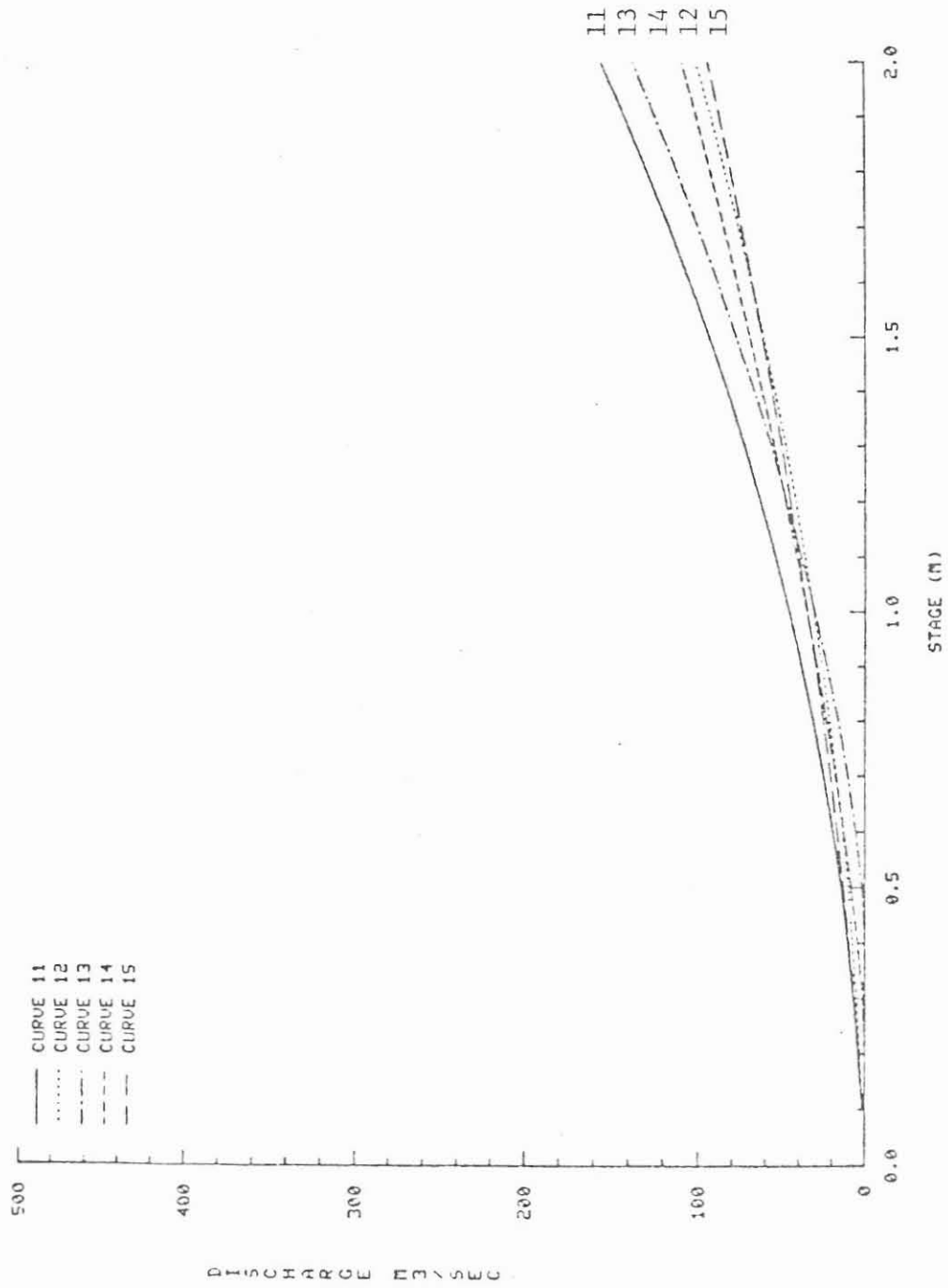


Figure 1.5.8 RATING CURVES FOR PALO DE CAJA

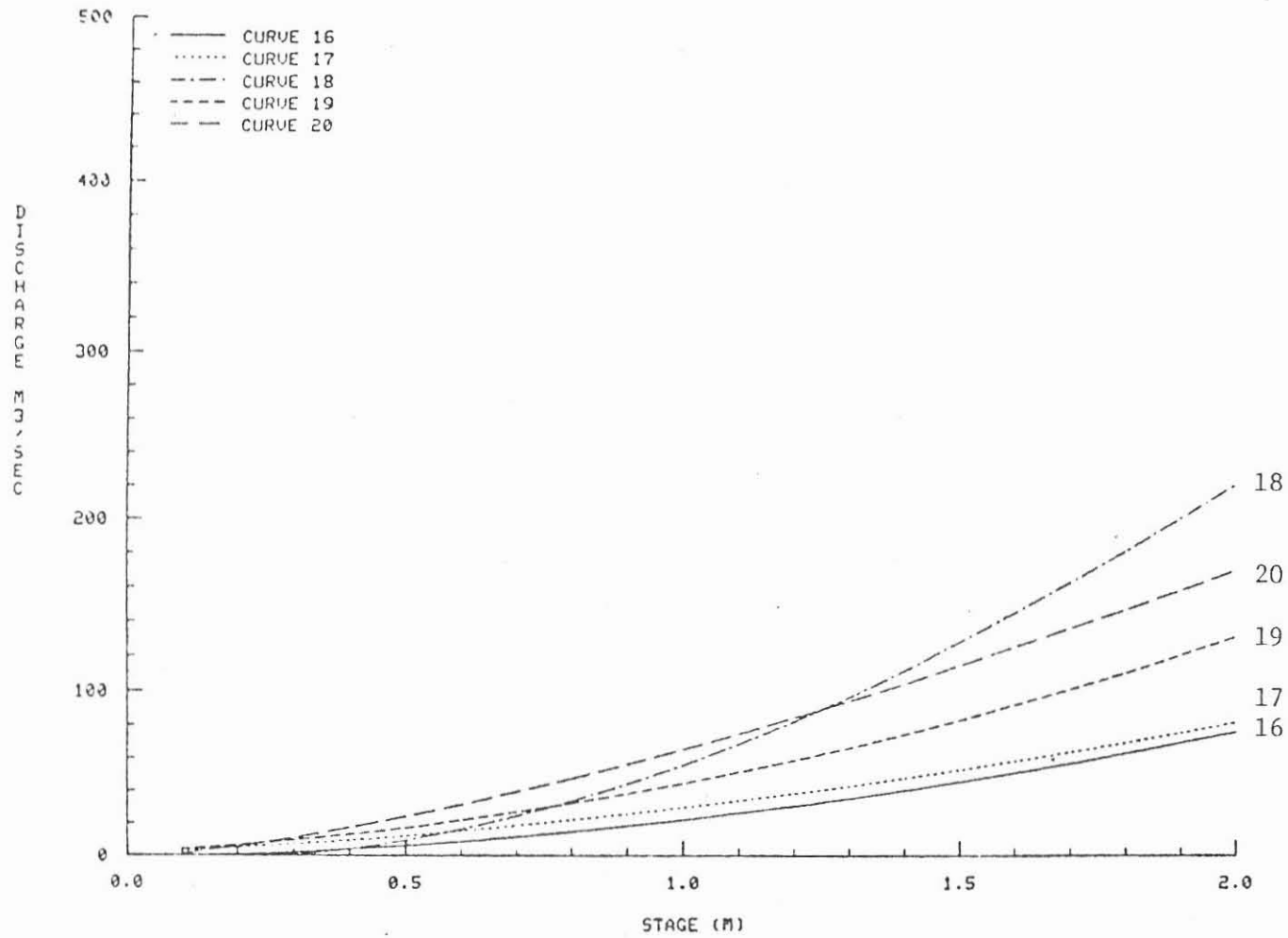


Figure 1.5.9 RATING CURVES FOR PALO DE CAJA

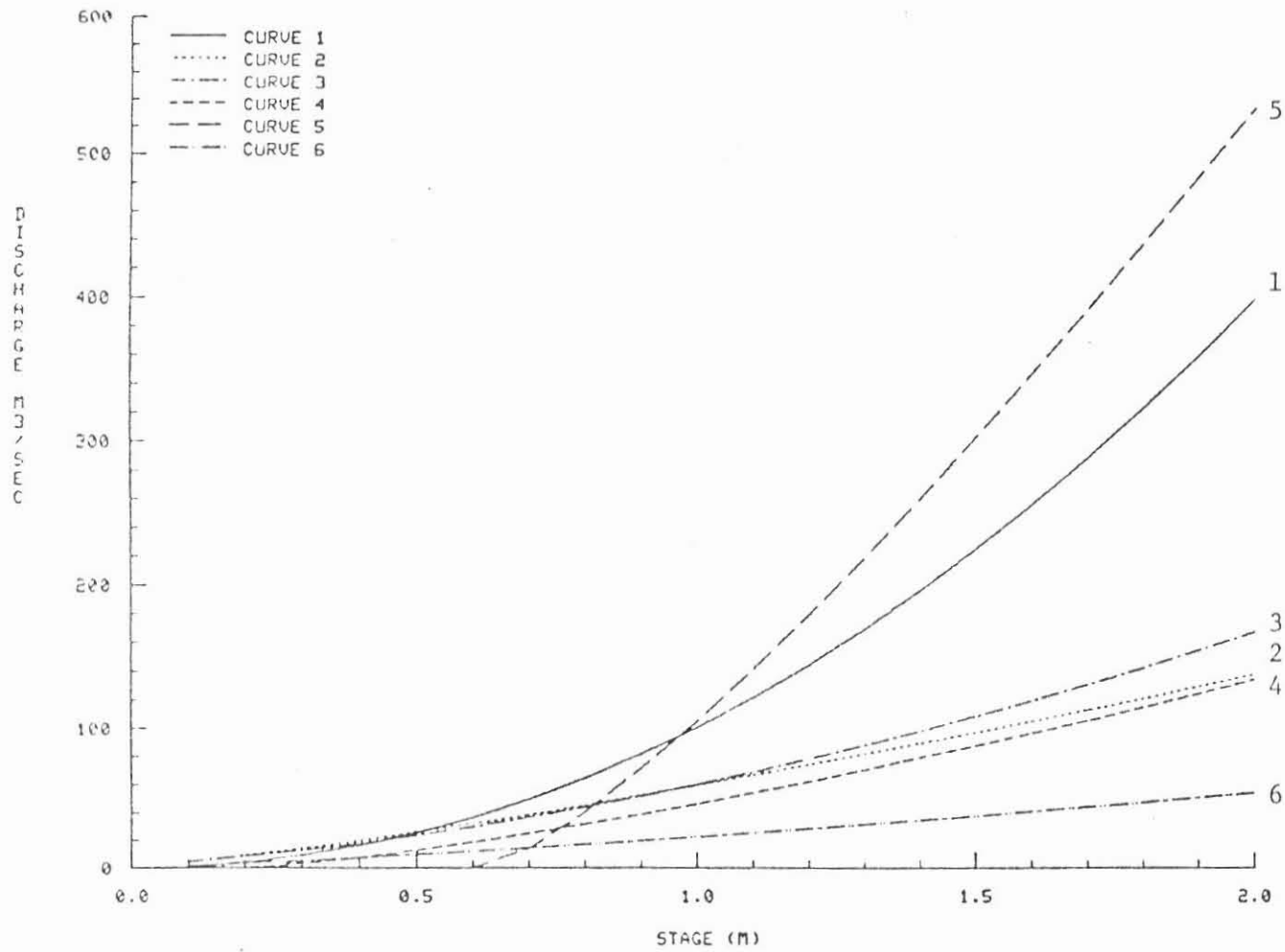


Figure 1.5.10 RATING CURVES FOR PASO DEL ERMITANO

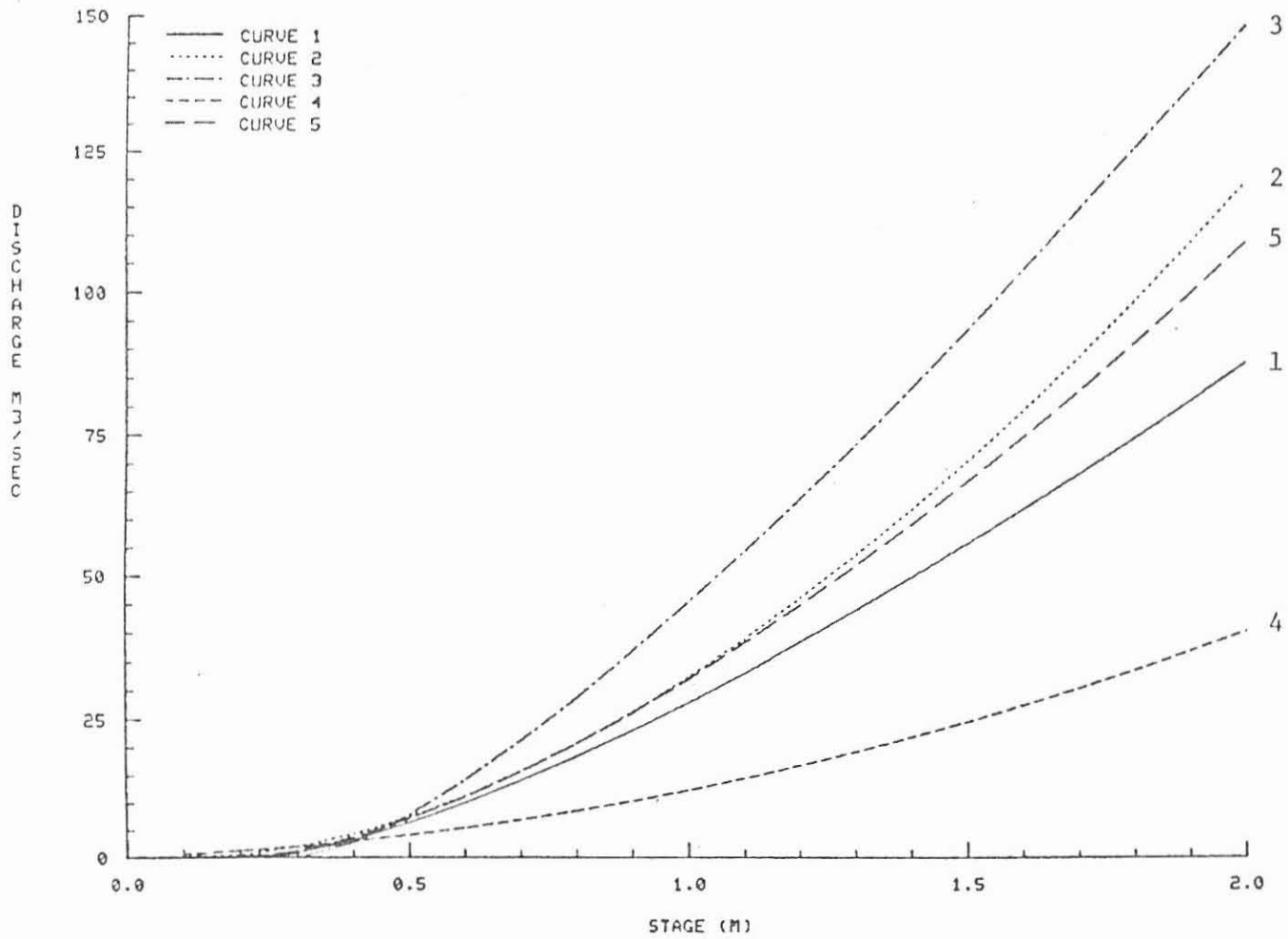


Figure 1.5.11 RATING CURVES FOR LA PENITA

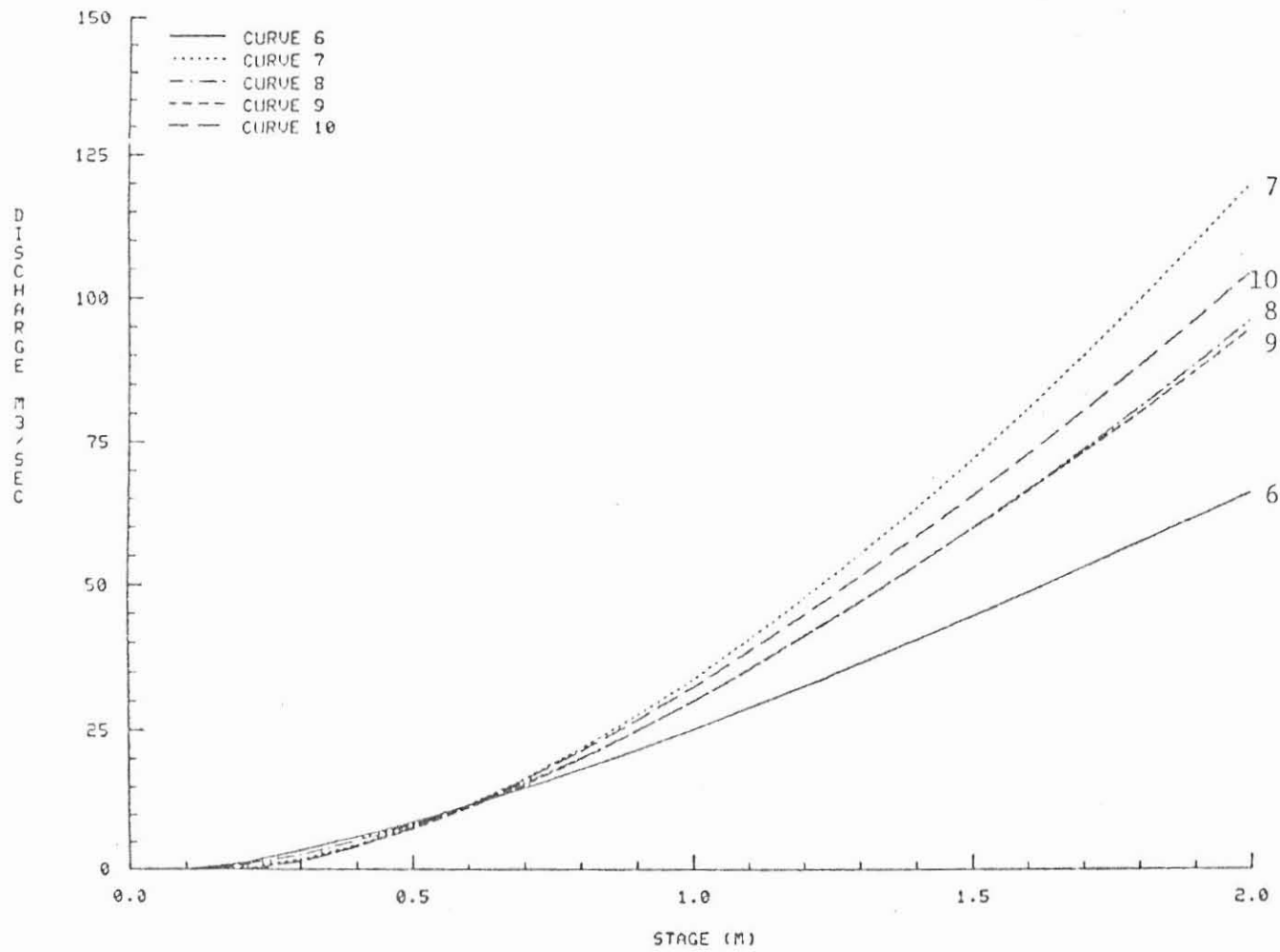


Figure 1.5.12 RATING CURVES FOR LA PENITA

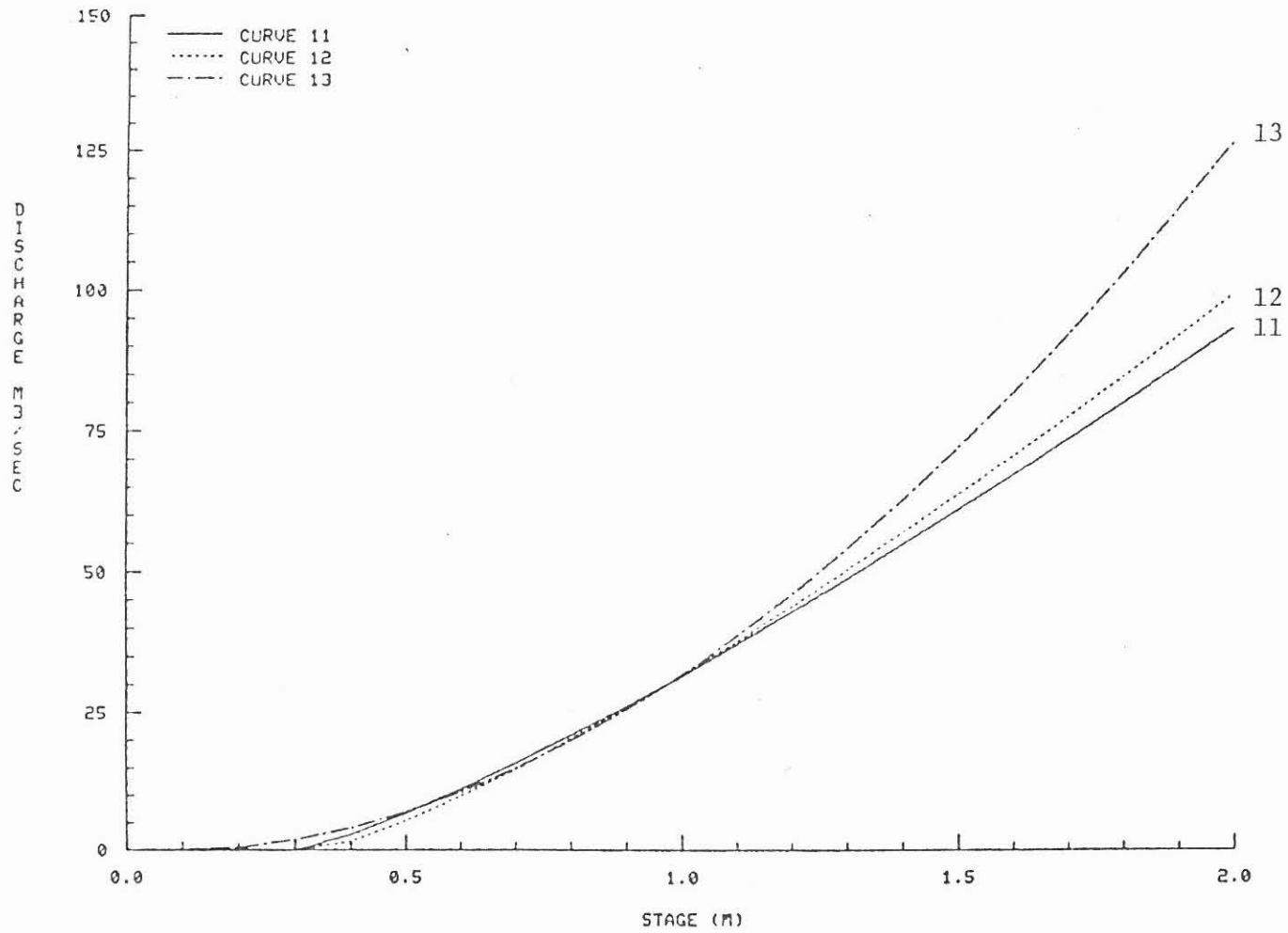


Figure 1.5.13 RATING CURVES FOR LA PENITA

2. May 6-20, 1981
3. April 10-20, 1983
4. September 13-14, 1985
5. October 23-26, 1985

For the two events in 1985, the data on operation of the spillway gates and turbines of Valdesia dam were also obtained. The two inflow hydrographs reconstructed using these data were found to be very valuable in rainfall-runoff model calibration.

1.5.5 Quality of Data

During the course of the study several obvious inconsistencies in rainfall and streamflow data were detected. The quality of data was of utmost concern since majority of modeling work to be carried out depended heavily on the accuracy of data. In general, it was felt that the entire data recording and processing procedures of INDRHI needs a careful review. Some specific problems are described below.

The first problem was encountered when the stage hourly data were transformed into discharges using the stage discharge relations provided by INDRHI. The curve numbers shown in Figures 1.5.4 to 1.5.13 refer to different equations corresponding to different time interval in which each is applicable. The equation parameters and dates are shown in Tables 1.5.6 to 1.5.11.

Very large differences are observed in some cases between consecutive curves; such is the case of curves 3 and 4 at La Penita, 5 and 6 at Paso del Ermitano, 4 and 5 at Estrechura and 3 and 4 at Palo De Caja. Curves 3 and 4 at La Penita are applicable at two consecutive periods with break point at May 22, 1977. This date coincides with storm M in the hydrograph classification and could mean that the flood

wave modified the cross section at the station. But curve 4 gives discharges more than three times smaller than those obtained with curve 3, and that is very unlikely to be true. Usually a big flood causes scour which should increase discharges for a given level. The same case is present in curves 4 and 5 at La Estrechura, with break point at August 8, 1974, which coincides with storm G.

After these problems were detected, the information on the stage-discharge data used to obtain these curves were received. They were analyzed in conjunction with cross section data to come out with the new stage-discharge relations. While developing the new stage-discharge curves, some inconsistencies were found with the data. All of them are explained in Appendix 1.5.A dealing with the development of the new stage-discharge relations, but a special case will be pointed out here. In the data received for station El Ermitano, there are two stages measured on December 11, 1970. The values in the stage-discharge data give readings of 1.75 m and 1.93 m. However, storm D has data for the same day and the maximum observed stage is only 1.55 m. The same case is observed on the data from December 14, 1970. Two stages 1.81 and 1.89 m are observed on that date, but the hydrograph data show a maximum of 1.43 m. On March 13, 1972 a stage of 0.47 m is observed while the hydrograph of storm E shows a minimum of 0.64 m for that day.

Even after the new stage-discharge hydrographs were developed, some unrealistic situations still exist. For example, consider the three hydrographs corresponding to storm F (see plot in Section 1.7). The basin area upstream El Ermitano is 800 km^2 and the basin area upstream Palo De Caja is 535 km^2 . The volume under the hydrograph at El Ermitano is 123.2 MCM while at Palo De Caja is 28.5 MCM. This means that the

subbasins downstream Palo De Caja, that is an area of 265 km², must contribute to the total flow with 94.7 MCM, or 357 mm of equivalent excess rainfall depth, but the observed total storm depths have a maximum of only 231 mm at station Valdesia.

From the 18 storms for which hourly streamflow data are available, only four storms were selected for model calibration. In many other cases the runoff appears before the rainfall stations start recording any data. For example, compare the plots corresponding to storm C. The only recording station available inside the basin is Nizao and the only observed streamflow station is El Ermitano. If we superimpose both plots we see that the precipitation recorded in Nizao has no effect in the hydrograph at Ermitano, unless there's a timing error in one of the records. If the precipitation record is lagged 15 hours or more to the back the correct response in the hydrograph is observed. In this case the solution of lagging one series could be used, but when several rainfall and runoff stations are present the problem is more complex. A similar case is observed in storm D.

The hydrographs at storm N shows a very peculiar case. The volume under the hydrograph at station Palo De Caja is larger than the one at La Penita, even though very significant precipitation was observed in the lower subbasins.

When the precipitation data used in the development of the DAD curves was compiled, several discrepancies between the hourly data and the daily data at the same station were detected. In many cases the hourly values added to values completely different from the daily ones.

The results of the HEC-1 calibration presented in Section 1.7 also point out some very severe problems in the timing of the hydrographs.

This was confirmed after checking the calibrated model with two hydrographs reconstructed from reservoir levels. In these two cases the timing seems to be correct, but in the cases of storms A, B, F, and M, we observe sharp differences in the timing of the hydrographs.

APPENDIX 1.5.ACalibration of Stage-Discharge Curves

Reliable stage-discharge curves are essential tools to convert stage readings into flow discharges, since we only normally measure variations of flow stages. Unfortunately the data used to correlate stage with flow discharge were collected at low flows. Thus a great deal of effort was used to construct and extend stage-discharge curves for gaging stations at La Estrechura, Palo de Caja, La Penita and Ermitano. The procedures of developing these stage-discharge curves are described below:

Step 1: Number of curves to be used for each station:

After plotting all stage-discharge data on logarithm papers for each station, (see Figures 1.5.A.1 through 1.5.A.4) the number of curves to describe the stage-discharge relationship for each station were decided by visual inspection.

Step 2: Separation of the low and high stage data:

For each curve chosen in step 1, we separate the data in two parts. The first part corresponds to stages lower than or equal to the stage of an expected break point which corresponds to the upper end of the low flow cross section. The second part includes stages higher than that of the above break point. For sake of illustration of the curve separation criterion a sketch is shown in Figure 1.5.A.5.

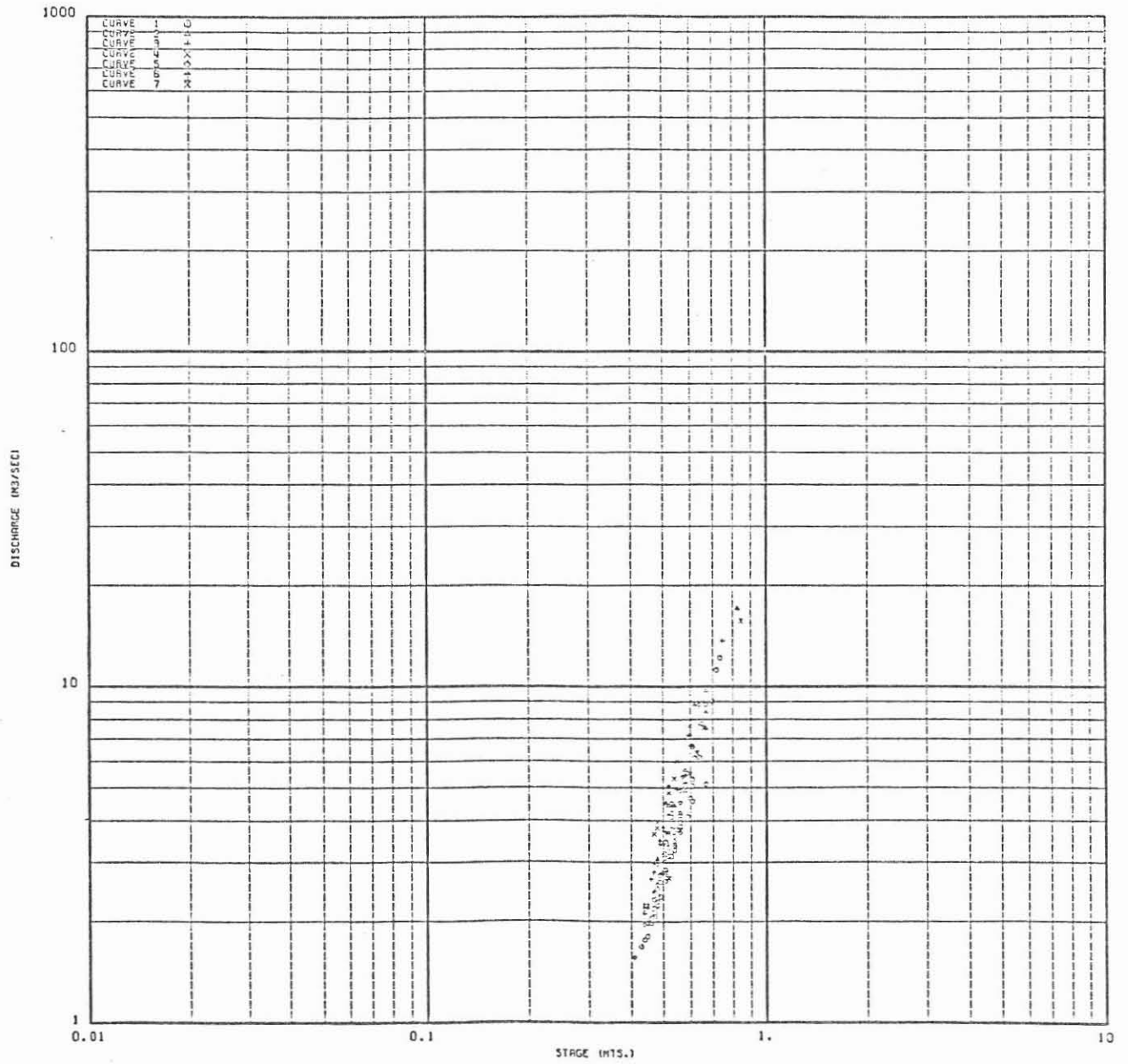


Figure 1.5.A.1. La Estrechura stage-discharge calibration points.

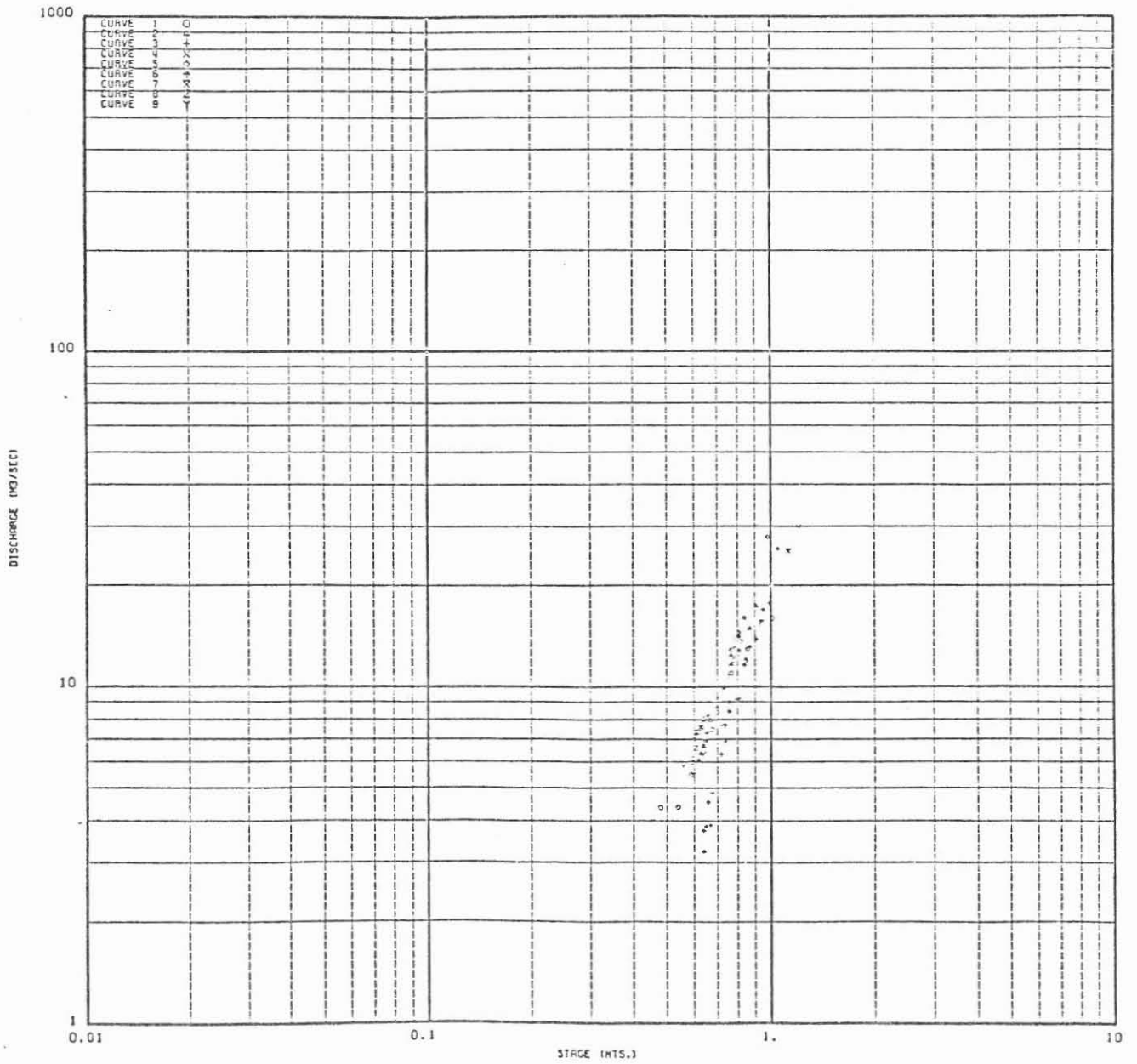


Figure 1.5.A.2. Palo de Caja stage-discharge calibration points.

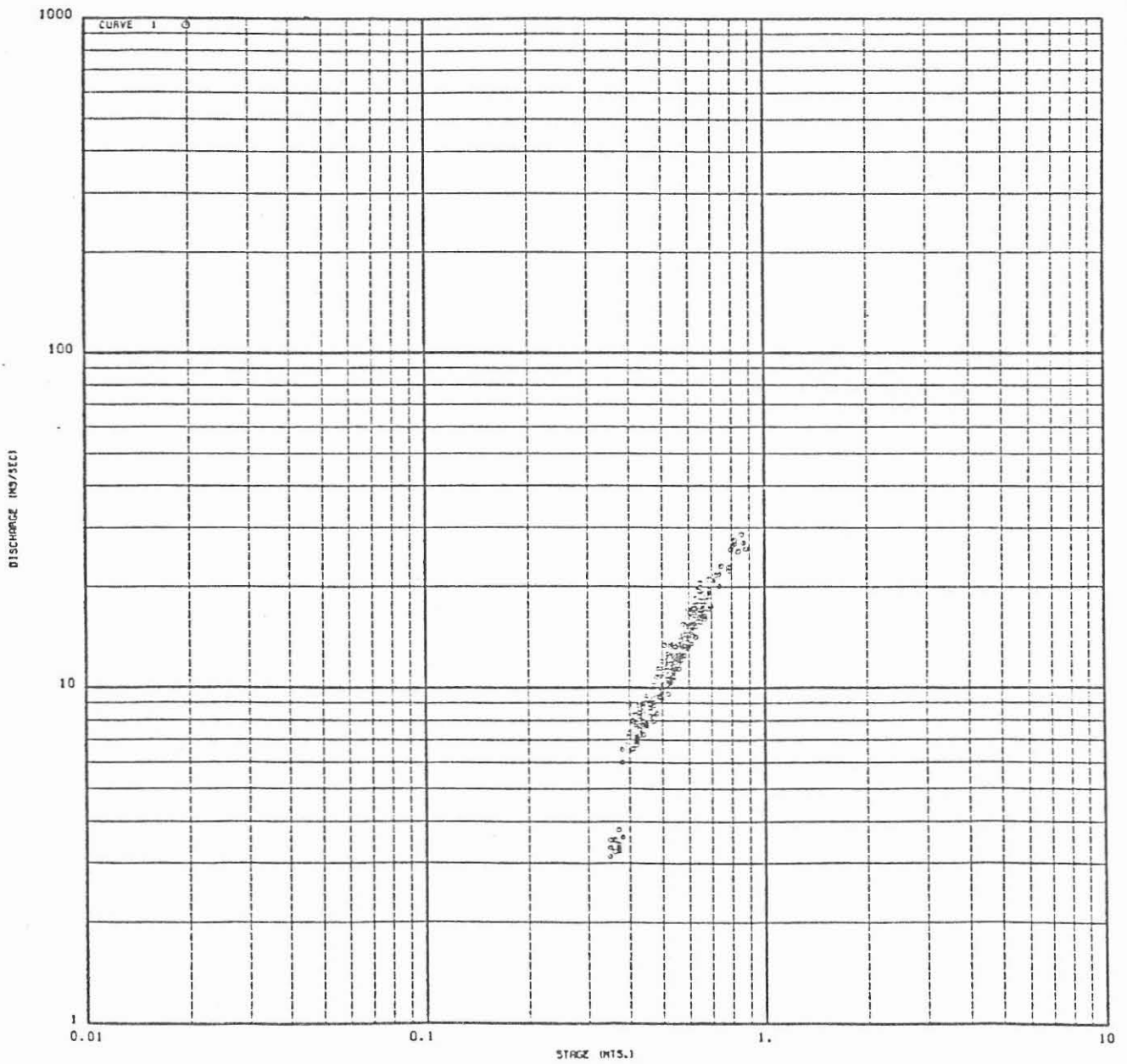


Figure 1.5.A.3. La Penita stage-discharge calibration points.

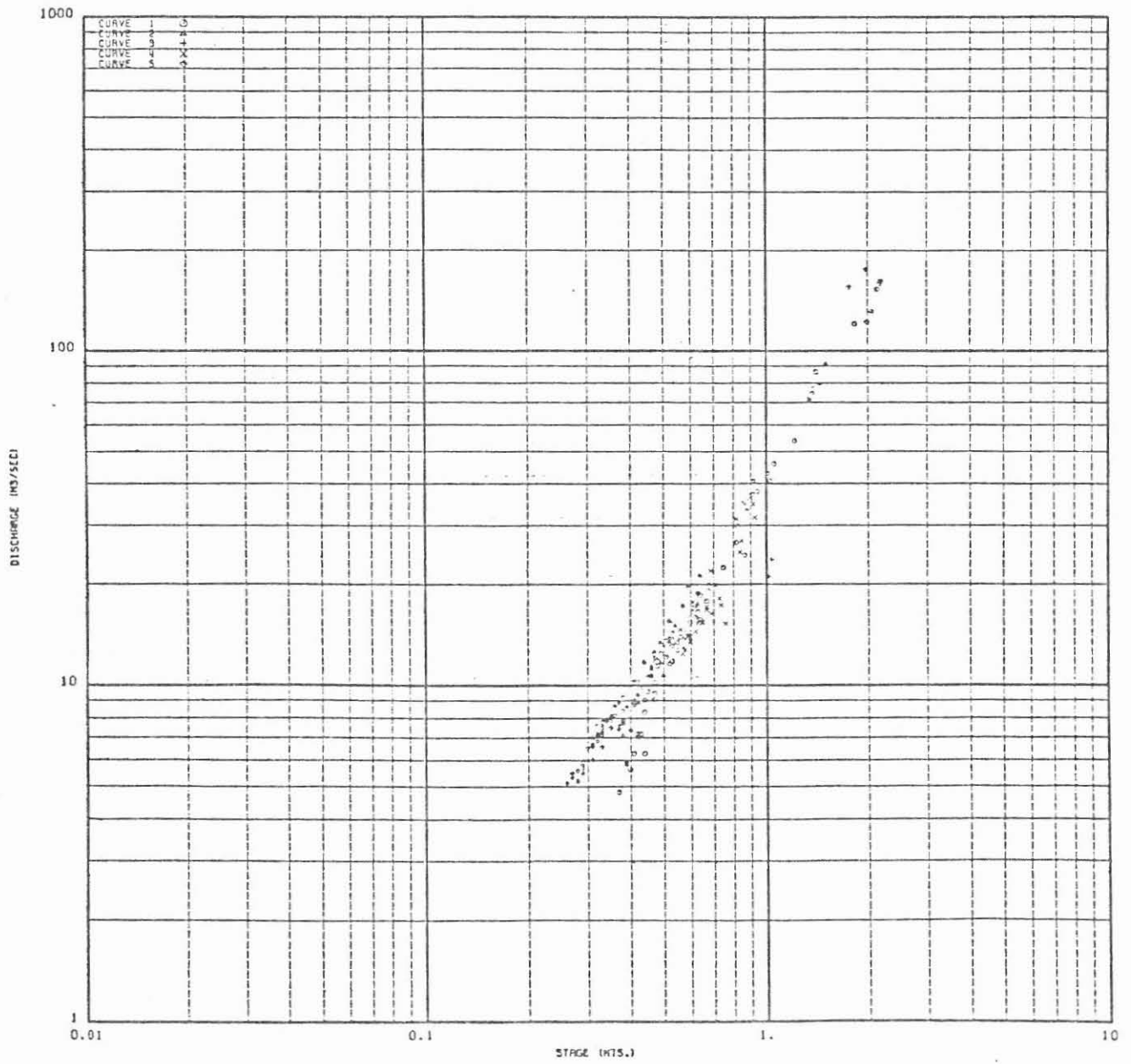


Figure 1.5.A.4. Paso del Ermitano stage-discharge calibration points.

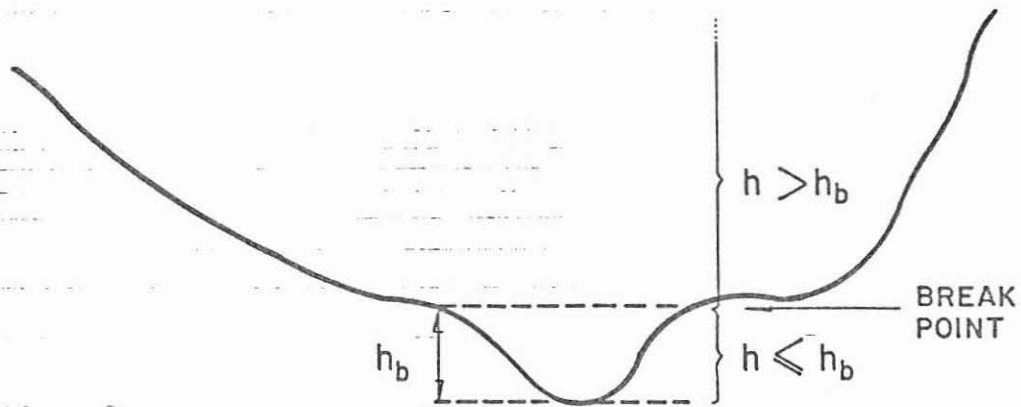


Figure 1.5.A.5 Possible gaging stations cross section shape.

Sometimes we cannot separate the data in two parts because all the stages are either lower or higher than h_b . Therefore, we fit only one curve for such the data.

Since ^{we dispose} only of cross sections of May 1985 are available, a rough estimate of the h_b was deduced from the shape of these cross sections.

Step 3: Development of the stage-discharge curves for low stages:

By using the stage and discharge data at low stages (as defined in step 2), and an optimization scheme of the polynomial function, we obtained the values of h_0 , c and m corresponding to each low-discharge curve. Same procedure was used for all stages greater than h_b .

Step 4: Decision on the number of curves to be used for stages above h_b :

First, the discharges Q versus $(h+h_0)^X$ were plotted on a log-log paper, where h_0 corresponded to the values obtained in step 3. If all points showed a tendency to scatter around a single straight line, at high stages only one curve will be used. Otherwise several straight lines will be used for each group of points.

Step 5: Development of stage-discharge curve for higher stages:

A single straight line, at high stages, in the logarithmic plot of Q versus $(h+h_0)$ usually indicates that the upper part of the cross section ($h>h_b$) can be considered as stable. In this case all the cross section rating curves will have the same parameters of the polynomial function except the value of h_0 . Each of the rating curves will have its respective h_0 .

More than one straight line, at high stages, in the logarithmic plot of Q versus $(h+h_0)$ might be caused by a nonstable upper part of the cross section or a change of the site of the cross section.

Once the high stage-discharge curves are fitted, they may be used for purpose of extension. It should be stressed that the extension or extrapolation is valid only for the range of the observed hourly stages at a given station and storm.

Results

Some inconsistency has been detected in the stage-discharge data of La Estrechura, La Penita and Ermitano. The inconsistent or questionable data were either deleted or corrected.

All the stations, except La Penita, have shown an instability in the lower part of the cross section and a stability in the upper part of the cross section. Therefore, more than one rating curve was used in the low stages and a single expression was used for the upper part of the Q versus $(h+h_0)$ logarithmic plot. The value of h_0 differs from one rating curve to another. The actual break points between the upper curve and each of the low stage discharge curves are determined by simultaneous solution of both polynomial equations corresponding to each of the lower and the upper portion of the rating curve.

La Penita logarithmic plot of Q versus $(h+h_0)$ shows a single straight line for the whole range of data, except the inconsistent points which have been deleted. This indicates that a single control is effective for the complete range of discharge. The stability of La Penita station might be due to its location just upstream the Valdesia reservoir.

A detailed description of the above results are given below for each station.

La Estrechura

Inconsistent data: The reliability of data measured on 11-03-69 with $h=0.36$ m and a corresponding discharge of $5.14 \text{ m}^3/\text{s}$ is doubtful. The error seems to be more related to an inadequate stage height reading rather than an inadequate measurement of the flow discharge.

Correction: Since there is no objective evidence regarding the origin of the error and since we have enough data to fit the

observations of curve 1 (which includes the above questionable data) we have decided to delete the above doubtful observation.

Fitting and extrapolation of the stage-discharge curves: After a close study of the data, we have decided to use seven rating curves. Three of the latter have different polynomial expressions in each of their low and high portion. Each of the four other rating curves have a single polynomial expression. We are only confident the extrapolation of the above seven curves only in the range of the observed hourly stages measured during the storm events.

Palo De Caja

Inconsistent data: None.

Fitting and extrapolation of the stage-discharge curves: Nine rating curves have been selected. Four of these curves have different polynomial expressions in for low and high flows. The remaining five curves have one unique polynomial expression for all flows. The extrapolation of the nine curves is valid only in the range of the hourly stages recorded during the storm events.

La Penita

Inconsistency: Both the plot of Q versus h on arithmetic scale and the logarithmic plot of Q versus $(h+h_0)$ show that the 14 data points observed for the time periods 03-03-77 to 03-15-77 were questionable. The latter doubtful data points would be due to inaccurate readings of the rate of revolution of the propeller-type meter. If the discharges corresponding to the above data are multiplied by 2, these 14 adjusted data points would join the same straight line as determined by all the other data. Next a logarithmic plot of Q versus $(h+h_0)$ was performed. Again the 392 data points (except the same 14 data points) defined a

single curve with about 30 data points having a stage height between 0.40 and 0.50 m (range of the stage height of the 14 doubtful data points). It was then decided to discard these 14 questionable stage-discharge observations.

Fitting and extrapolation of the stage-discharge curve: Only one curve is used for the whole range of data since a plot of Q versus $(h+h_0)$ of the above 392 data points follow a single straight line. The extrapolation of the latter curve is guaranteed only for the range of the hourly stages corresponding to the observed storms.

Ermitano

Inconsistency: After a comparison of the stages of the stage-discharge measurements with those of the hourly storms observations we found that the stage-discharge data measured on the 12-11-70, the 12-14-70 and the 03-13-72 might be subject to inaccurate stage height reading. For the two rating curve measurements of the 12-11-70 the stages were set respectively to 1.75 and 1.89 m but, the maximum hourly stage observed during the same day was only 1.55 m. The same remark can be done for the 12-14-70 measurements where the maximum hourly observed stages is 1.43 m and the two stage-discharge data were set respectively to 1.81 m and 1.89 m. The 03-13-72 stage of the rating curve was 0.47 m, whereas the minimum hourly observed stage on the same day was 0.64 m. All the above five questionable stage discharge observations were deleted. A plot of Q versus $(h+h_0)$ has shown that the stage-discharge data measured from 05-22-72 to 05-24-72 were doubtful. The error might be due to incorrect rate of revolution of the propeller-type meter. It was decided to correct the above observation on account of the need for higher stage-discharge data in the period between 05-22-72 to 09-14-72.

Correction: The questionable stage-discharge data measured from 05-22-72 to 05-24-72 were corrected by dividing their corresponding discharges by 2.

Fitting and extrapolation of the stage-discharge curves: Five rating curves have been selected. One rating curve as a single polynomial expression for both low and high stages. Each of the other four curves has two polynomial expressions, one for the low portion and the other for the higher portion. The extrapolation of each of the rating curves is reliable only for the range of the hourly stages recorded during the storm events.

The polynomial parameters of all the rating curves, the date of validity, and some observations are shown in a tabular form for each station in Tables 1.5.A.1 through 1.5.A.4.

Table 1.5.A.1 Rating Curve Equations $Q = c(h + h_o)^m$ for La Estrechura

Station	Date	h_o in m	C	m	Stage in m	Observation
La Estrechura	10/20/67 to 7/15/70	0.31	38.75	3.51	$h \geq 0.10$	
	12/9/70 to 12/16/72	0.36	46.05	3.75	$h \leq 0.13$	
		0.36	38.75	3.51	$h > 0.13$	
	1/12/73 to 11/15/73	0.30	38.75	3.51	$h \geq 0.10$	
	12/11/73 to 7/9/74	0.25	38.75	3.51	$h \geq 0.15$	
	8/10/74 to 10/8/76	0.34	56.66	4.30	$h \leq 0.27$	
		0.34	38.75	3.51	$h > 0.27$	
	11/9/76 to 7/12/78	0.35	38.75	3.51	$h \geq 0.05$	
	8/23/78 to 3/7/79	0.45	40.20	4.11	$h \leq 0.49$	
		0.45	38.75	3.51	$h > 0.49$	

Table 1.5.A.2 Rating Curves for Palo De Caja

Station	Date	h_o in m	C	m	Stage in m	Observation
Palo De Caja	6/22/71 to 9/12/73	0.26	20.26	2.20	$h \geq 0.20$	
	4/4/74 to 6/11/74	0.31	22.06	2.62	$h \leq 0.50$	
		0.31	20.26	2.20	$h > 0.50$	
	7/5/74 to 7/29/74	0.44	19.29	3.49	$h \leq 0.60$	
		0.44	20.26	2.20	$h \geq 0.60$	
	8/6/74 to 8/30/74	0.22	20.26	2.20	$h \geq 0.48$	
	9/6/74 to 10/16/74	0.04	20.26	2.20	$h \geq 0.60$	
	10/28/74 to 10/22/75	0.27	38.39	5.35	$h \leq 0.55$	
		0.27	20.26	2.20	$h \geq 0.55$	
	2/26/76 to 9/17/76	0.30	19.80	2.11	$h \leq 0.47$	
0.30		20.26	2.20	$h > 0.47$		
11/9/76 to 5/12/77	0.09	20.26	2.20	$h \geq 0.50$		
6/23/77 to 3/22/79	0.33	20.26	2.20	$h \geq 0.20$		

Table 1.5.A.3 Rating Curves for La Penita

Station	Date	h_o in m	C	m	Stage in m	Observation
La Penita	11/30/76 to 8/9/79	-0.14	40.00	1.63	$h \geq 0.40$	The 14 points observed between 3/3/76 and 3/15/77 are deleted. The rating curve will be adjusted to get the observed hydrograph.
		-0.08	36.81	1.79	" $h \geq 0.40$	The above 14 points are corrected by multiplying their discharges by

Table 1.5.A.4 Rating Curves for Ermitano

Station	Date	h_0 in m	C	m	Stage in m	Observation
Ermitano	8/27/68 to 12/14/70	0.25	42.42	1.999	$h \leq 0.80$	Curve 1.
		0.25	43.11	1.70	$h > 0.80$	The data corresponding to $h=1.75, 1.91, 1.81,$ and 1.81 are deleted
	1/14/71 to 5/16/72	-0.01	43.11	1.70	$h \geq 0.40$	The data corresponding to $h =$ is deleted.
	5/22/72 to 9/14/72	0.30	20.73	2.98	$h \leq 1.47$	The discharge corresponding to $h \geq 0.95$ are multiplied by 2.
		0.30	43.11	1.70	$h > 1.47$	
	10/3/72 to 8/12/74	-0.22	20.86	1.48	$h \leq 0.25$	
-0.22		43.11	1.70	$h > 0.25$		
1/22/75 to 10/17/75	-0.07	43.11	1.70	$h > 0.50$		

1.6 DESIGN STORMS

1.6.1 Historic Storms

Analysis of critical historic storms is a prerequisite to the development of Depth-Area-Duration (DAD) curves which are required to develop hypothetical floods such as the standard project flood (SPF). The past records of hourly rainfall and daily runoff were examined to single out a critical storm for every year. A preliminary analysis of maximum 1-day, 2-day, and 3-day rainfall data at each station and the inspection of daily runoff plots allowed a rough determination of the dates of occurrence of the critical storms. More than one critical storm were included for certain years. The initial selection included the rainfall due to hurricane David (August 30, 1979), and the tropical storm Frederic (September 5, 1979). Then a careful inspection of the hourly rainfall records at many gages enabled the selection of exact dates and times of occurrence of the storms which are to be analyzed further. Table 1.6.1 below presents the historic storms selected for further analysis. The mass curves of rainfall for these storms, are shown in Appendix 1.6.A.

1.6.2 Isohyetal Mapping

To derive the depth-area-duration curves, the first step is to analyze each storm for its isohyetal pattern. The isohyets are computed by a spatial interpolation technique known as multiquadratic interpolation. In multiquadratic interpolation, the influence of each sampling point is represented by quadric cones as a function of the coordinates of these points. The estimate for a given point with coordinates (x_0, y_0) is thus obtained by the sum of the contributions from all those quadric cones. This is mathematically expressed as

TABLE 1.6.1 HISTORIC STORMS SELECTED FOR ANALYSIS

<u>Beginning Date and Time</u>				<u>Ending Date and Time</u>				
Year	Month	Date	Hour	Year	Month	Date	Hour	
1	1963	Oct.	1	20	1963	Oct.	5	5
2	1964	Aug.	6	8	1964	Aug.	7	3
3	1965	May	2	8	1965	May	5	8
4	1966	May	25	8	1966	May	27	8
5	1966	Sept.	28	20	1966	Sept.	30	8
6	1967	Sept.	10	8	1967	Sept.	13	8
7	1968	Aug.	8	8	1968	Aug.	10	8
8	1969	July	19	8	1969	July	20	8
9	1970	Aug.	22	8	1970	Aug.	23	8
10	1971	Feb.	19	8	1971	Feb.	21	8
11	1972	May	20	8	1972	May	23	8
12	1973	Oct.	14	8	1973	Oct.	21	8
13	1974	Aug.	30	8	1974	Aug.	31	8
14	1975	Sept.	16	8	1975	Sept.	18	8
15	1976	Oct.	10	8	1976	Oct.	12	8
16	1977	May	21	8	1977	May	24	8
17	1977	Dec.	28	8	1978	Jan.	1	8
18	1978	Aug.	3	8	1978	Aug.	6	8
19	1979	Aug.	30	8	1979	Sept.	2	8*
20	1979	Sept.	5	8	1979	Sept.	8	3**
21	1980	Aug.	4	8	1980	Aug.	7	8
22	1981	May	8	8	1981	May	11	8
23	1982	May	9	8	1982	May	13	8
24	1983	April	12	8	1983	April	13	8
25	1984	Aug.	1	8	1984	Aug.	3	8

*Hurricane David

**Tropical Storm Frederick

$$h_o = \sum_{i=1}^n c_i d_{oi} \quad (1.6.1)$$

where h_o is an estimate of rainfall process at any point (x_o, y_o) , c_i is the multiquadric coefficient of sampling point with coordinates (x_i, y_i) , d_{oi} is the distance between point (x_o, y_o) , and (x_i, y_i) , and n is the number of sampling points. The distance d_{oi} is computed from the formula:

$$d_{oi} = \sqrt{(x_o - x_i)^2 + (y_o - y_i)^2} \quad (1.6.2)$$

The estimate h_o at any point (x_o, y_o) can be represented by a weighted linear combination of the observed values h_j at each sampling point (x_j, y_j) as

$$h_o = \sum_{j=1}^n w_j h_j \quad (1.6.3)$$

where w_j is the weight at sampling point j . To estimate the coefficients c_i and express Eq. (1.6.1) in terms of the weights as in Eq. (1.6.3), we do the following.

Let h_j of each sampling point (x_j, y_j) assume Eq. (2.1) as

$$h_j = \sum_{i=1}^n c_i d_{ji} \quad \text{for } j = 1, 2, \dots, n$$

Then the coefficients c_i are determined by

$$c_i = \sum_{j=1}^n \delta_{ij} h_j \quad \text{for } i = 1, 2, \dots, n \quad (1.6.4)$$

where δ_{ij} is an element of the inverse of the $n \times n$ interstation distance matrix with element d_{ji} , $j = 1, \dots, n$ and $i = 1, \dots, n$. Substitution of Eq. (1.6.4) in Eq. (1.6.1) yields

$$h_o = \sum_{i=1}^n d_{oi} \sum_{j=1}^n \delta_{ij} h_j$$

or upon rearranging the numeration terms,

$$h_o = \sum_{j=1}^n \left[\sum_{i=1}^n \delta_{ij} d_{oi} \right] h_j$$

Thus, the interpolation equation (1.6.3) has weights

$$w_j = \sum_{i=1}^n \delta_{ij} d_{oi} \quad \text{for } j = 1, \dots, n \quad (1.6.5)$$

For this study, Eqs. (1.6.3) and (1.6.5) are used to compute the isohyetal pattern of each storm at any point (x_o, y_o) in the study area

using the data h_j at each sampling point (x_j, y_j) available from $j=1, \dots, n$ stations.

From the previous explanation it is noted that rainfall interpolates at any point in the basic area of interest are solely function of the distances between such point and the observation points (rainfall stations) available in the area. By virtue of this method, rainfall pattern anomalies due to orographic effects or bias in rainfall information due to topography are not accounted for in the interpolation. The Hydrology Group at INDRHI strongly suggested that perhaps such rainfall pattern anomalies should be considered in the derivation of the isohyetal patterns. In this connection, two approaches are tried which are briefly described below.

First is the adoption of the precipitation weighing method given by the U.S. Corps of Engineers in the HEC-1 Flood Hydrograph Package which was likewise suggested by the INDRHI Hydrology Group. This method is based on the weighting equation given by

$$h_o = h_{B_o} \frac{\sum_{j=1}^n h_j w_j}{\sum_{j=1}^n \bar{h}_j w_j} \quad (1.6.6)$$

where h_o is the rainfall interpolate at any point in the area, h_{B_o} is the interpolated (using optimal interpolation) normal annual precipitation at any point in the area, h_j is the total storm precipitation at sampling station j and \bar{h}_j is the j th station normal annual precipitation. The weight w_j of station j can likewise be obtained using the multiquadric interpolation technique such that

$$w_j = \sum_{i=1}^n \delta_{ij} d_{oi} \quad \text{for } j=1, \dots, n \quad (1.6.7)$$

where d_{oi} is the distance between the point with coordinates (x_o, y_o) and j th station point with coordinates (x_j, y_j) , and δ_{ij} is an element of the

inverse of an $n \times n$ interstation distance matrix with elements d_{ij} , $j=1, \dots, n$ rows and $i=1, \dots, n$ columns.

As indicated by the U.S. Corps of Engineers, the above approach could correct rainfall estimation bias associated to elevation effects which is accounted for by the station normal annual precipitation term \bar{h}_j . However, this claim may be rather dubious since the elevation is not explicitly parameterized in the weighting scheme and that any adjustments for bias affected by incorporating either or both terms h_B and \bar{h}_j can be associated to rainfall anomalies other than elevation effects.

In view of this, the second approach tried accounts for orographic effects which explicitly parameterized the basin elevation. This approach is based on representing the rainfall by a polynomial function written as

$$h_o = a_o + \sum_{k=1}^m a_k E_o^k + H_o \quad (1.6.8)$$

where h_o is the rainfall estimate at any point (x_o, y_o) , the a 's are polynomial coefficients, E_o is the elevation at point (x_o, y_o) and H_o is the elevation-free rainfall values. Similarly, the observed rainfall values at the available station points can be represented as in the above equation as

$$h_j = a_o + \sum_{k=1}^n a_k E_j^k + H_j \quad (1.6.9)$$

where w_j is the j th station weight obtained by the multiquadric interpolation technique. Finally, the rainfall interpolate in the actual domain can be obtained using Equation (1.6.8) given the elevation E_o .

Note that an elevation map is required in the above approach for interpolating over an area. The multiquadric interpolation technique is used also to derive the elevation map for the basin.

The two approaches above were tried in this study followed by developing a new set of depth-area-duration (DAD) curves. A comparison was made using the two approaches as well as the previously obtained DAD curves based on rainfall isohyetal patterns without considering orographic effects.

From results obtained, it is found that using the second approach in which the elevation of the basing is explicitly parameterized gave the most reasonable and constant rainfall isohyetal pattern. The elevation map required in this approach on multiquadric interpolation is based on more than 200 elevation data points. A first-order polynomial is found sufficient to represent the rainfall-elevation anomaly function such that $m=1$ in Equation (1.6.8). The listing of programs used is given in Appendix 1.6.B.

1.6.3 Development of Depth-Area-Duration (DAD) Curves

Given the rainfall isohyetal patterns described in the previous section and the Mass Curves shown in Appendix 1.6.A, the procedure to develop the Depth-Area-Duration curves can be summarized as follows.

First, define class intervals based on the observed range of precipitation depths. Based on the 25 storms selected (see Table 1.6.1) the range is taken as 0 to 625 mm. Twenty five classes of class widths 25 mm each (i.e., class 1 is defined as 600-625, class 2 as 575-600, etc.) were selected. From the derived isohyetal pattern map of a given storm, the total area and average depth for each class are computed using the following equations:

$$A_i = \Delta A \sum_{j=1}^{NG} I_j \quad (1.6.10)$$

and

$$D_i = \frac{\sum_{j=1}^{NG} d_j I_j}{\sum_{j=1}^{NG} I_j} \quad (1.6.11)$$

where:

A_i = area corresponding to class interval i

ΔA = unit area of one grid point

NG = total number of grids

D_i = average depth of class interval i

d_j = depth at grid point j

$I_j = \begin{cases} 1 & \text{if } L_i \leq d_j < U_i \\ 0 & \text{otherwise} \end{cases}$

L_i = lower class limit of class i

U_i = upper class limit of class i

After obtaining the area and average depth corresponding to each class interval cumulative areas and corresponding average depths were computed using the following equations:

$$AC_i = \sum_{j=1}^i A_j \quad (1.6.12)$$

$$DC_i = \frac{\sum_{j=1}^i A_j D_j}{AC_i} \quad (1.6.12)$$

where

AC_i = cumulative area of classes greater than or equal to class i

DC_i = weighted average depth corresponding to classes greater than

or equal to class i

Note that the cumulative areas and average depths computed above correspond to what is referred to as "extended class" where each extended class always has an upper limit of 625 mm and a lower limit equal to the lower class limit of class i defined earlier.

The above defines the Depth-Area curve for the total duration of the given storm. Now, the Depth-Area curves for shorter durations are derived by using mass curves of rainfall. To do this, one has to compute the weights that will be used in obtaining an average mass curve for each class interval.

The weight given to each recording station vary according to the distribution of the area assigned to a certain class interval. Since we

know which grid points belong to each class interval, we can compute the distance between a point and the recording stations and determine the closest one. Then we can count what fraction of the total area was assigned to each station and compute the weights accordingly, using the equation below.

$$W_i(k) = \frac{\sum_{j=1}^{NG} S_j(k)}{\sum_{j=1}^{NG} I_j} \quad (1.6.14)$$

where

$W_i(k)$ = Weight assigned to station k in class interval i

$$S_j(k) = \begin{cases} 1 & \text{if } L_i \leq d_j < U_i \text{ and } k \text{ is the nearest station to grid point } j \\ 0 & \text{otherwise} \end{cases}$$

The cumulative weights corresponding to each extended class is computed by using the equation.

$$WC_i = \frac{\sum_{j=1}^i W_j(k)A_j}{AC_i} \quad (1.6.15)$$

Using the cumulative weights an average mass curve for each cumulative class is computed. From this average mass curve one obtains the maximum precipitation recorded at different durations and compute the fraction of the total storm depth for each duration. These fractions are multiplied by the depth at the corresponding class interval to obtain the depths for different storm durations. The procedure is repeated for all the extended classes. The individual DAD curves for the 25 selected storms as well as the enveloping curves obtained by picking the maximum observed depth for each area and duration are included in Appendix 1.6.C.

As explained in the previous section, two methods were tried to account for topographic effects in the rainfall interpolation program. The results from both trials were used to derive two new sets of DAD curves. The performance of the two methods were judged based on the

TABLE 1.6.2 CLASSIFICATION OF STORMS USED IN DAD
CURVES COMPUTATION

STARTING DATE	CLASSIFICATION	
8/1/84	non-hurricane	
4/12/83	non-hurricane	
5/9/82	non-hurricane	
5/8/81	non-hurricane	
8/4/80	hurricane	Hurricane Allen
9/5/79	hurricane	Tropical Storm Frederick
8/30/79	hurricane	Hurricane David
8/3/78	non-hurricane	
12/28/77	non-hurricane	
5/21/77	non-hurricane	
10/10/76	non-hurricane	
9/16/75	hurricane	Hurricane Eloise
8/30/74	hurricane	Hurricane Carmen
10/14/73	hurricane	Tropical Storm Gilda
5/20/72	non-hurricane	
2/19/71	non-hurricane	
8/22/70	hurricane	Tropical Storm Dorothy
7/19/69	non-hurricane	
8/8/68	non-hurricane	
9/10/67	hurricane	Hurricane Beulah
9/28/66	hurricane	Hurricane Inez
5/25/66	non-hurricane	
5/2/65	non-hurricane	
8/6/64	non-hurricane	
10/1/63	hurricane	Hurricane Flora

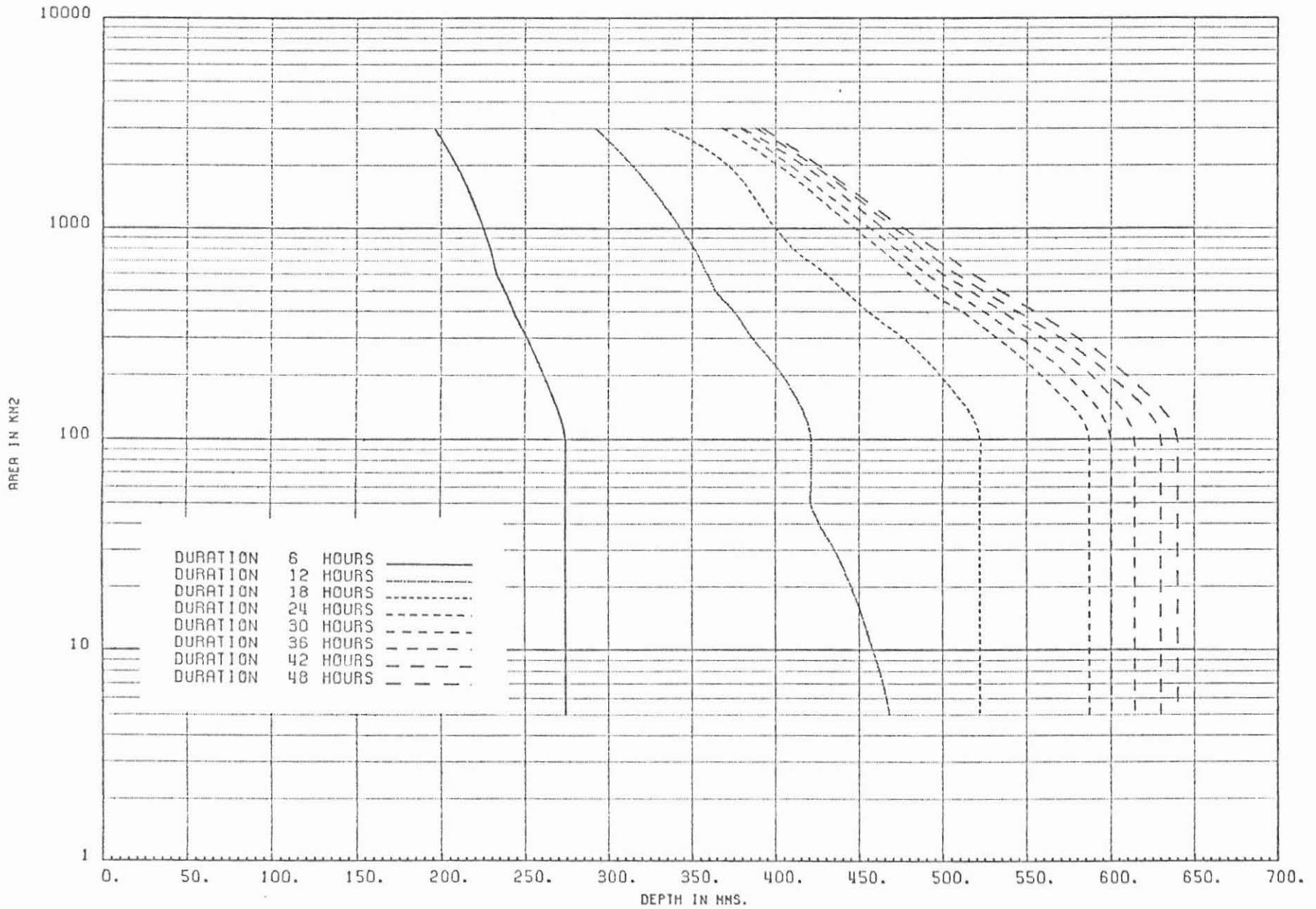


Figure 1.6.1. Hurricane depth-area-duration curves for Nizao basin.

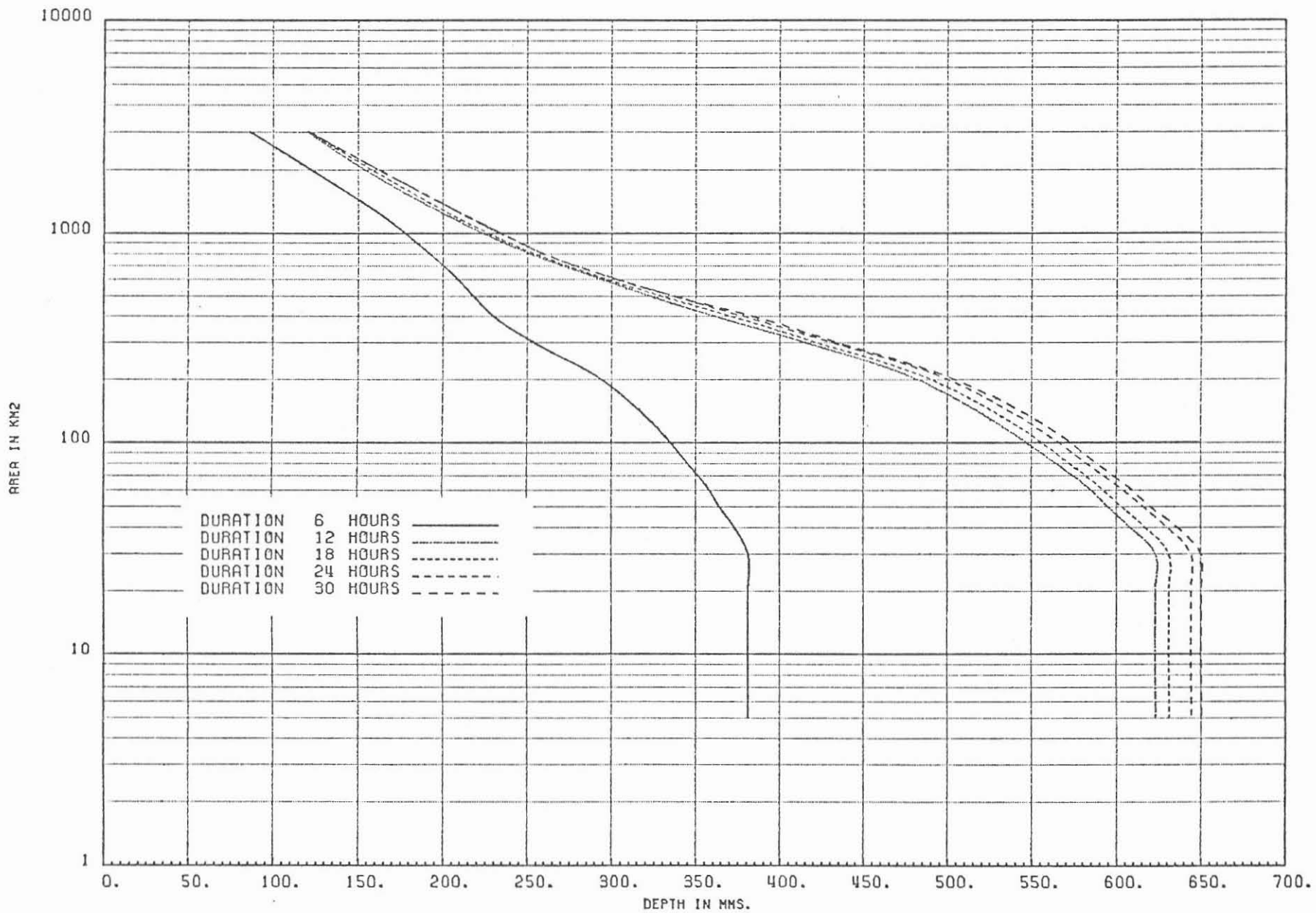


Figure 1.6.2. Non-hurricane depth-area duration curves for Nizao basin.

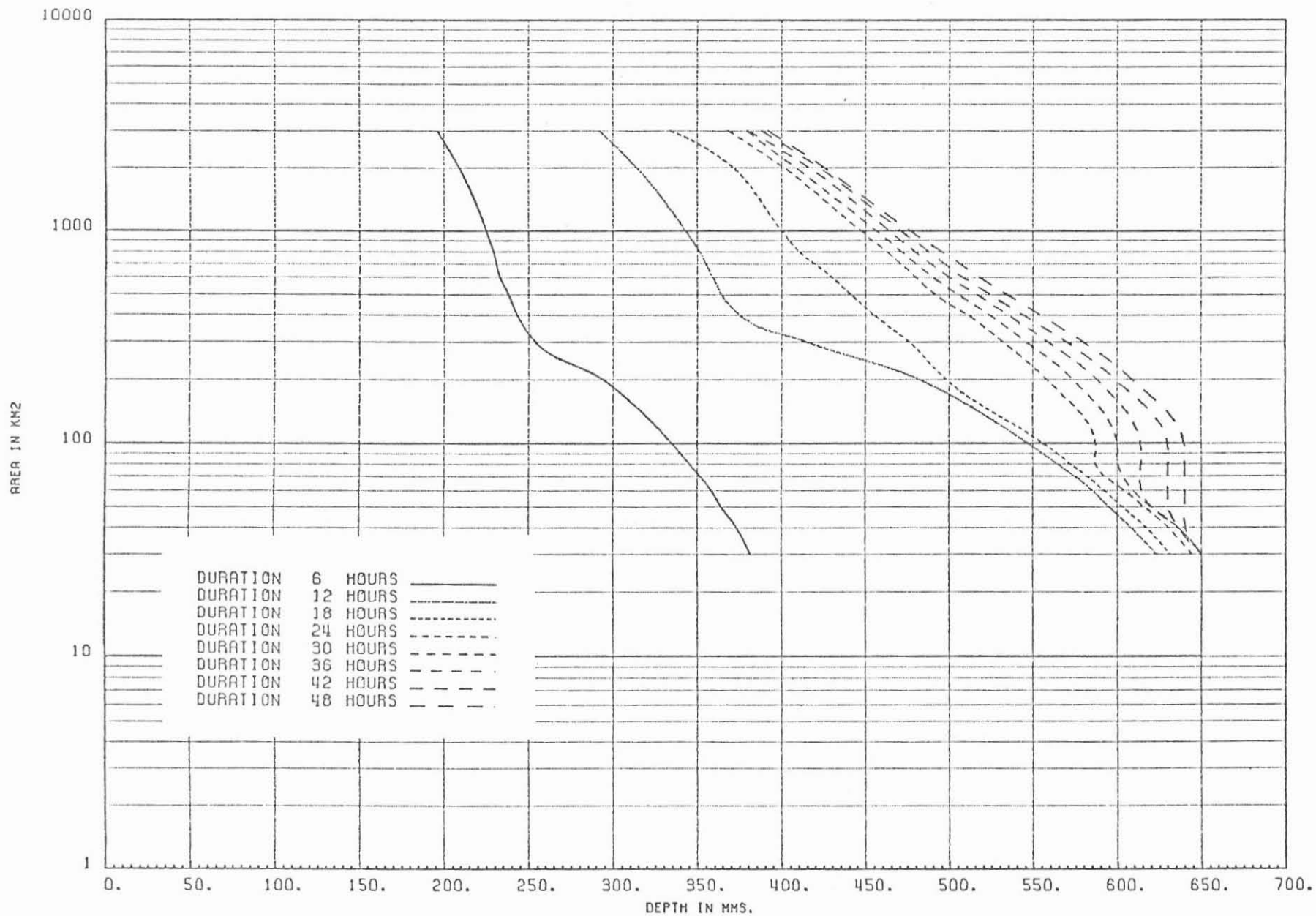


Figure 1.6.3. Enveloping depth-area-duration curves for Nizao basin.

comparison of the isohyetal patterns with those obtained without considering the topographic effects as well as by observing the DAD curves obtained with each method. Based on these it was decided that parameterizing the elevation in the interpolation function gave the most realistic results. The DAD curves obtained with the two methods are also included in Appendix 1.6.C.

Before obtaining the enveloping DAD curve from the individual curves computed with the selected method, the 25 storms used were divided into two groups: (a) hurricane and (b) non-hurricane. Table 1.6.2 shows this classification. Then the "hurricane" DAD curves and the "non-hurricane" DAD curves were derived from the corresponding groups. Based on these, a third set of curves that represent the worst conditions observed in the basin was derived. The "hurricane", "non-hurricane", and "enveloping" DAD curves are shown in Figures 1.6.1 to 1.6.3.

1.6.4 Standard Project Storm (SPS)

From the three sets of DAD curves shown in Figures 1.6.1 through 1.6.3, the standard project storms corresponding to duration of 24 hours and 48 hours for an area of 820 sq. km. (Nizao basin upstream of Valdesia dam) were derived. The total precipitation magnitudes of these standard projects storms are given below:

TABLE 1.6.3 Standard Project Storm (Hurricane, Non-Hurricane) and Enveloping) Precipitation Depth

DURATION (hrs)	PRECIPITATION DEPTH (mm)		
	HURRICANE	NON-HURRICANE	ENVELOPING
24	460	255	460
48	493	260	493

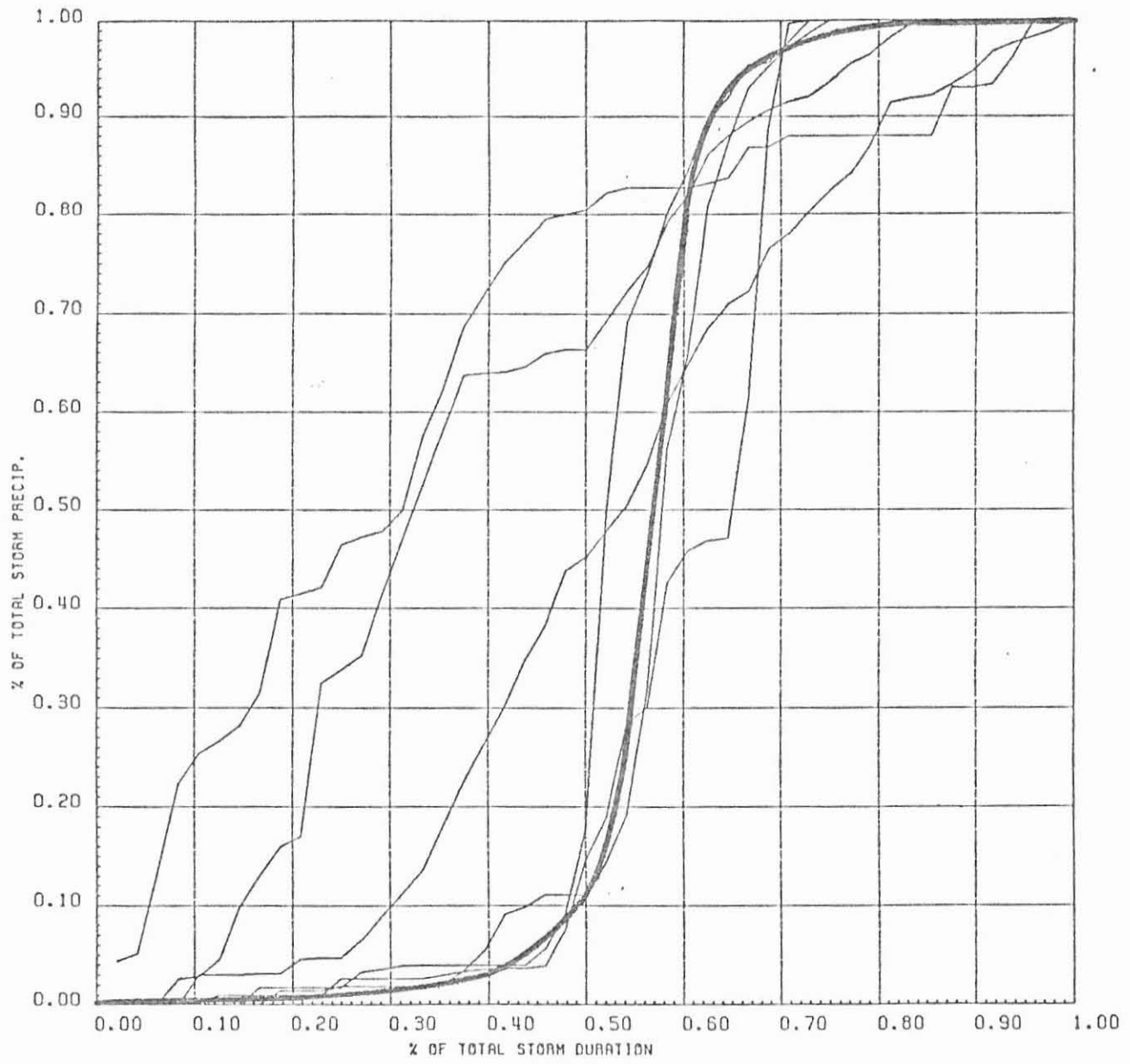


Figure 1.6.4. Historic rainfall patterns and selected design temporal distributions.

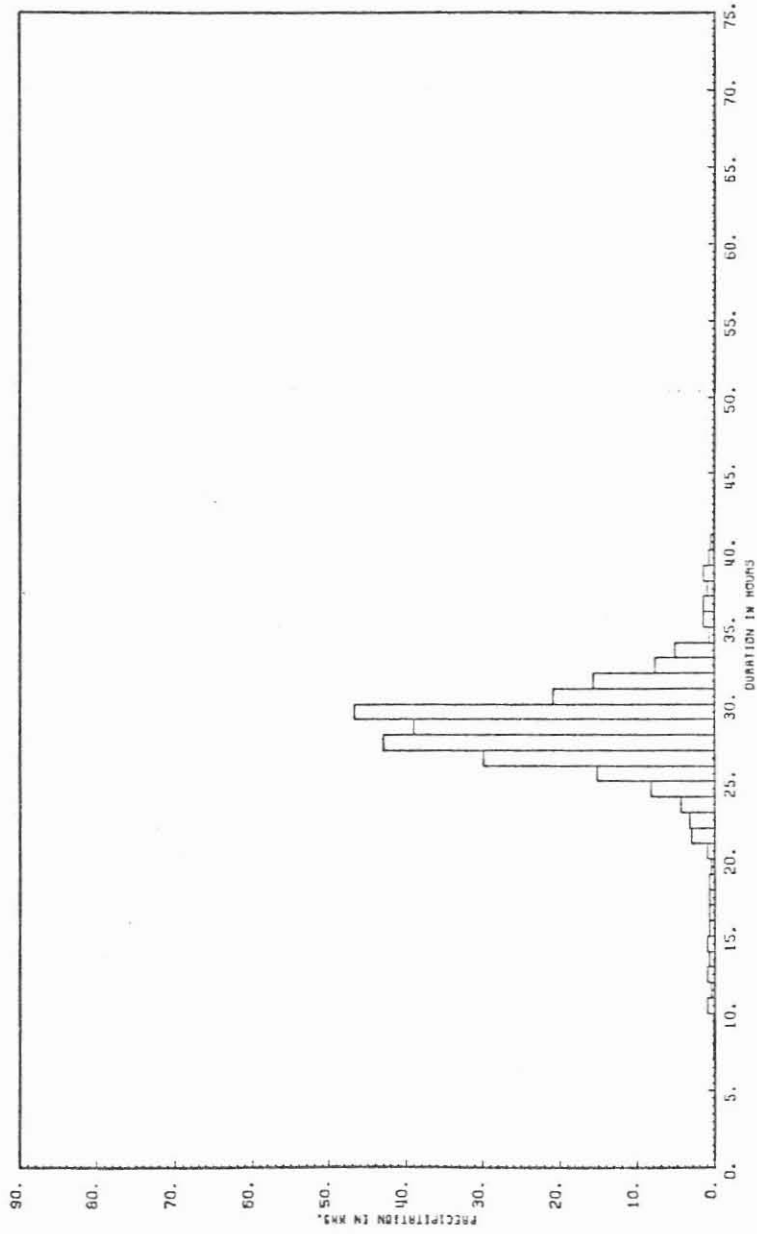


Figure 1.6.5. Non-hurricane standard project storm (48 hours).

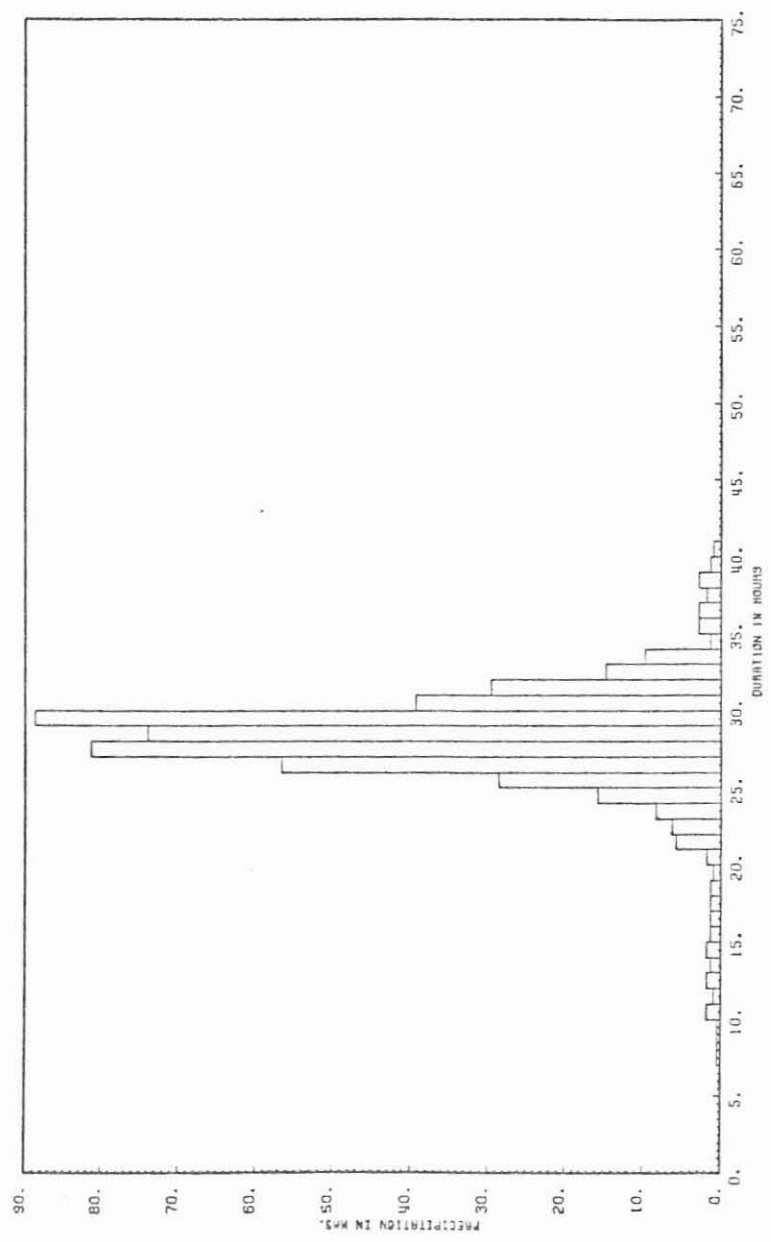


Figure 1.6.6. Hurricane standard project storm (48 hours).

It is seen that precipitation magnitude corresponding to hurricane conditions is almost twice as big as the corresponding depth for non-hurricane conditions from the same storm duration.

1.6.5 Temporal Distribution of SPS

During the course of the study two different criteria have been used for temporal distribution of the total precipitation magnitudes reported in Table 1.6.3. First approach is to select two critical patterns from the 25 historic storms selected for detailed analysis in the derivation of DAD curves. Specifically, the temporal distributions corresponding to hurricane David and tropical storm Frederick were selected. After recognizing the subjective nature of this first approach, a second approach which proved to give a more critical temporal distribution was used as follows. The percentage magnitude versus percentage duration plots were made for all 48 hour storms selected earlier. Then an enveloping curve, which lies below all the curves at small percentage duration, and above all curves for larger durations, was plotted. This exercise is illustrated in Figure 1.6.4.

Combining the enveloping temporal distribution shown in Figure 1.6.4 and the standard project storm magnitudes in Table 1.6.3, the 48 hour standard project storm isohyetal patterns were generated. These design storms are presented in Figure 1.6.5 and 1.6.6.

1.6.6 Probable Maximum Precipitation (PMP)

Given the location of the Nizao basin, a hurricane is most likely to produce the PMP. This is supported by the almost twice the precipitation depths obtained for hurricane conditions than for non-hurricane conditions for a given area and a duration (see DAD analysis). The Hurricane Model of U.S. Weather Bureau (1961) has been used by INDRHI/CDE to produce a hurricane PMP pattern for the Nizao basin. The original model has been modified for conditions existing in the

Dominican Republic and has been applied to compute PMP for the Tavera-Bao Watershed (CDE, personal communication). The precipitation pattern of average PMP over Nizao watershed obtained from the Hurricane model is presented in Figure 1.6.7.

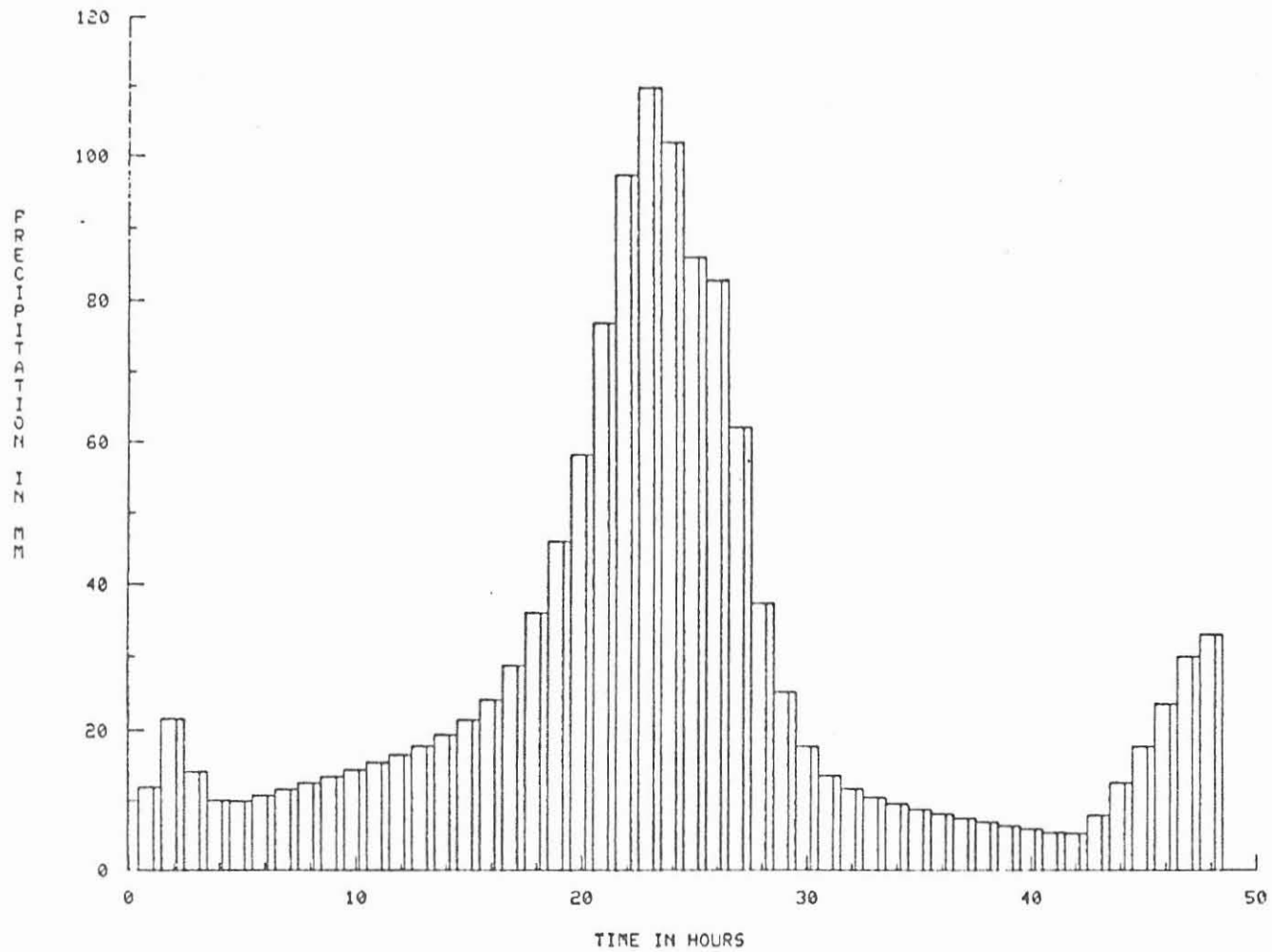
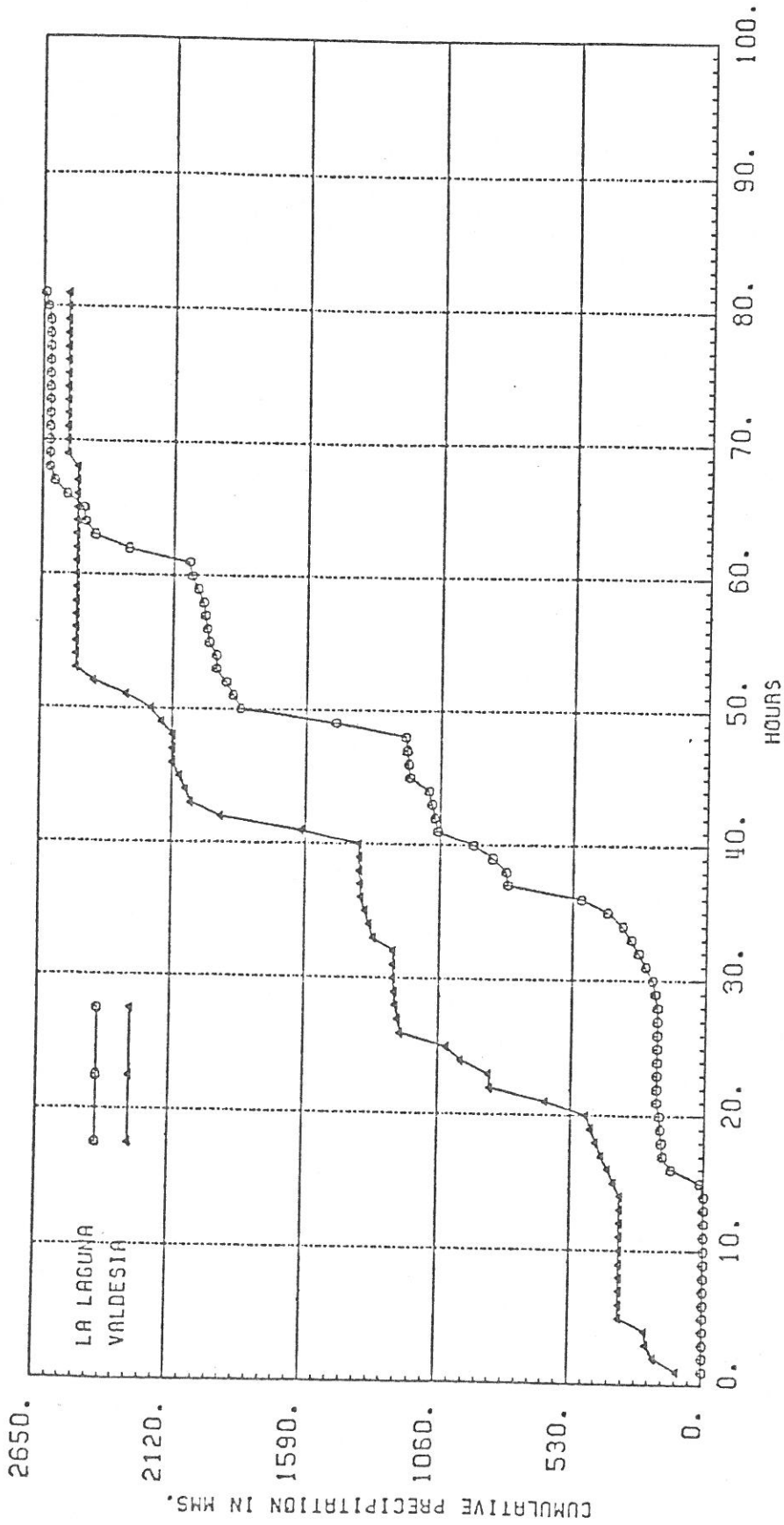


Figure 1.6.7. Precipitation pattern of average Probable Maximum Precipitation over Nizao watershed.

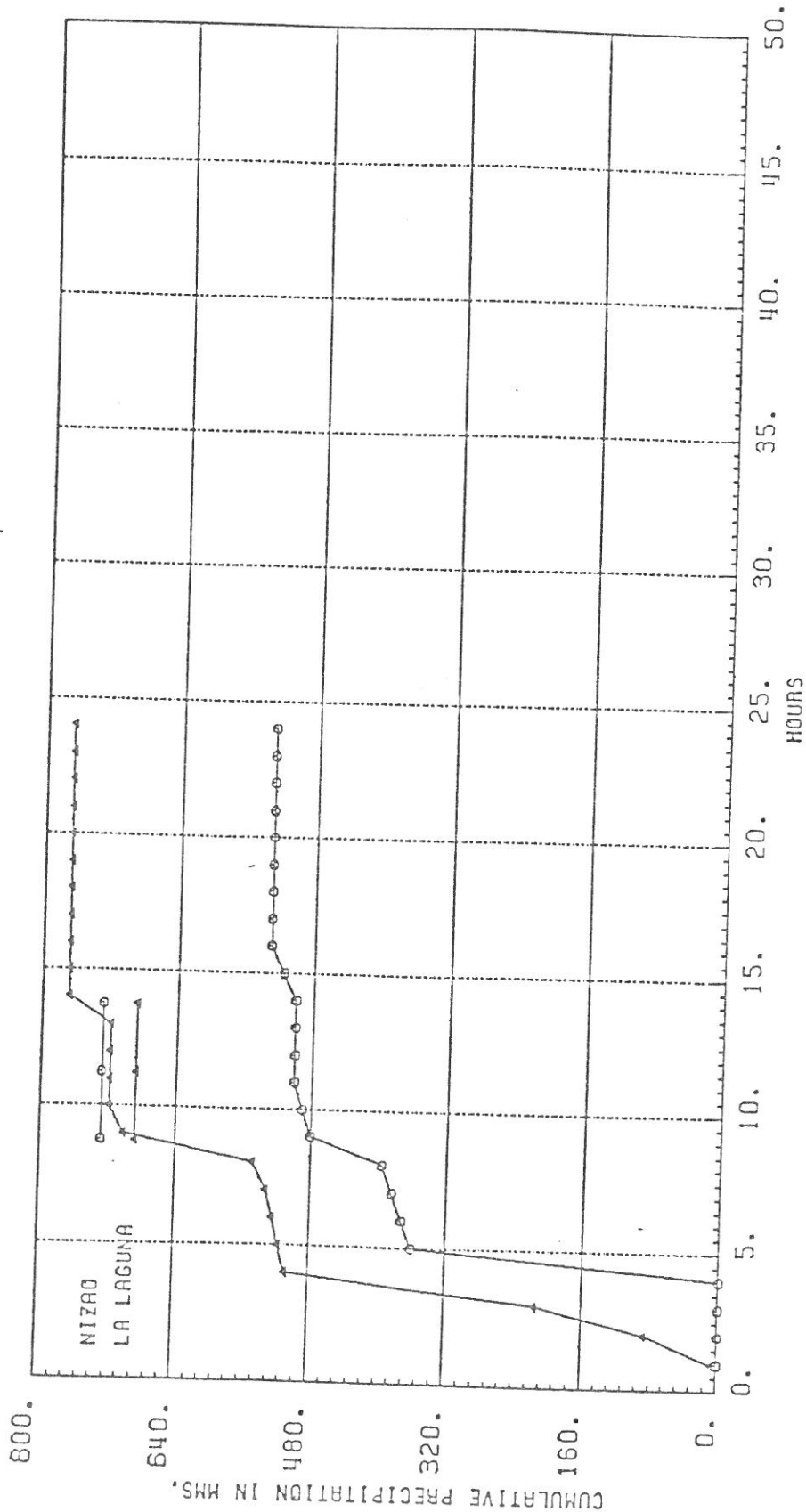
APPENDIX 1.6.A

MASS CURVES OF RAINFALL FOR SELECTED STORMS



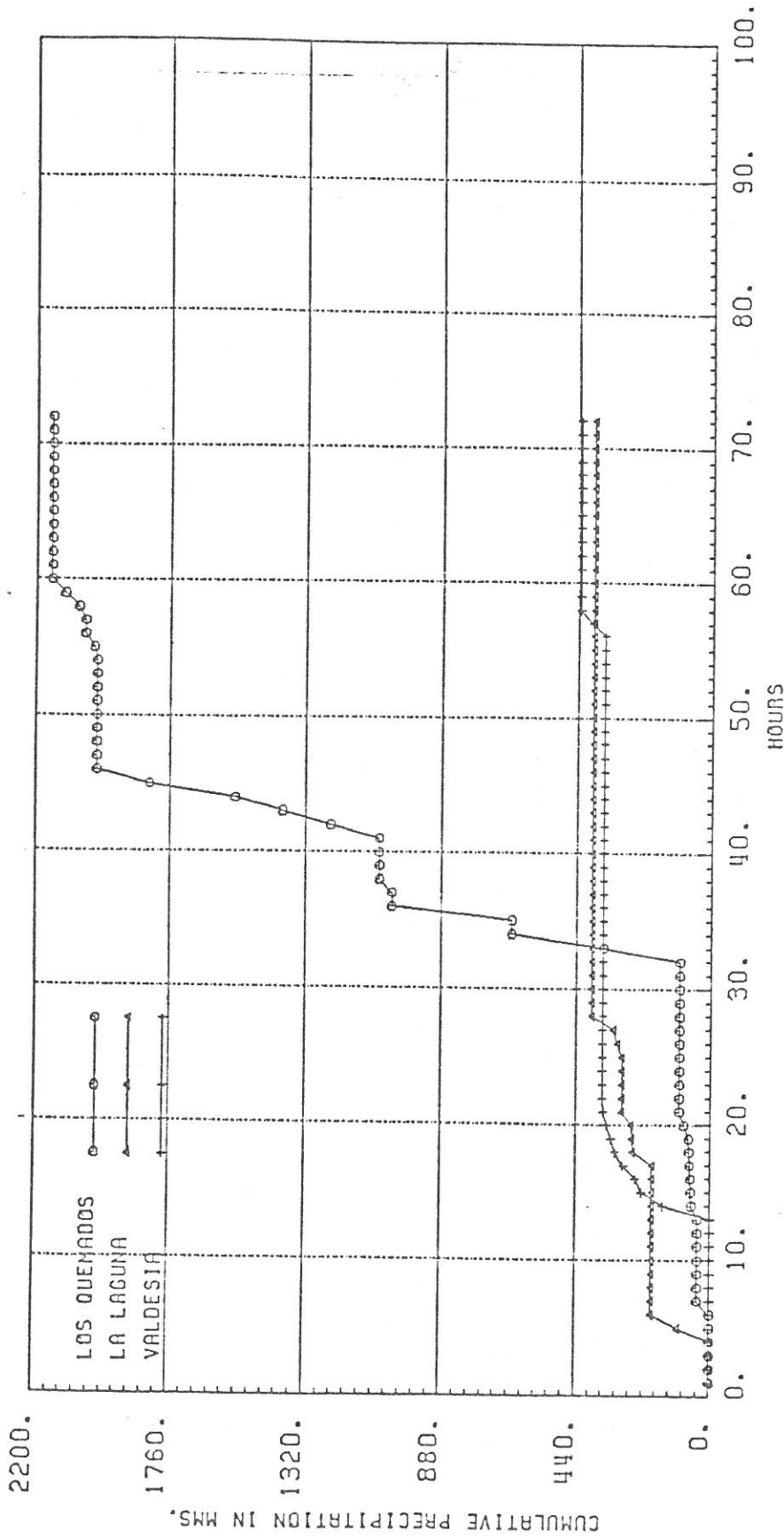
MASS CURVES STORM STARTING 63 OCT 1 20 ENDING 63 OCT 5 5

Figure 1.6.A.1.



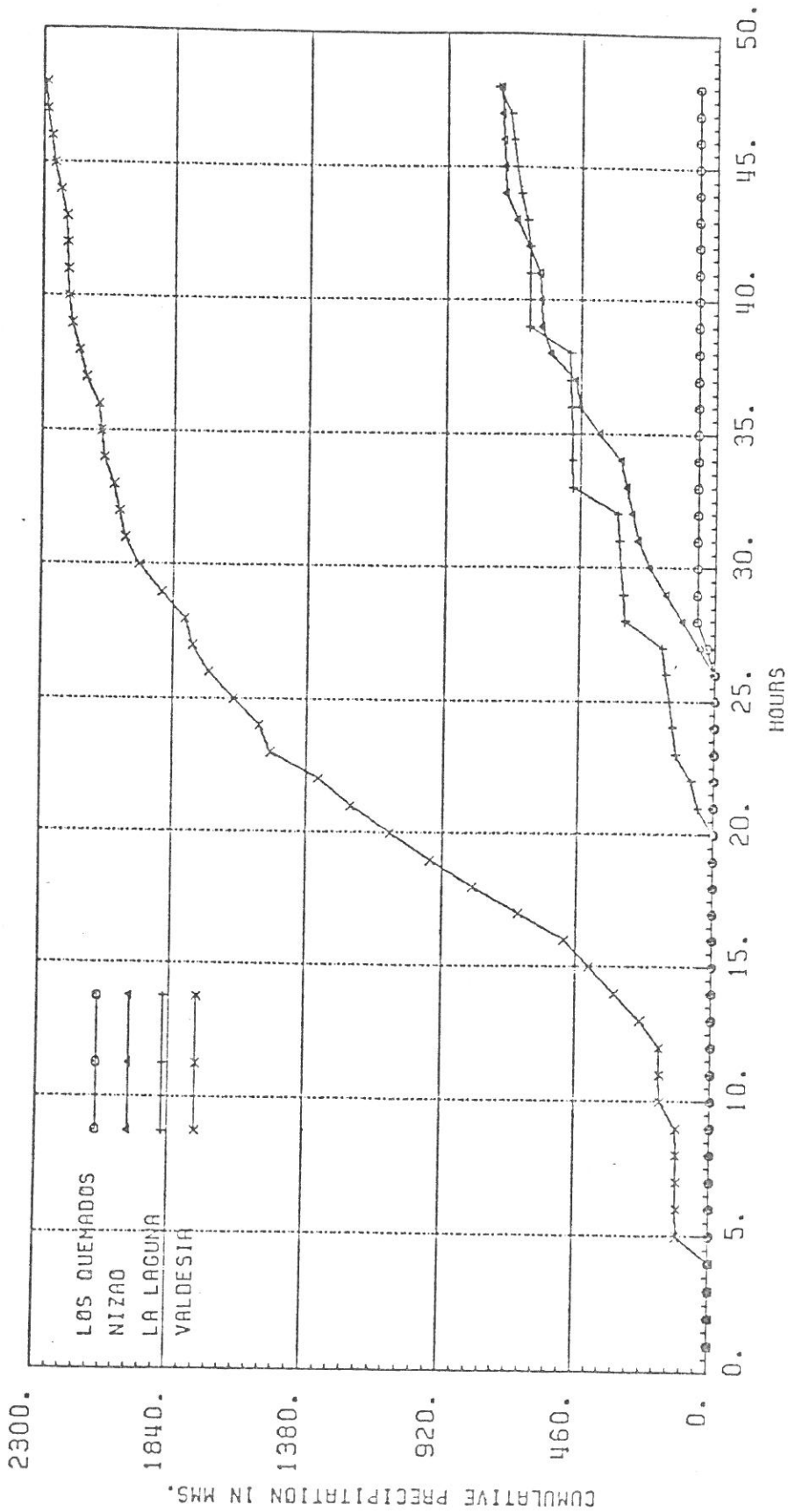
MASS CURVES STORM STARTING 64 AGO 6 8 ENDING 64 AGO 7 8

Figure 1.6.A.2



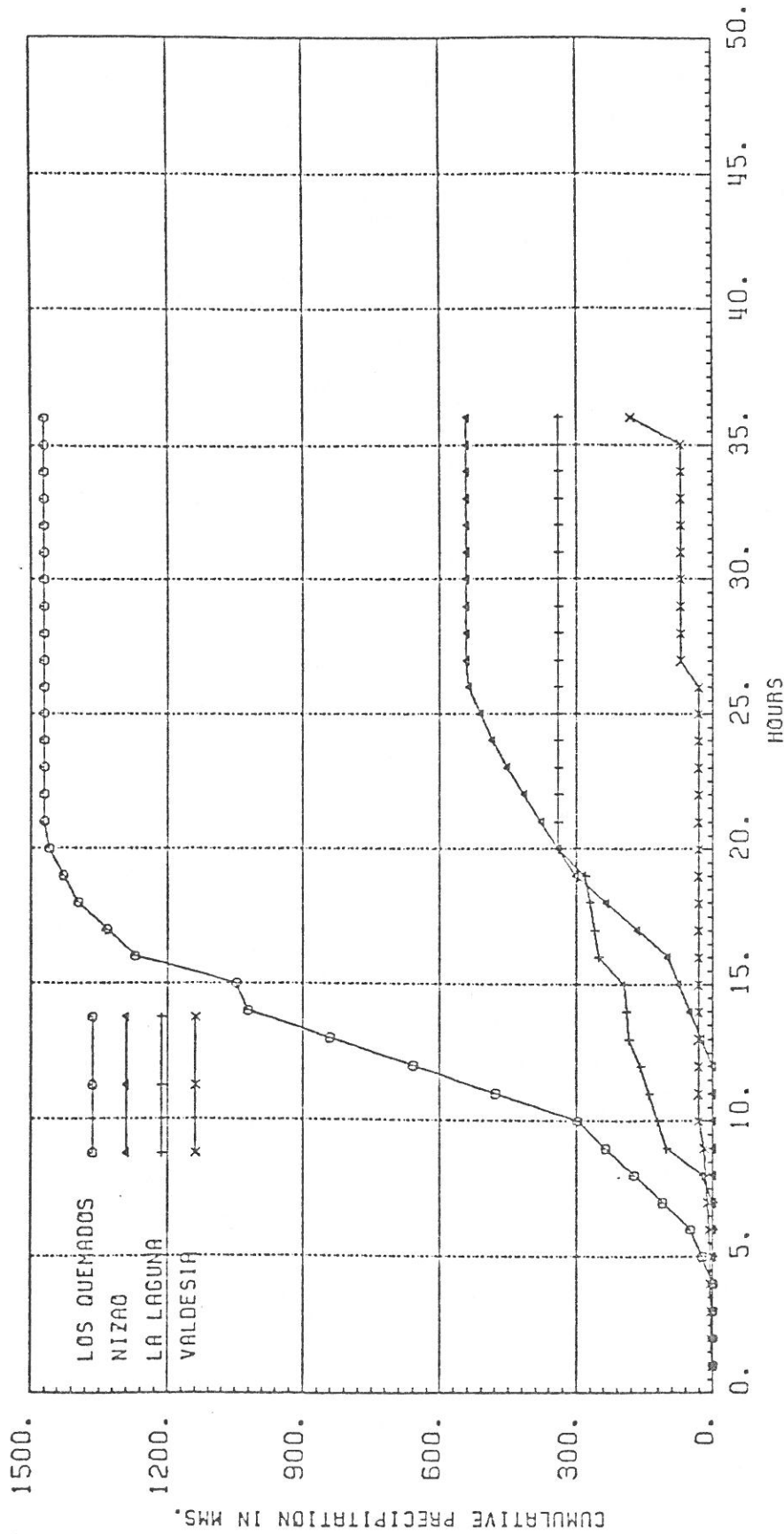
MASS CURVES STORM STARTING 65 MAY 2 8 ENDING 65 MAY 5 8

Figure 1.6.A.3.



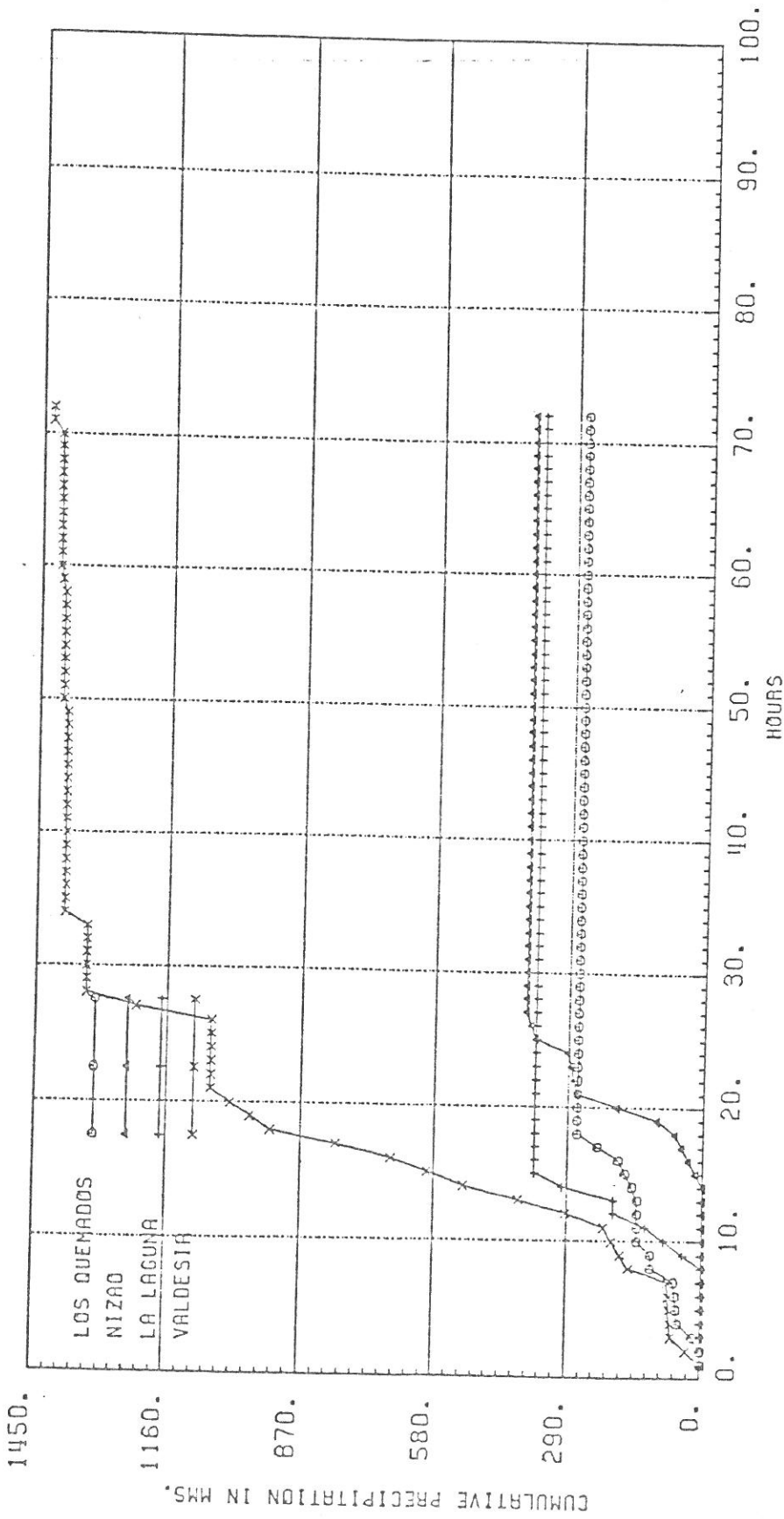
MASS CURVES STORM STARTING 66 MAY 25 8 ENDING 66 MAY 27 8

Figure 1.6.A.4



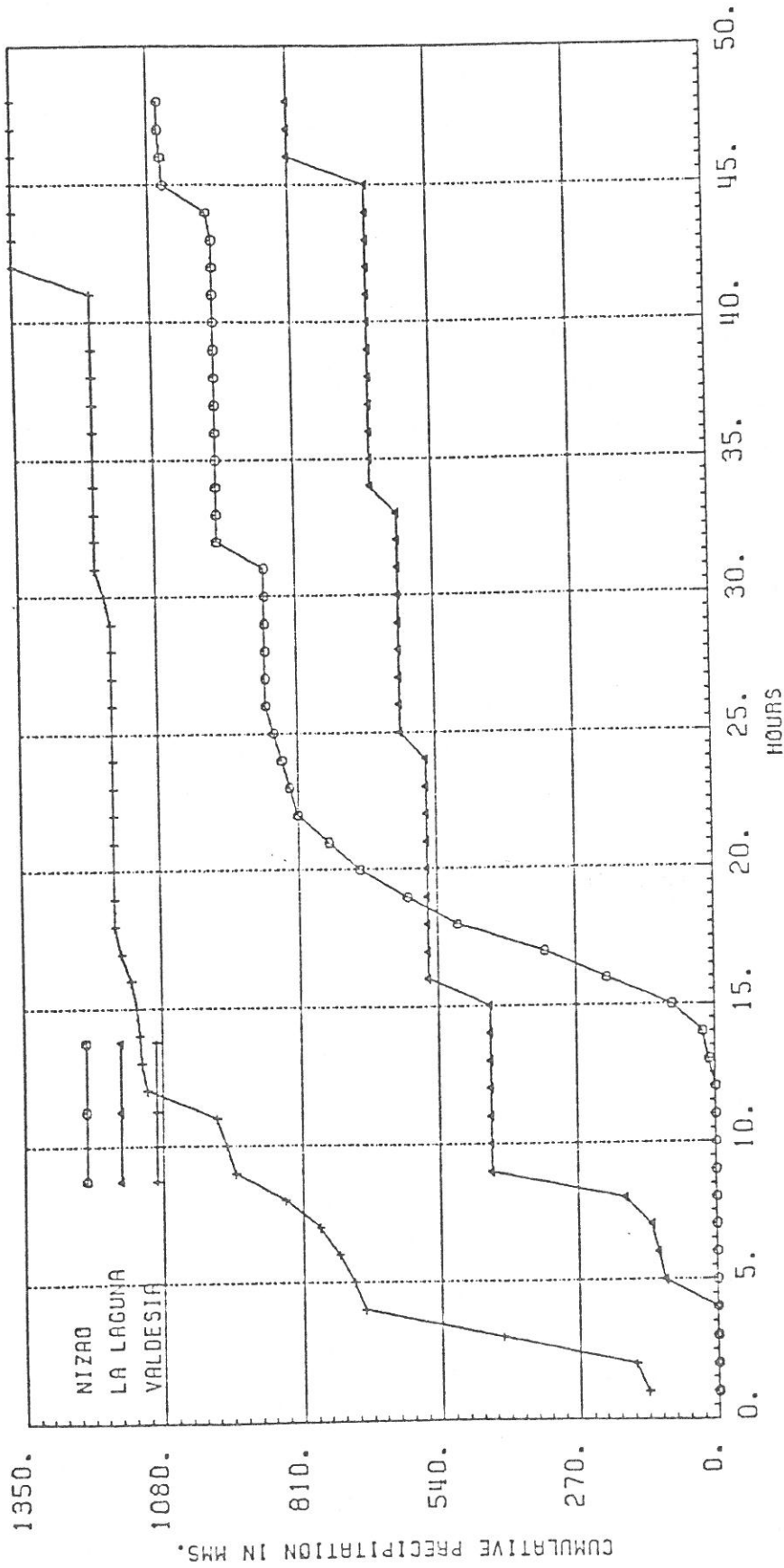
MASS CURVES STORM STARTING 66 SEP 28 20 ENDING 66 SEP 30 8

Figure 1.6.A.5



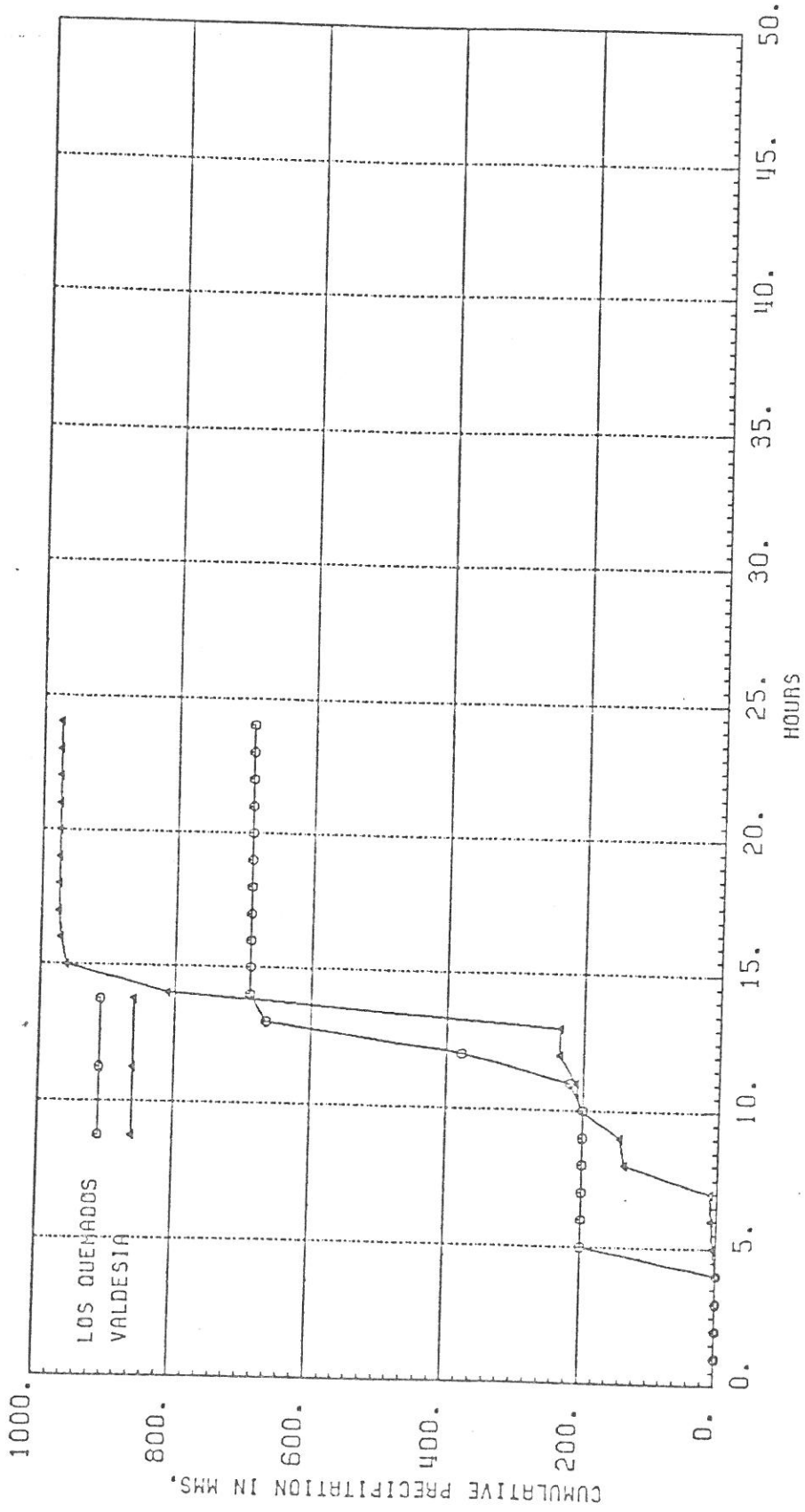
MASS CURVES STORM STARTING 67 SEP 10 8 ENDING 67 SEP 13 8

Figure 1.6.A.6



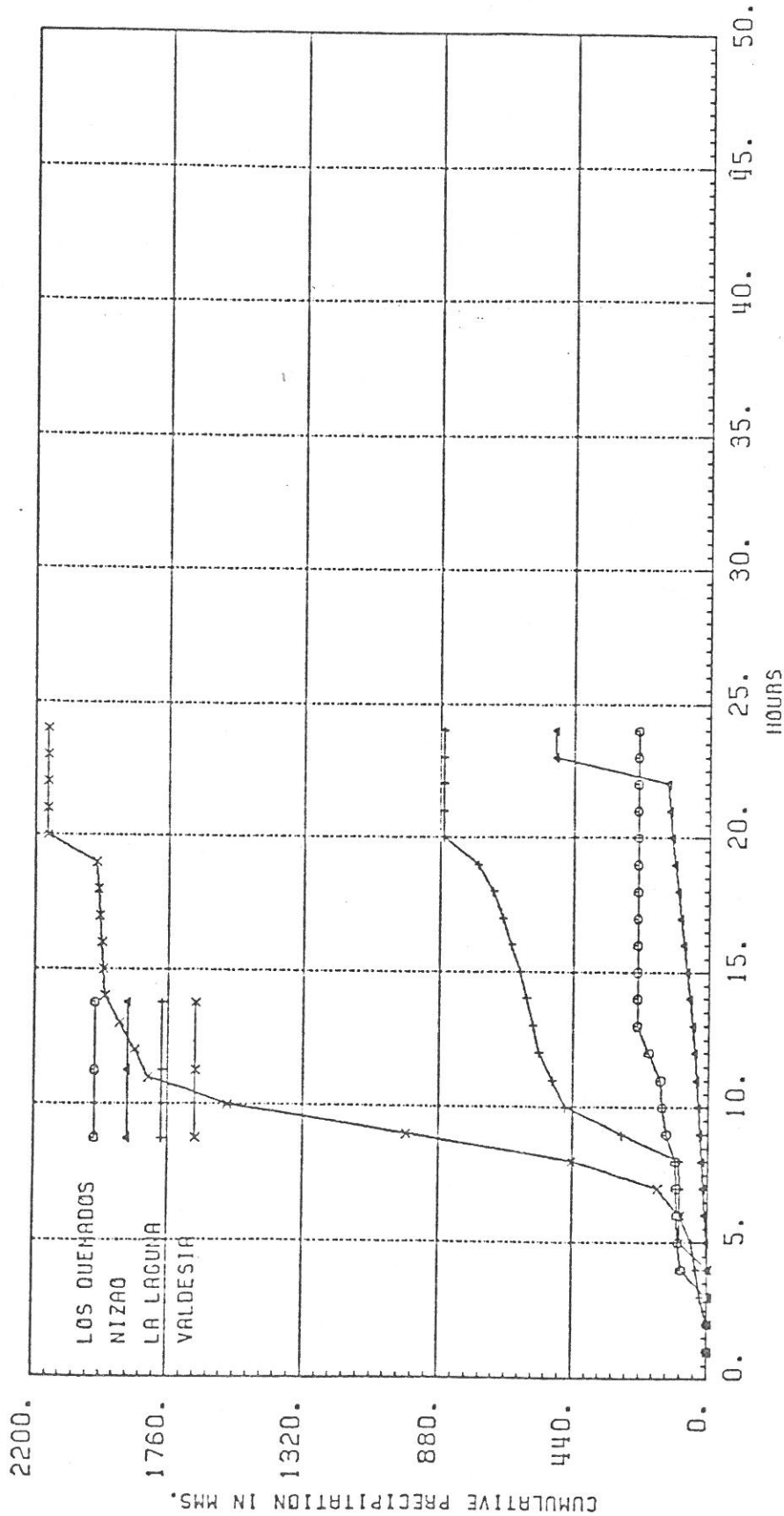
MASS CURVES STORM STARTING 68 AGO 8 8 ENDING 68 AGO 10 8

Figure 1.6.A.7



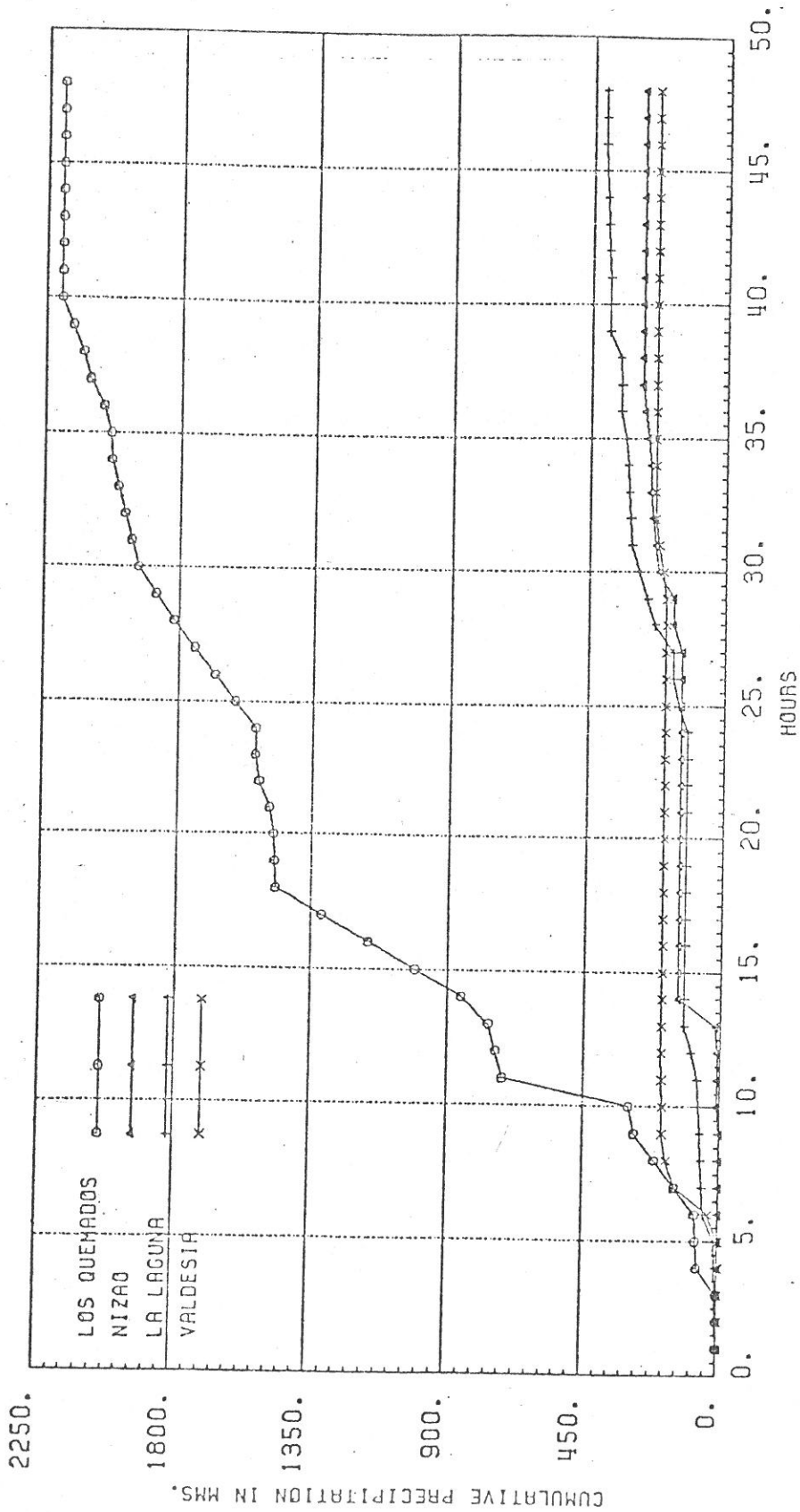
MASS CURVES STORM STARTING 69 JUL 19 8 ENDING 69 JUL 20 8

Figure 1.6.A.8



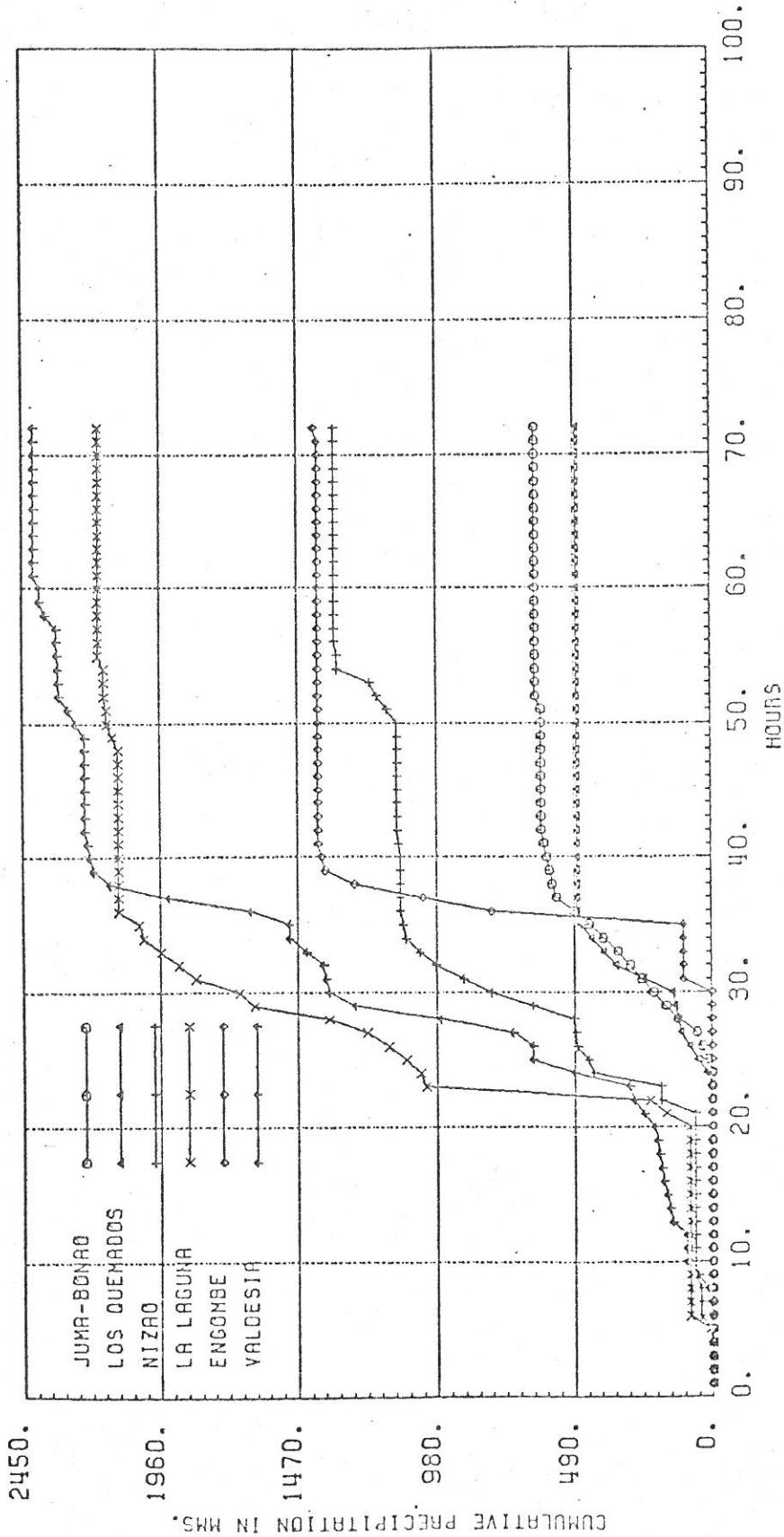
MASS CURVES STORM STARTING 70 AGO 22 8 ENDING 70 AGO 23 8

Figure 1.6.A.9



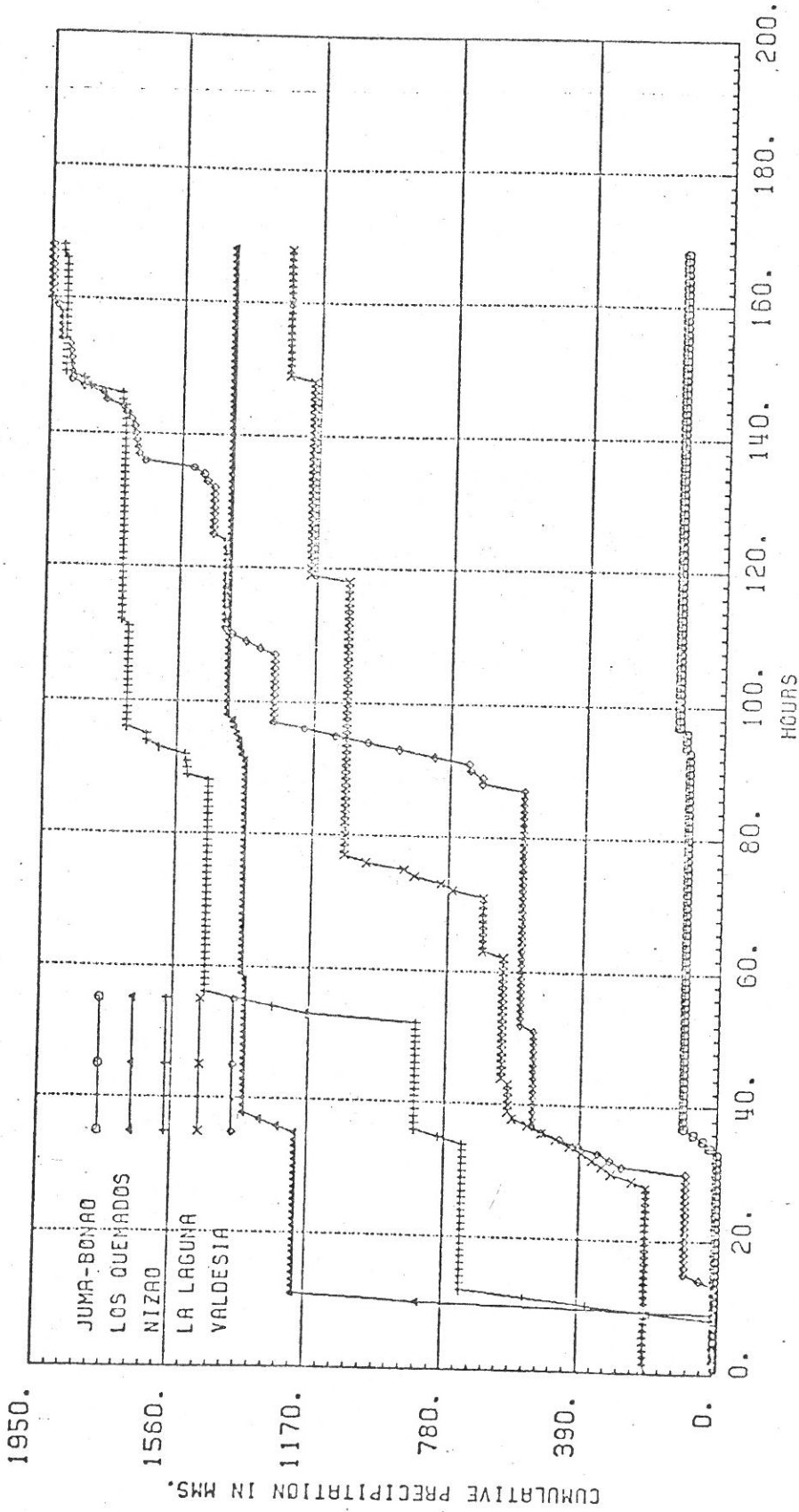
MASS CURVES STORM STARTING 71 FEB 19 8 ENDING 71 FEB 21 8

Figure 1.6.A.10



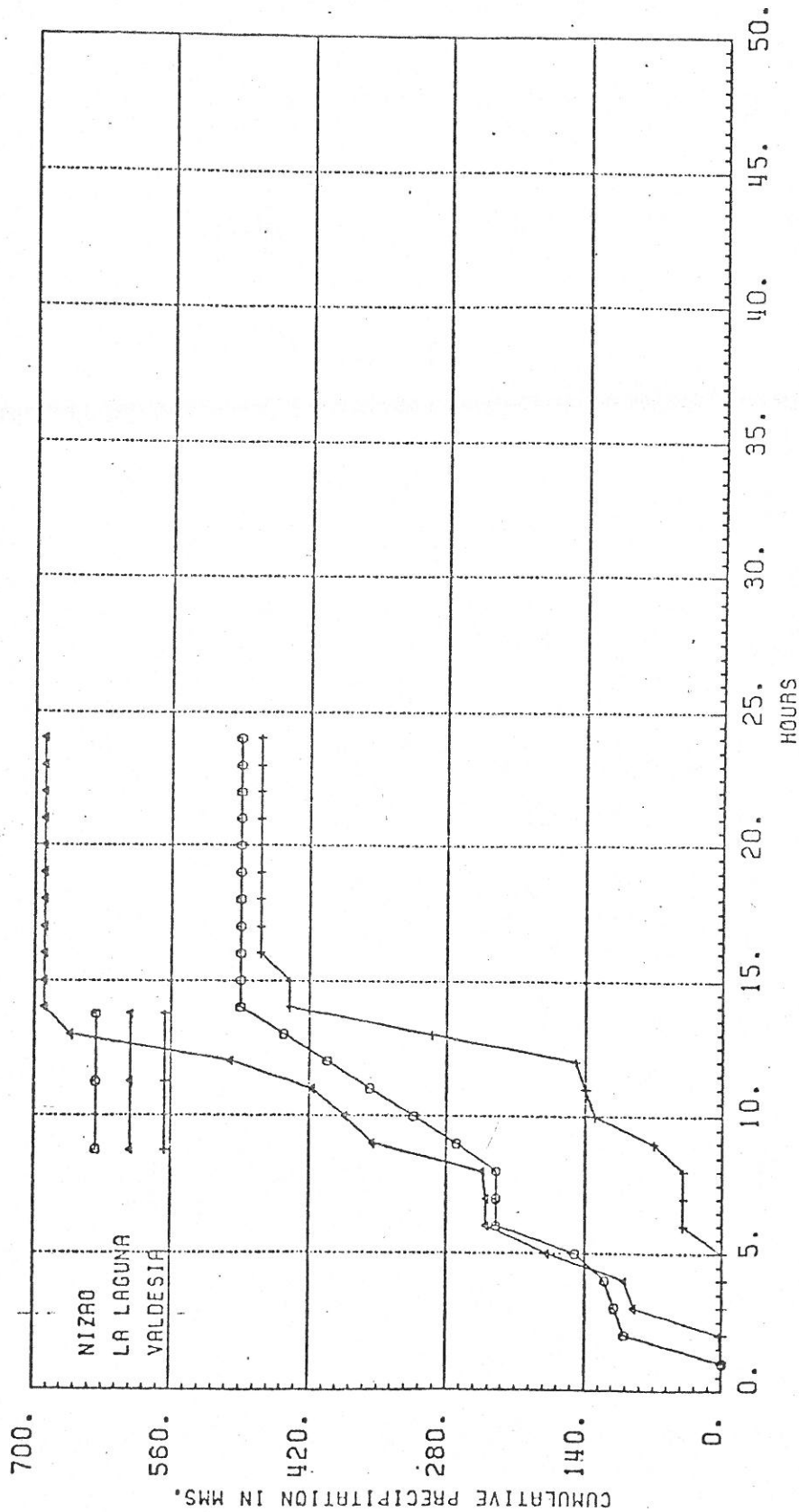
MASS CURVES STORM STARTING 72 MAY 20 8 ENDING 72 MAY 23 8

Figure 1.6.A.11



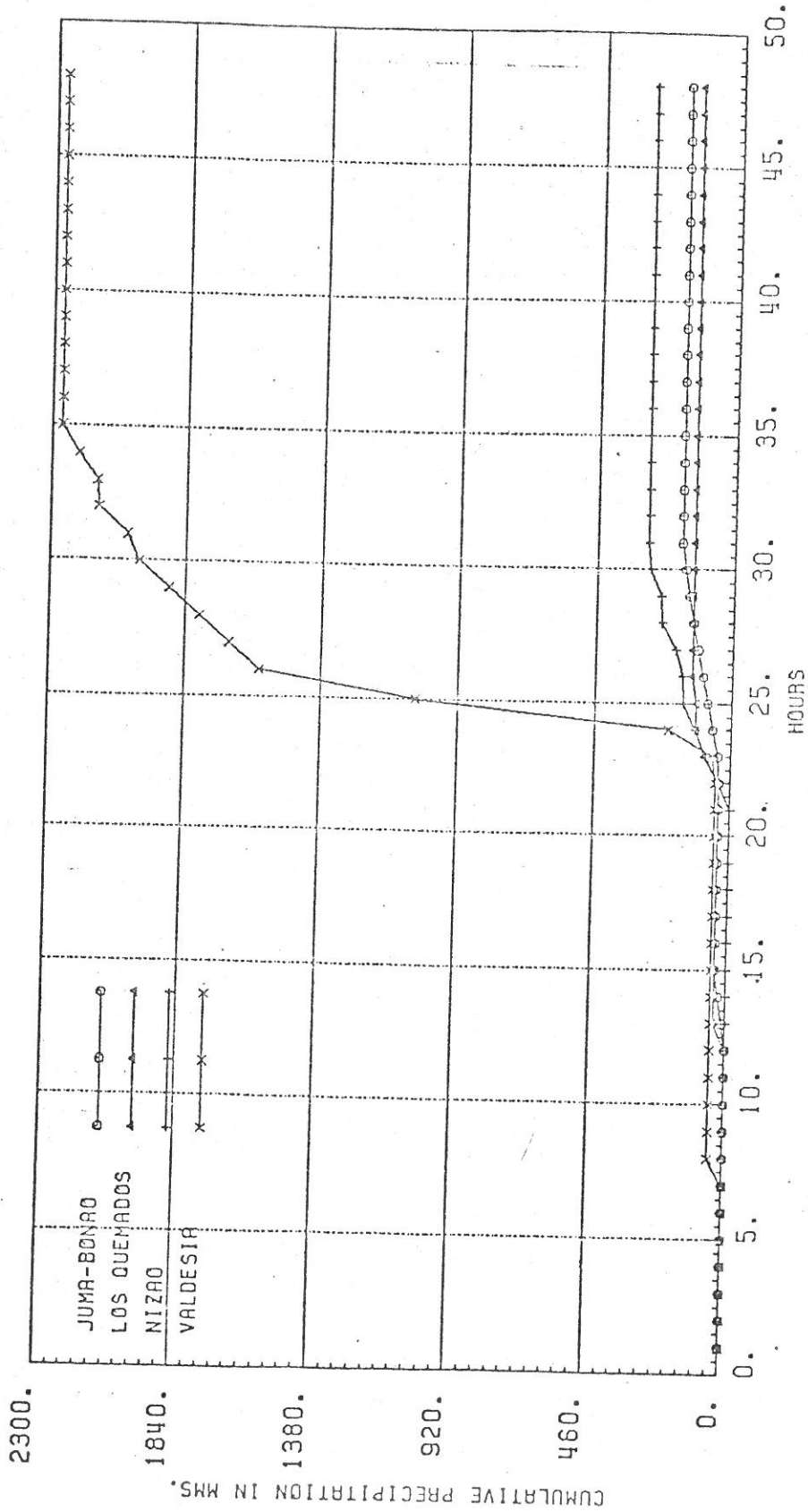
MASS CURVES STORM STARTING 73 OCT 14 8 ENDING 73 OCT 21 8

Figure 1.6.A.12



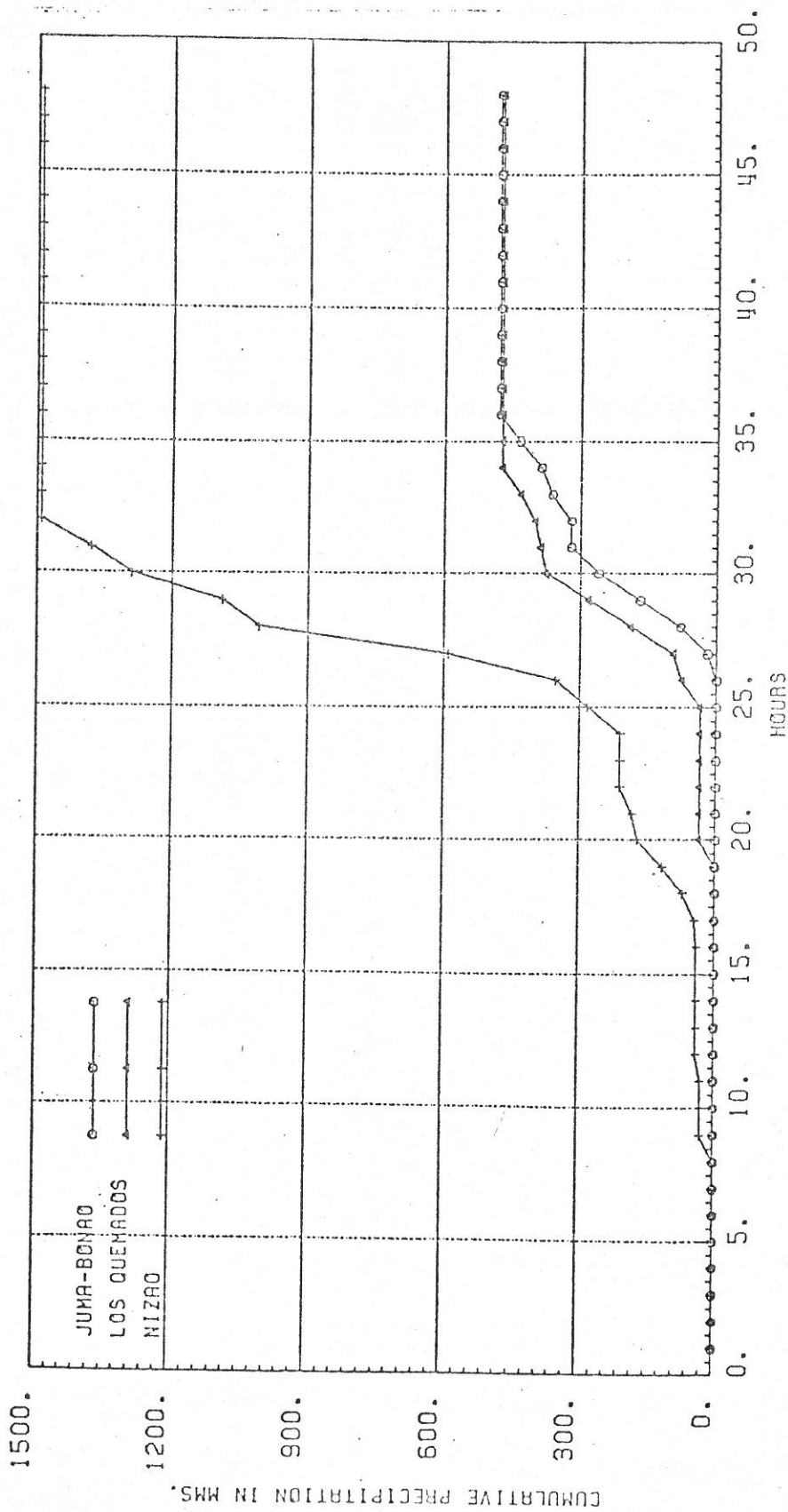
MASS CURVES STORM STARTING 74 AGO 30 8 ENDING 74 AGO 31 8

Figure 1.6.A.13



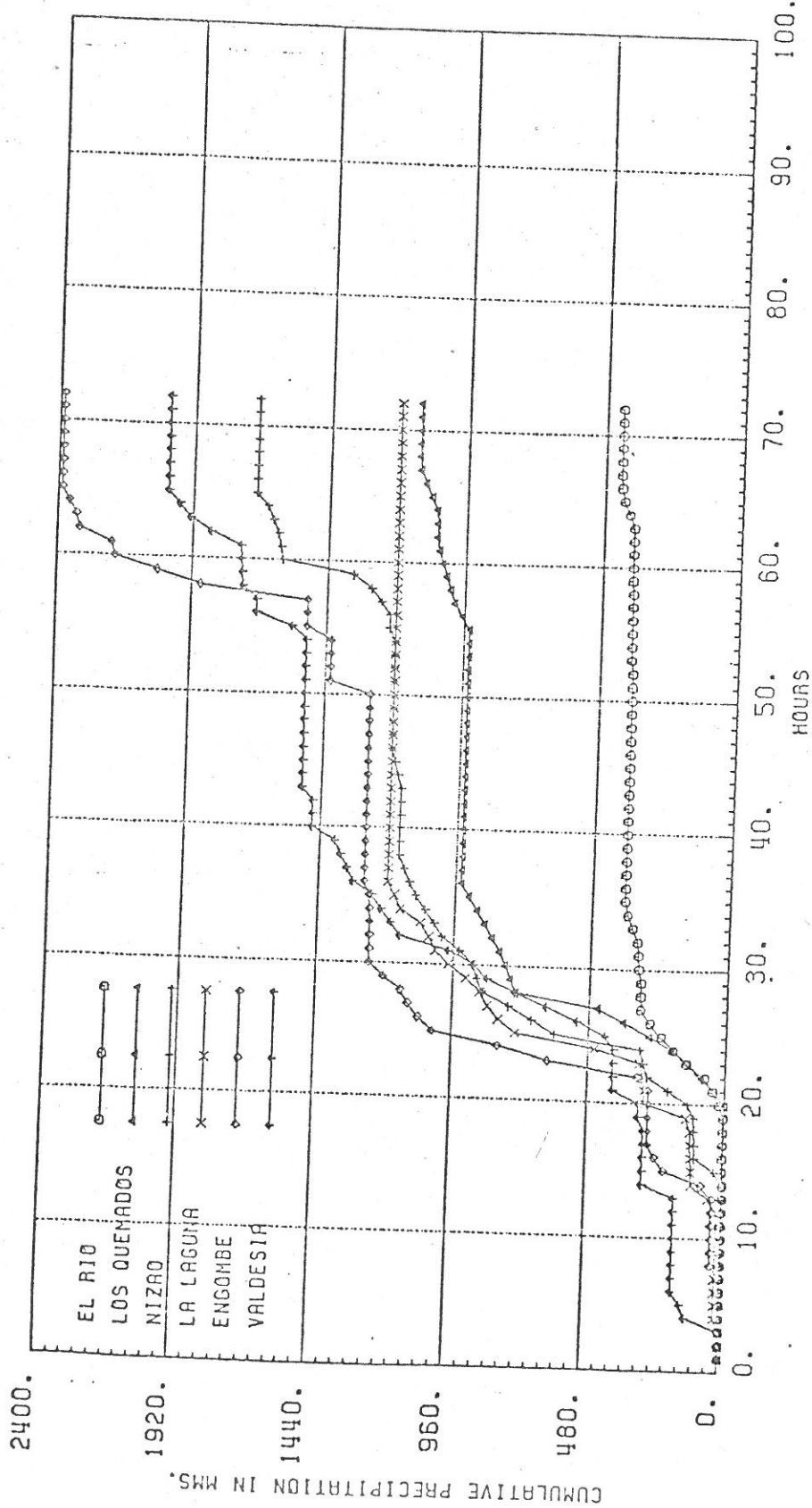
MASS CURVES STORM STARTING 75 SEP 16 8 ENDING 75 SEP 18 8

Figure 1.6.A.14



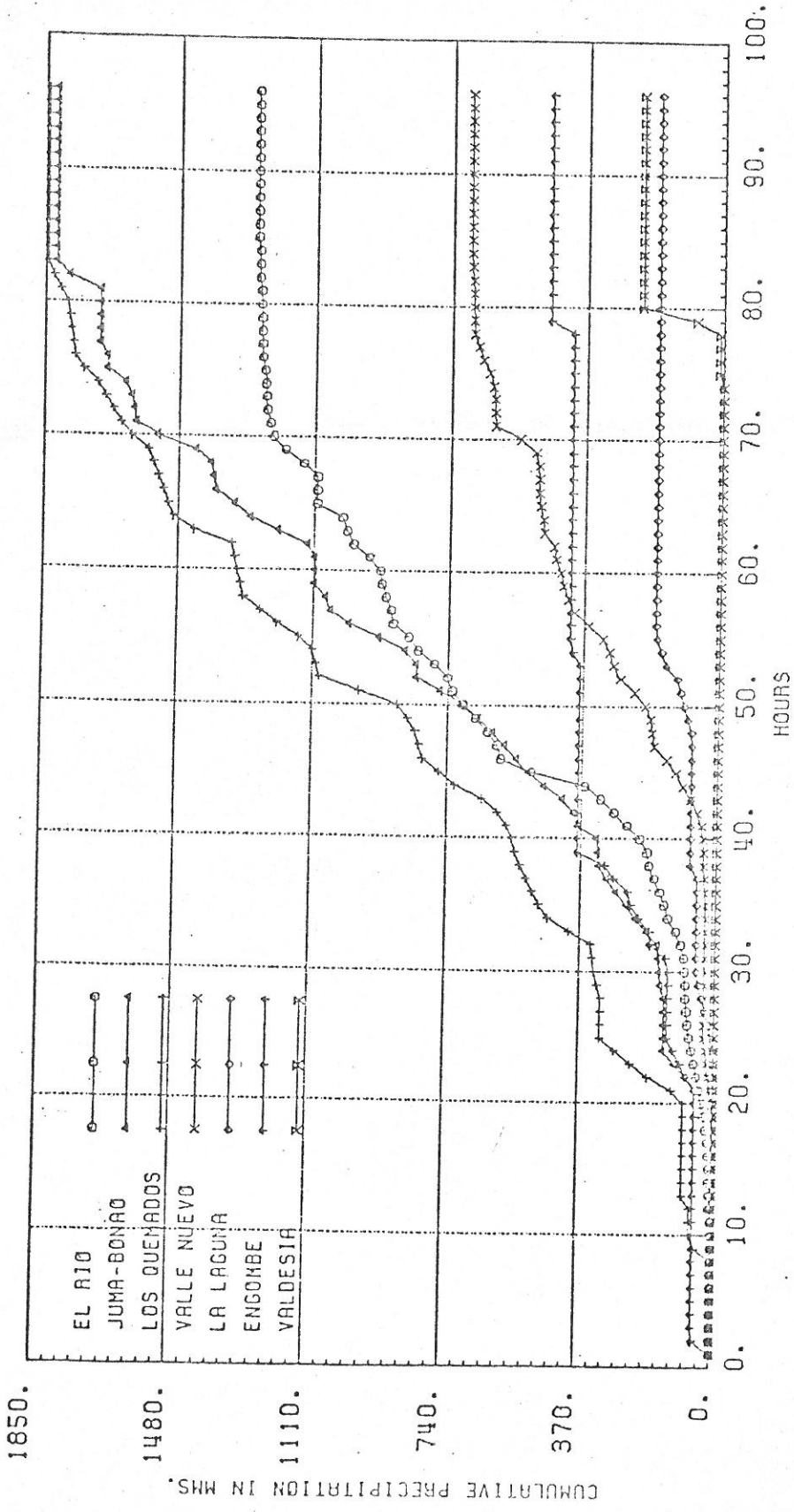
MASS CURVES STORM STARTING 76 OCT 10 8 ENDING 76 OCT 12 8

Figure 1.6.A.15



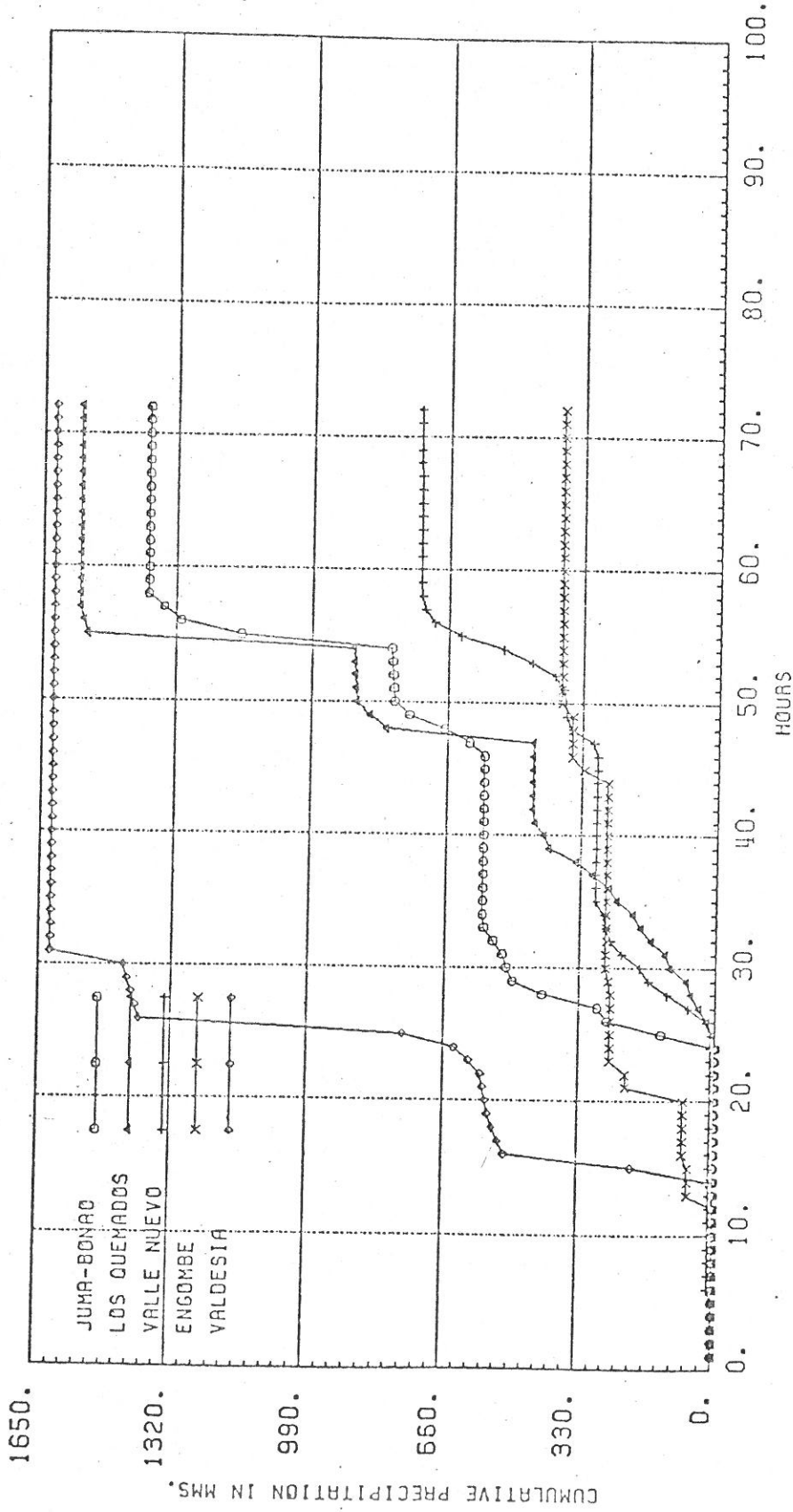
MASS CURVES STORM STARTING 77 MAY 21 8 ENDING 77 MAY 24 8

Figure 1.6.A.16



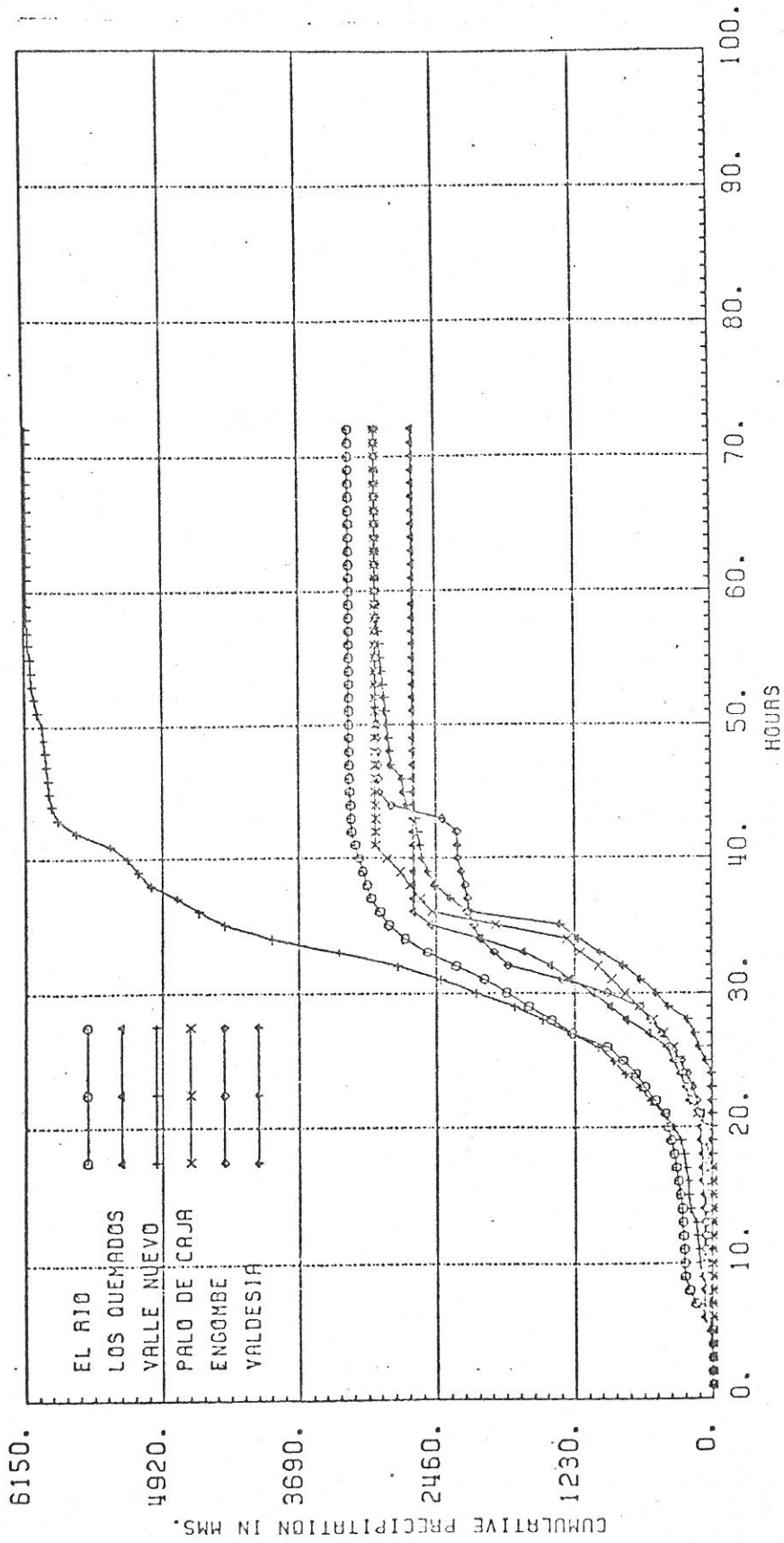
MASS CURVES STORM STARTING 77 DEC 28 8 ENDING 78 ENE 1 8

Figure 1.6.A.17



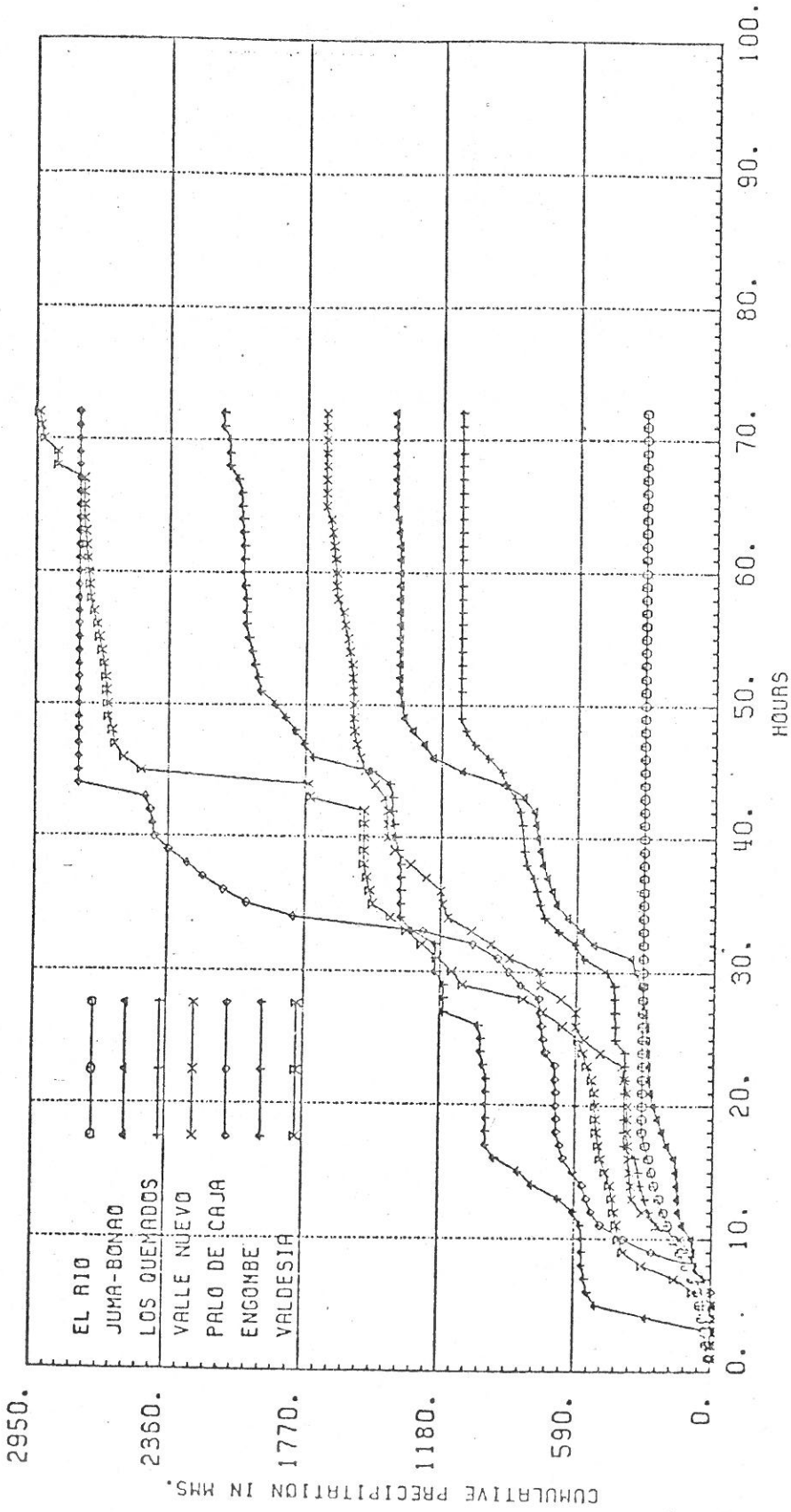
MASS CURVES STORM STARTING 78 AGO 3 8 ENDING 78 AGO 6 8

Figure 1.6.A.18



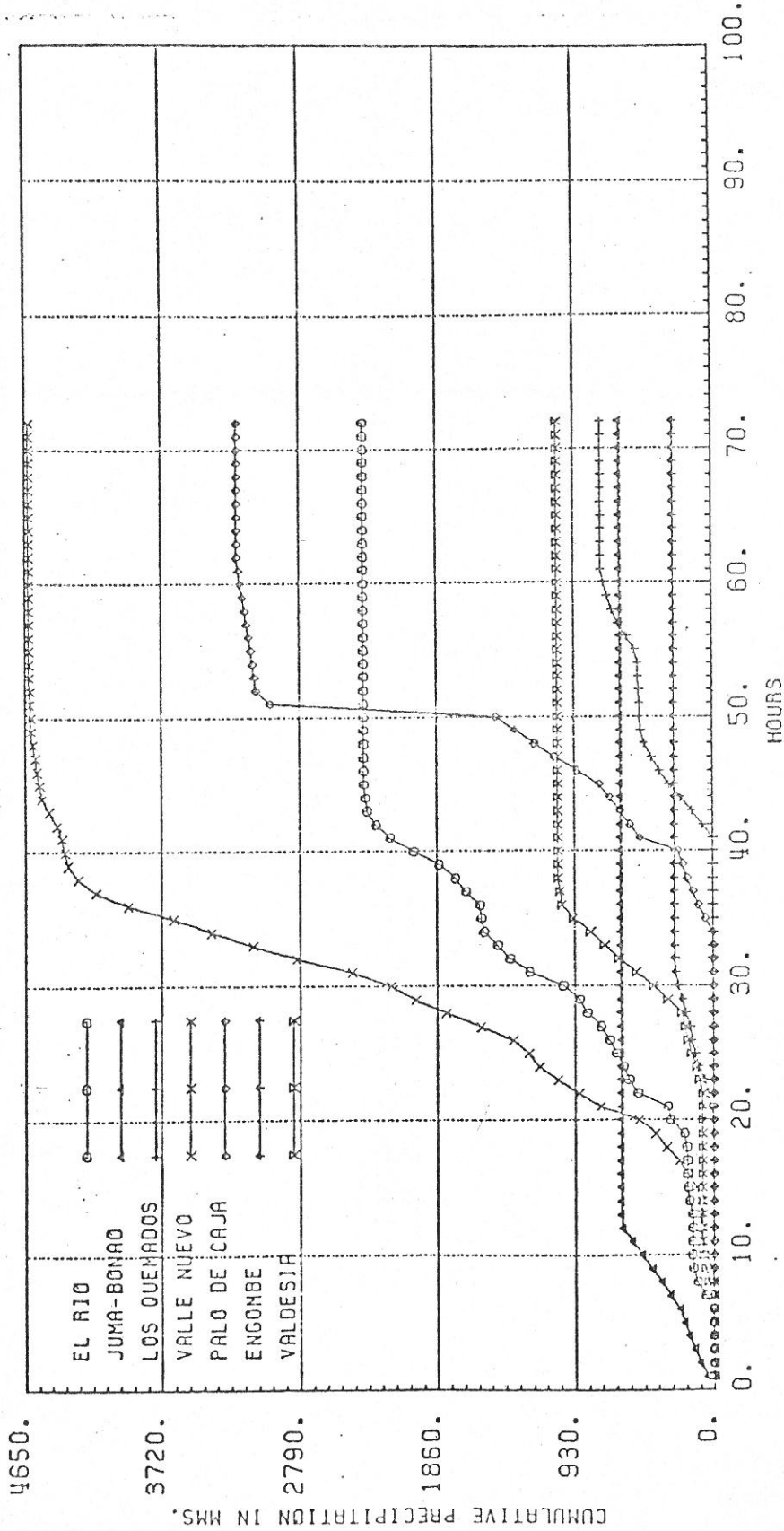
MASS CURVES STORM STARTING 79 AGO 30 8 ENDING 79 SEP 2 8

Figure 1.6.A.19



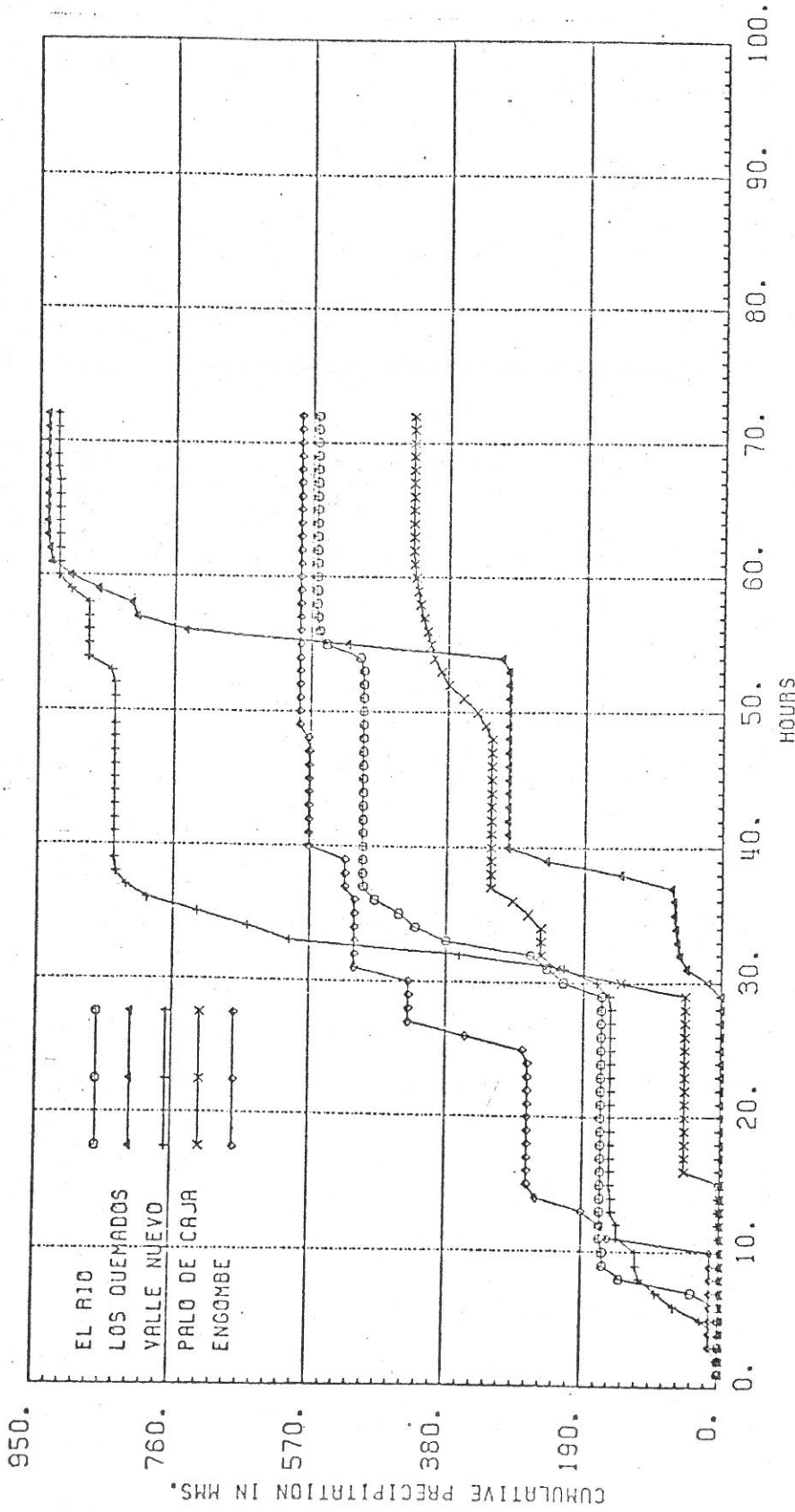
MASS CURVES STORM STARTING 79 SEP 5 8 ENDING 79 SEP 8 8

Figure 1.6.A.20



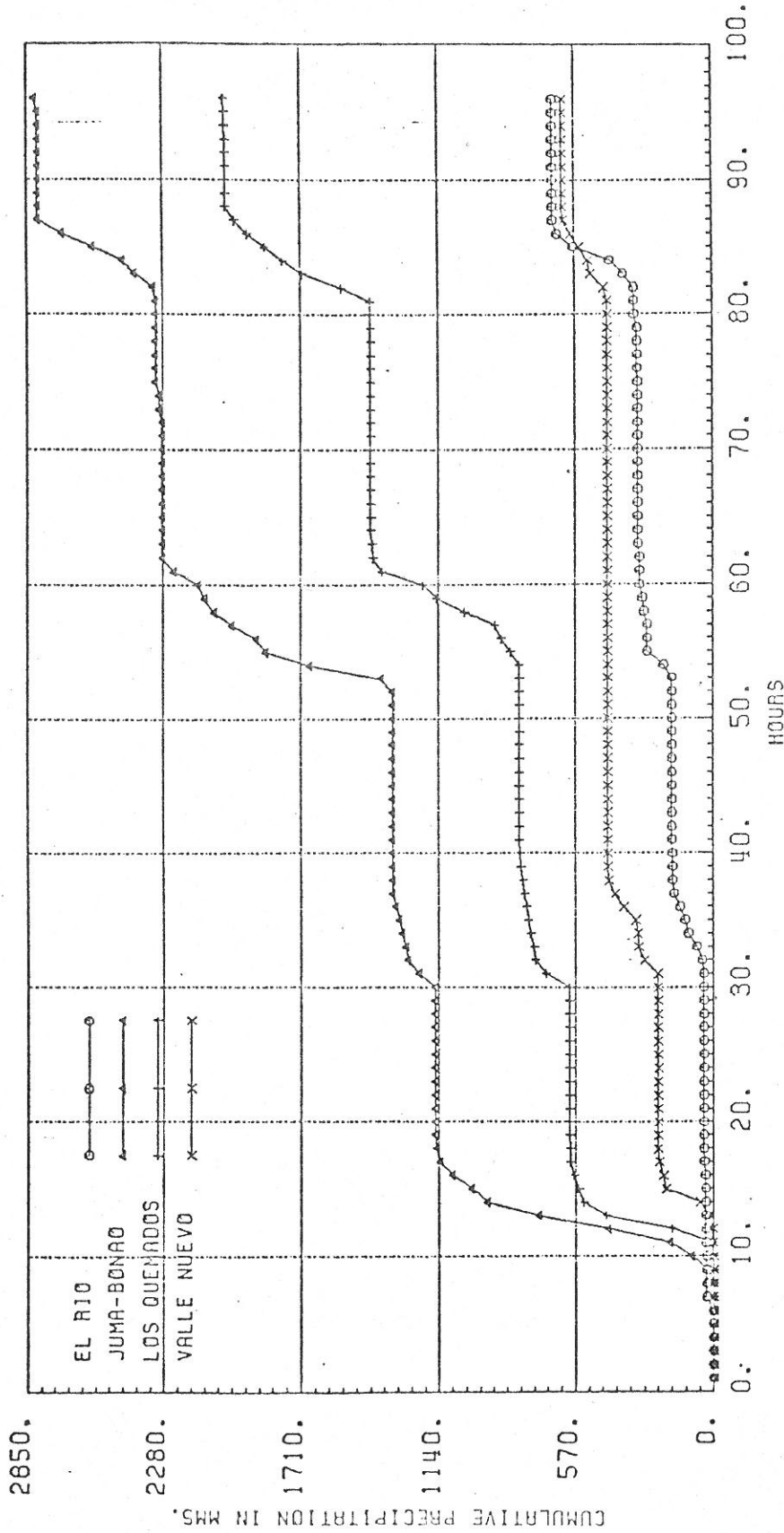
MASS CURVES STORM STARTING 80 AGO 4 8 ENDING 80 AGO 7 8

Figure 1.6.A.21



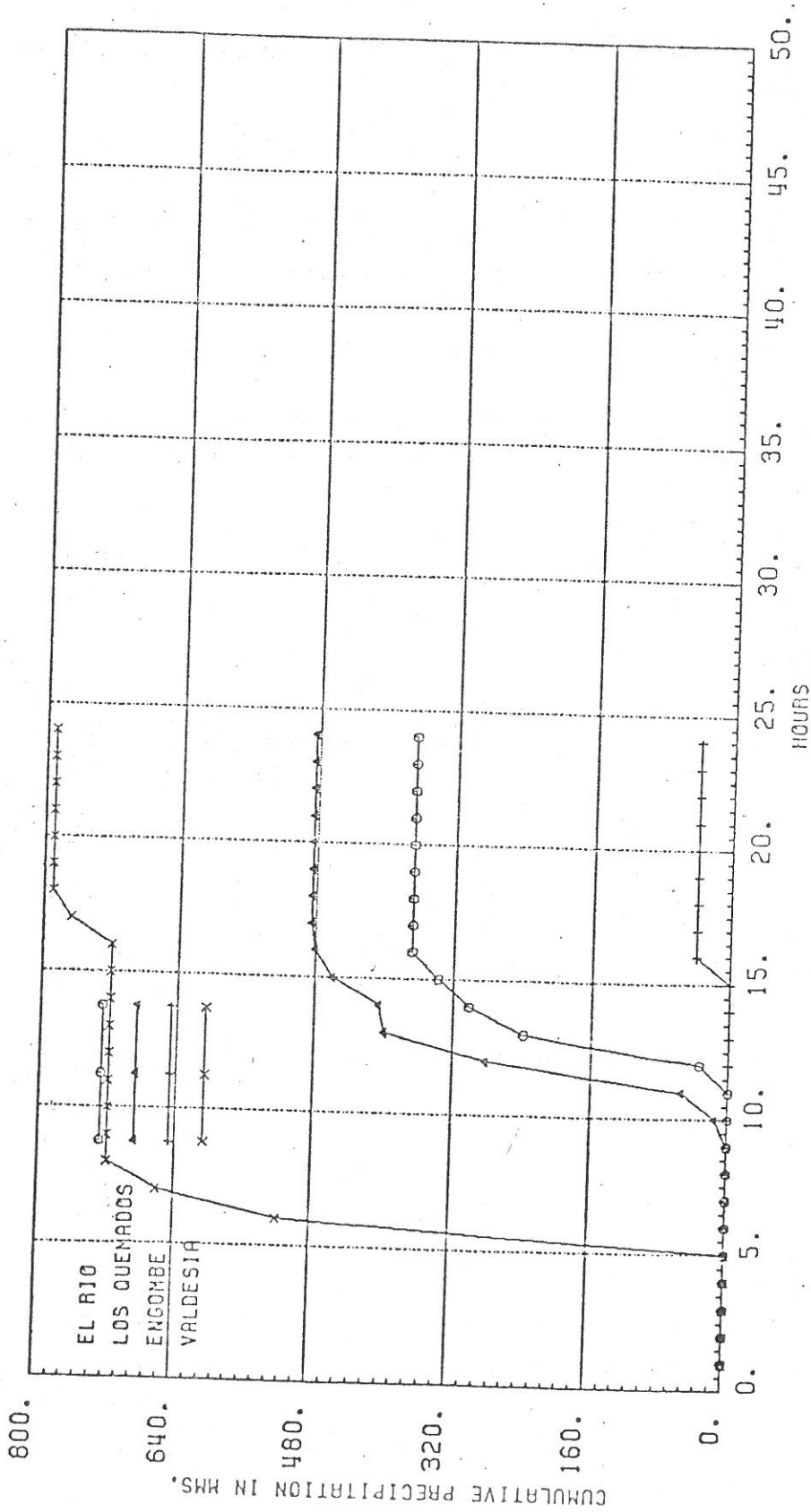
MASS CURVES STORM STARTING 81 MAY 8 8 ENDING 81 MAY 11 8

Figure 1.6.A.22



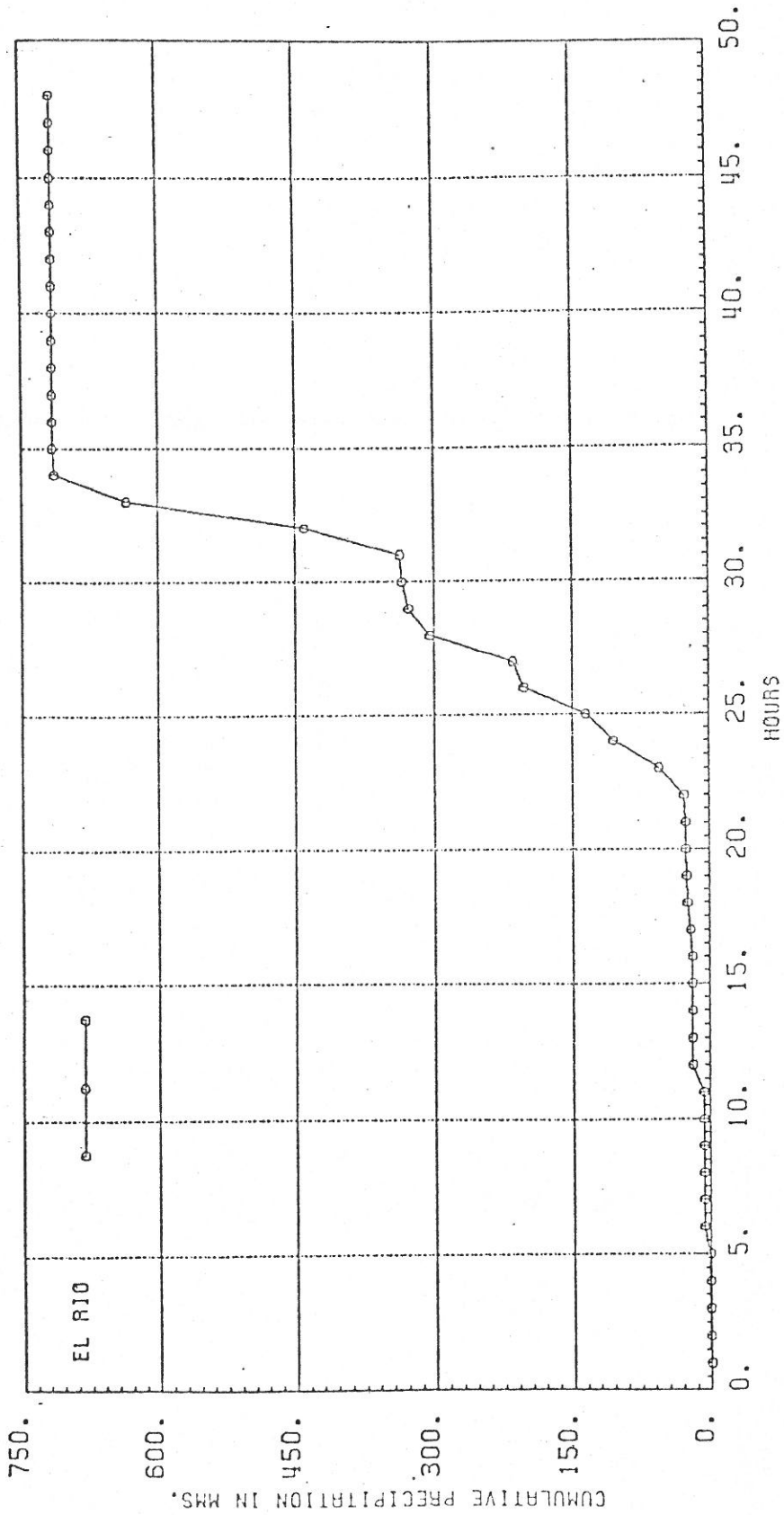
MASS CURVES STORM STARTING 82 MAY 9 8 ENDING 82 MAY 13 8

Figure 1.6.A.23



MASS CURVES STORM STARTING 83 ABR 12 8 ENDING 83 ABR 13 8

Figure 1.6.A.24



MASS CURVES STORM STARTING 84 AGO 1 8 ENDING 84 AGO 3 8

Figure 1.6.A.25

APPENDIX 1.6.B

PROGRAM LISTING FOR RAINFALL ISOHYETH
COMPUTATIONS

PROGRAM PCMAP (INPUT,OUTPUT,TAPE5=INPUT,TAPE6=OUTPUT,
1 TAPE10)

C
C
C

PROGRAM FOR MAPPING, INTERPOLATION AND AREAL AVERAGING

CHARACTER*8 LIST,ALFA,PROC,TITLE,SNAME,BL,BLANK,ENDER,SKIP

CHARACTER*1 STN

CHARACTER*8 SDNAME

DIMENSION

ALFA(10) , PROC(23) , BL(3) ,

1 IND(5)

COMMON /SYSID/ STORE(400,101)

COMMON /SYSID2/ TITLE(10) , SNAME(50)

COMMON /CONST/ IDSC , NS , NX ,

1 NY , DX , DY , XMAX ,

2 XMIN , YMAX , YMIN , HMAX ,

3 HMIN , SMAX , SMIN ,

COMMON /SYSTAT/ AMU , STD , STE ,

1 RC , ATC , CRSCO(50) , SPAC(50,50) ,

COMMON /CVAR/ KVR , OMEGA , ALPHA ,

1 SVAR(50) , SMU(50) ,

COMMON /SYSDAT/ X(50) , Y(50) , H(50) ,

1 Z(50) ,

COMMON /SBOUND/ NB , XB(100) , YB(100) ,

COMMON /SBOUND2/ STN(50,8)

COMMON /AREAL/ AREA , TAR ,

COMMON /CODE/ ID , VR(5) , KCOR ,

1 KERR

COMMON /CORP/ NCR , FA , FB ,

1 AXM , AYL , XM ,

COMMON /PTP/ NP , STP(50) , XP(50) ,

1 YP(50) , KDP(50) ,

COMMON /DEPTH1/ SDNAME(50)

COMMON /DEPTH2/ NRS , XRS(50) , YRS(50)

DATA PROC/'TITLE','RANGE','STATION','BOUNDARY','DATA','STATISTI',

1 'THIESSEN','HARMONIC','OPTIMAL','POLYNOMI','INVERSE','QUADRIC',

2 'KRIGING','STOP','TAPE','RETITLE','MASK','SPACOR','VARIOGRA',

3 'PTDATA','DEPTH','STAMEANS','STADEPTH'/

DATA BLANK/' /,ENDER/'END'/

IDSC = 200

RTXY=8.0

NIG = 23

100 REWIND 10

KCOR = 0

KVR = 0

NC = 0

NB=0

NS=0

NRS=0

IND(1) = 0

IND(2) = 0

IND(3) = 0

IND(4) = 0

IND(5) = 0

110 READ (5,940,ERR=120,END=120) LIST,ALFA

NC = NC + 1

WRITE (10,930) ALFA

```
IF (LIST.EQ.PROC(14)) GO TO 120
GO TO 110
120 IF (NC.EQ.0) STOP
WRITE (10,820) PROC(14)
REWIND 10
WRITE (6,800)
NR = 0
130 NR = NR + 1
IF (NR.GT.NC) GO TO 400
READ (10,940) LIST,ALFA
DO 140 IG = 1,NIG
    IF (LIST.EQ.PROC(IG)) GO TO 150
140 CONTINUE
PNT = NR/1000.
WRITE (6,950) PNT,ALFA
GO TO 130
150 GO TO (160,180,190,220,300,360,370,370,370,370,370,370,370,380,400,390
1,170,250,260,270,280,370,300,375), IG
160 IND(1) = 1
170 PNT = NR/1000.
WRITE (6,960) PNT,ALFA
NR = NR + 1
READ (10,930) ALFA
PNT = NR/1000.
WRITE (6,970) PNT,ALFA
GO TO 130
180 IND(2) = 1
PNT = NR/1000.
WRITE (6,960) PNT,ALFA
NR = NR + 1
READ (10,930) ALFA
PNT = NR/1000.
WRITE (6,970) PNT,ALFA
GO TO 130
190 IND(3) = 1
PNT = NR/1000.
WRITE (6,960) PNT,ALFA
200 NR = NR + 1
IF (NR.GT.NC) GO TO 350
NS = NS + 1
READ (10,940) LIST,ALFA
PNT = NR/1000.
IF (LIST.EQ.ENDER) GO TO 210
WRITE (6,970) PNT,ALFA
GO TO 200
210 WRITE (6,970) PNT,ALFA
NS = NS - 1
GO TO 130
220 PNT = NR/1000.
WRITE (6,960) PNT,ALFA
230 NR = NR + 1
IF (NR.GT.NC) GO TO 350
NB = NB + 1
READ (10,940) LIST,ALFA
PNT = NR/1000.
IF (LIST.EQ.ENDER) GO TO 240
```

```
WRITE (6,970) PNT,ALFA
GO TO 230
240 WRITE (6,970) PNT,ALFA
NB = NB - 1
GO TO 130
250 PNT = NR/1000.
WRITE (6,1010) PNT,ALFA
GO TO 130
260 PNT = NR/1000.0
WRITE (6,960) PNT,ALFA
NR = NR + 1
READ (10,930) ALFA
PNT = NR/1000.0
WRITE (6,970) PNT,ALFA
GO TO 130
270 PNT = NR/1000.
WRITE (6,960) PNT,ALFA
NR = NR + 1
READ (10,930) ALFA
PNT = NR/1000.
WRITE (6,970) PNT,ALFA
GO TO 130
280 PNT = NR/1000.
WRITE (6,1010) PNT,ALFA
BACKSPACE 10
READ (10,1110) KPT
IF (KPT.EQ.0) GO TO 130
290 NR = NR + 1
IF (NR.GT.NC) GO TO 350
READ (10,940) LIST,ALFA
PNT = NR/1000.
WRITE (6,970) PNT,ALFA
IF (LIST.EQ.ENDER) GO TO 130
GO TO 290
300 KND = 4
IF (IG.NE.5) KND = 5
NDT = 0
IND(KND) = 1
PNT = NR/1000.
WRITE (6,960) PNT,ALFA
310 NR = NR + 1
IF (NR.GT.NC) GO TO 350
NDT = NDT + 1
READ (10,940) LIST,ALFA
PNT = NR/1000.
IF (LIST.EQ.ENDER) GO TO 320
WRITE (6,970) PNT,ALFA
GO TO 310
320 NDT = NDT - 1
WRITE (6,970) PNT,ALFA
IF (NDT.EQ.NS) GO TO 130
IND(KND) = 0
IF (KND.EQ.4) WRITE (6,980)
IF (KND.EQ.5) WRITE (6,990)
330 WRITE (6,1000)
NR = NR + 1
```

```
DO 340 KF = NR,NC
  READ (10,940) LIST,ALFA
  PNT = KF/1000.
340 WRITE (6,970) PNT,ALFA
  GO TO 790
350 WRITE (6,1070) PROC(IG)
  WRITE (6,1000)
  WRITE (6,1060)
  GO TO 790
360 PNT = NR/1000.
  WRITE (6,960) PNT,ALFA
  NR = NR + 1
  READ (10,930) ALFA
  PNT = NR/1000.
  WRITE (6,970) PNT,ALFA
  GO TO 130
370 PNT = NR/1000.
  WRITE (6,1010) PNT,ALFA
  GO TO 130
375 PNT = NR/1000.
  WRITE (6,1010) PNT, ALFA
376 NR = NR + 1
  IF(NR.GT.NC) GO TO 350
  READ (10,940) LIST, ALFA
  PNT = NR/1000.
  IF (LIST.EQ.ENDR) GO TO 377
  WRITE (6,970) PNT, ALFA
  GO TO 376
377 WRITE (6,970) PNT, ALFA
  GO TO 130
380 PNT = NR/1000.
  WRITE (6,1010) PNT,ALFA
  BACKSPACE 10
  READ (10,1110) ID
  IF (ID.NE.1) GO TO 130
  IF (IND(5).EQ.1) GO TO 130
  WRITE (6,1050)
  GO TO 330
390 PNT = NR/1000.
  WRITE (6,1020) PNT
  GO TO 130
400 KERR = 0
  DO 410 II = 1,4
    IF (IND(II).EQ.1) GO TO 410
    KERR = 1
    WRITE (6,1040) PROC(II)
410 CONTINUE
  IF (KERR.NE.1) GO TO 420
  WRITE (6,1060)
  GO TO 790
420 PNT = NR/1000.
  WRITE (6,1030) PNT
  IF (LIST.NE.PROC(14)) WRITE (6,1080)
  REWIND 10
  NR = 0
430 NR = NR + 1
```



```

IF (NR.GT.NC) GO TO 790
KERR = 0
READ (10,810) LIST, ID, VR
DO 440 IG = 1, NIG
  IF (LIST.EQ.PROC(IG)) GO TO 450
440 CONTINUE
  GO TO 430
450 IF (IG.LE.6) GO TO 480
  IF (ISIG.EQ.1) GO TO 480
  ISIG = 1
  CALL CHECK (RTXY)
  IF (KERR.EQ.1) GO TO 790
  CALL BOUND (INAME)
  GO TO 480
460 CALL TAPERWR (PROC(ITP), STORE, 0)
  GO TO 430
480 GO TO (490, 500, 510, 530, 550, 570, 600, 610, 620, 630, 640, 650, 660, 790, 680
  1, 490, 730, 740, 770, 780, 670, 580, 675), IG
490 NR = NR + 1
  READ (10, 820) TITLE
  GO TO 430
500 NR = NR + 1
  INAME = ID
  READ (10, 810) LIST, ID, VR
  ITP = 4
  ISIG = 0
  NCR = MAX0(ID, 0)
  NCR = MIN0(NCR, 4)
  XMAX = VR(1)
  XMIN = VR(2)
  YMAX = VR(3)
  YMIN = VR(4)
  GO TO 430
510 DO 520 K = 1, NS
  NR = NR + 1
520 READ (10, 830) SNAME(K), (STN(K, L), L = 1, 8), X(K), Y(K), Z(K)
  READ (10, 1100) SKIP
  NR = NR + 1
  GO TO 430
530 DO 540 L = 1, NB
  NR = NR + 1
540 READ (10, 840) XB(L), YB(L)
  READ (10, 1100) SKIP
  NR = NR + 1
  GO TO 430
550 AMU = 0.0
  STD = 0.0
  STE = 0.0
  DO 560 K = 1, NS
  NR = NR + 1
  READ (10, 850) H(K)
  AMU = AMU + H(K)
560 STD = STD + H(K) * H(K)
  STD = SQRT((STD - AMU * AMU/NS)/(NS - 1.))
  AMU = AMU/NS
  READ (10, 1100) SKIP

```

```
NR = NR + 1
GO TO 430
570 NR = NR + 1
  READ (10,860) ID,VR,BL
  IF (BL(1).NE.BLANK) AMU = VR(1)
  IF (BL(2).NE.BLANK) STD = VR(2)
  IF (BL(3).NE.BLANK) STE = VR(3)
  GO TO 430
580 DO 590 K = 1,NS
      NR = NR + 1
590 READ (10,850) SMU(K)
  READ (10,1100) SKIP
  NR = NR + 1
  GO TO 430
600 ITP = 7
  CALL THIESN
  KTP = ID
  GO TO 430
610 ITP = 8
  CALL HARMON
  KTP = ID
  GO TO 430
620 ITP = 9
  CALL OPTIM
  KTP = ID
  GO TO 430
630 ITP = 10
  CALL POLYN
  KTP = ID
  GO TO 430
640 ITP = 11
  CALL INDISM
  KTP = ID
  GO TO 430
650 ITP = 12
  CALL QUAD
  KTP = ID
  GO TO 430
660 ITP = 13
  CALL KRIGING
  KTP = ID
  GO TO 430
670 CALL DEPTH
  GO TO 430
675 NRS=0
676 NR=NR+1
  NRS=NRS+1
  READ (10,815) SDNAME(NRS), XRS(NRS), YRS(NRS)
  IF(SDNAME(NRS).NE.ENDER) GO TO 676
  NRS=NRS-1
  GO TO 430
680 IF (ITP.NE.0) GO TO 690
  WRITE (6,870)
  GO TO 430
690 IF (ITP.NE.4) GO TO 700
  GO TO 460
```

```

700 CALL TAPERWR (PROC(IITP),STORE,0)
WRITE (6,880) PROC(IITP),KTP
NIP = IITP
IITP = 0
IF (KCOR.EQ.0) GO TO 430
IF (NIP.LE.8) GO TO 430
CALL TAPERWR (PROC(IITP),STORE,1)
GO TO 430
730 IF (IITP.EQ.8) GO TO 430
WRITE (6,1090) PROC(IITP),TITLE
CALL MASK
GO TO 430
740 READ (10,860) ID,VR,BL
RHO = 0.0
IF (BL(1).NE.BLANK) RAD = VR(1)
IF (BL(2).NE.BLANK) RHO = VR(2)
KCOR = MAX0(ID, - 1)
KCOR = MIN0(KCOR,2)
RC = RAD
IF (RHO * (RHO - 1.).GE.0.) GO TO 760
IF (KCOR.EQ.2) GO TO 750
RC = - RAD/ALOG(RHO)
GO TO 760
750 RC = RHO * RAD/(1.0 - RHO)
760 CALL SPACOR
GO TO 430
770 READ (10,810) LIST, ID, VR
CALL VARIO
GO TO 430
780 CALL PTDATA (ID)
GO TO 430
790 IF (LIST.EQ.PROC(14)) GO TO 100
STOP

```

C
C
C

```

800 FORMAT (1H1,/T18,1H+,68(1H-),1H+/,T18,1H!,T20,68H  HYDROLOGIC MAP
LPING, INTERPOLATION AND AREAL AVERAGING SYSTEM !,/T18,1H+,68(1H
2-),1H+)
810 FORMAT (A8,2X,I5,4X,5F10.0)
815 FORMAT (A8,11X,2F10.0)
820 FORMAT (10A8)
830 FORMAT (A8,T1,8A1,T20,3F10.0)
840 FORMAT (T20,2F10.0)
850 FORMAT (T20,F10.0)
860 FORMAT (T14,I2,T20,5F10.0,T20,3(2X,A8))
870 FORMAT (1H1,/T20,17HDATA SET IS EMPTY)
880 FORMAT (1H-,T20,26HMAPPED VALUES ON TAPE - ,A8,6HCODE =,I3)
890 FORMAT (1H1,/T20,A8,I3,3X,10A8/T22,2I6,2F15.5)
900 FORMAT (10F10.3)
910 FORMAT (8HSTDERROR,2X,10A8)
920 FORMAT (10I10)
930 FORMAT (10A8)
940 FORMAT (A8,T1,10A8)
950 FORMAT (/T5,1HS,F4.3,T25,10A8/T5,49H***** INVALID STATEMENT/PROCEED
LURE NOT FOUND *****,/)

```

```

960 FORMAT (/T5,1HS,F4.3,T19,6HINPUT ,10A8)
970 FORMAT (T5,1HS,F4.3,T25,10A8)
980 FORMAT (/T5,81H***** STATION INPUT DATA IS GREATER OR LESS THAN T
THE NUMBER OF STATION(S) *****)
990 FORMAT (/T5,81H***** STATION INPUT MEANS IS GREATER OR LESS THAN T
THE NUMBER OF STATION(S) *****)
1000 FORMAT (/T5,58H***** INPUT STREAM IS FLUSH UP TO STOP/END OF FILE
1 *****)
1010 FORMAT (/T5,1HS,F4.3,T20,5HEROC ,10A8)
1020 FORMAT (T5,1HS,F4.3,T20,25HTAPE PRECEEDING PROCEDURE)
1030 FORMAT (/T5,1HS,F4.3,T20,4HSTOP)
1040 FORMAT (//T5,30HNOTE: INPUT REQUIREMENTS FOR ,A8,17H IS NOT SATIS
IFIED)
1050 FORMAT (/T5,63HNOTE: INPUT REQUIREMENTS FOR KRIGING OPTION 1 IS N
OT SATISFIED)
1060 FORMAT (/T5,35H***** NO FURTHER PROCESSING *****)
1070 FORMAT (/T5,13H***** END OF ,A8,1X,16HNOT FOUND *****)
1080 FORMAT (1H+,T25,16H* STMT GENERATED)
1090 FORMAT (1H1,T20,16HMASKED VALUES - ,A8,3H : ,10A8/)
1100 FORMAT (A8)
1110 FORMAT (T14,I2)
END
SUBROUTINE CHECK (RTXY)

```

C
C
C

DIAGNOSTIC CHECK OF COORDINATE SYSTEM

```

CHARACTER*8 TITLE, SNAME
CHARACTER*1 STN
COMMON /SYSID/          STORE(400,101)
COMMON /SYSID2/        TITLE(10), SNAME(50)
COMMON /CONST/         IDSC      , NS      , NX      ,
1      NY      , DX      , DY      , XMAX     ,
2      XMIN     , YMAX     , YMIN     , HMAX     ,
3      HMIN     , SMAX     , SMIN     ,
COMMON /SYSDAT/        X(50)     , Y(50)     , H(50)     ,
1      Z(50)
COMMON /SYSTAT/        AMJ      , STD      , STE      ,
1      RC      , ATC      , CRSCO(50) , SPAC(50,50)
COMMON /SBOUND/        NB      , XB(100)  , YB(100)
COMMON /SBOUN2/        STN(50,8)
COMMON /AREAL/         AREA     , TAR
COMMON /CODE/          ID      , VR(5)   , KCOR     ,
1      KERR
COMMON /CORP/          NCR      , FA      , FB      ,
1      AXM      , AYL      , XM
BIG = 9999999.0
TAR = 0.0
FA = 3.141592654/180.
FB = 111.
WRITE (6,230) TITLE
IF (NCR.EQ.0) GO TO 100
AXL = ANGLE(XMIN,1)
AXH = ANGLE(XMAX,1)
AXM = (AXL + AXH)/2.
AYL = ANGLE(YMIN,2)
AYH = ANGLE(YMAX,2)

```

```

XL = XMIN
XH = XMAX
YL = YMIN
YH = YMAX
CL = COS(FA * AYL)
CH = COS(FA * AYH)
XMAX = FB * (AXH - AXL) * AMAX1(CL, CH)
XMIN = 0.0
XM = XMAX/2.
YMAX = FB * (AYH - AYL)
YMIN = 0.
WRITE (6,240) XL, YL, XMIN, YMIN, XH, YH, XMAX, YMAX
GO TO 110
100 WRITE (6,250) XMIN, YMIN, XMAX, YMAX
110 NX = 101
DX = (XMAX - XMIN)/100.0
DY = DX * 10./RTXY
NY = IFIX((YMAX - YMIN)/DY) + 1
IF (NS.LE.0) GO TO 220
DO 150 K = 1, NS
  IF (NCR.EQ.0) GO TO 120
  AX = ANGLE(X(K), 1)
  AY = ANGLE(Y(K), 2)
  XT = X(K)
  YT = Y(K)
  X(K) = XM + FB * COS(FA * AY) * (AX - AXM)
  Y(K) = FB * (AY - AYL)
  WRITE (6,260) K, SNAME(K), XT, YT, X(K), Y(K), Z(K)
  GO TO 130
120 WRITE (6,270) K, SNAME(K), X(K), Y(K), Z(K)
130 IF ((X(K) - XMIN) * (X(K) - XMAX).GT.0.) GO TO 140
  IF ((Y(K) - YMIN) * (Y(K) - YMAX).GT.0.0) GO TO 140
  WRITE (6,280)
  GO TO 150
140 WRITE (6,290)
150 CONTINUE
KERR = 0
IF (NB.GT.0) GO TO 170
NB = 4
TAR = BIG
XB(1) = XMIN
XB(2) = XMAX
XB(3) = XMAX
XB(4) = XMIN
YB(1) = YMIN
YB(2) = YMIN
YB(3) = YMAX
YB(4) = YMAX
DO 160 L = 1, NB
160 WRITE (6,300) L, XB(L), YB(L)
RETURN
170 DO 210 L = 1, NB
  IF (NCR.EQ.0) GO TO 180
  AX = ANGLE(XB(L), 1)
  AY = ANGLE(YB(L), 2)
  XT = XB(L)

```

```

YT = YB(L)
XB(L) = XM + FB * COS(FA * AY) * (AX - AXM)
YB(L) = FB * (AY - AYL)
WRITE (6,310) L,XT,YT,XB(L),YB(L)
GO TO 190
180 WRITE (6,320) L,XB(L),YB(L)
190 IF ((XB(L) - XMIN) * (XB(L) - XMAX).GT.0.) GO TO 200
IF ((YB(L) - YMIN) * (YB(L) - YMAX).GT.0.) GO TO 200
WRITE (6,280)
GO TO 210
200 WRITE (6,290)
KERR = 1
210 CONTINUE
IF (KERR.EQ.1) WRITE (6,330)
RETURN
220 KERR = 1
WRITE (6,340)
RETURN

```

C
C
C

```

230 FORMAT (1H1,/,T12,40HDIAGNOSTIC CHECK OF COORDINATE SYSTEM : ,10A8/
1/T38,9HSFERICAL,T68,11HRECTANGULAR,/,T10,14HCONTROL POINTS,T30,11
2HDEG.MIN.SEC,T45,11HDEG.MIN.SEC,T60,12HX-COORDINATE,T75,12HY-COORD
3INATE,T90,11HZ-ELEVATION,T105,11HDIAGNOSTICS,/,T31,9HLONGITUDE,T47,
48HLATITUDE,/)
240 FORMAT (T10,7HMINIMUM,T30,F10.0,T45,F10.0,T60,F10.3,T75,F10.3/T10,
17HMAXIMUM,T30,F10.0,T45,F10.0,T60,F10.3,T75,F10.3/)
250 FORMAT (T10,7HMINIMUM,T60,F10.3,T75,F10.3/T10,7HMAXIMUM,T60,F10.3,
1T75,F10.3/)
260 FORMAT (T10,3HSTA,I4,2X,A8,T30,F10.0,T45,F10.0,T60,F10.3,T75,F10.3
1,T90,F10.3)
270 FORMAT (T10,3HSTA,I4,2X,A8,T60,F10.3,T75,F10.3,T90,F10.3)
280 FORMAT (1H+,T105,15HINSIDE OF RANGE)
290 FORMAT (1H+,T103,18H*OUTSIDE OF RANGE*)
300 FORMAT (T10,22HDEFAULT BOUNDARY POINT,I3,T60,F10.3,T75,F10.3,T105,
115HINSIDE OF RANGE)
310 FORMAT (T10,14HBOUNDARY POINT,I3,T30,F10.0,T45,F10.0,T60,F10.3,T75
1,F10.3)
320 FORMAT (T10,14HBOUNDARY POINT,I3,T60,F10.3,T75,F10.3)
330 FORMAT (/T40,55HNO FURTHER PROCESSING - BOUNDARY POINT OUTSIDE OF
1RANGE,/)
340 FORMAT (T105,17HNO INPUT STATIONS,/)
END
FUNCTION ANGLE(A,KL)

```

C
C
C
C
C
C
C
C
C

```

CONVERSION FROM DEGREE-MINUTES-SECONDS TO DEGREE DECIMALS
OR FROM DEGREE DECIMALS DO DEGREE-MINUTES-SECONDS
NCR=1, DEG-MIN-SEC TO DEG DECIMALS (EAST)
NCR=1, DEG-MIN-SEC TO DEG DECIMALS (WEST)
NCR=3, DEG DECIMALS TO DEG-MIN-SEC (EAST)
NCR=4, DEG DECIMALS TO DEG-MIN-SEC (WEST)

```

```

COMMON /CORP/          NCR          , FA          , FB          ,
1 AXM          , AYL          , XM
GO TO (100,100,110,110) , NCR

```

```

100 IA = IFIX(A)
    IDEG = IA/10000
    IMIN = IA/100 - 100 * IDEG
    ISEC = IA - 10000 * IDEG - 100 * IMIN
    ANG = FLOAT(IDEG) + FLOAT(IMIN)/60. + FLOAT(ISEC)/3600.
    IF (NCR.EQ.1) GO TO 120
    IF (KL.EQ.2) GO TO 120
    ANG = 180.0 - ANG
    GO TO 120
110 ANG = A
    IA = A
    DEG = IA
    A = (A - DEG) * 60.0
    IA = A
    AMIN = IA
    A = (A - AMIN) * 60.0
    IA = A
    ASEC = IA
    A = DEG * 10000.0 + AMIN * 100.0 + ASEC
    IF (NCR.EQ.3) GO TO 120
    IF (KL.EQ.2) GO TO 120
    ANG = 180.0 - ANG
120 ANGLE = ANG
    RETURN
    END
SUBROUTINE BOUND (INAME)

```

C
C
C

DETERMINATION OF BASIN BOUNDARY POINTS

```

CHARACTER*8 TITLE, SNAME
CHARACTER*1 STN, PLOT(200,101), BLK, PLUS, PERD
COMMON /SYSID/          STORE(400,101)
COMMON /SYSID2/        TITLE(10), SNAME(50)
COMMON /CONST/        IDSC          , NS          , NX
1      NY          , DX          , DY          , XMAX
2      XMIN        , YMAX        , YMIN        , HMAX
3      HMIN        , SMAX        , SMIN
COMMON /SYSDAT/        X(50)        , Y(50)        , H(50)
1      Z(50)
COMMON /SBOUND/        NB          , XB(100)        , YB(100)
COMMON /SBOUND2/       STN(50,8)
COMMON /AREAL/         AREA        , TAR
DATA BLK, PLUS, PERD/ ' ', '+', '.' /
BIG = 9999999.0
IDR = 0
IF (TAR.LT.BIG) GO TO 110
DO 100 I = 1, NY
DO 100 J = 1, NX
100 STORE(I, J) = 1.0
GO TO 240
110 DO 120 I = 1, NY
DO 120 J = 1, NX
120 STORE(I, J) = 0.0
DO 230 L = 1, NB
    XL = XB(L)
    YL = YB(L)

```

```

LB = L + 1
IF (LB.LE.NB) GO TO 130
LB = 1
130 X2 = XB(LB)
Y2 = YB(LB)
IX1 = IFIX((X1 - XMIN)/DX) + 1
IX2 = IFIX((X2 - XMIN)/DX) + 1
IF (IX2 - IX1) 140,170,200
140 IX2 = IX2 + 1
DO 160 IX = IX2,IX1
    XT = XMIN + DX * (IX - 1.0)
    YT = Y2 + (Y1 - Y2) * (XT - X2)/(X1 - X2)
    IYM = IFIX((YT - YMIN)/DY) + 1
    DO 150 IY = 1,IYM
150     STORE(IY,IX) = STORE(IY,IX) + 1.0
160     CONTINUE
    GO TO 230
170 IF (Y2.GE.Y1) GO TO 180
    TP = Y1
    Y1 = Y2
    Y2 = TP
180 IY1 = IFIX((Y1 - YMIN)/DY) + 1
    IY2 = IFIX((Y2 - YMIN)/DY) + 1
    DO 190 IY = IY1,IY2
190     STORE(IY,IX1) = STORE(IY,IX1) + 1.0
    GO TO 230
200 IX1 = IX1 + 1
    DO 220 IX = IX1,IX2
        XT = XMIN + DX * (IX - 1.0)
        YT = Y1 + (Y2 - Y1) * (XT - X1)/(X2 - X1)
        IYM = IFIX((YT - YMIN)/DY)
        IF (IYM.LE.0) GO TO 220
        DO 210 IY = 1,IYM
210         STORE(IY,IX) = STORE(IY,IX) - 1.0
220     CONTINUE
230 CONTINUE
240 TAR = 0.0
    DO 270 IY = 1,NY
        IDS3 = IY + IDR
        DO 260 IX = 1,NX
            IF (STORE(IY,IX).LE.0.0) GO TO 250
            STORE(IY,IX) = 0.0
            TAR = TAR + 1.0
            PLOTG(IDS3,IX) = BLK
            GO TO 260
250         PLOTG(IDS3,IX) = PERD
            STORE(IY,IX) = - BIG
260     CONTINUE
270 CONTINUE
AREA = TAR * DX * DY
DO 340 K = 1,NS
    IF ((X(K) - XMIN) * (X(K) - XMAX).GT.0.0) GO TO 340
    IF ((Y(K) - YMIN) * (Y(K) - YMAX).GT.0.0) GO TO 340
    IX = IFIX((X(K) - XMIN)/DX) + 1
    IY = IFIX((Y(K) - YMIN)/DY) + 1
    IY = IY + IDR

```



```

PLOTG(IY,IX) = PLUS
IF (INAME.EQ.1) GO TO 340
IY = IY - IDR
IXP = IX - 3
280 IF (IXP.GT.0) GO TO 290
IXP = IXP + 1
GO TO 280
290 IT = IX + 4
300 IF (IT.LE.NX) GO TO 310
IT = IT - 1
IXP = IXP - 1
GO TO 300
310 IYP = IY + 1
IF (IYP.LE.NY) GO TO 320
IYP = IY - 1
320 IYP = IYP + IDR
KX = 0
DO 330 IX = IXP,IT
KX = KX + 1
330 PLOTG(IYP,IX) = SIN(K,KX)
340 CONTINUE
WRITE (6,350) TITLE
CALL GRAPH (PLOTG)
UA = DX * DY
WRITE (6,360) AREA,UA,DX,DY
RETURN

```

C
C
C

```

350 FORMAT (1HL,T22,31HDELINEATION OF BASIN BOUNDARY: ,10A8//)
360 FORMAT (// T32,12HTOTAL AREA =,F10.3/T32,11HUNIT AREA =,F10.3/T3
12,9HDELTA X =,F10.3/T32,9HDELTA Y =,F10.3//T32,7HLEGEND: ,/T33,21HI
2.I OUTSIDE OF BASIN,/T33,20HI I INSIDE OF BASIN,/T33,13HI+I ST
3ATION)
END
SUBROUTINE GRAPH (PLOTG)

```

C
C
C

```

MAPPING OF CHARACTER INFORMATION

CHARACTER*8 TITLE, SNAME
CHARACTER*1 PLOTG(200,101)
DIMENSION XAX(11)
COMMON /SYSID/ STORE(400,101)
COMMON /SYSID2/ TITLE(10), SNAME(50)
COMMON /CONST/ IDSC , NS , NX
1 NY , DX , DY , XMAX ,
2 XMIN , YMAX , YMIN , HMAX ,
3 HMIN , SMAX , SMIN

IDR = 0
WRITE (6,140)
AX = DX * 10
XAX(1) = XMIN
DO 100 K = 2,11
100 XAX(K) = XAX(K - 1) + AX
DO 110 II = 1,NY
110 I = NY + 1 - II

```

```

      YAX = YMIN + DY * (I - 1.0)
      IDS = I + IDR
110 WRITE (6,120) YAX, (PLOTG(IDS,J), J = 1, NX)
      WRITE (6,140)
      WRITE (6,130) XAX
      RETURN

```

C
C
C

```

120 FORMAT (5X,F8.2,1X,1HI,101A1,1HI)
130 FORMAT (9X,3HY/X,F6.2,10(3X,F7.2))
140 FORMAT (14X,2HI+,10(9(1H-),1H+),1HI)
      END
      SUBROUTINE MAP. (KDSK)

```

C
C
C

MAPPING OF CONTOURS

```

CHARACTER*8 TITLE, SNAME
CHARACTER*1 PLOTG(200,101), PT, BLK, PERD, PLUS
COMMON /SYSID/          STORE(400,101)
COMMON /SYSID2/        TITLE(10), SNAME(50)
DIMENSION              HG(21), PT(21)
COMMON /CONST/        IDSC, NS, NX,
1 NY, DX, DY, XMAX,
2 XMIN, YMAX, YMIN, HMAX,
3 HMIN, SMAX, SMIN
COMMON /SYSDAT/       X(50), Y(50), H(50)
1 Z(50)
COMMON /CODE/        ID, VR(5), KCOR
1 KERR
DATA BLK, PERD, PLUS/ ' ', '.', '+'/
DATA PT/'0','1','2','3','4','5','6','7','8','9','A','B','C','D',
1 'E','F','G','H','I','J','K'/
BIG = 9999999.0
KD = IDSC * (KDSK - 1)
IDR = 0
HMX = HMAX
HMN = HMIN
IF (KDSK.EQ.1) GO TO 100
HMX = SMAX
HMN = SMIN
100 RH = ABS(HMX - HMN)
IF (RH.EQ.0.0) RETURN
IF (ID.NE. - 99) GO TO 110
DH = RH/20.0
TOL = DH * VR(1)
GO TO 150
110 BS = 0.001
120 DO 130 K = 1,9
      DH = BS * K
      IF (RH/DH.LE.18.0) GO TO 140
130 CONTINUE
      BS = BS * 10.0
      GO TO 120
140 NL = IFIX(HMN/DH) - 1
      HMN = NL * DH

```

```

TOL = DH * 0.25
150 DO 160 K = 1,21
    HG(K) = HMN + DH * (K - 1.0)
    HH = HG(K) * 10000.
    HT = HH - IFIX(HH)
    IF (HT.LT.6.) HG(K) = HG(K) + .0003
160 CONTINUE
    DO 230 I = 1,NY
        KDD = KD + I
        IDS3 = IDR + I
        DO 220 J = 1,NX
            PL = STORE(KDD,J)
            PLOTG(IDS3,J) = BLK
            IF (STORE(I,J).EQ.-BIG) GO TO 220
            K = IFIX((PL - HMN)/DH) + 1
            IF (K - 1) 170,190,180
170         PLOTG(IDS3,J) = PT(1)
            GO TO 220
180         IF (K.LT.21) GO TO 190
            PLOTG(IDS3,J) = PT(21)
            GO TO 220
190         IF ((PL - HG(K) + TOL) * (PL - HG(K) - TOL).LE.0.0) GO TO 20
1          0
            K = K + 1
            IF ((PL - HG(K) + TOL) * (PL - HG(K) - TOL).GT.0.0) GO TO 21
1          0
200         PLOTG(IDS3,J) = PT(K)
            GO TO 220
210         PLOTG(IDS3,J) = PERD
220         CONTINUE
230 CONTINUE
    IF (NS.EQ.0) GO TO 250
    DO 240 K = 1,NS
        IF ((X(K) - XMIN) * (X(K) - XMAX).GT.0.0) GO TO 240
        IF ((Y(K) - YMIN) * (Y(K) - YMAX).GT.0.0) GO TO 240
        J = IFIX((X(K) - XMIN)/DX) + 1
        I = IFIX((Y(K) - YMIN)/DY) + 1
        IDS3 = I + IDR
        PLOTG(IDS3,J) = PLUS
240 CONTINUE
250 CONTINUE
    CALL GRAPH (PLOTG)
    WRITE (6,270)
    DO 260 K = 1,21
260 WRITE (6,280) PT(K),HG(K),TOL
    RETURN
C
C
C
270 FORMAT (// T27,8HLEGEND : ,//T28,7HSYMBOLS,T40,8HCONTOURS,T54,9HTO
    LLERANCE,/ )
280 FORMAT (T31,A1,T37,F10.3,T53,3H+/-,T58,F6.3)
    END
    SUBROUTINE GEPCON (AI,N)
C
C    MATRIX INVERSION USING GAUSSIAN ELIMINATION

```

C WITH PIVOTAL CONDENSATION

C

COMMON /CODE/ ID , VR(5) , KCOR
 1 KERR
 DIMENSION AI(52,52) , A(52,104) , W(52)

C

```

KERR = 0
NM = N - 1
IM = N + 1
NDL = N + N
DO 110 I = 1,N
  DO 100 J = 1,N
    A(I,J) = AI(I,J)
    M = N + J
100  A(I,M) = 0.0
    M = N + I
110  A(I,M) = 1.0
    DO 170 I = 1,NM
      MX = I
      AMX = ABS(A(I,I))
      IN = I + 1
      DO 120 K = IN,N
        IF (ABS(A(K,I)).LE.AMX) GO TO 120
        MX = K
        AMX = ABS(A(K,I))
120  CONTINUE
      IF (MX.EQ.I) GO TO 140
      DO 130 L = I,NDL
        TT = A(I,L)
        A(I,L) = A(MX,L)
130  A(MX,L) = TT
140  DV = A(I,I)
      IF (ABS(DV).LE.0.00000001) GO TO 220
      DO 160 J = IN,N
        IF (A(J,I).EQ.0.0) GO TO 160
        CN = - A(J,I)/DV
        DO 150 L = I,NDL
150  A(J,L) = A(J,L) + CN * A(I,L)
160  CONTINUE
170  CONTINUE
      IK = 0
      IF (ABS(A(N,N)).LE.0.00000001) GO TO 220
      DO 210 II = IM,NDL
        W(N) = A(N,II)/A(N,N)
        DO 190 I = 1,NM
          K = N - I
          JK = K + 1
          WW = A(K,II)
          DO 180 J = JK,N
180  WW = WW - A(K,J) * W(J)
          IF (ABS(A(K,K)).LE.0.00000001) GO TO 220
190  W(K) = WW/A(K,K)
          IK = IK + 1
          DO 200 L = 1,N
200  AI(L,IK) = W(L)
210  CONTINUE

```

```

RETURN
220 KERR = 1
WRITE (6,230)
RETURN
C
C
230 FORMAT (/5X, 44HPROCESSING STOPPED - MATRIX DETERMINANT ZERO,/5X,
120HIN SUBROUTINE GEPCON,/)
END
SUBROUTINE THIESN
C
C
C
THIESSEN METHOD AND AREAL AVERAGING

CHARACTER*8 TITLE, SNAME
CHARACTER*1 PLOT(200,101), PT, BLK, PLUS
DIMENSION DST(50) , W(50) , ISM(50) ,
1 PT(30)
COMMON /SYSID/ STORE(400,101)
COMMON /SYSID2/ TITLE(10) , SNAME(50)
COMMON /CONST/ IDSC , NS , NX ,
1 NY , DX , DY , XMAX ,
2 XMIN , YMAX , YMIN , HMAX ,
3 HMIN , SMAX , SMIN ,
COMMON /SYSDAT/ X(50) , Y(50) , H(50) ,
1 Z(50)
COMMON /SYSTAT/ AMU , STD , STE ,
1 RC , ATC , CRSCO(50) , SPAC(50,50) ,
COMMON /AREAL/ AREA , TAR ,
COMMON /CODE/ KOD , VR(50) , KCOR ,
1 KERR
DATA BLK, PLUS/' ', '+'/
DATA PT/'0','1','2','3','4','5','6','7','8','9','A','B','C','D',
1 'E','F','G','H','I','J','K','L','M','N','O','P','Q','R','S','T'/
WRITE (6,380) TITLE
BIG = 9999999.0
KOD1 = MAX0(KOD,1)
KOD1 = MIN0(KOD1,6)
IF (KOD1.NE.5) GO TO 100
IF (KCOR.NE.0) GO TO 100
WRITE (6,490)
RETURN
100 HM = 0.0
ATM = 0.0
DO 110 K = 1,NS
ATM = ATM + H(K)
110 W(K) = 0.0
TOL = SQRT(DX * DX + DY * DY)
IF (KOD1.EQ.5) TOL = ABS(ATM/NS) * TOL
IF (KOD1.EQ.6) TOL = STD * TOL/RC
DO 310 I = 1,NY
YT = YMIN + DY * (I - 1.0)
DO 300 J = 1,NX
IF (STORE(I,J).EQ.-BIG) GO TO 290
XT = XMIN + DX * (J - 1.0)
DO 120 K = 1,NS
ISM(K) = K

```

```

120     DST(K) = BIG
      DO 220 K = 1, NS
130         GO TO (130,140,150,160,170,190), KOD1
          DST(K) = SQRT((XT - X(K)) * * 2 + (YT - Y(K)) * * 2)
          GO TO 220
140         DST(K) = (XT - X(K)) * * 2 + (YT - Y(K)) * * 2
          GO TO 220
150         DST(K) = ABS(XT - X(K)) + ABS(YT - Y(K))
          GO TO 220
160         DST(K) = AMAX1(ABS(XT - X(K)), ABS(YT - Y(K)))
          GO TO 220
170         DIS = SQRT((XT - X(K)) * * 2 + (YT - Y(K)) * * 2)
          IF (DIS.GT.0.0) GO TO 180
          DST(K) = 0.0
          GO TO 220
180         IF (H(K).EQ.0.0) GO TO 220
          DST(K) = ABS(H(K)) * DIS
          GO TO 220
190         DIS = SQRT((XT - X(K)) * (XT - X(K)) + (YT - Y(K)) * (YT
1          - Y(K)))
          IF (KCOR.EQ.2) GO TO 200
          RHO = EXP(- DIS/RC)
          GO TO 210
200         RHO = 1.0/(1.0 + DIS/RC)
210         DST(K) = - ABS(RHO * (H(K) - AMU))
220     CONTINUE
      IF (NS.GT.2) GO TO 230
      K = 1
      IF (NS.EQ.1) GO TO 270
      IF (DST(1).GT.DST(2)) GO TO 260
      K = 2
      GO TO 260
230     DO 250 K = 1, 2
          IS = K + 1
          DO 240 L = IS, NS
              IF (DST(K).LE.DST(L)) GO TO 240
              TMP = DST(K)
              DST(K) = DST(L)
              DST(L) = TMP
              IT = ISM(K)
              ISM(K) = ISM(L)
              ISM(L) = IT
240     CONTINUE
250     CONTINUE
      K = ISM(1)
260     IF (ABS(DST(2) - DST(1)).GT.TOL) GO TO 270
          PLOT(I,J) = BLK
          GO TO 280
270     PLOT(I,J) = PT(K)
280     W(K) = W(K) + 1.0
          STORE(I,J) = H(K)
          HM = HM + H(K)
          GO TO 300
290     PLOT(I,J) = BLK
300     CONTINUE
310     CONTINUE

```

```

DO 320 K = 1, NS
  IF ((X(K) - XMIN) * (X(K) - XMAX).GT.0.0) GO TO 320
  IF ((Y(K) - YMIN) * (Y(K) - YMAX).GT.0.0) GO TO 320
  J = 1 + IFIX((X(K) - XMIN)/DX)
  I = 1 + IFIX((Y(K) - YMIN)/DY)
  I = I + IDR
  PLOT(I, J) = PLUS
320 CONTINUE
  CALL GRAPH (PLOT)
  WRITE (6, 390)
  DO 330 K = 1, NS
    W(K) = W(K)/TAR
    ART = W(K) * AREA
330 WRITE (6, 400) K, SNAME(K), PT(K), W(K), ART, H(K)
    HM = HM/TAR
    IF (KCOR.NE.0) GO TO 340
    WRITE (6, 480) HM
    GO TO 370
340 SD2 = STD * STD
    SE2 = STE * STE
    SE = AIC * SD2
    DO 360 K = 1, NS
      WT = W(K)
      SE = SE + WT * WT * SE2 - 2. * WT * CRSCO(K) * SD2
    DO 350 L = 1, NS
350 SE = SE + WT * W(L) * SD2
360 CONTINUE
    IF (SE.LE.0.0) SE = 0.0
    SE = SQRT(SE)
    WRITE (6, 410) HM, SE
370 WRITE (6, 500) KOD1
    RETURN

```

C

```

380 FORMAT (1H1, T20, 28HMAPPING BY THIESSEN POLYGON:, 10A8//)
390 FORMAT (// T20, 7HSTA.NO., 6X, 9HSTA. NAME, 6X, 6HSYMBOL, 4X, 15HTHIESSE
  IN WEIGHT, 5X, 12HPOLYGON AREA, T96, 9HSTA. DATA,/)
400 FORMAT (1H , T21, I3, 10X, A8, 8X, A1, 7X, F10.5, 8X, F10.2, T95, F9.3)
410 FORMAT (// T40, 23HAREAL AVERAGE OF DATA =, F16.4/T40, 24HERROR OF A
  LREAL AVERAGE =, F15.4/)
480 FORMAT (// T40, 23HAREAL AVERAGE OF DATA =, F16.4/)
490 FORMAT (1H1, //T40, 41HTHIESSEN METHOD - OPTION 5 NOT APPLICABLE, //T
  140, 42HSPATIAL CORRELATION FUNCTION NOT SPECIFIED)
500 FORMAT (/20X, 'THIESSEN METHOD OPTION CODE =', I3//
  1 23X, '1 MIN STRAIGHT DISTANCE'//
  2 23X, '2 MIN SQUARE DISTANCE'//
  3 23X, '3 MIN ORTHOGONAL DISTANCE'//
  4 23X, '4 MIN MAX LEG DISTANCE'//
  5 23X, '5 MIN ABS (DATA) X STRAIGHT DISTANCE'//
  6 23X, '6 MAX ABS (DEVIATE) X SPATIAL CORRELATION COEFF'//
  END
  SUBROUTINE QUAD

```

C

C

C

MULTIQUADRIC INTERPOLATION

CHARACTER*8 TITLE, SNAME

DIMENSION

C(50)

, A(52,52)

, RAD(50)

```

1      COR(50)      , W(50)
COMMON /SYSID/      STORE(400,101)
COMMON /SYSID2/     TITLE(10) , SNAME(50)
COMMON /CONST/     IDSC      , NS      , NX      ,
1      NY          , DX      , DY      , XMAX     ,
2      XMIN        , YMAX     , YMIN     , HMAX     ,
3      HMIN        , SMAX     , SMIN     ,
COMMON /SYSDAT/    X(50)     , Y(50)     , H(50)     ,
1      DAS(50)
COMMON /SYSTAT/    AMU      , STD      , STE      ,
1      RC          , ATC      , CRSCO(50) , SPAC(50,50) ,
COMMON /AREAL/    AREA     , TAR      ,
COMMON /CODE/     KOD      , VR(5)    , KCOR     ,
1      KERR
BIG = 9999999.0
SD2 = STD * STD
SE2 = STE * STE
KD = MAX0(1,KOD)
KD = MIN0(KD,4)
WRITE (6,370) TITLE
A(1,1) = 0.
DO 160 K = 2,NS
  X1 = X(K)
  Y1 = Y(K)
  KL = K - 1
  A(K,K) = 0.
  DO 150 J = 1,KL
    X2 = X(J)
    Y2 = Y(J)
    DSX = ABS(X1 - X2)
    DSY = ABS(Y1 - Y2)
    GO TO (100,110,120,130) , KD
100   DST = SQRT(DSX * DSX + DSY * DSY)
    GO TO 140
110   DST = DSX * DSX + DSY * DSY
    GO TO 140
120   DST = DSX + DSY
    GO TO 140
130   DST = AMAX1(DSX,DSY)
140   A(K,J) = DST
    A(J,K) = DST
150   CONTINUE
160   CONTINUE
    IF (KD.NE.2) GO TO 190
    DO 180 K = 1,NS
      IF (DAS(K).LE.0.0) GO TO 180
      RG2 = DAS(K)/6.2832
      DO 170 J = 1,NS
170   A(K,J) = A(K,J) + RG2
180   CONTINUE
190   NSH = NS/2
    DO 210 K = 1,NSH
      IS = NS - K + 1
      DO 200 J = 1,NS
        TEMP = A(K,J)
        A(K,J) = A(IS,J)

```



```

200   A(IS,J) = TEMP
210   CONTINUE
      CALL GEPCON (A,NS)
      HM = 0.0
      IF (KERR.EQ.1) WRITE (6,380) KD,HM
      IF (KERR.EQ.1) RETURN
      DO 230 J = 1,NS
          SUM = 0.
          DO 220 K = 1,NS
              IS = NS + 1 - K
220     SUM = SUM + A(J,K) * H(IS)
230   C(J) = SUM
      HMAX = - BIG
      HMIN = BIG
      SMAX = - BIG
      SMIN = BIG
      DO 360 I = 1,NY
          YT = YMIN + DY * (I - 1.0)
          IDS2 = I + IDSC
          DO 350 J = 1,NX
              IF (STORE(I,J).EQ.-BIG) GO TO 350
              XT = XMIN + DX * (J - 1.0)
              HH = 0.
              SW = 0.
              DO 280 K = 1,NS
                  DSX = ABS(XT - X(K))
                  DSY = ABS(YT - Y(K))
                  GO TO (240,250,260,270), KD
240     RAD(K) = SQRT(DSX * DSX + DSY * DSY)
                  GO TO 280
250     RAD(K) = DSX * DSX + DSY * DSY
                  GO TO 280
260     RAD(K) = DSX + DSY
                  GO TO 280
270     RAD(K) = AMAX1(DSX,DSY)
280     CONTINUE
          DO 300 K = 1,NS
              SUM = 0.
              IS = NS + 1 - K
              DO 290 N = 1,NS
290     SUM = SUM + RAD(N) * A(N,IS)
              W(K) = SUM
              SW = SW + SUM
300     HH = HH + C(K) * RAD(K)
          STORE(I,J) = HH
          HMAX = AMAX1(HMAX,HH)
          HMIN = AMIN1(HMIN,HH)
          HM = HM + HH
          IF (KCOR,EQ.0) GO TO 350
          DO 320 K = 1,NS
              DS = SQRT((XT - X(K)) * * 2 + (YT - Y(K)) * * 2)
              IF (KCOR.EQ.2) GO TO 310
              COR(K) = EXP(- DS/RC)
              GO TO 320
310     COR(K) = 1./(1. + DS/RC)
320     CONTINUE

```

```

SE = ((1. - SW) * AMU) * * 2 + SD2
DO 340 K = 1,NS
  WT = W(K)
  SE = SE + WT * WT * SE2 - 2. * WT * COR(K) * SD2
  DO 330 IS = 1,NS
330    SE = SE + WT * W(IS) * SPAC(K,IS) * SD2
340    CONTINUE
      IF (SE.LE.0.0) SE = 0.0
      SE = SQRT(SE)
      STORE(IDS2,J) = SE
      SMAX = AMAX1(SE,SMAX)
      SMIN = AMIN1(SE,SMIN)
350    CONTINUE
360    CONTINUE
      HM = HM/TAR
      CALL MAP (1)
      WRITE (6,380) KD,HM
      IF (KCOR.EQ.0) RETURN
      WRITE (6,390) TITLE
      CALL MAP (2)
      RETURN

C
C
C
370 FORMAT (1H1,/T20,28HMULTIQUADRIC INTERPOLATION :,10A8//)
380 FORMAT (// T20,26HMULTIQUADRIC OPTION CODE =,I4/T23,28H1 SQRT((X
1-XS)**2+(Y-YS)**2),/T23,24H2 (X-XS)**2 + (Y-YS)**2,/T23,24H3 ABS
2(X-XS) + ABS(Y-YS),/T23,27H4 MAX(ABS(X-XS),ABS(Y-YS)),//T20,35HAR
3EAL MEAN OF INTERPOLATED VALUES =,F12.5)
390 FORMAT (1H1,/T20,23HERROR OF INTERPOLATION: ,10A8//)
      END
      SUBROUTINE DEPTH

C
C
C
      DEPTH-AREA CURVE COMPUTATIONS

      DIMENSION          CSA(50)      , CSB(50) , WT(50,50)
      CHARACTER*8 TITLE, SNAME, SDNAME
      COMMON /SYSID/      STORE(400,101)
      COMMON /SYSID2/     TITLE(10) , SNAME(50)
      COMMON /CONST/     IDSC          , NS          , NX          ,
1      NY          , DX          , DY          , XMAX          ,
2      XMIN        , YMAX        , YMIN        , HMAX          ,
3      HMIN        , SMAX        , SMIN        ,
      COMMON /AREAL/     AREA          , TAR
      COMMON /CODE/      KOD          , VR(5)          , KCOR          ,
1      KERR

      COMMON /DEPTH1/ SDNAME(50)
      COMMON /DEPTH2/ NRS, XRS(50), YRS(50)
      BIG = 9999999.0
      NC = MIN0(KOD,50)
      NC = MAX0(2,NC)
      CMIN = VR(1)
      CMAX = VR(2)
      IF (CMAX.NE.CMIN) GO TO 100
      CMIN = HMIN
      CMAX = HMAX

```

```

100 DC = (CMAX - CMIN)/NC
    DO 110 K = 1,NC
        CSA(K) = 0.0
        CSB(K) = 0.0
        IF(NRS.EQ.0) GO TO 110
        DO 105 N=1,NRS
105 WT(N,K) = 0.0
110 CONTINUE
    NCL = NC - 1
    DO 150 I = 1,NY
        YT = YMIN + DY * (I-1.0)
        DO 140 J = 1,NX
            IF (STORE(I,J).EQ.-BIG) GO TO 140
            XT = XMIN + DX * (J-1.0)
            DO 120 K = 1,NCL
                KC = K
                CL = CMIN + (K - 1.0) * DC
                CH = CL + DC
                CV = STORE(I,J)
                IF (CV.GE.CL.AND.CV.LT.CH) GO TO 130
120 CONTINUE
                KC = NC
                IF (CV.LT.CMIN.AND.CV.GT.CMAX) GO TO 140
130 CSA(KC) = CSA(KC) + 1.0
                CSB(KC) = CSB(KC) + CV
                IF (NRS.EQ.0) GO TO 140
                DSM = BIG
                ISM = 1
                DO 135 N=1,NRS
                    DST = SQRT ((XRS(N) - XT)**2 + (YRS(N) - YT) **2)
                    IF (DST.GE.DSM) GO TO 135
                    DSM = DST
                    ISM = N
135 CONTINUE
                    WT(ISM,KC) = WT(ISM,KC) + 1.0
140 CONTINUE
150 CONTINUE
    WRITE (6,170) TITLE
    UAREA = DX * DY
    SA = 0.0
    SB = 0.0
    DO 160 L = 1,NC
        K = NC - L + 1
        CL = CMIN + (K - 1.0) * DC
        SA = SA + CSA(K)
        SAP = SA * UAREA
        SB = SB + CSB(K)
        SBP = 0.0
        IF (SA.NE.0.0) SBP = SB/SA
160 WRITE (6,180) K,CL,CMAX,SAP,SBP
        IF (NRS.EQ.0) RETURN
        WRITE (6,185) (SDNAME(N),N=1,NRS)
        WRITE(6,186)
        SA= 0.0
        DO 165 L=1,NC
            K=NC-L+1

```

```

CL= CMIN+DC*(K-1.0)
SA=SA+CSA(K)
IF (SA.LE.0.0) SA=1.0
DO 166 N=1,NRS
IF(K.NE.NC) WT(N,K)=WT(N,K)+WT(N,K+1)
166 CSB(N)=WT(N,K)/SA
165 WRITE (6,190) K,CL,CMAX,(CSB(N),N=1,NRS)
RETURN

```

C
C
C

```

170 FORMAT (1H1/5X,31HDEPTH-AREA CURVE COMPUTATIONS: ,10A8//10X,6HNUMB
1ER,4X,11HLOWER LIMIT,4X,11HUPPER LIMIT,4X,8HCUM.AREA,4X,9HAVE.DEPT
2H,/)
180 FORMAT (11X,I3,6X,F10.3,5X,F10.3,2X,F10.2,3X,F10.3)
185 FORMAT (///5X,'STATION WEIGHTS USING THIESSEN POLYGON METHOD (STRA
1IGHT DISTANCE FORMULA) '//6X,'NUMBER',2X,'LOWER LIMIT',2X,'UPPER LI
2MIT',2X,8(2X,A8)/2X,12(2X,A8)/2X,12(2X,A8))
186 FORMAT(/)
190 FORMAT (7X,I3,4X,F10.3,3X,F10.3,3X,8F10.4/2X,12F10.4/2X,12F10.4)
END
SUBROUTINE SPACOR

```

C
C
C

DETERMINATION OF SYSTEM STATISTICS

```

DIMENSION          DST(50,50)
CHARACTER*8 TITLE, SNAME
COMMON /SYSID2/ TITLE(10), SNAME(50)
COMMON /SYSID/ STORE(400,101)
COMMON /CONST/
1   NY              , DX              , NS              , NX              ,
2   XMIN            , YMAX            , DY              , XMAX            ,
3   HMIN            , SMAX            , YMIN            , HMAX            ,
COMMON /SYSDAT/
1   Z(50)           , X(50)           , Y(50)           , H(50)           ,
COMMON /SYSTAT/
1   RC              , ATC              , STD              , STE              ,
COMMON /SBOUND/
COMMON /AREAL/
COMMON /CODE/
1   KERR            , ID              , VR(5)           , KCOR            ,
BIG = 9999999.0
KRTN = 0
IF (KCOR.NE. - 1) GO TO 100
KRTN = - 1
READ (5,290) KCOR,ATC
READ (5,280) (CRSCO(I),I = 1,NS)
100 DO 130 K = 1,NS
    DO 120 L = 1,NS
        RAD = SQRT((X(L) - X(K)) * * 2 + (Y(L) - Y(K)) * * 2)
        DST(K,L) = RAD
        IF (KCOR.EQ.2) GO TO 110
        SPAC(K,L) = EXP(- RAD/RC)
        GO TO 120
110   SPAC(K,L) = 1.0/(1.0 + RAD/RC)
        DST(K,K) = 0.0

```

```

120  CONTINUE
130  SPAC(K,K) = 1.0
     IF (KRTN.EQ. - 1) RETURN
     SX = 0.0
     SY = 0.0
     DO 150 I = 1,NY
       YT = YMIN + DY * (I - 1.0)
       DO 140 J = 1,NX
         IF (STORE(I,J).EQ.-BIG) GO TO 140
         XT = XMIN + DX * (J - 1.0)
         SX = SX + XT
         SY = SY + YT
140  CONTINUE
150  CONTINUE
     XC = SX/TAR
     YC = SY/TAR
     DO 190 K = 1,NS
       SUM = 0.0
       DO 180 I = 1,NY
         YT = YMIN + DY * (I - 1.0)
         DO 170 J = 1,NX
           IF (STORE(I,J).EQ.-BIG) GO TO 170
           XT = XMIN + DX * (J - 1.0)
           RAD = SQRT((XT - X(K)) * * 2 + (YT - Y(K)) * * 2)
           IF (KCOR.EQ.2) GO TO 160
           SUM = SUM + EXP(- RAD/RC)
           GO TO 170
160  SUM = SUM + 1.0/(1.0 + RAD/RC)
170  CONTINUE
180  CONTINUE
190  CRSCO(K) = SUM/TAR
     SUM = 0.0
     DO 220 I = 1,NY
       YT = YMIN + DY * (I - 1.0)
       DO 210 J = 1,NX
         IF (STORE(I,J).EQ.-BIG) GO TO 210
         XT = XMIN + DX * (J - 1.0)
         RAD = SQRT((XT - XC) * (XT - XC) + (YT - YC) * (YT - YC))
         IF (KCOR.EQ.2) GO TO 200
         SUM = SUM + EXP(- RAD/RC)
         GO TO 210
200  SUM = SUM + 1./(1. + RAD/RC)
210  CONTINUE
220  CONTINUE
     CROC = SUM/TAR
     CROC2 = CROC * CROC
     PERM = 0.0
     DO 240 L = 1,NB
       X1 = XB(L)
       Y1 = YB(L)
       LB = L + 1
       IF (LB.LE.NB) GO TO 230
       LB = 1
230  X2 = XB(LB)
       Y2 = YB(LB)
240  PERM = PERM + SQRT((X2 - X1) * (X2 - X1) + (Y2 - Y1) * (Y2 - Y1))

```

```

REQ = 2. * AREA/PERM
RR = REQ/RC
R2 = RR * RR
R3 = R2 * RR
R4 = R3 * RR
IF (KCOR.EQ.2) GO TO 250
ATCLO = 4.0 * ( - EXP( - RR)/RR + (1.0 - EXP( - RR))/R2) * * 2
ATCHI = 8.0 * (1./ (3. * RR) - .5/R2 - EXP( - RR)/R3 + (1. - EXP( -
1 RR))/R4)
GO TO 260
250 ATCLO = 4.0 * ( - 2./ (3. * RR) - 1./R2 + 1./ (6. * R4)) * ALOG(1. +
1 2. * RR) + 4.0 * (2./ (3. * RR) + 1./R2 - 1./ (3. * R4)) * ALOG(1.
2+ RR) + 4. * (2./ (3. * RR) + 1./ (6. * R2))
ATCHI = 8. * (1./ (3. * RR) + .5/R2 - 1./ (6. * R4)) * ALOG(1. + RR)
1 + 8. * ( - 4./ (9. * RR) - 1./ (12. * R2) + 1./ (6. * R3))
260 ATC = AMINI(1., (ATCHI + ATCLO)/2.)
WRITE (6,300) AMJ,STD,STE,ATC,ATCHI,ATCLO,CROC2,CROC,AREA,PERM,REQ
1,XC,YC
WRITE (6,340)
IF (KCOR.EQ.1) WRITE (6,370)
WRITE (6,350)
IF (KCOR.EQ.2) WRITE (6,370)
WRITE (6,360) RC
WRITE (6,310) (SNAME(K),K = 1,NS)
DO 270 K = 1,NS
WRITE (6,320) SNAME(K),CRSCO(K), (SPAC(K,L),L = 1,NS)
270 WRITE (6,330) (DST(K,L),L = 1,NS)
RETURN
C
C
C
280 FORMAT (10F10.0)
290 FORMAT (I2,F10.0)
300 FORMAT (1H1,T30,34H S Y S T E M S T A T I S T I C S : ,/T25,38HSPA
TIAL MEAN ..... ,F10.3/T25,38HSPATIAL STANDARD
2DEVIATION ..... ,F10.3/T25,38HSTANDARD MEASUREMENT ERROR ....
3..... ,F10.3//T25,38HAREAL AUTOCORRELATION COEFFICIENT ..... ,F10.5
4/T25,38H UPPER LIMIT ..... ,F10.5/T25,38H
5HIGH LOWER LIMIT ..... ,F10.5/T25,38H LOW LOWER LIM
6T ..... ,F10.5/T25,31HCENTROID-AREA CROSS CORRELATION,/T
725,38H COEFFICIENT ..... ,F10.5//T25,38HBASIN A
8REA ..... ,F10.3/T25,38HBASIN PERIMETER .....
9..... ,F10.3/T25,38HBASIN EQUIVALENT RADIUS .....
0... ,F10.3/T25,14HBASIN CENTROID,/T25,38H X - AXIS .....
1..... ,F10.3/T25,38H Y - AXIS ..... ,F
210.3)
310 FORMAT (/T6,22HSTATION AREAL CROSS,T45,36H STATION-STATION CR
LOSS CORRELATION,/T6,22H NAME CORR. COEFF.,T39,T52,26HCOEFFICIE
2NTS AND DISTANCES,/(T31,10(2X,A8)))
320 FORMAT (/T5,A8,T16,F10.5,(T31,10F10.5))
330 FORMAT (T5,11H(DISTANCES),(T31,10F10.3))
340 FORMAT (1H1,T30,61H S Y S T E M C O R R E L A T I O N C O E F F I
1 C I E N T S : ,/T28,65H ( S T A T I O N C O R R E L A T I O N C O
2 E F F I C I E N T S ),//T35,29HSPATIAL CORRELATION FUNCTION: ,/T38
3,21HRHO = EXP(-RADIUS/RC))
350 FORMAT (1H ,T38,25HRHO = 1.0/(1.0+RADIUS/RC))

```

```

360 FORMAT (1H ,T35,28H CHARACTERISTIC RADIUS, RC =,F10.3)
370 FORMAT (1H+,T34,3H=>)
END
SUBROUTINE OPTIM

```

C
C
C

```

OPTIMAL INTERPOLATION AND AREAL AVERAGING

DIMENSION          W(52)          , COR(52)          , B(52)          ,
1      A(52,52)
CHARACTER*8 TITLE, SNAME
COMMON /SYSID2/ TITLE(10), SNAME(50)
COMMON /SYSID/ STORE(400,101)
COMMON /CONST/      IDSC          , NS          , NX          ,
1      NY          , DX          , DY          , XMAX          ,
2      XMIN          , YMAX          , YMIN          , HMAX          ,
3      HMIN          , SMAX          , SMIN          ,
COMMON /SYSDAT/      X(50)          , Y(50)          , H(50)          ,
1      Z(50)
COMMON /SYSTAT/      AMU          , STD          , STE          ,
1      RC          , ATC          , CRSCO(50)          , SPAC(50,50)
COMMON /AREAL/      AREA          , TAR          ,
COMMON /CODE/      KOD          , VR(5)          , KCOR          ,
1      KERR
BIG = 9999999.0
KOPT = MAX0(KOD,1)
KOPT = MIN0(KOPT,2)
IF (KOPT.EQ.2) GO TO 100
WRITE (6,350) (TITLE(LT),LT = 1,9)
GO TO 110
100 WRITE (6,360) (TITLE(LT),LT = 1,9)
110 CONTINUE
HMAX = - BIG
HMIN = BIG
SMAX = - BIG
SMIN = BIG
SD2 = STD * STD
SE2 = STE * STE
MS = NS + 2 - KOPT
HM = 0.0
DO 130 K = 1,NS
    HMAX = AMAX1(H(K),HMAX)
    HMIN = AMIN1(H(K),HMIN)
    DO 120 L = 1,NS
120    A(K,L) = SD2 * SPAC(K,L)
130    A(K,K) = A(K,K) + SE2
    NSS = NS + 1
    IF (KOPT.EQ.2) GO TO 150
    DO 140 K = 1,NS
        A(K,NSS) = 1.0
140    A(NSS,K) = 1.0
        A(NSS,NSS) = 0.0
        B(NSS) = 1.0
150    CALL GEPCON (A,MS)
    IF (KERR.EQ.1) RETURN
    DO 260 I = 1,NY
        YI = YMIN + DY * (I - 1.0)

```

```

IDS2 = I + IDSC
DO 250 J = 1, NX
  IF (STORE(I,J).EQ.-BIG) GO TO 250
  XT = XMIN + DX * (J - 1.0)
  DO 170 K = 1, NS
    RAD = SQRT((XT - X(K)) * * 2 + (YT - Y(K)) * * 2)
    IF (KCOR.EQ.2) GO TO 160
    COR(K) = EXP(- RAD/RC)
    GO TO 170
160   COR(K) = 1.0/(1.0 + RAD/RC)
170   B(K) = SD2 * COR(K)
    IF (KOPT.EQ.2) GO TO 190
    WL = 0.0
    DO 180 K = 1, MS
180   WL = WL + A(MS,K) * B(K)
190   SUM = 0.0
    DO 210 K = 1, NS
      W(K) = 0.0
      DO 200 L = 1, MS
200     W(K) = W(K) + A(K,L) * B(L)
210   SUM = SUM + W(K)
    IF (KOPT.EQ.2) GO TO 220
    HH = 0.0
    SE = SD2 - WL
    GO TO 230
220   HH = (1.0 - SUM) * AMU
    SE = SD2
230   DO 240 K = 1, NS
      HH = HH + W(K) * H(K)
240   SE = SE - W(K) * SD2 * COR(K)
    STORE(I,J) = HH
    IF (SE.LE.0.0) SE = 0.0
    SE = SQRT(SE)
    STORE(IDS2,J) = SE
    SMAX = AMAX1(SE, SMAX)
    SMIN = AMIN1(SE, SMIN)
    HM = HM + HH
250   CONTINUE
260   CONTINUE
    HM = HM/TAR
    CALL MAP (1)
    WRITE (6,410) HM
    WRITE (6,370) (TITLE(LT), LT = 1,9)
    CALL MAP (2)
    DO 270 K = 1, NS
270   B(K) = SD2 * CRSCO(K)
    IF (KOPT.EQ.2) GO TO 290
    WL = 0.0
    DO 280 K = 1, MS
280   WL = WL + A(MS,K) * B(K)
290   SUM = 0.0
    DO 310 K = 1, NS
      W(K) = 0.0
      DO 300 L = 1, MS
300     W(K) = W(K) + A(K,L) * B(L)
310   SUM = SUM + W(K)

```



```

IF (KOPT.EQ.2) GO TO 320
HH = 0.0
SE = SD2 * ATC - WL
GO TO 330
320 HH = (1.0 - SUM) * AMJ
SE = SD2 * ATC
330 DO 340 K = 1,NS
      HH = HH + W(K) * H(K)
340 SE = SE - W(K) * SD2 * CRSCO(K)
      HM = HH
      IF (SE.LE.0.0) SE = 0.0
      SE = SQRT(SE)
      WRITE (6,380)
      WRITE (6,400) (K,SNAME(K),H(K),W(K),K = 1,NS)
      WRITE (6,390) SUM,HM,SE
      RETURN

```

C
C
C

```

350 FORMAT (1H1,/T14,46HOPTIMAL INTERPOLATION BASED ON STRAIGHT DATA
1,9A8//)
360 FORMAT (1H1,/T17,41HOPTIMAL INTERPOLATION BASED ON DEVIATES: ,9A8/
1/)
370 FORMAT (1H1,T17,41HOPTIMAL STANDARD ERROR OF INTERPOLATION: ,9A8//
1)
380 FORMAT (// T35,23HOPTIMAL AREAL AVERAGING, //T20,7HSTA.NO.,T30,9HS
ITA. NAME,T43,10HINEUT DATA,T58,14HOPTIMAL WEIGHT,/)
390 FORMAT (// T25,23HSUM OF OPTIMAL WEIGHT =,F10.5/T25,23HOPTIMAL AR
LEAL MEAN =,F10.3/T25,26HSTD. ERROR OF AREAL MEAN =,F10.4/)
400 FORMAT (T22,I2,T31,A8,T44,F8.3,T59,F8.4)
410 FORMAT (// T15,35HAREAL MEAN OF INTERPOLATED VALUES =,F12.4)
      END

```

~eor

```

PROGRAM DAD2 (INPUT,OUTPUT,TAPE5=INPUT,TAPE6=OUTPUT,TAPE10=TAPE10)
CHARACTER TITL*80
DIMENSION A(30),DUR(200),DADD(30,200),W(7),XM(8,200),SN(7),
1  IFOR(11),WM(200)
DATA IFOR /4H(13X,9*1H ,6H,F8.0)/
DO 1000 IC=1,25
  READ (10,10) TITL
  WRITE (6,10) TITL
10  FORMAT (A80)
  READ (5,20) DUM
20  FORMAT (A1)
  READ (5,30) (SN(I),I=1,7)
30  FORMAT (16X,7F8.3)
  DO 100 I=1,7
    IF (SN(I).EQ.0.) GO TO 110
100  CONTINUE
110  NST=I-1
  NCPL=I
  J=1
  DO 200 I=1,9
    W(I)=0.
    IF(I.NE.SN(J)) GO TO 200
    W(I)=1.
    J=J+1
200  CONTINUE
  DO 210 I=2,10
    IF(W(I-1).EQ.0.) GO TO 220
    IFOR(I)=5H,F6.0
    GO TO 210
220  IFOR(I)=5H,6X
210  CONTINUE
  DO 230 I=1,2
230  READ (10,20) DUM
  DO 300 I=1,200
    READ (10,IFOR) (XM(J,I),J=1,NCPL)
    IF (XM(NCPL,I).EQ.-100.) GO TO 310
300  CONTINUE
310  CONTINUE
  NPIS = I-1
  DO 400 NA=1,30
    READ (5,40) D,A(NA), (W(I),I=1,NST)
    IF (D.EQ.-100.) GO TO 500
  DO 320 I=1,NPIS
    WM(I)=0.
    DO 320 J=1,NST
320  WM(I)=WM(I)+W(J)*XM(J,I)
  CALL MDEPTH (WM,NPIS,DUR)
  DO 330 I=1,NPIS
330  DADD(NA,I) = D*DUR(I)
400  CONTINUE
500  IAMAX=NA-1
  IT=(NPIS-1)/12 + 1
  DO 600 I=1,IT
    INIT = 12*(I-1) + 1
    IEND = INIT + 11
    IF (IEND.GT.NPIS) IEND=NPIS

```

```

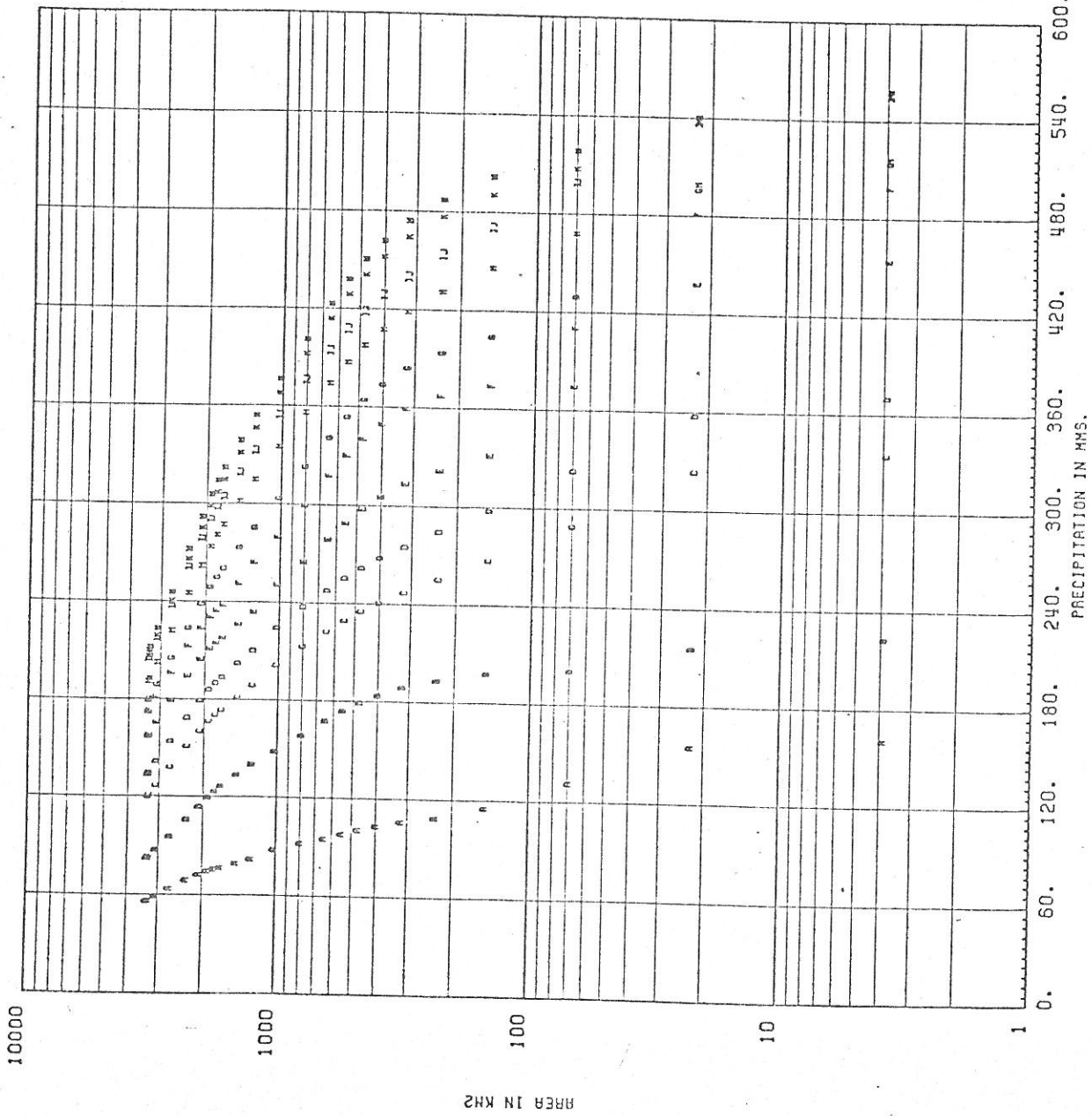
        WRITE (6,50) (J,J=INIT,IEND)
50  FORMAT('  AREA      ',12(' DUR.',I3,' H'))
        DO 410 J=1,IAMAX
410  WRITE (6,60) A(J), (DADD(J,K),K=INIT,IEND)
600  CONTINUE
1000 CONTINUE
        60  FORMAT (13F10.3)
        40  FORMAT (9F8.2)
        END
        SUBROUTINE MDEPTH (WM,NPTS,DUR)
        DIMENSION WM(200), DUR(200)
        DO 100 I=1,NPTS
100  WM(I)=WM(I)/WM(NPTS)
        IM=NPTS-1
        DO 430 I=1,IM
        AVMAX=WM(I)
        IMI=NPTS-I
        DO 440 J=1,IMI
        SUM=WM(I+J)-WM(J)
440  AVMAX=AMAX1(AVMAX,SUM)
430  DUR(I) = AVMAX
        DUR(NPTS)=1.
        RETURN
        END

```

~eor

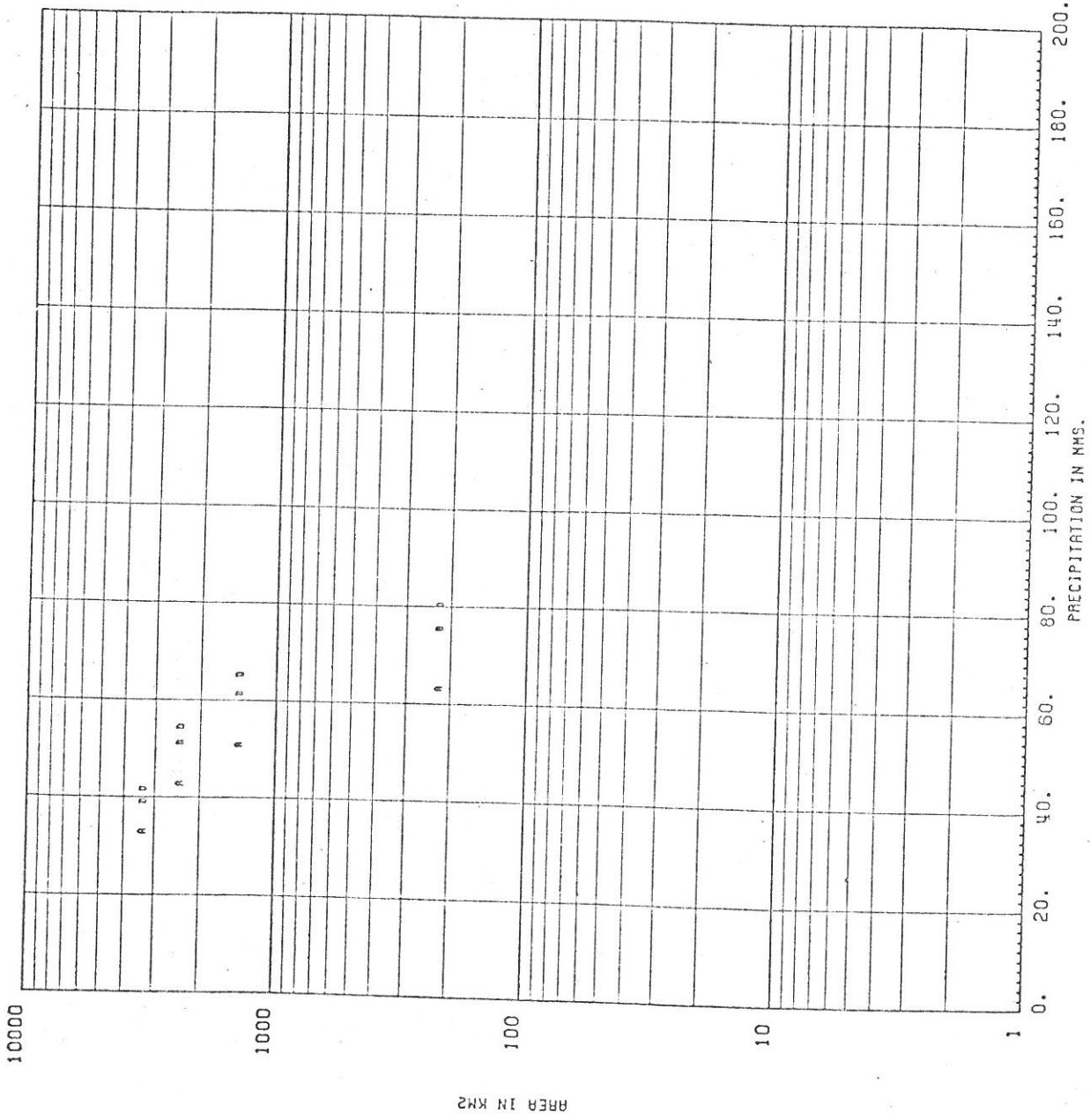
APPENDIX 1.6.C

DEPTH-AREA-DURATION CURVES FOR SELECTED STORMS



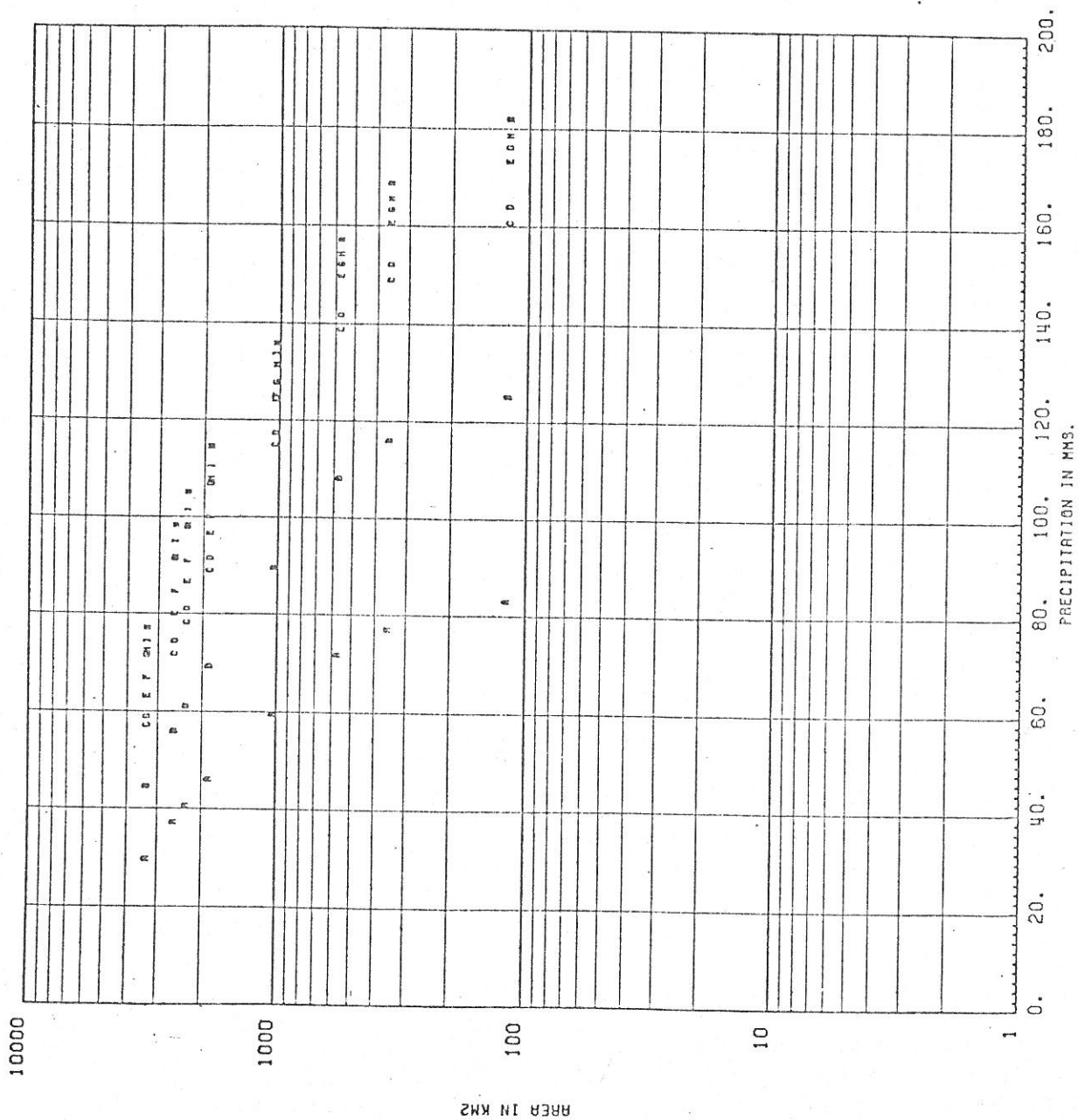
DEPTH AREA DURATION STORM STARTING 63 OCT 1 20 ENDING 63 OCT 5 5

Figure 1.6.C.1



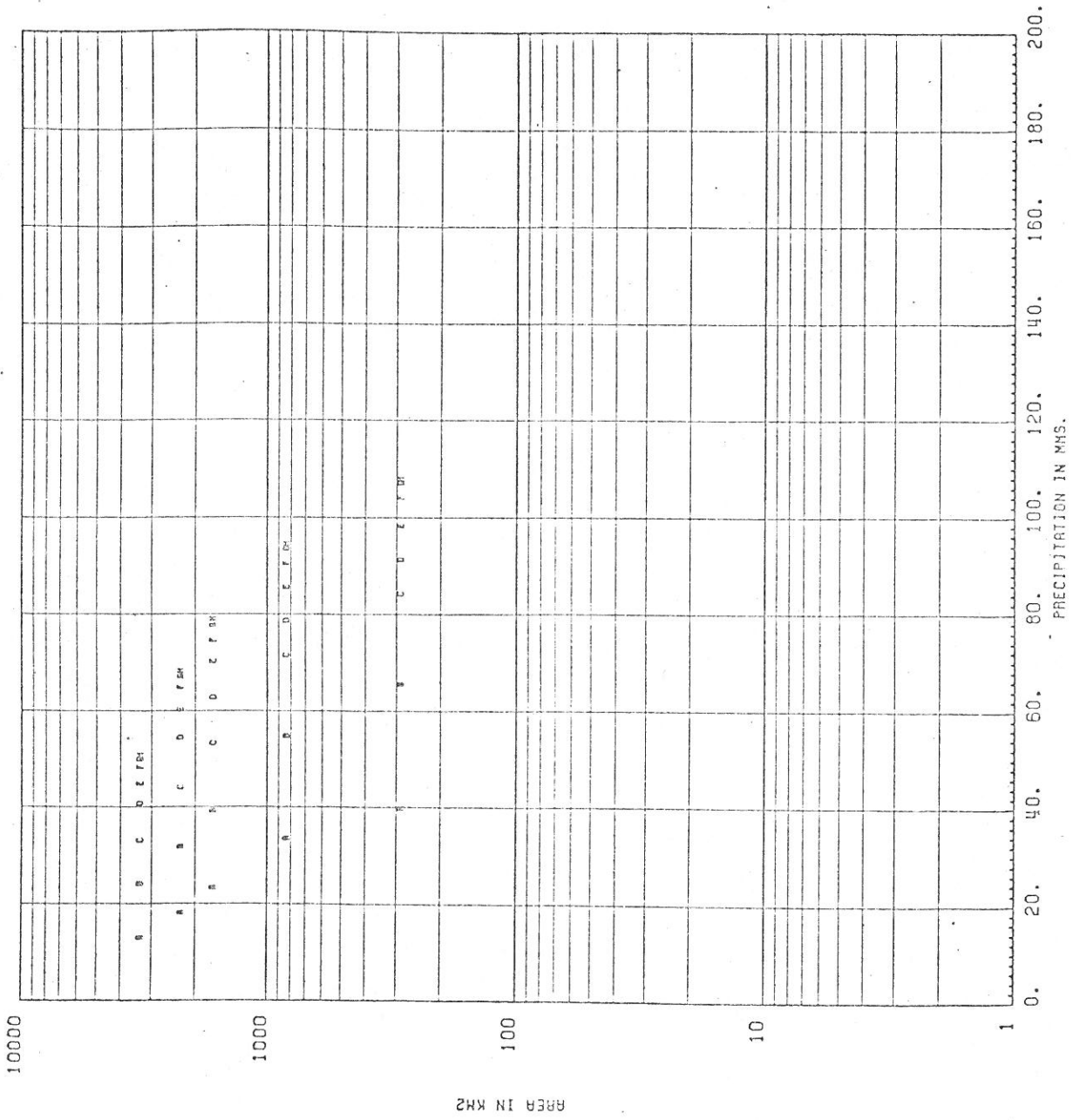
DEPTH AREA DURATION STORM STARTING 64 AGO 6 8 ENDING 64 AGO 7 8

Figure 1.6.C.2.



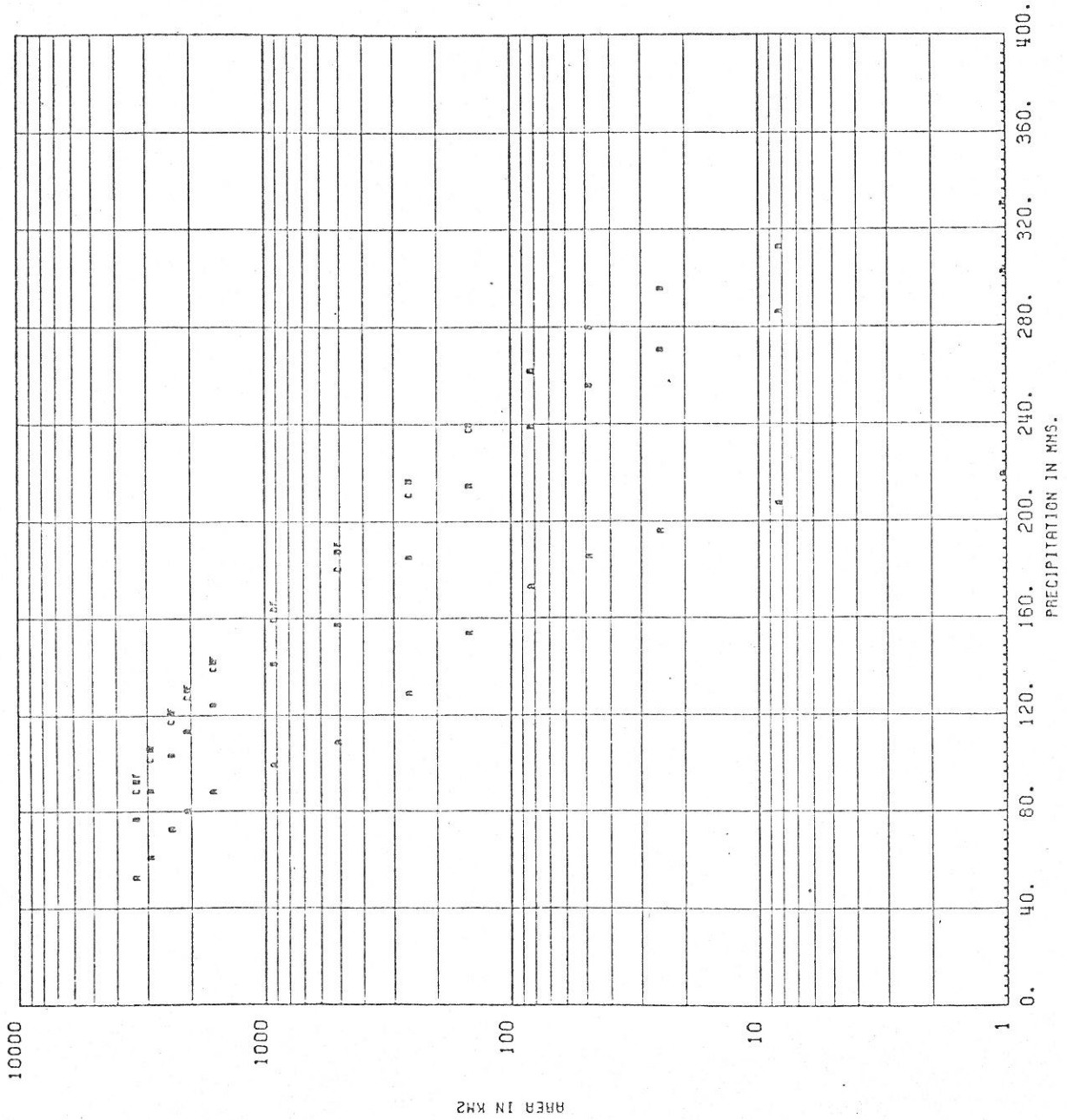
DEPTH AREA DURATION STORM STARTING 65 MAY 2 8 ENDING 65 MAY 5 8

Figure 1.6.C.3.



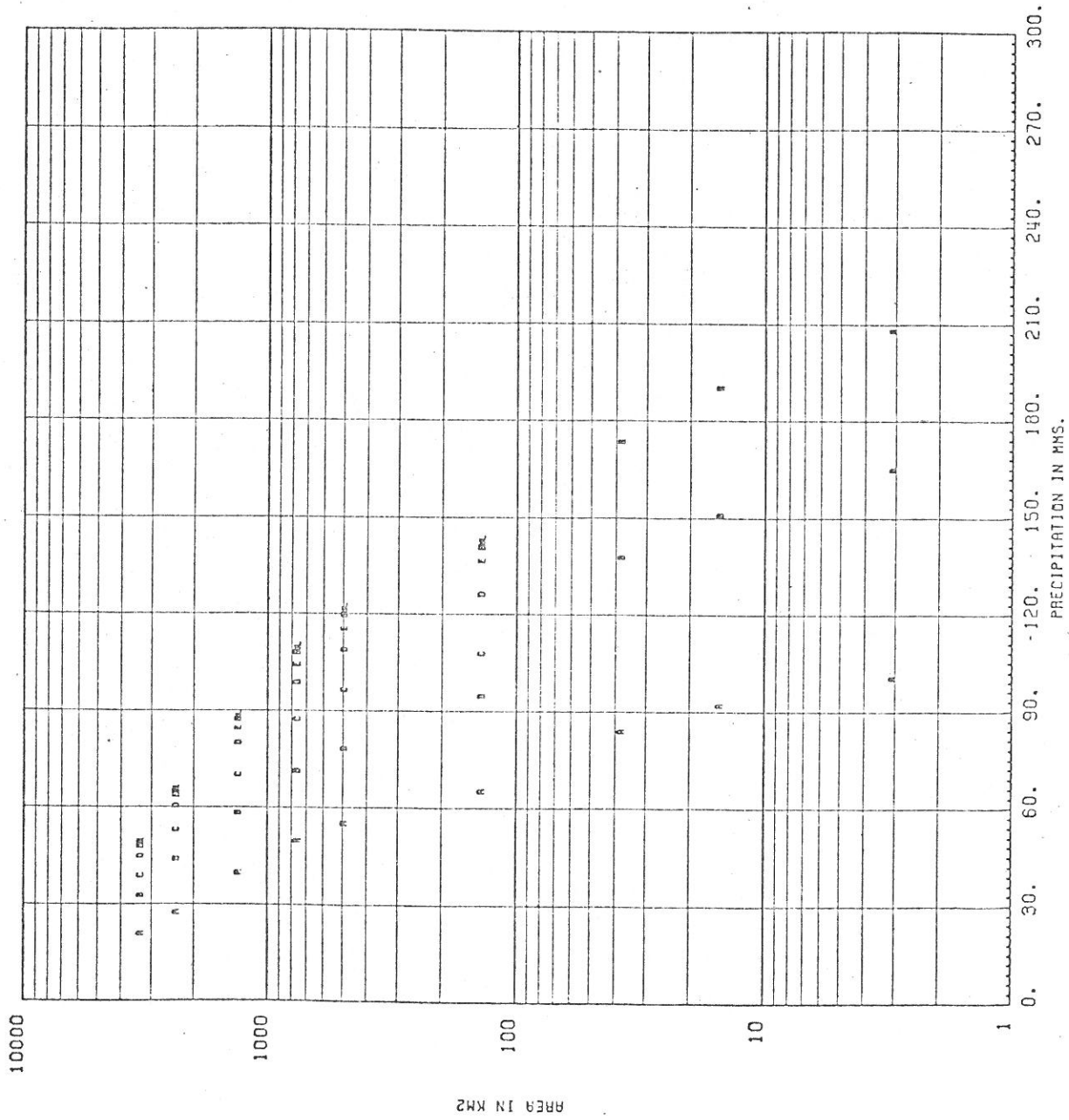
DEPTH AREA DURATION STORM STARTING 66 MAY 25 8 ENDING 66 MAY 27 8

Figure 1.6.C.4.



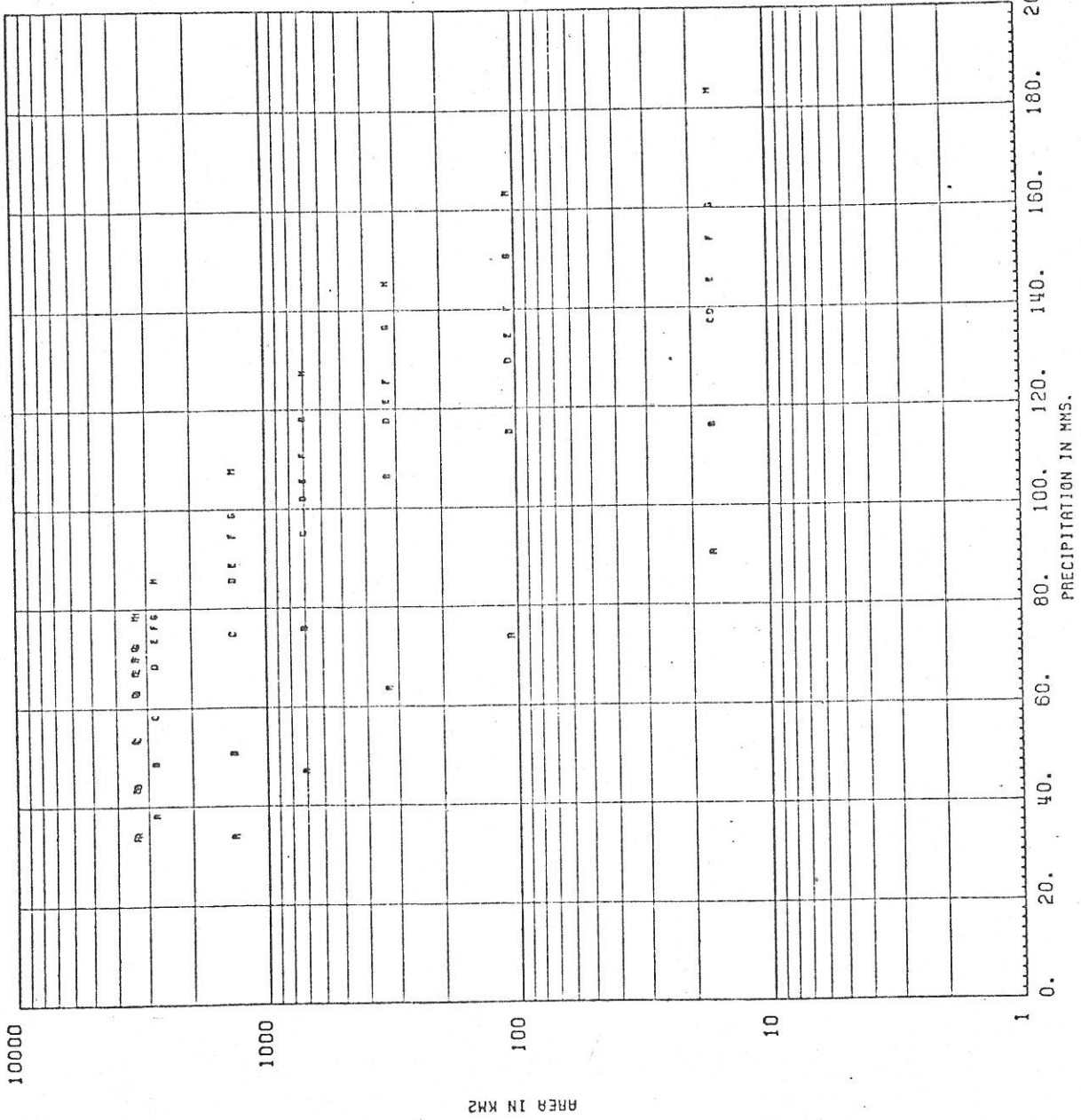
DEPTH AREA DURATION STORM STARTING 66 SEP 28 20 ENDING 66 SEP 30 8

Figure 1.6.C.5.



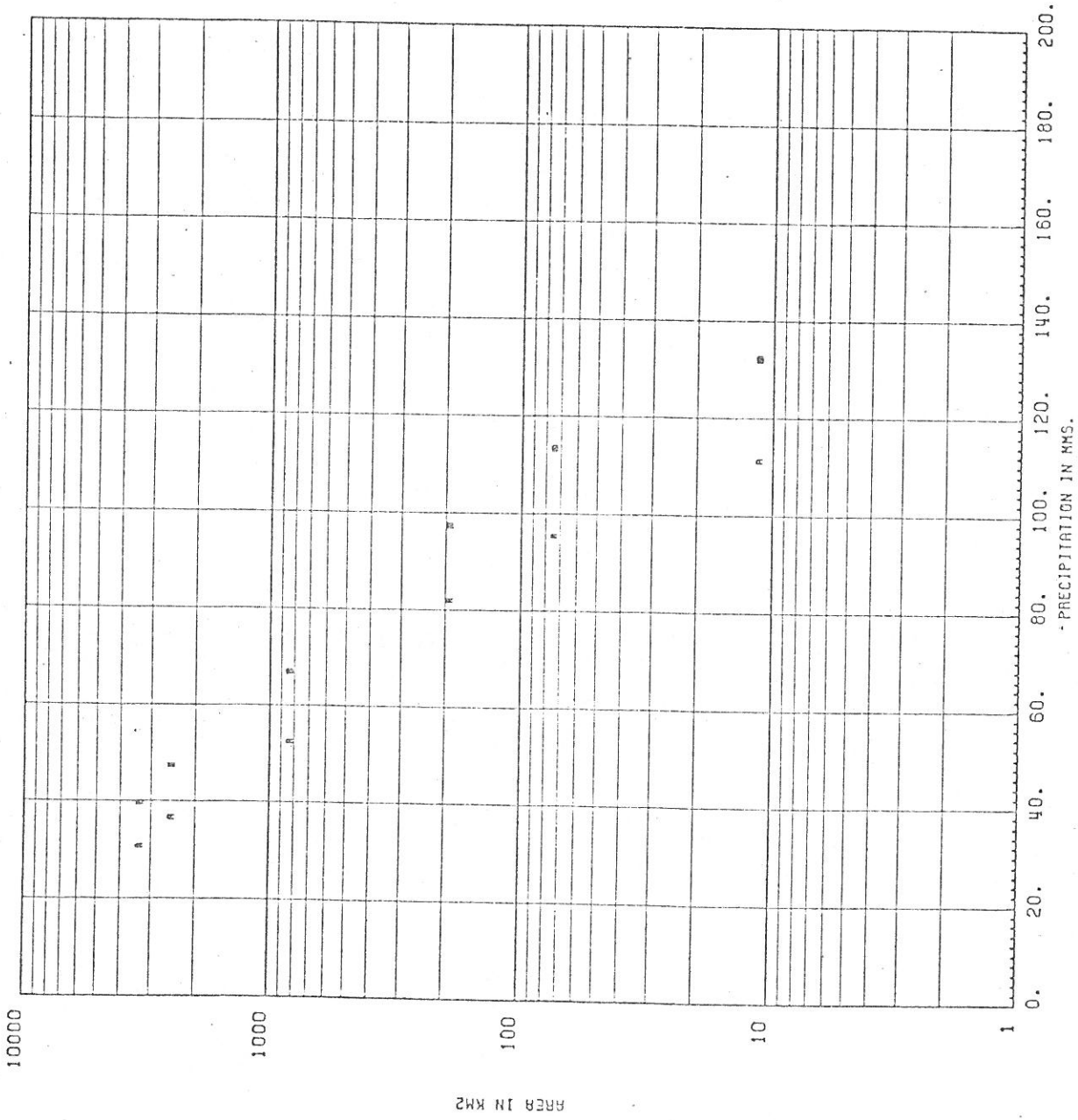
DEPTH AREA DURATION STORM STARTING 67 SEP 10 8 ENDING 67 SEP 13 8

Figure 1.6.C.6.



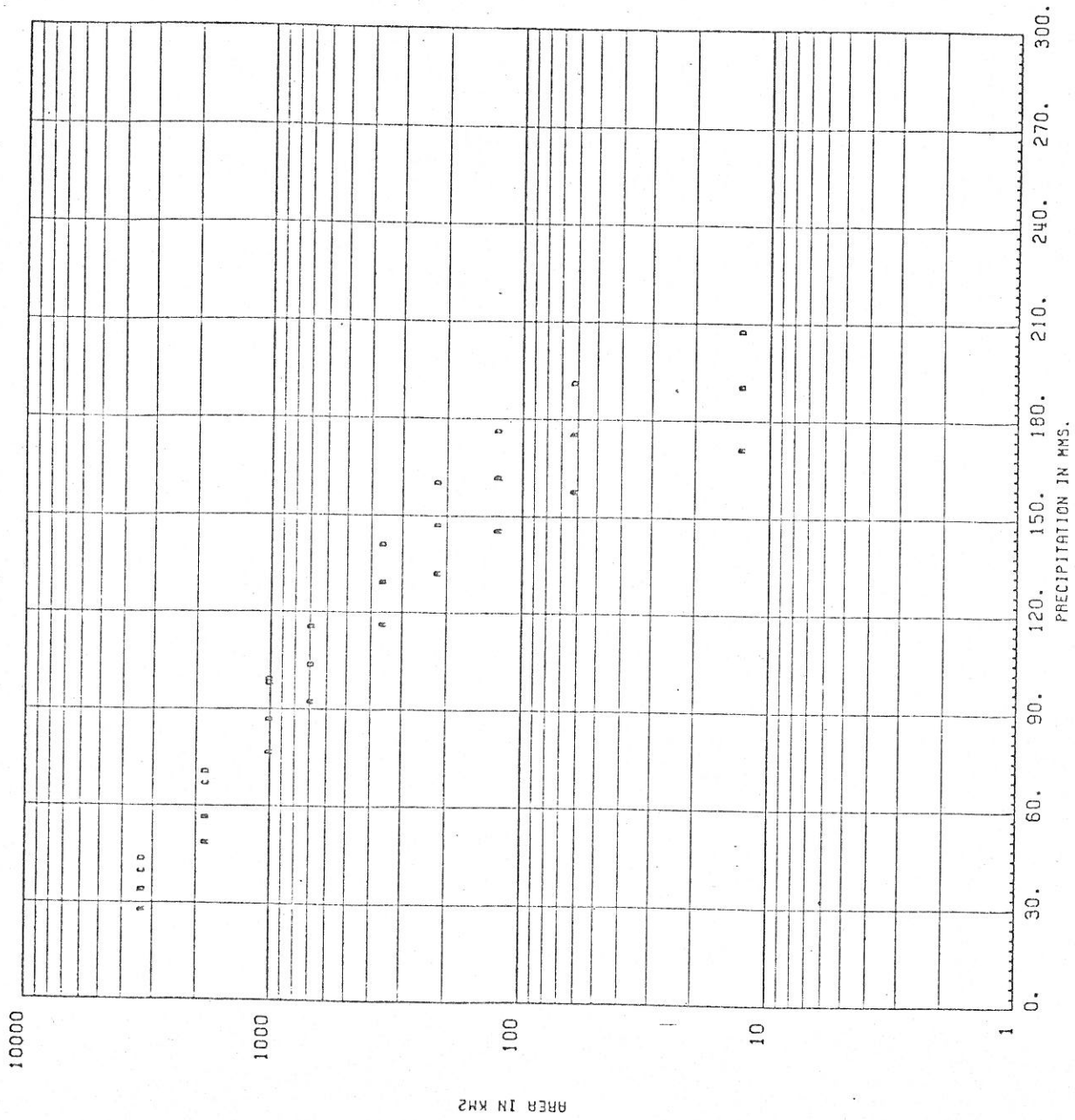
DEPTH AREA DURATION STORM STARTING 68 AGO 8 8 ENDING 68 AGO 10 8

Figure 1.6.C.7.



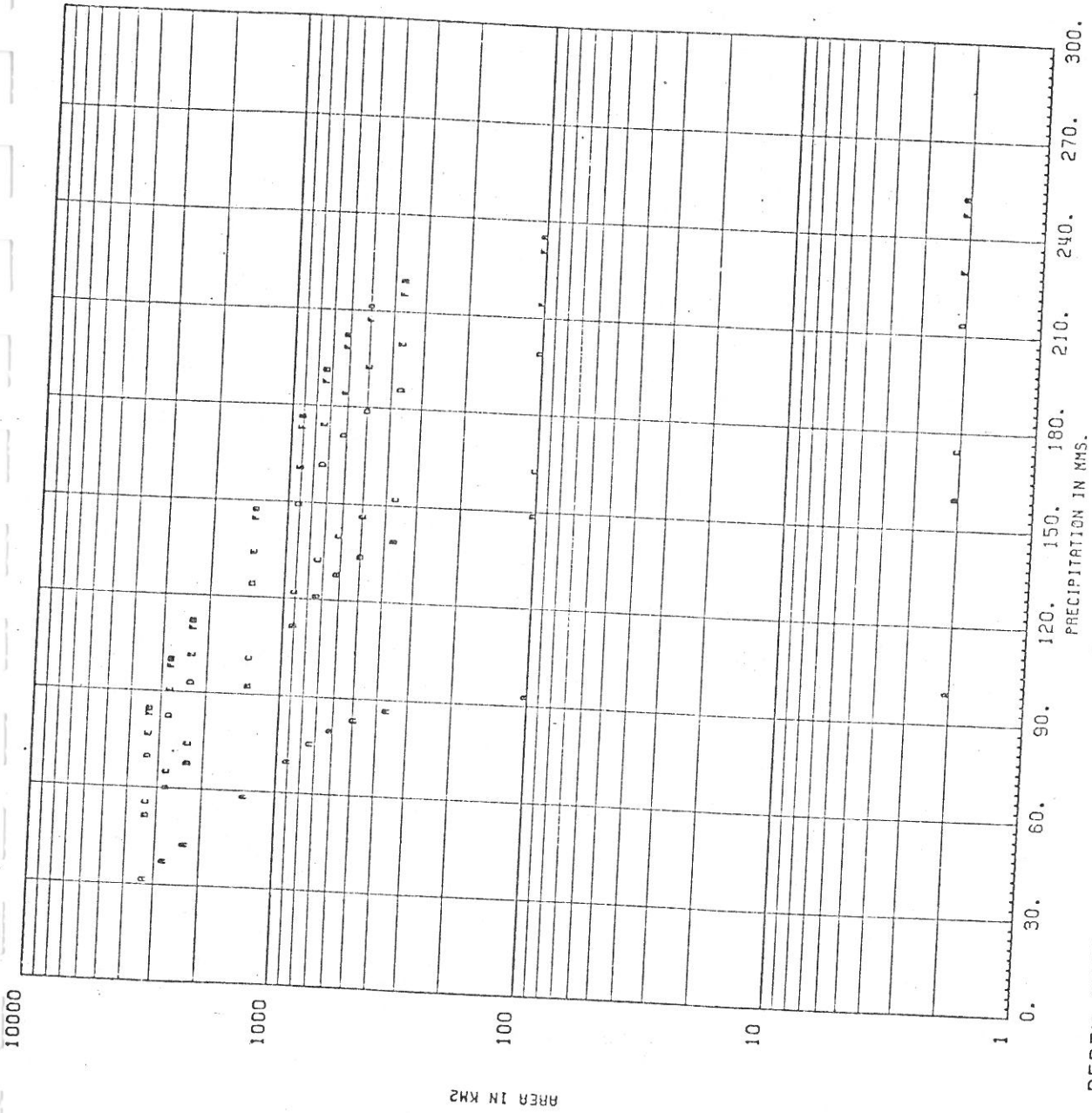
DEPTH AREA DURATION STORM STARTING 69 JUL 19 8 ENDING 69 JUL 20 8

Figure I.6.C.8.



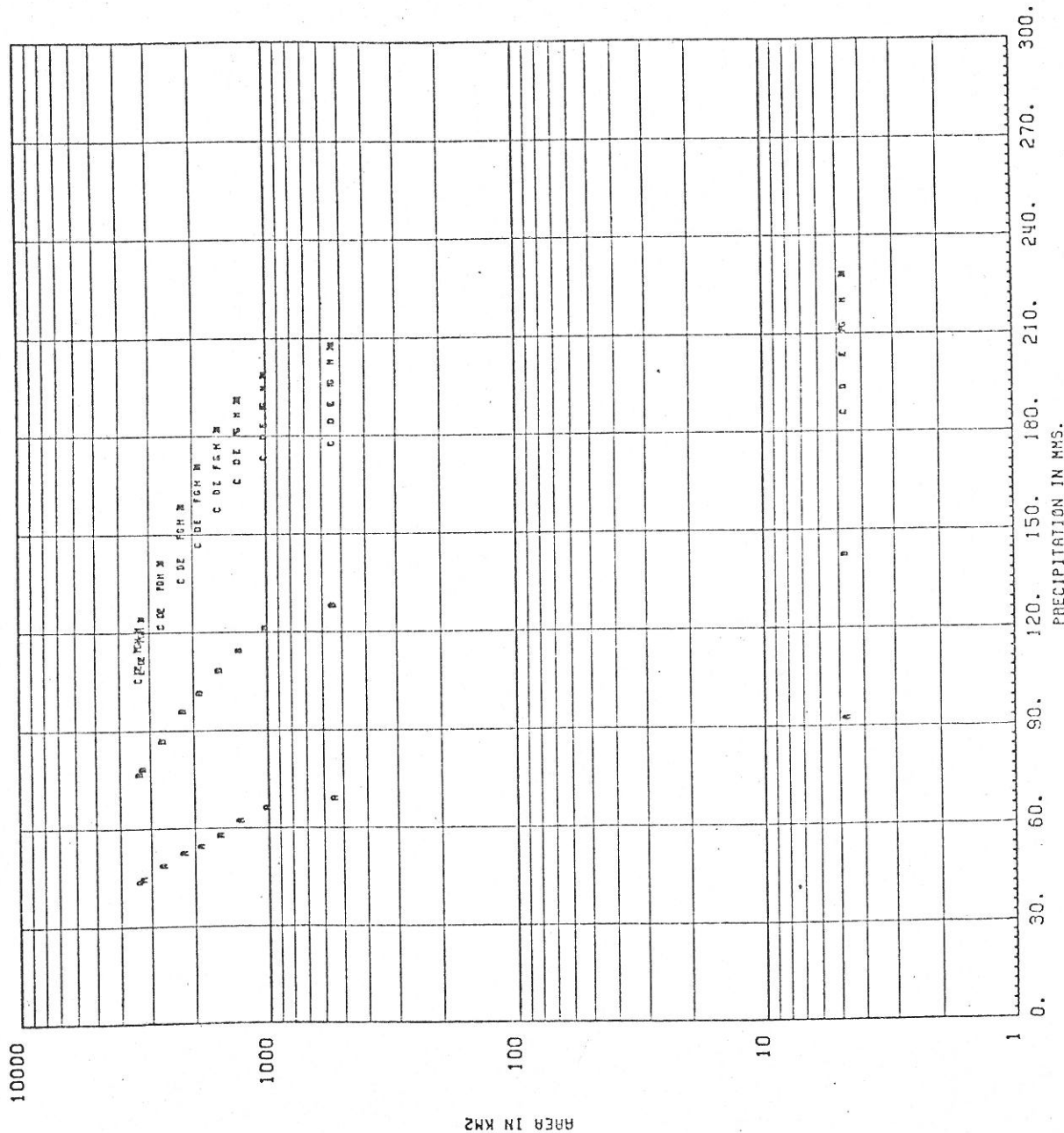
DEPTH AREA DURATION STORM STARTING 70 AGO 22 8 ENDING 70 AGO 23 8

Figure 1.6.C.9.



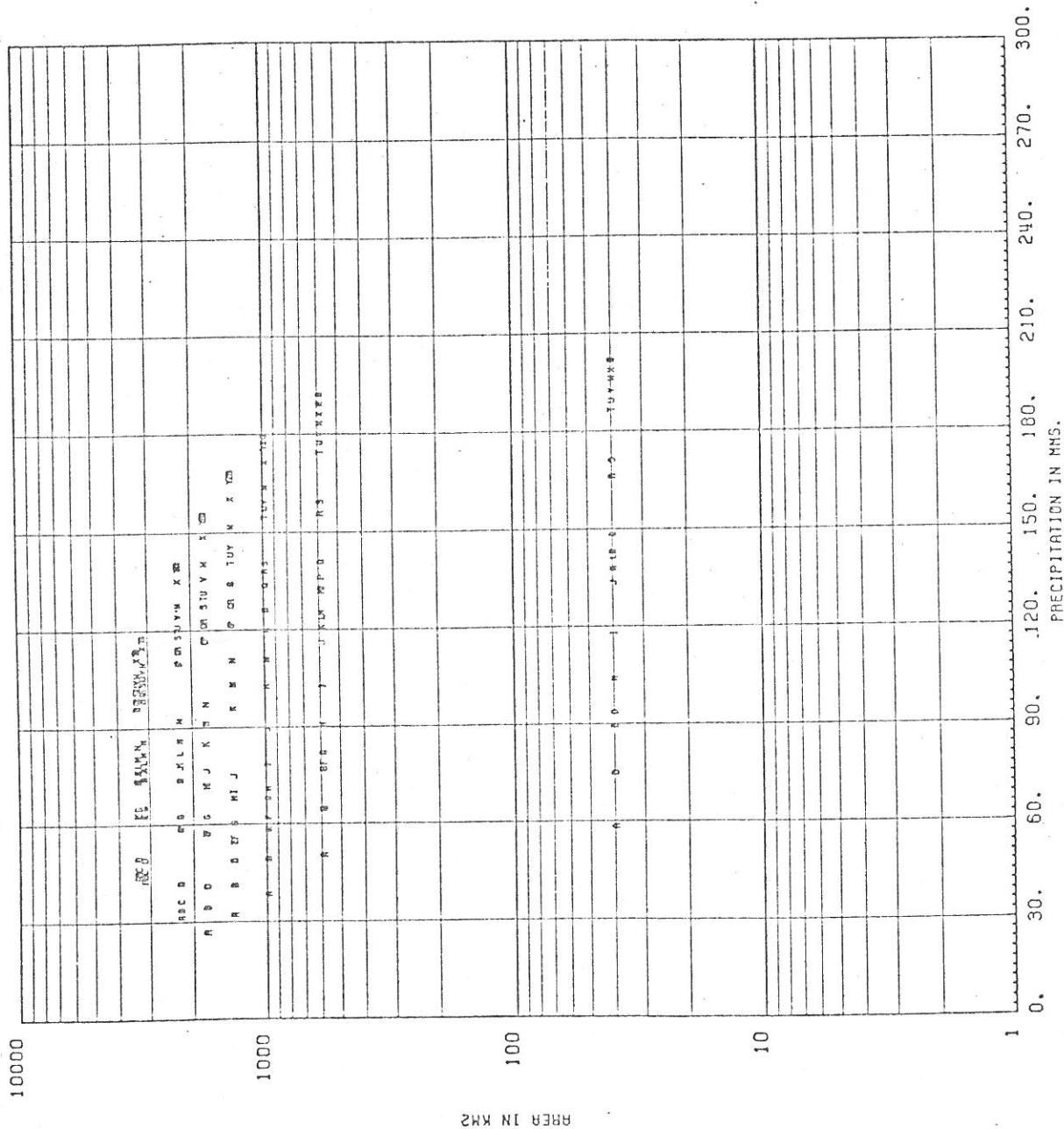
DEPTH AREA DURATION STORM STARTING 71 FEB 19 8 ENDING 71 FEB 21 8

Figure 1.6.C.10



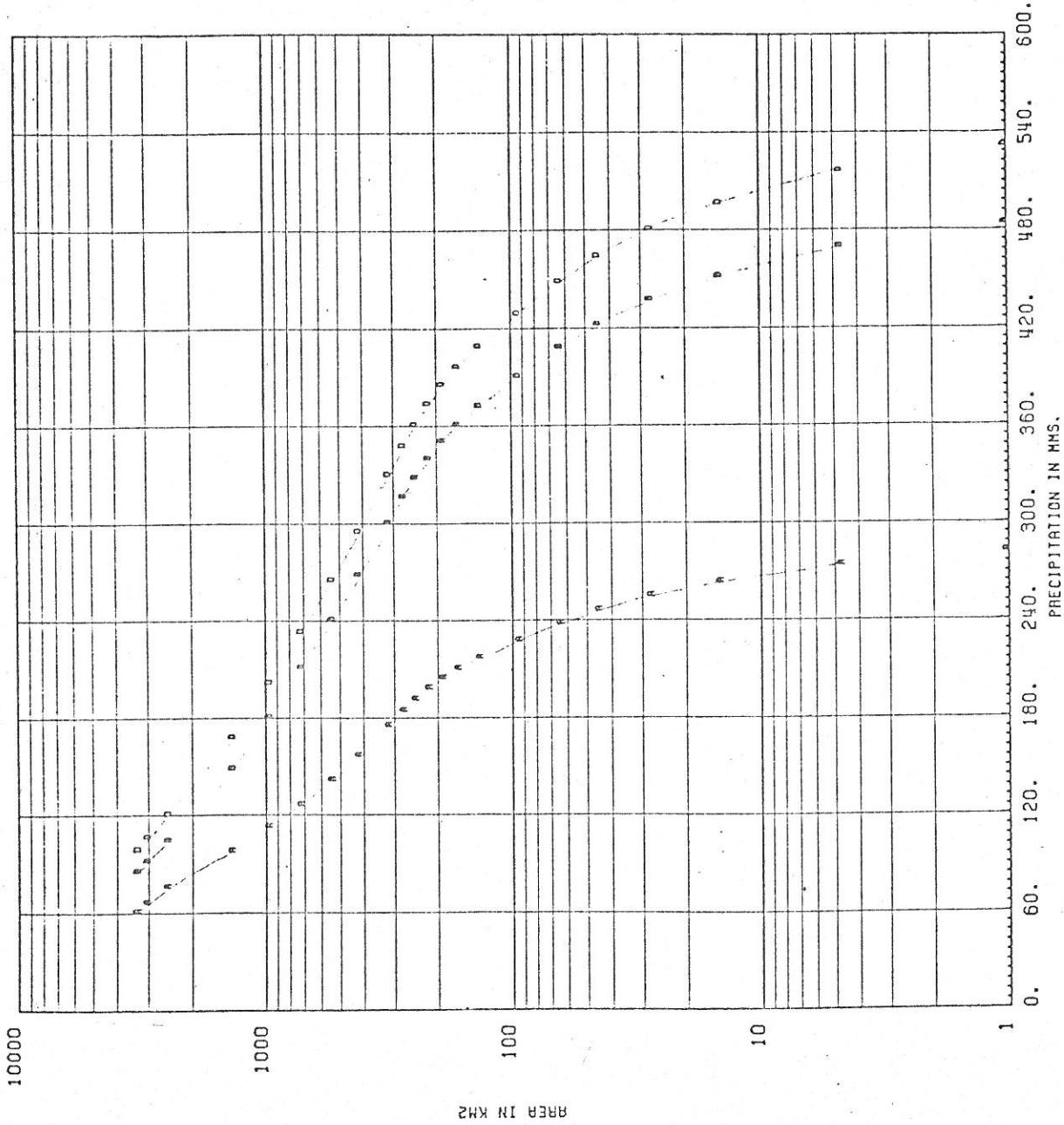
DEPTH AREA DURATION STORM STARTING 72 MAY 20 8 ENDING 72 MAY 23 8

Figure 1.6.C.11.



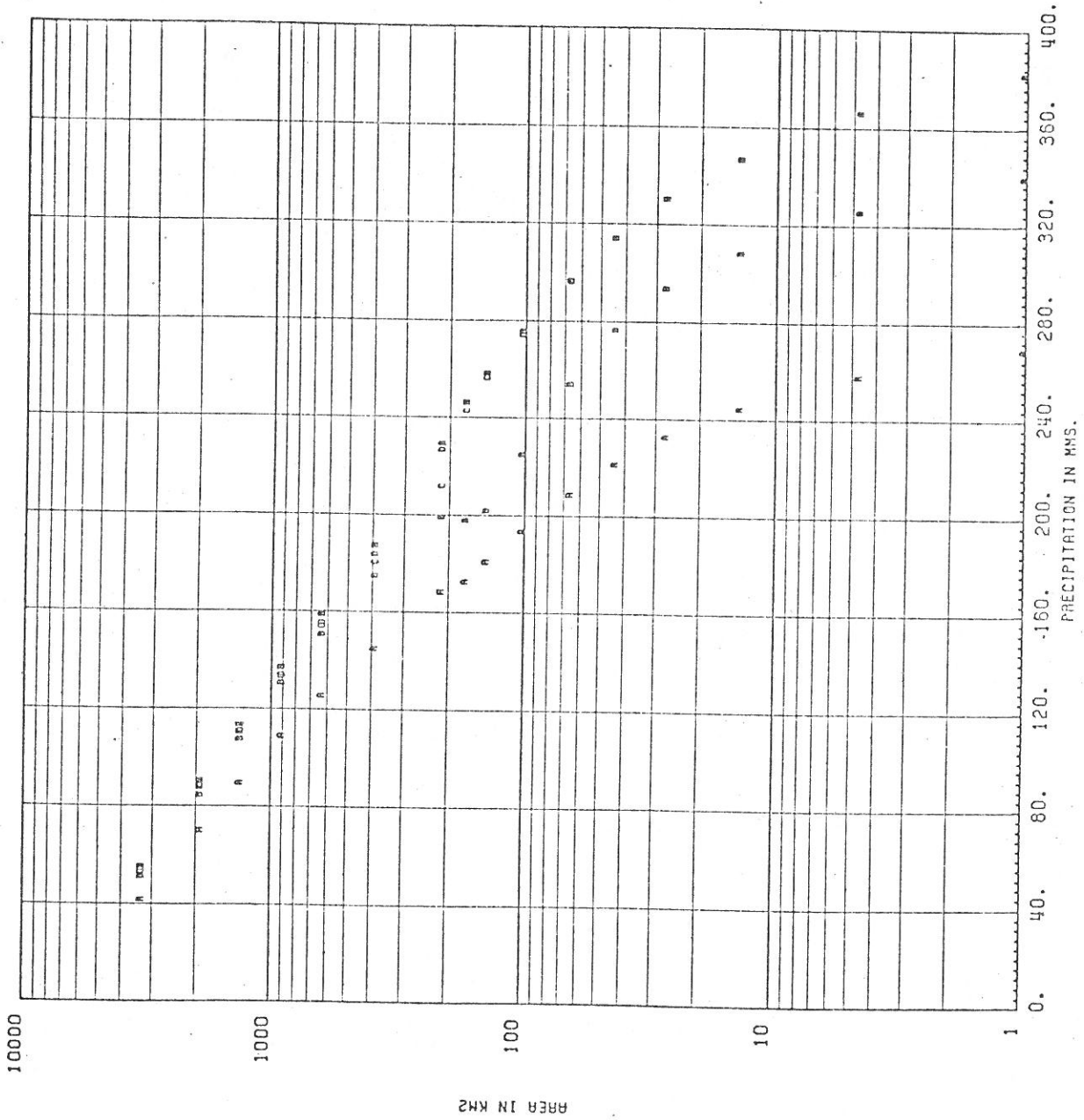
DEPTH AREA DURATION STORM STARTING 73 OCT 14 8 ENDING 73 OCT 21 8

Figure 1.6.C.12



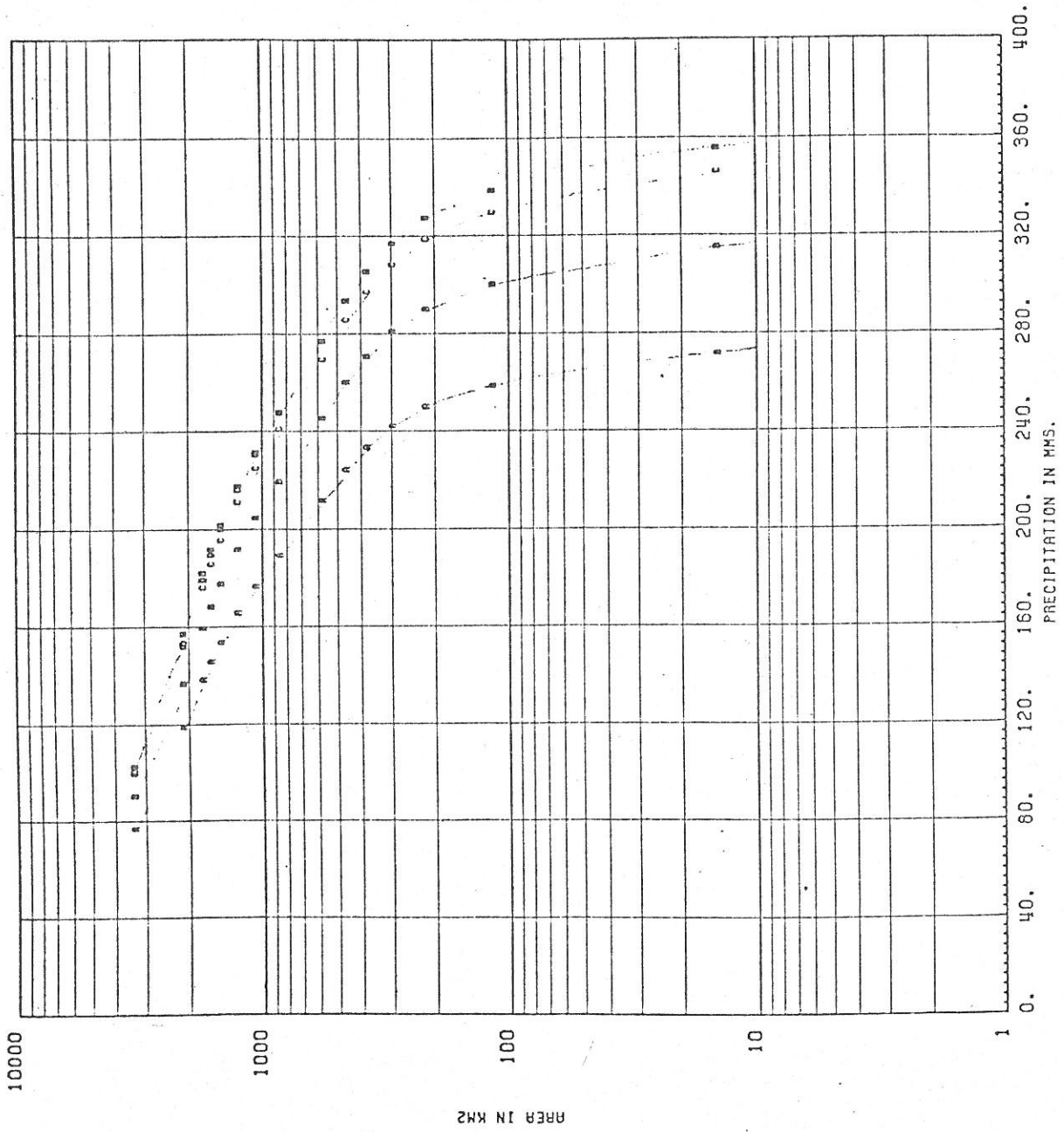
DEPTH AREA DURATION STORM STARTING 74 AGO 30 8 ENDING 74 AGO 31 8

Figure 1.6.C.13.



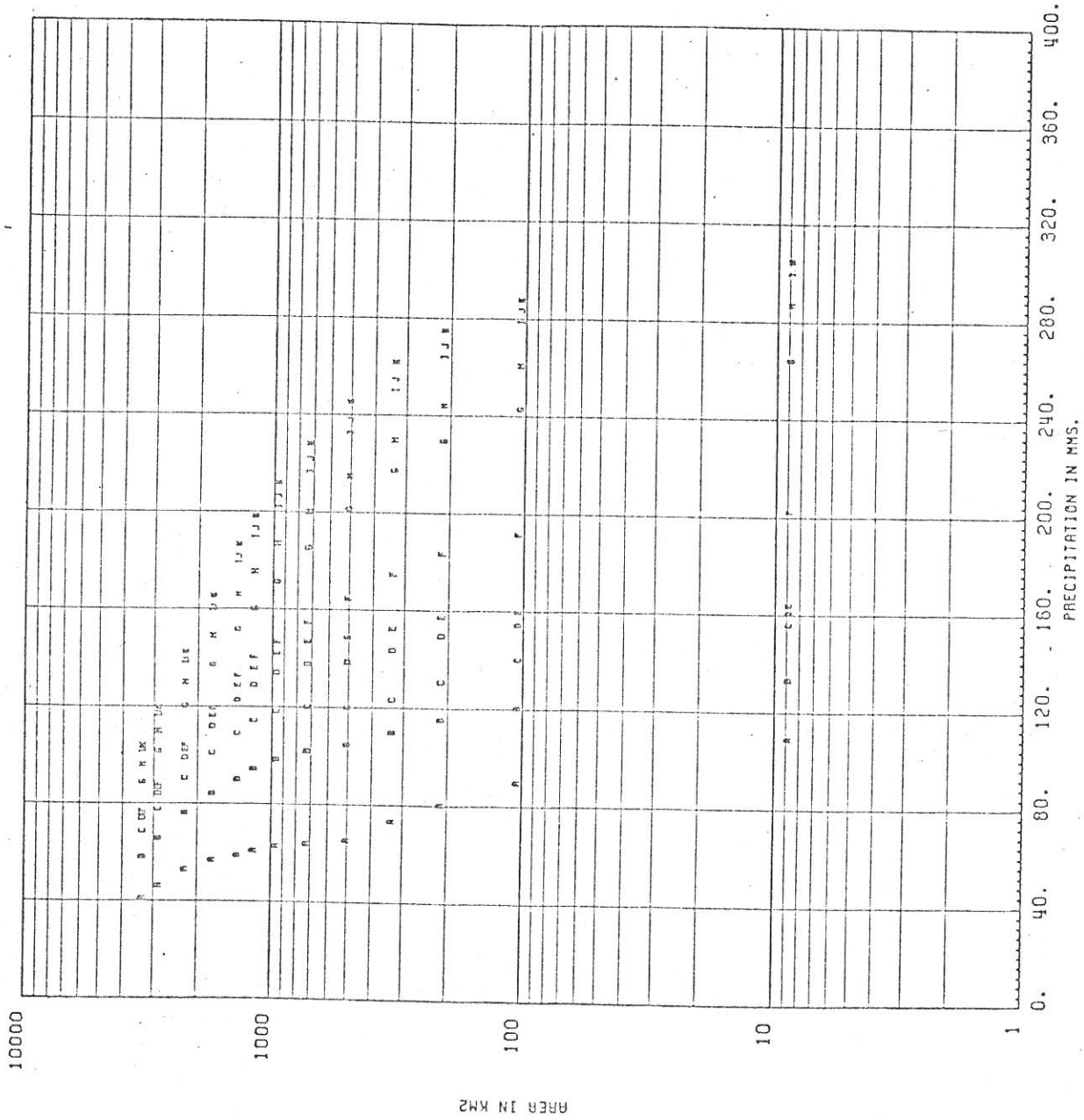
DEPTH AREA DURATION STORM STARTING 75 SEP 16 8 ENDING 75 SEP 18 8

Figure 1.6.C.14



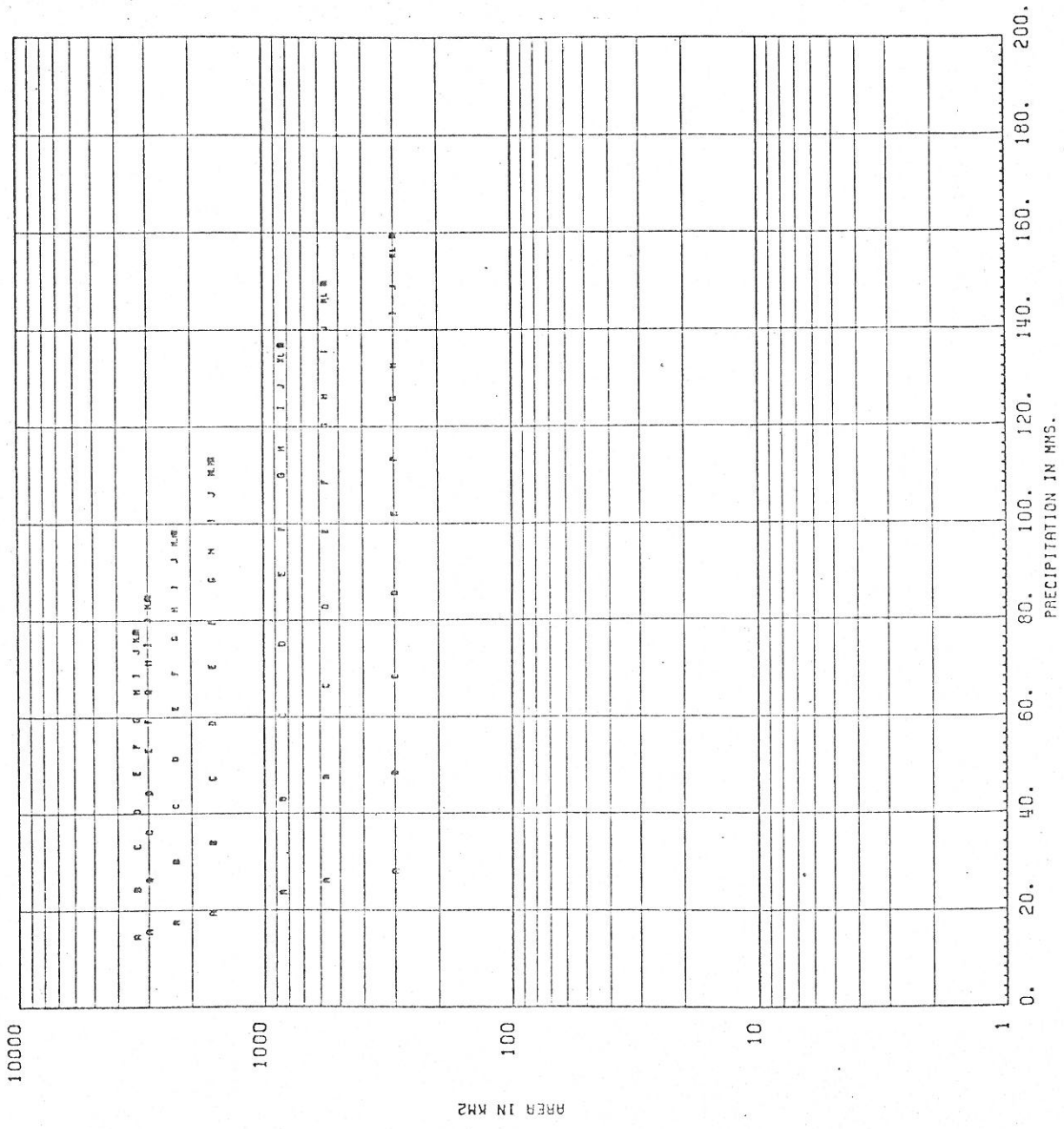
DEPTH AREA DURATION STORM STARTING 76 OCT 10 8 ENDING 76 OCT 12 8

Figure 1.6.C.15



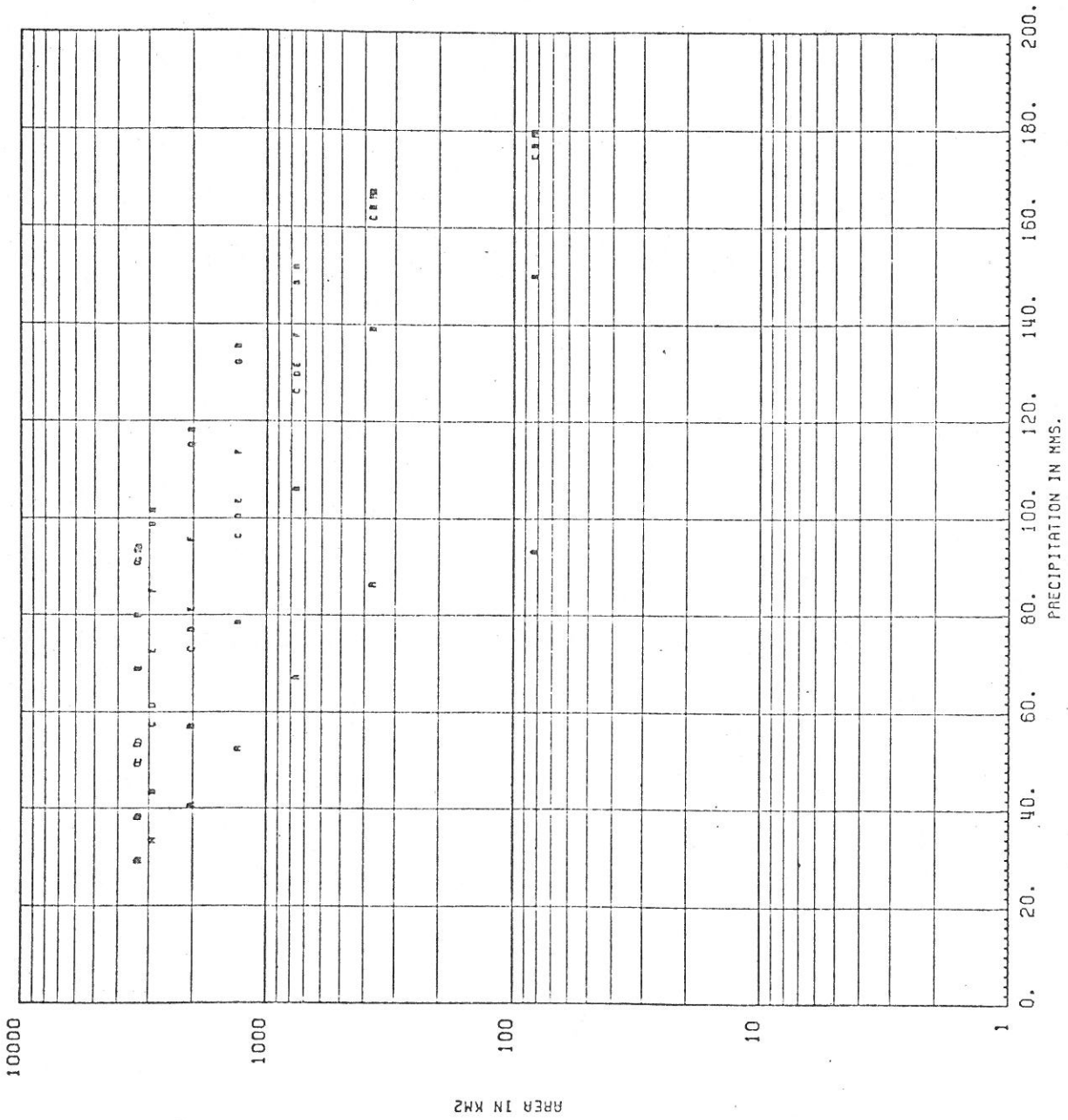
DEPTH AREA DURATION STORM STARTING 77 MAY 21 8 ENDING 77 MAY 24 8

Figure 1.6.C.16



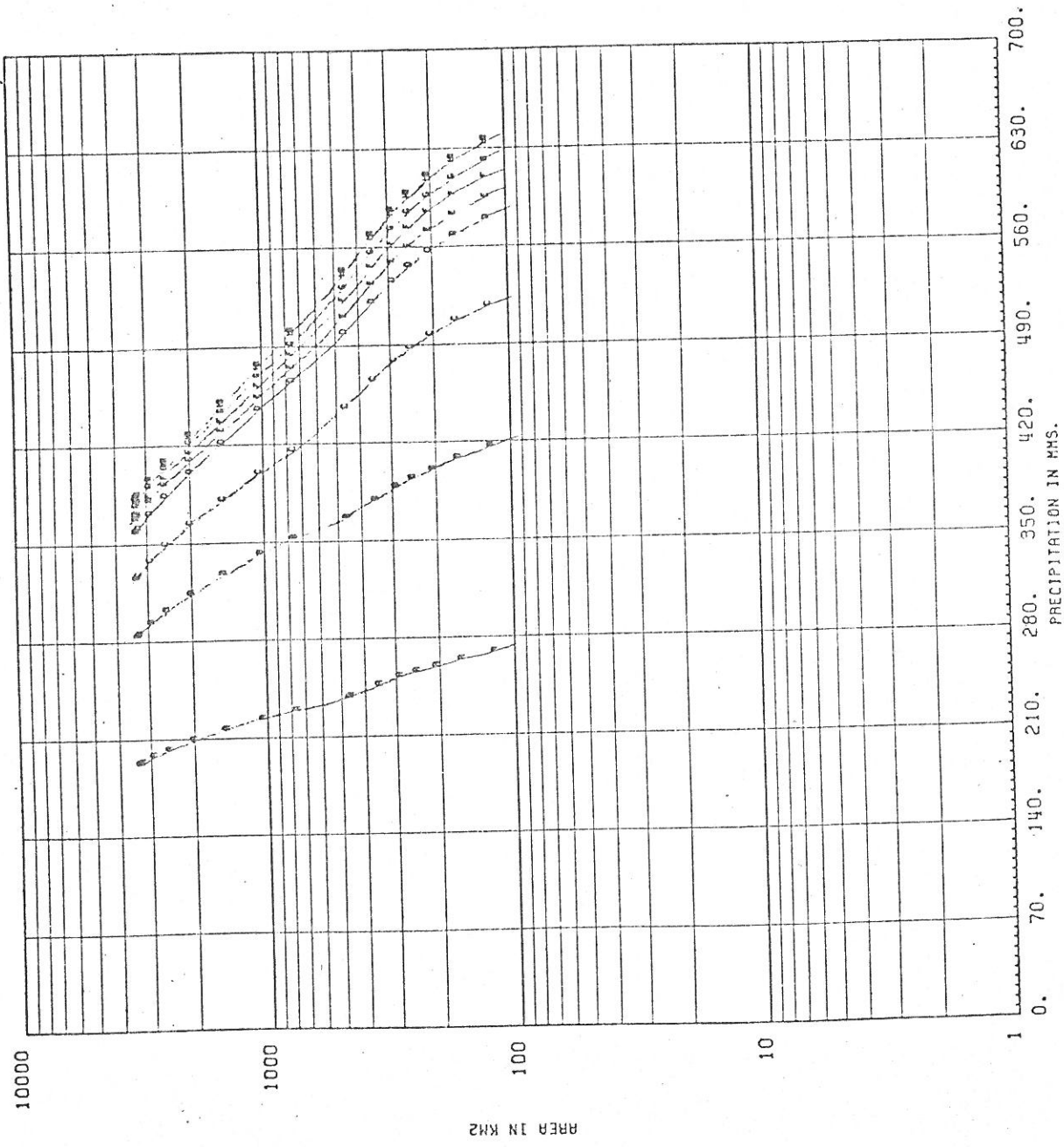
DEPTH AREA DURATION STORM STARTING 77 DEC 28 8 ENDING 78 ENE 1 8

Figure 1.6.C.17



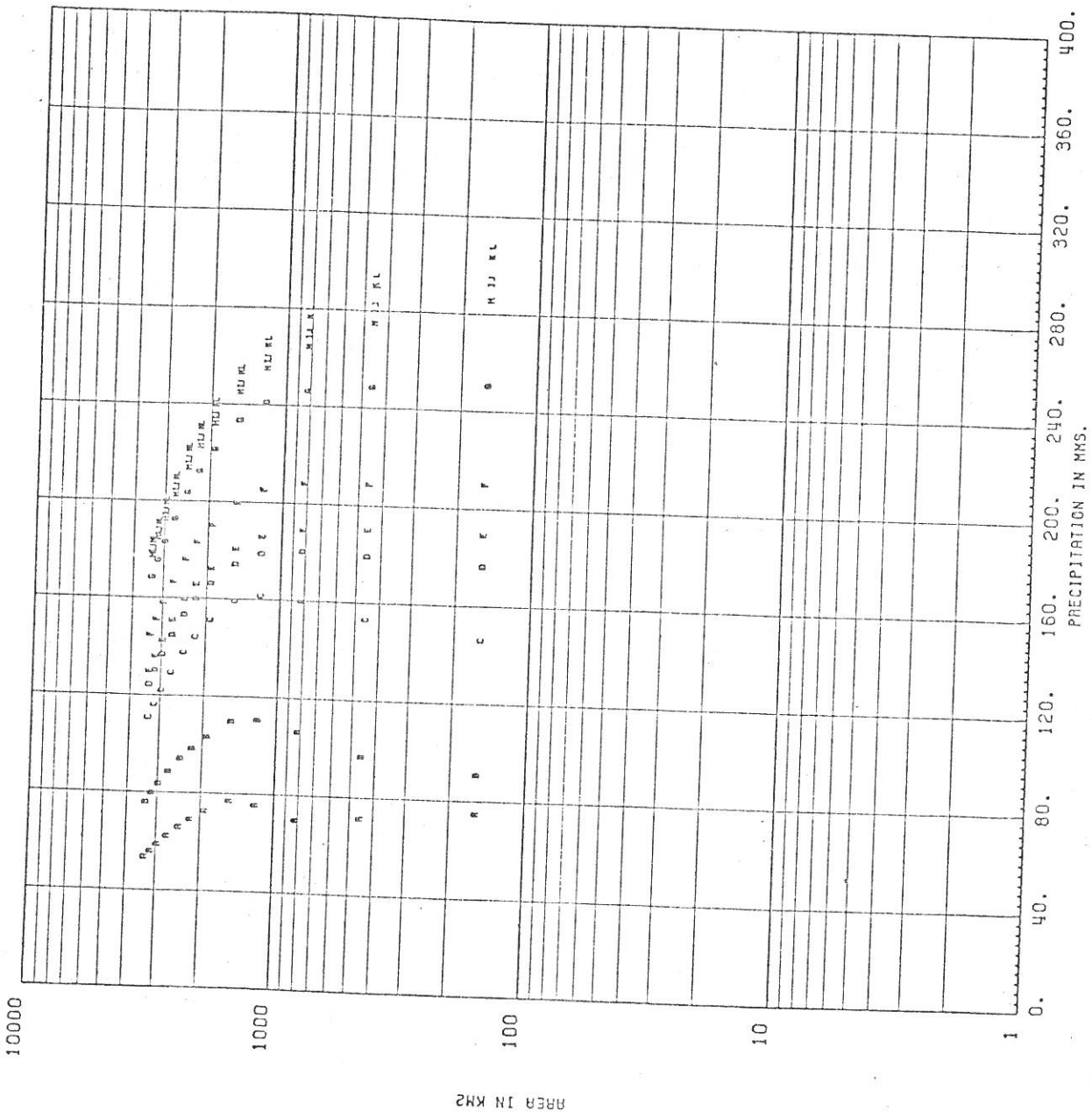
DEPTH AREA DURATION STORM STARTING 78 AGO 3 8 ENDING 78 AGO 6 8

Figure 1.6.C.18



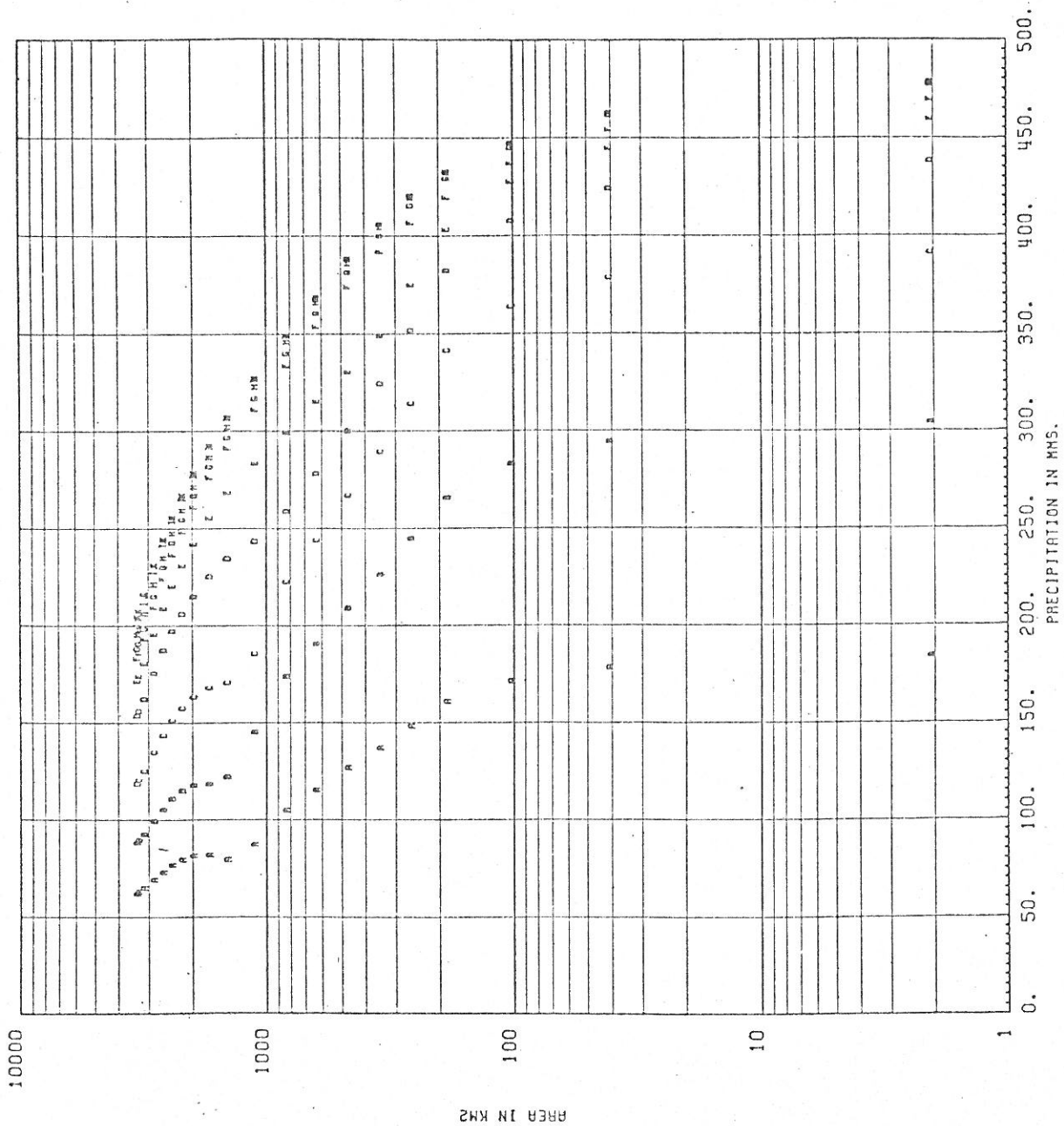
DEPTH AREA DURATION STORM STARTING 79 AGO 30 8 ENDING 79 SEP 2 8

Figure 1.6.C.19



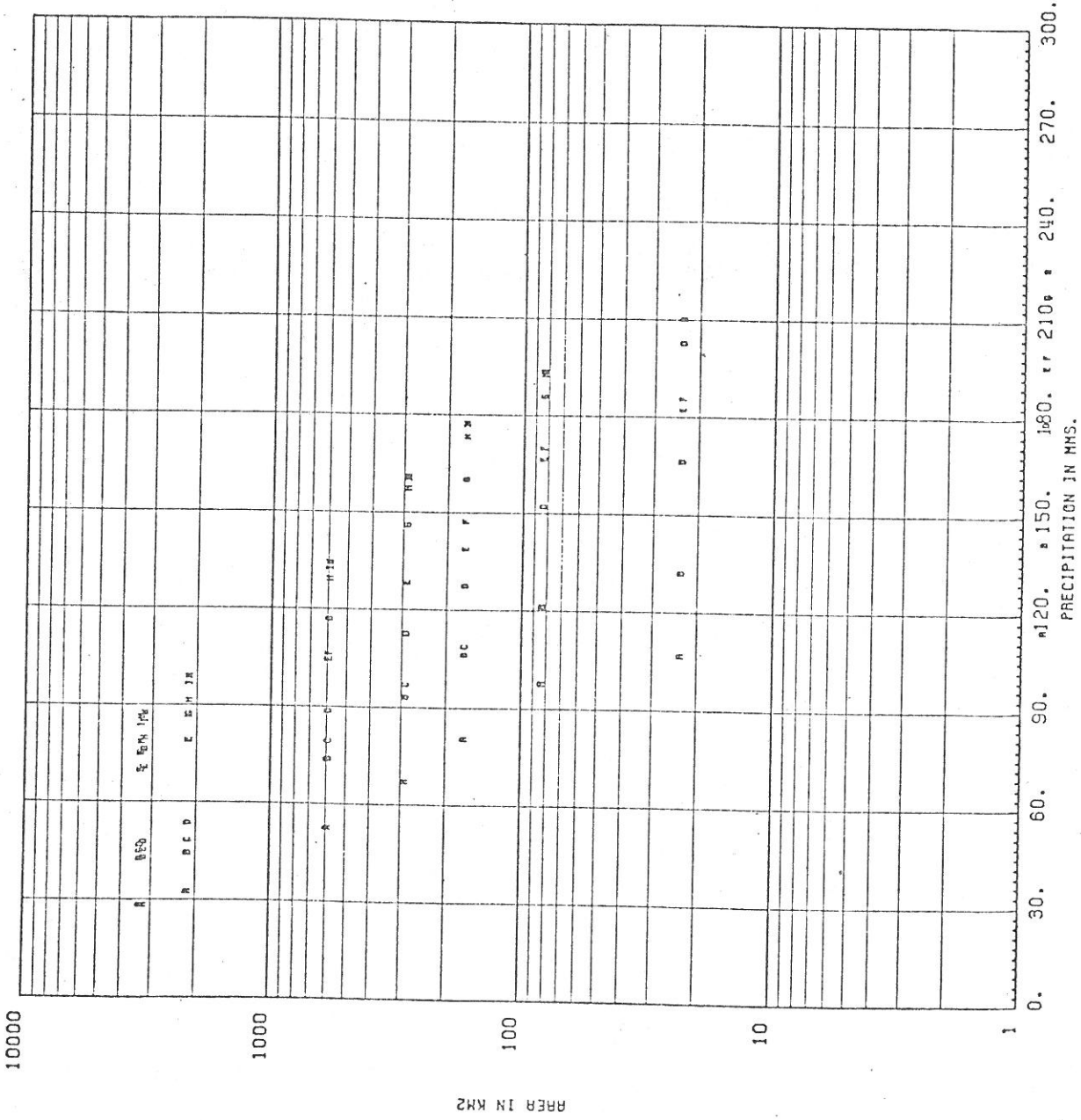
DEPTH AREA DURATION STORM STARTING 79 SEP 5 8 ENDING 79 SEP 8 8

Figure 1.6.C.20



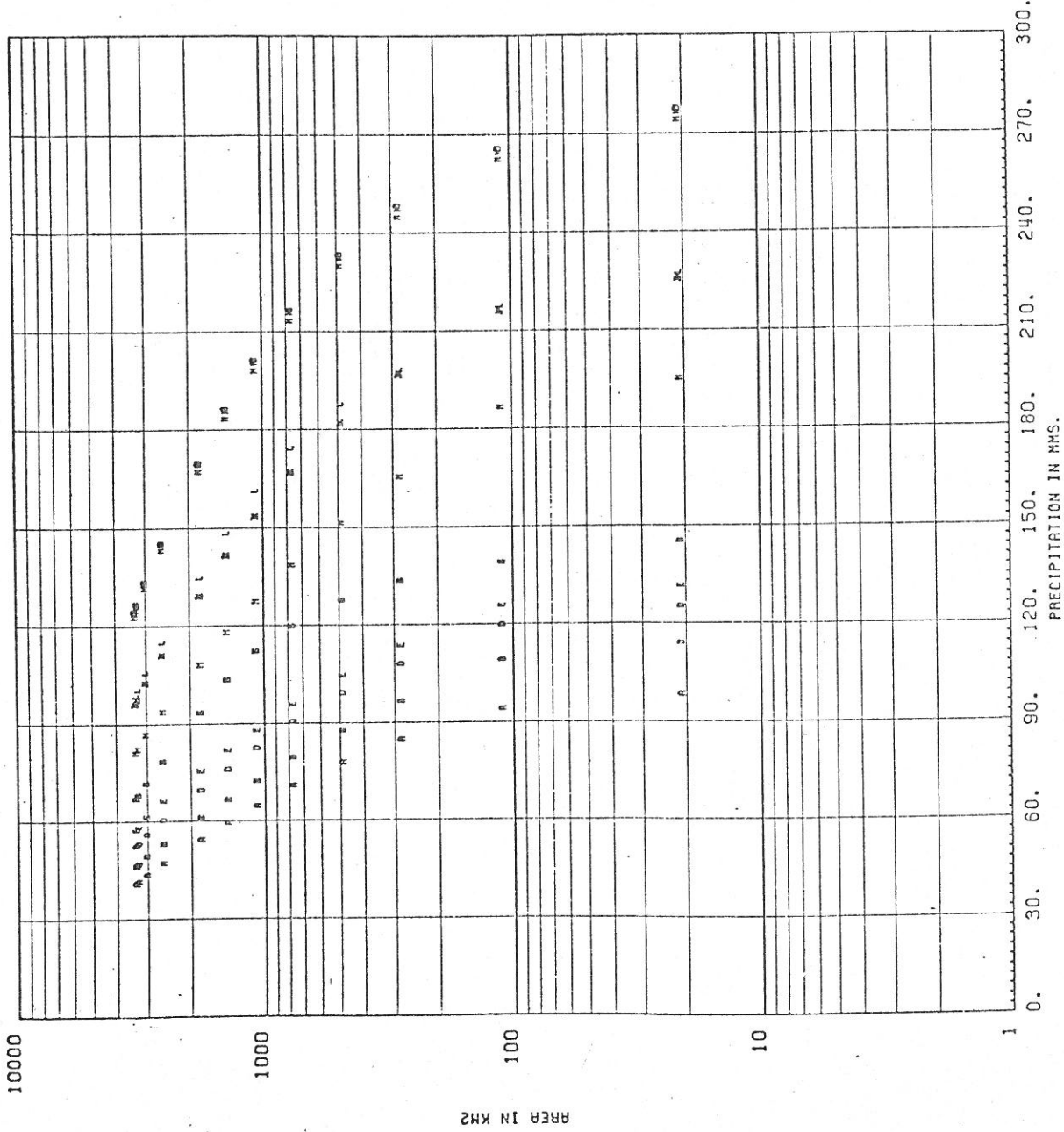
DEPTH AREA DURATION STORM STARTING 80 AGO 4 8 ENDING 80 AGO 7 8

Figure I.6.C.21



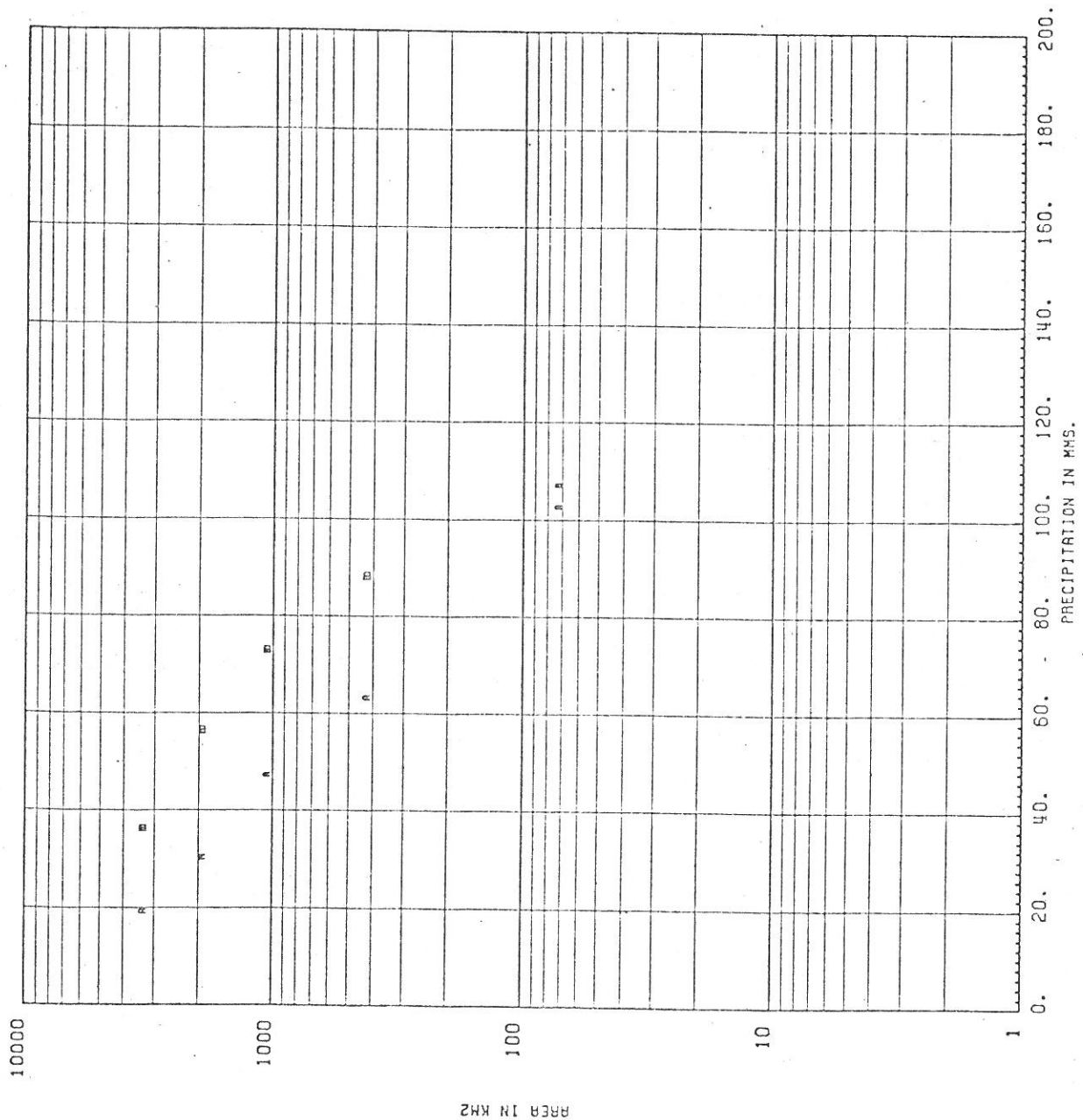
DEPTH AREA DURATION STORM STARTING 81 MAY 8 8 ENDING 81 MAY 11 8

Figure 1.6.C.22



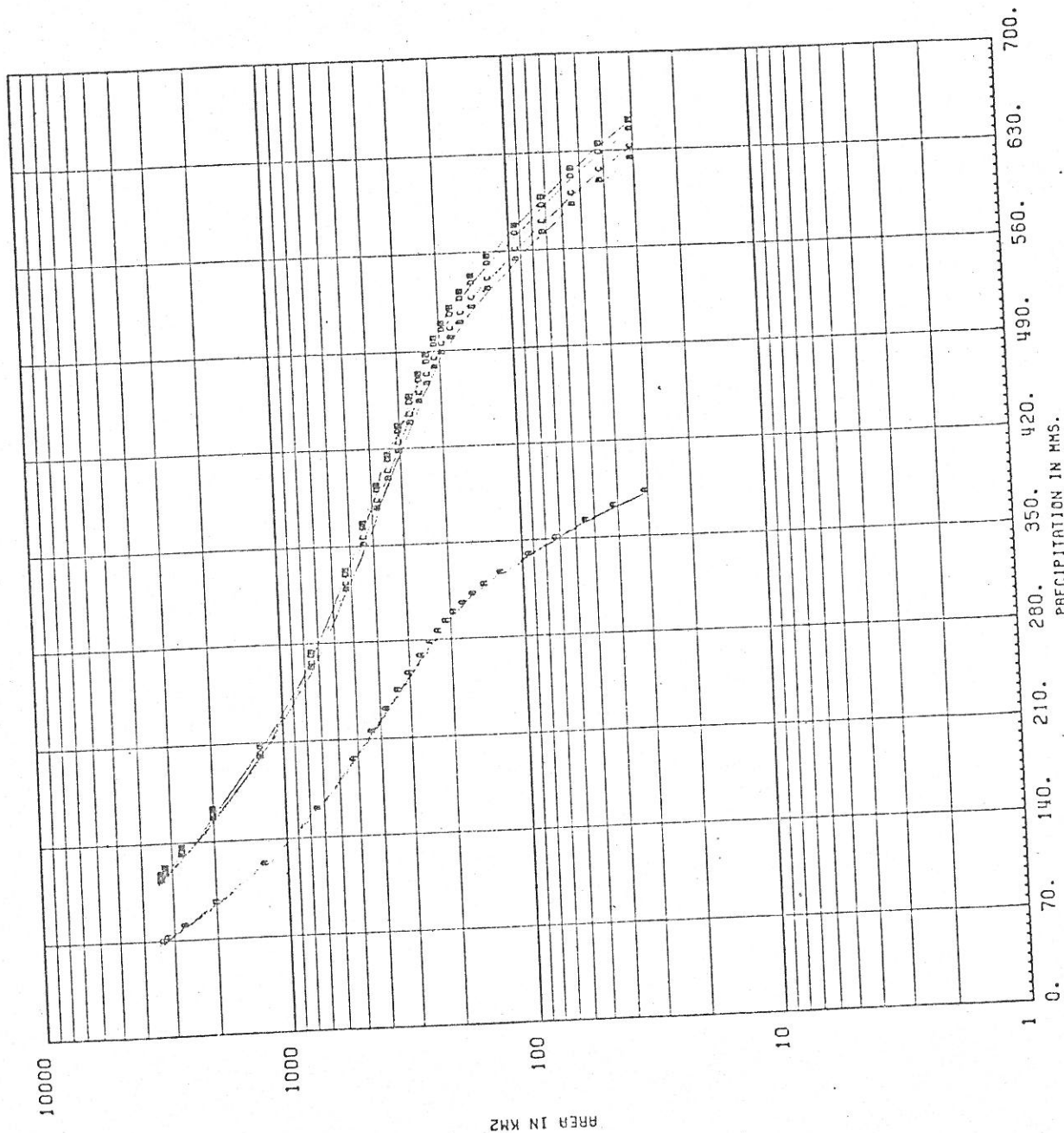
DEPTH AREA DURATION STORM STARTING 82 MAY 9 8 ENDING 82 MAY 13 8

Figure 1.6.C.23



DEPTH AREA DURATION STORM STARTING 83 ABR 12 8 ENDING 83 ABR 13 8

Figure 1.6.C.24



DEPTH AREA DURATION STORM STARTING 84 AGO 1 8 ENDING 84 AGO 3 8

Figure 1.6.C.25

1.7 Rainfall-Runoff Modeling

1.7.1 Selection of Model

Several existing simulation models were considered for various applications in this project. In particular, a close scrutiny was exercised for models Flood Hydrograph Package (HEC-1) of U.S. Army Corps of Engineers, National Weather Service River Forecast System (NWSRFS), Hydrologic Simulation Program-Fortran (HSPF), and Simulation of Flood Control and Conservation Systems (HEC5). Most of these models obtained by Colorado State University require a large computer memory and storage and they cannot be implemented readily in the IBM System 34 or in the personal computers available at INDRHI or CDE. Moreover, the available data for Nizao basin do not satisfy the requirements of some of the models.

Since flood hydrographs entering Valdesia reservoir are to be simulated from historic and hypothetical storms, an event type rainfall-runoff model is needed. From all the event models considered, the HEC-1 was chosen based on the data availability and ease of implementation. A version of HEC-1 which fits into an IBM personal computer (PC) has been obtained and installed in the IBM-PC at INDRHI.

1.7.2 HEC-1 Model Calibration

Clark Unit Hydrograph and Muskingum Routing: The HEC-1 rainfall-runoff model has many options for computing rainfall excess, watershed routing, and channel routing. Many of these options can be used with the optimization model which computes the "best" parameters on basis of "best fit" of the observed and computed flood hydrographs. A common approach is to use the HEC-1 exponential loss rate function with the Clark unit hydrograph to compute subbasin flood hydrographs and then to

use Muskingum method to route flood hydrographs to the basin outlet. Naturally, this was our initial approach to calibrate HEC-1 for Nizao watershed. Unfortunately, the results obtained by this method were unacceptable. The poor quality of rainfall and runoff data used in our optimization runs was partly responsible for this poor initial calibration results. The variation of parameters from one storm event to another was too large to be acceptable (see Table 1.7.1).

The particular topography of Nizao basin (high slopes with no significant flood plains) is thought to be another reason for the failure of storage routing techniques such as Clark Unit hydrograph and Muskingum method. Consequently, the kinematic wave approach of overland flow routing and channel routing was deemed to be more appropriate for rainfall-runoff modeling in Nizao basin.

Kinematic wave model: The first step taken in the Kinematic Wave Model formulation was to subdivide the basin into 10 subbasins, in an effort to simulate each important tributary of the Nizao basin independently. Also, one additional subdivision was included at gaging station La Estrechura, to be able to compare simulated and observed hydrographs at that point. Other two subbasin limits coincide with gaging stations Palo de Caja and Paso del Ermitano. The general configuration is shown in Figure 1.7.1.

The physical characteristics of the basin, such as channel lengths, slopes, widths, areas, etc., as well as the land use was estimated from available maps of the area. Mannings's N was obtained by calibration.

The loss rate method that was chosen was the SCS (Soil Conservation Service) loss rate function. Since very little data on soil type was available it was judged convenient to reduce the number of unknown

TABLE 1.7.1 INITIAL HEC-1 CALIBRATION FOR EXPONENTIAL LOSS RATE
AND CLARK UNIT HYDROGRAPH PARAMETERS

*	1st RUN	STRKR	DLTKR	RTIOL	TC	R	ERAIN
	Storm E	12.76	24.82	7.99	11.03	11.25	.44
	Storm F	14.69	95.04	2.76	1.09	12.35	.13
	Storm H	4.05	11.46	1.90	9.12	12.64	.81
	Storm M	19.53	96.78	4.33	16.83	18.09	.28
	Storm N	1.90	7.37	1.26	17.01	18.34	.10
**	2nd RUN						
	Storm E	13.70	26.89	1.98	11.04	10.80	.44
	Storm F	24.00	94.37	2.91	7.03	5.71	.22
	Storm H	4.53	12.59	1.98	6.00	18.01	.50
	Storm M	6.96	88.38	2.15	11.26	10.47	.13
	Storm N	3.20	11.04	1.50	16.83	18.08	.10
***	Storm E	10.75	5.89	1.39	1.36	32.06	.08
	Storm F	3.68	93.94	11.52	4.66	14.73	.06
	Storm H	4.26	10.59	1.73	10.12	12.74	.38
	Storm M	9.66	86.96	2.14	11.24	10.28	.13
	Storm N	6.03	16.77	1.67	11.55	13.91	.31

* Using all precipitation stations (recording and non-recording)
all hydrograph ordinates

** Range of optimization limited to period around peak outflow

*** Inconsistent precipitation data removed from record.

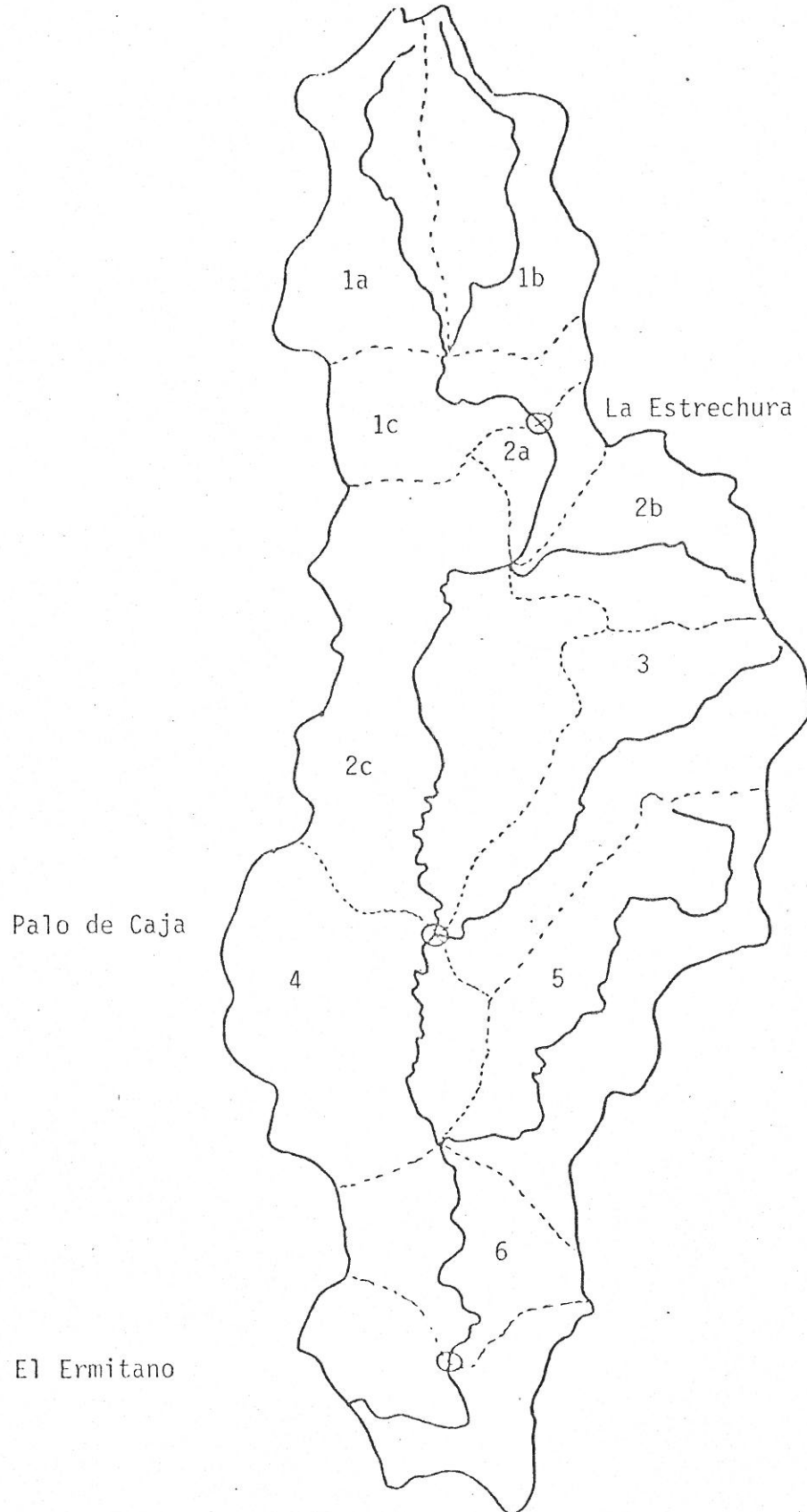


Figure 1.7.1. Nizao river basin configuration used in kinematic wave routing model.

parameters in the model to a minimum, and the SCS method only uses the Curve Number as input (once we assume that the initial abstraction is 20% of the storage capacity). The SCS method has proven to give very good results in similar situations.

The selection of the storms used for the calibration was a little difficult. After the revision of the stage-discharge relations and the computation of the new hydrographs, very carefully the plots of precipitation and runoff of all the available storms were inspected. Most of them showed dramatic inconsistencies in terms of timing and volume of water. In many cases the hydrographs show a peak before the stations show any record of precipitation. This is why only four storms were selected for calibration. The corresponding hydrographs are shown in Figures 1.7.2 through 1.7.5.

The creation of each storm file starts by assigning weights to each of the precipitation stations with available data for each storm. The HEC-1 model computes an average total storm depth based on given weights and storm values for each precipitation station. Then an average time distribution is computed based on given weights and precipitation patterns of the recording stations. The weights were assigned by using program PCMAPS to map the Thiessen polygons which were superimposed on the basin subdivision. In this way, each subbasin has an individual average storm depth and an individual time distribution pattern. The weights assigned to each station for each subbasin and storm are included in the HEC-1 input files printed as part of the output. A computer output of the Thiessen polygons is included in Appendix 1.7.A.

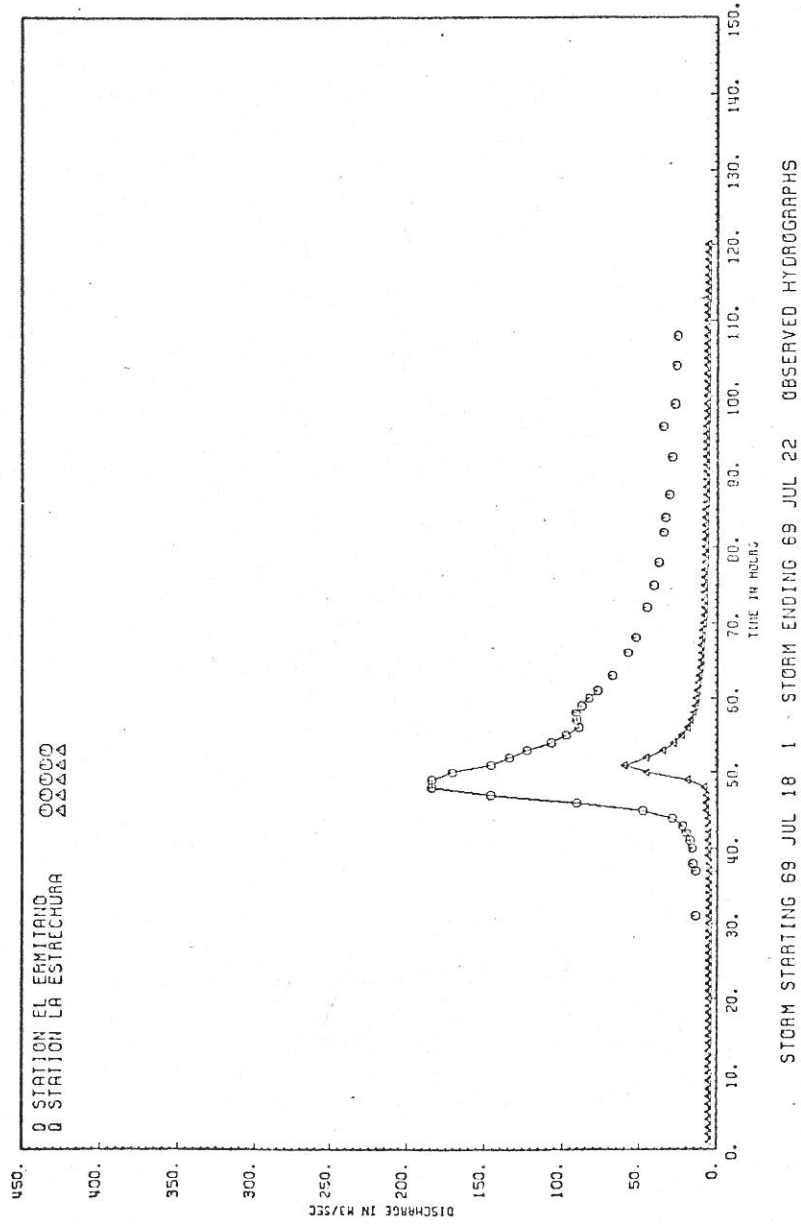


Figure 1.7.2. Streamflow hydrographs for period July 18-22, 1969 (storm A).

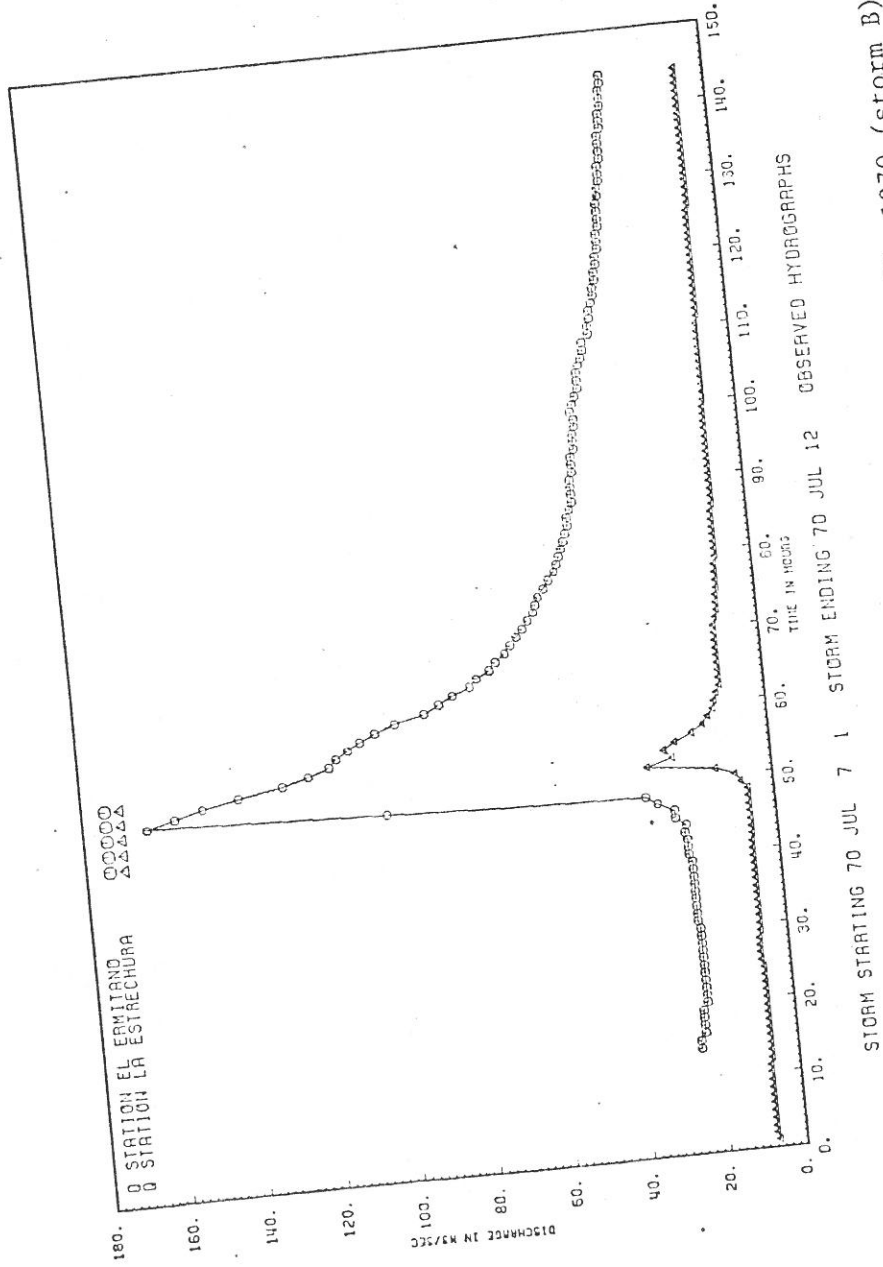


Figure 1.7.3. Streamflow hydrographs for period July 1-12, 1970 (storm B).

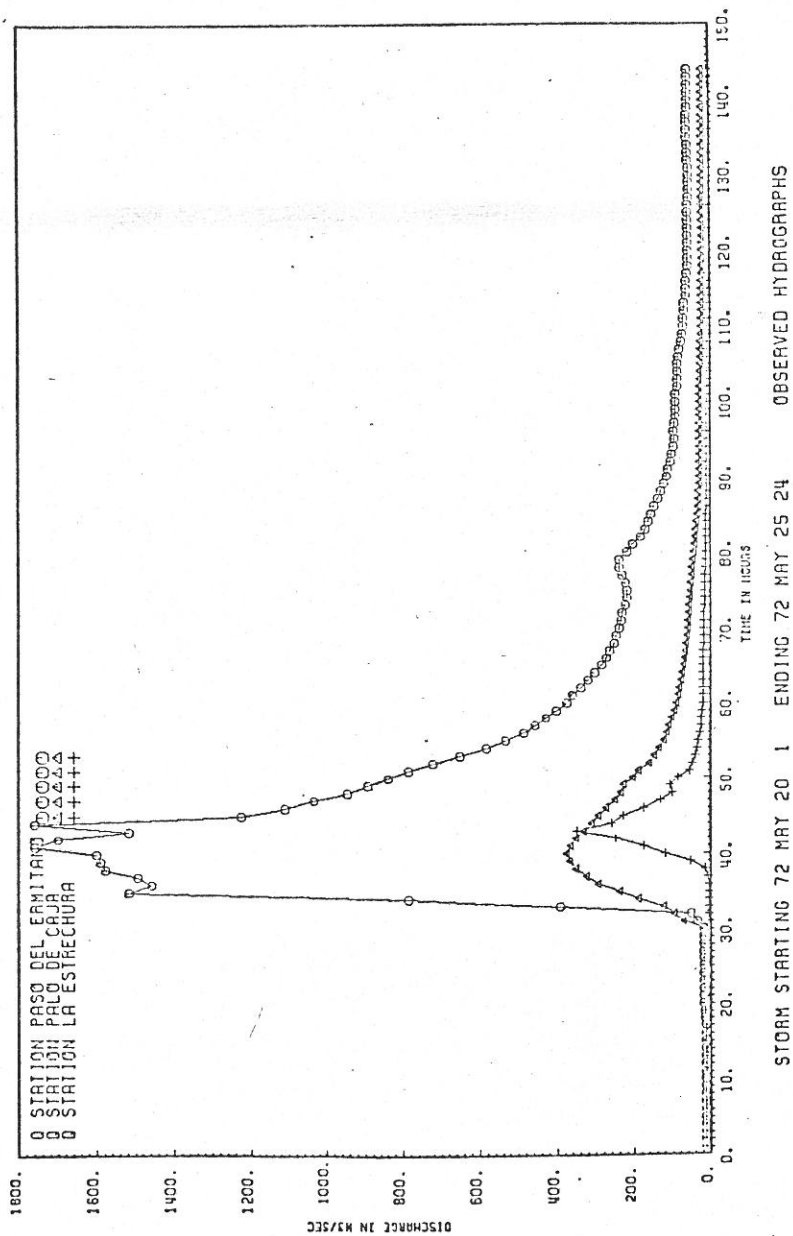


Figure 1.7.4. Streamflow hydrographs for period May 20-25, 1972 (storm F).

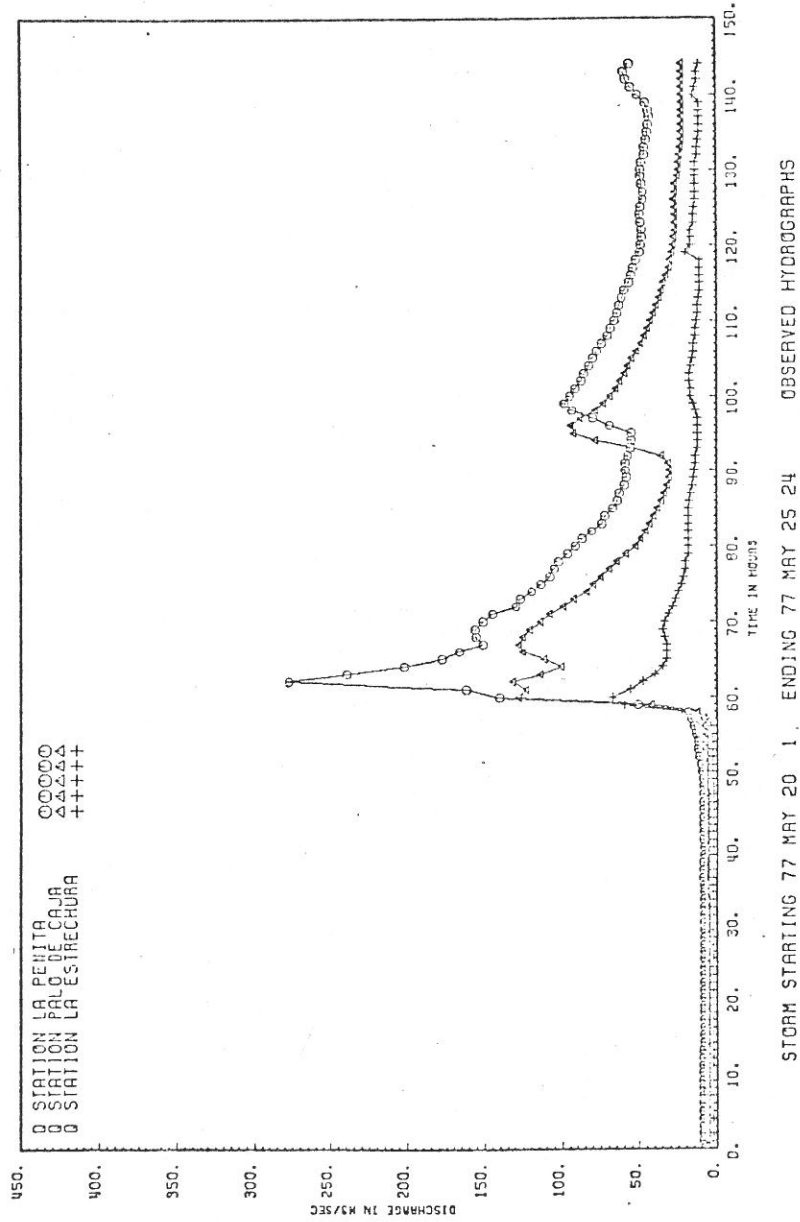


Figure 1.7.5. Streamflow hydrographs for period May 20-25, 1977 (storm M).

The calibration procedure starts by trying to adjust the volume under the computed hydrograph to the volume under the observed hydrograph. In most of the cases the curve number fall in the range 60 to 80 which is reasonable for the vegetation cover and expected soil type in the area. Some extremes, as low as 45 and high as 100 are observed in case of storms F and M. The low of 45 could be reasonable for a very dry antecedent condition but in the case of CN = 100 (see storm F and La Penita), i.e. with practically no infiltration, in the lower subbasins there is a deficit of 45% of the observed volume. It is concluded that some errors are present in the data.

Once the runoff volumes are matched, next step is to change the Manning's N, within reasonable values, to try to adjust the timing of the hydrographs. The best set of parameters so obtained is shown in Tables 1.7.2 and 1.7.3. Unfortunately, the computer outputs showed that the match is not too good in most of the cases. Since, in some cases, the observed peak comes before the computed, and in other cases is the opposite, it was not possible to arrive at a unique set of parameters that will fit all the storms. Suspecting errors in data, additional data consisting of reservoir levels and operation records of the Valdesia reservoir, during certain major storm events were requested. The data for two reasonably large storms in September and October 1985 were received which proved to be very useful. The calibration results for the September storm are illustrated in Figure 1.7.6. The timing of the reconstructed inflow hydrographs to the reservoir matches reasonably well those computed with the calibrated model.

TABLE 1.7.2 . KINEMATIC WAVE MODEL SUBBASIN CHARACTERISTICS

SUBBASIN	OVERLAND FLOW PLANE LENGTH	OVERLAND FLOW PLANE SLOPE	MANNING N	DRAINAGE AREA (sq km)	LAND USE
1a	2500	0.60	0.400	70	Forest
1b	2500	0.60	0.400	70	Forest
1c	4500	0.50	0.300	45	Pasture
2a	2000	0.40	0.300	24	Pasture/Forest
2b	3500	0.20	0.300	56	Agriculture/Forest
2c	4500	0.40	0.300	164	Pasture/Forest
3	2000	0.40	0.300	106	Pasture/Forest
4	5000	0.50	0.300	103	Pasture/Forest
5	3000	0.40	0.300	109	Forest
6	5000	0.60	0.300	53	Forest

TABLE 1.7.3 KINEMATIC WAVE MODEL CHANNEL CHARACTERISTICS

SUBBASIN	LENGTH (m)	SLOPE (m/m)	MANNING N	WIDTH	SIDE SLOPE (m/m)	UPSTREAM INFLOW
1a	16000	0.1050	0.060	45	1	no
1b	20000	0.0837	0.060	45	1	no
1c	8500	0.0133	0.040	85	1	yes
2a	7500	0.0108	0.040	85	1	yes
2b	13500	0.0452	0.050	45	1	no
2c	33500	0.0086	0.040	70	1	yes
3	33000	0.0293	0.050	45	1	no
4	18000	0.0148	0.040	80	1	yes
5	34000	0.0323	0.050	45	1	no
6	18000	0.0119	0.040	80	1	yes

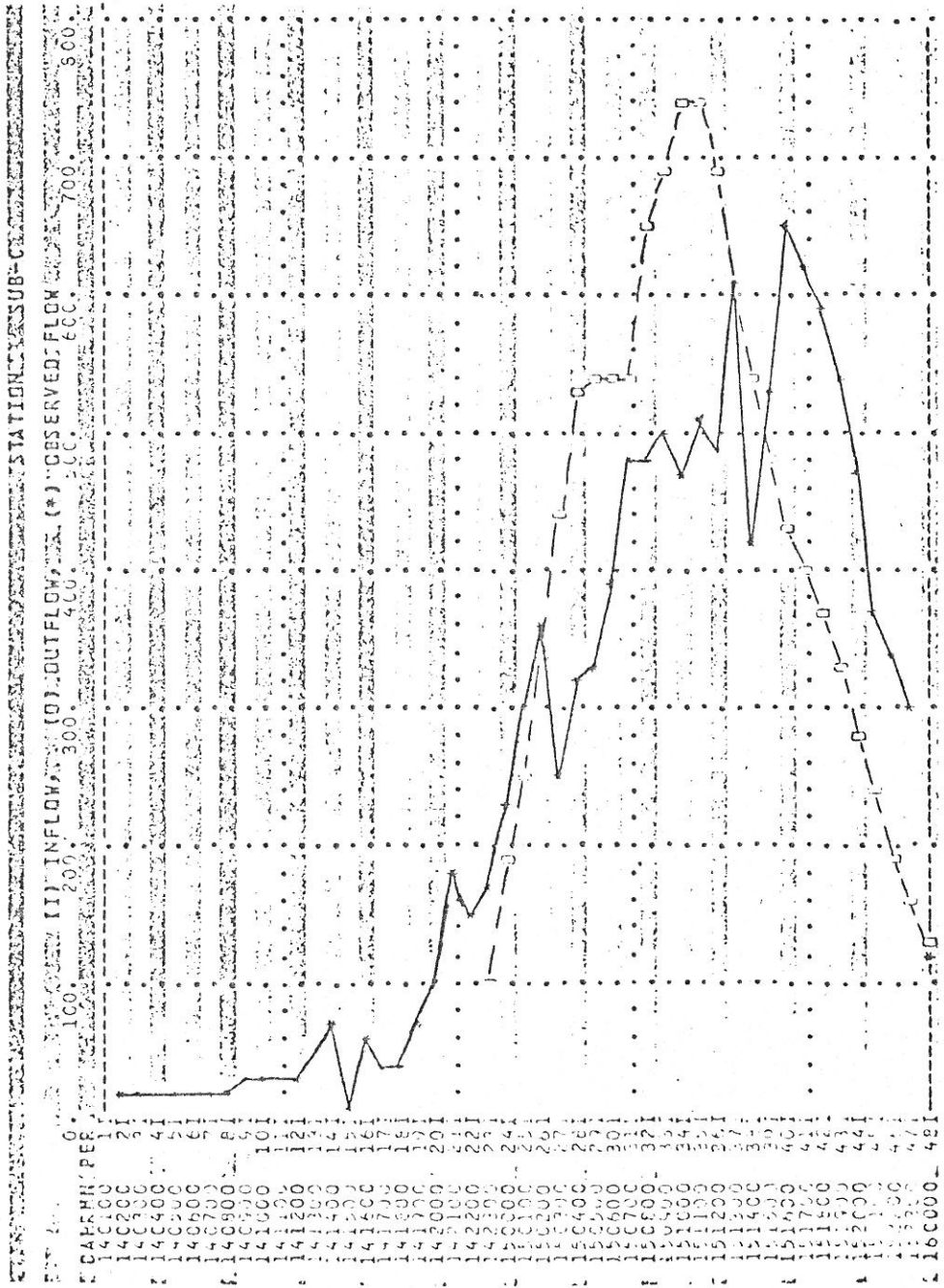


Figure 1.7.6. Comparison of computed (from reservoir levels) and calibrated (HEC-1) inflow hydrographs to Valdesia Reservoir due to storm on September 13-14, 1985.

1.7.3 Design Flood Hydrographs

The primary use of the calibrated HEC-1 model is in computing design flood hydrographs from hypothetical design storm events. The hypothetical storms under consideration are (a) Standard Project storm for non-hurricane conditions; (b) Standard Project storm for hurricane conditions; and (c) probable maximum precipitation. The calibrated HEC-1 model was used to compute hydrographs corresponding to all three design storms.

An important consideration in computing design flood hydrographs is the antecedent moisture condition of the basin under consideration. Since the SCS curve number method of HEC-1 model is being used for computing losses, the antecedent basin condition is reflected in the curve numbers corresponding to each subbasin. The actual curve number magnitudes depend on three types of basin conditions: (a) Antecedent Moisture Condition type-I (AMC-I). This corresponds to the dry soil conditions; (b) Antecedent Moisture Condition type-II (AMC II). This corresponds to average soil moisture conditions; and (c) Antecedent Moisture Condition type-III (AMC-III). This corresponds to the nearly saturated soil condition. The soil conservation service provides the Curve Numbers for AMC II conditions for a variety of land uses and four types of soil cover complexes. Based on a past study on soil types and land uses in Nizao basin (Perez, 1982), the curve numbers were identified for various subbasins in the watershed subdivision made for kinematic wave model. These curve numbers correspond to AMC II. The curve numbers for AMC I and AMC III were identified from another table provided by soil conservation service. These results are summarized in Table 1.7.4.

TABLE 1.7.4 CURVE NUMBERS USED IN HEC-1 MODEL

SUB-BASIN	AMC-I	AMC-II	AMC-III
1a	40	60	78
1b	40	60	78
1c	51	70	85
2a	51	70	85
2b	51	70	85
2c	45	65	82
3	45	65	82
4	45	65	82
5	45	65	82
6	40	60	78

Finally, by using the curve numbers corresponding to AMC I, AMC II and AMC III in the calibrated HEC-1 model three design flood hydrographs were computed for each of the three hypothetical design storms mentioned above. The resulting hydrographs are presented in Figures 1.7.7 through 1.7.9.

1.7.4 Reconstruction of Hydrographs for Hurricane DAVID

Hurricane DAVID constitutes the largest storm event recorded in Nizao. Unfortunately, the hydrograph corresponding to this event is not available due to failure of equipment during the hurricane. The HEC-1 model was used to reconstruct the hydrograph corresponding to recorded precipitation pattern of hurricane DAVID. Since the antecedent basin condition prior to the hurricane is unknown, the hydrograph corresponding to all three antecedent moisture conditions AMC I, AMC II, and AMC III were computed. These results are presented in Figure 1.7.10. It is noted that the peak flow corresponding to even AMC I is considerably larger than the peak discharge of about $3800 \text{ m}^3/\text{D}$ reported in some documents obtained from INDRHI. For the present study however, exact reproduction of the actual peak discharge of flood due to hurricane DAVID is not critical.

1.7.5 Effects of Natural Storages in the Watershed

It is noted here that the kinematic wave approach of HEC-1 does not have a facility to attenuate a flood hydrograph due to natural storages such as those present in flood plains. However, it is important to account for such storages wherever they exist since the hydrograph can be greatly modified by attenuation due to storage effects. A study was undertaken to investigate the presence of natural storage which may have

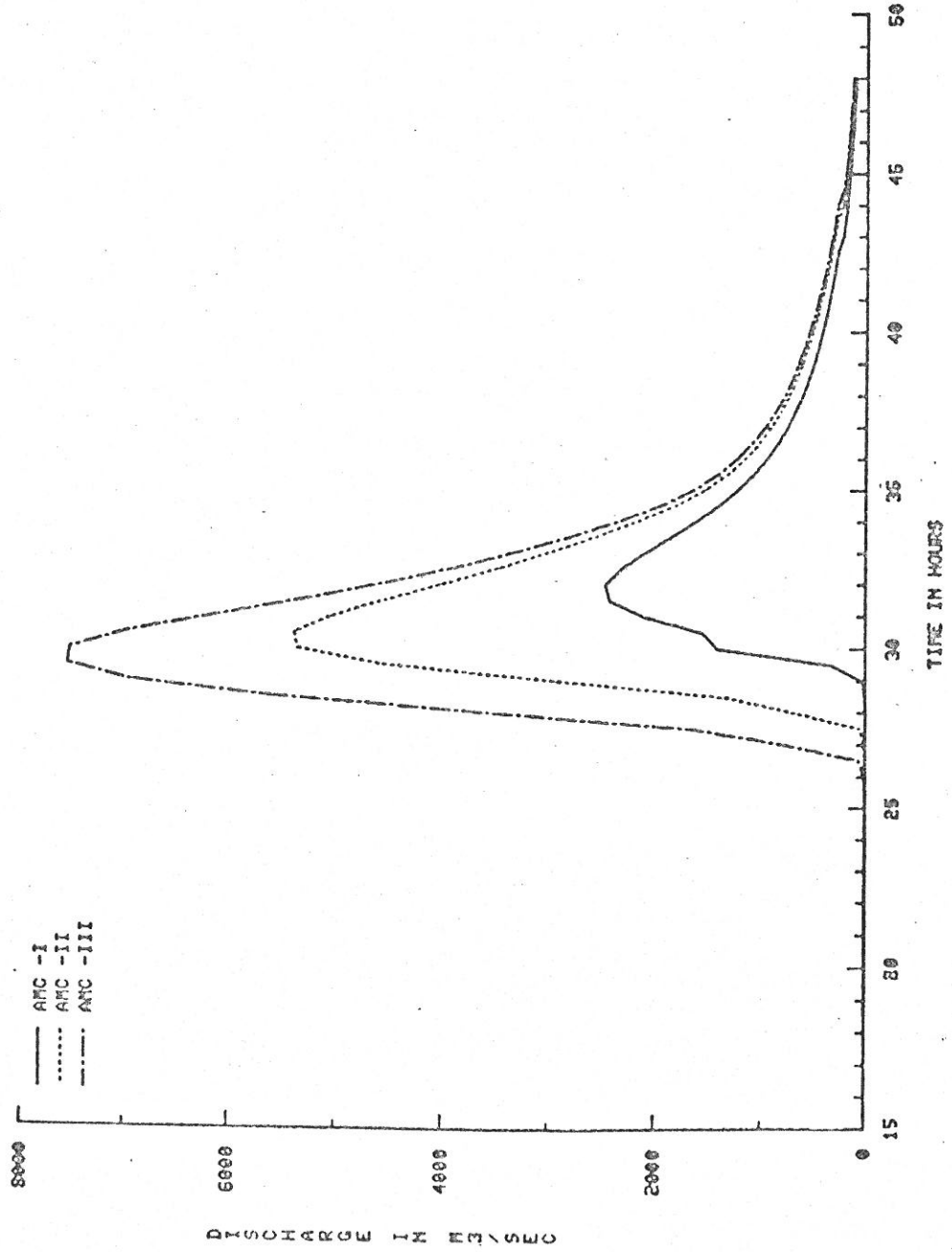


Figure 1.7.7. Non-hurricane standard project flood.

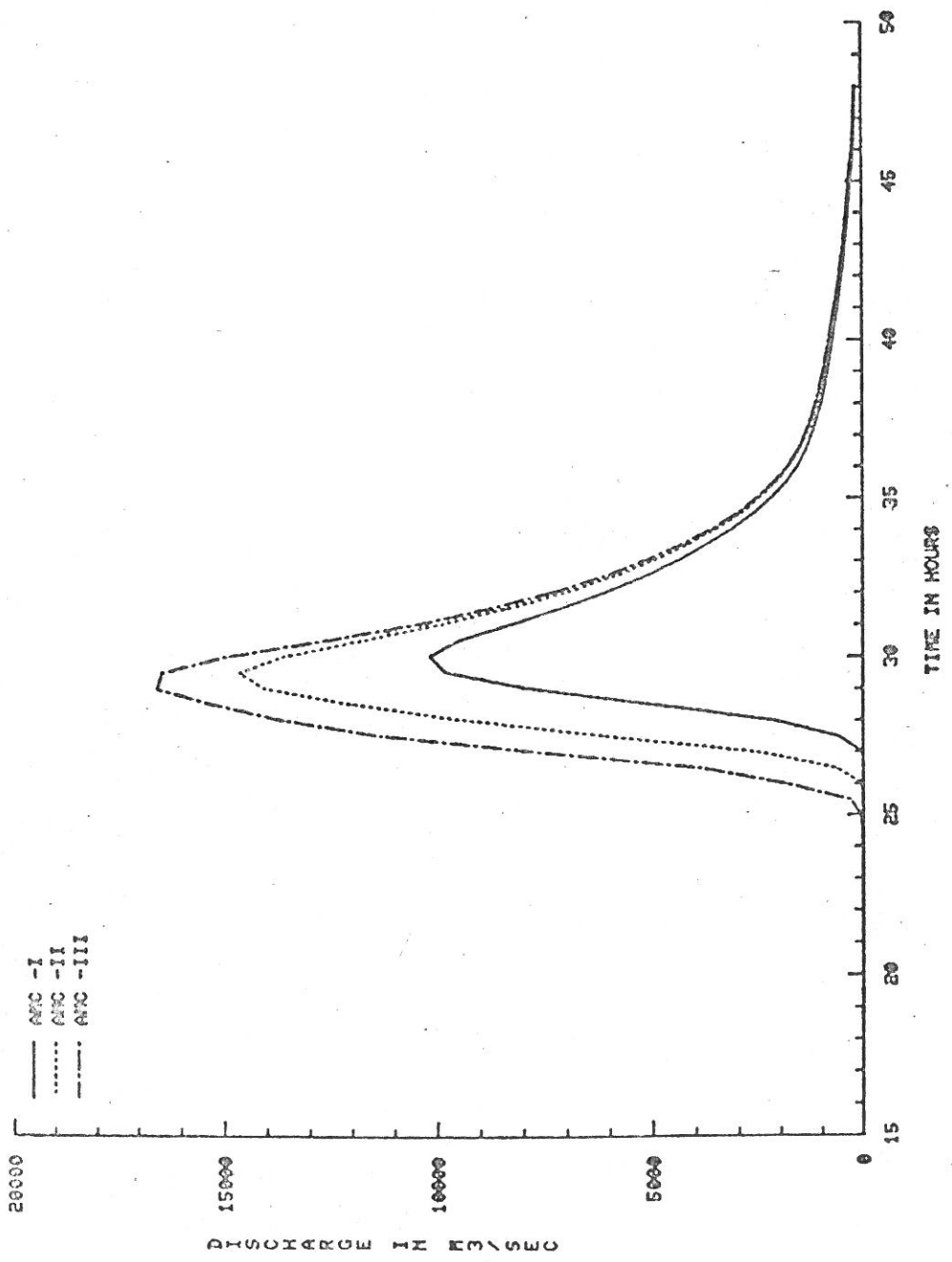


Figure 1.7.8. Hurricane standard project flood.

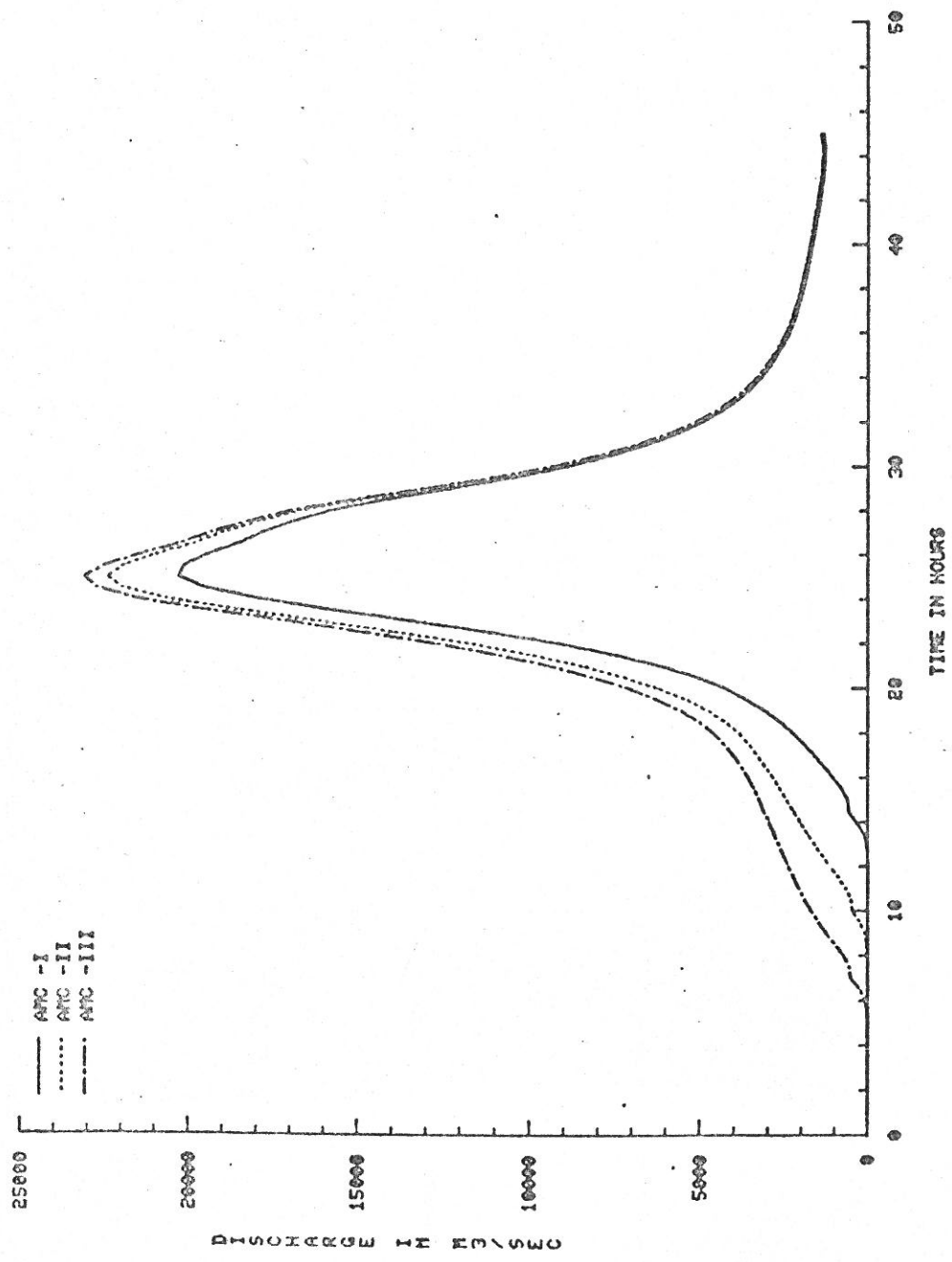


Figure 1.7.9. Simulated probable maximum flood.

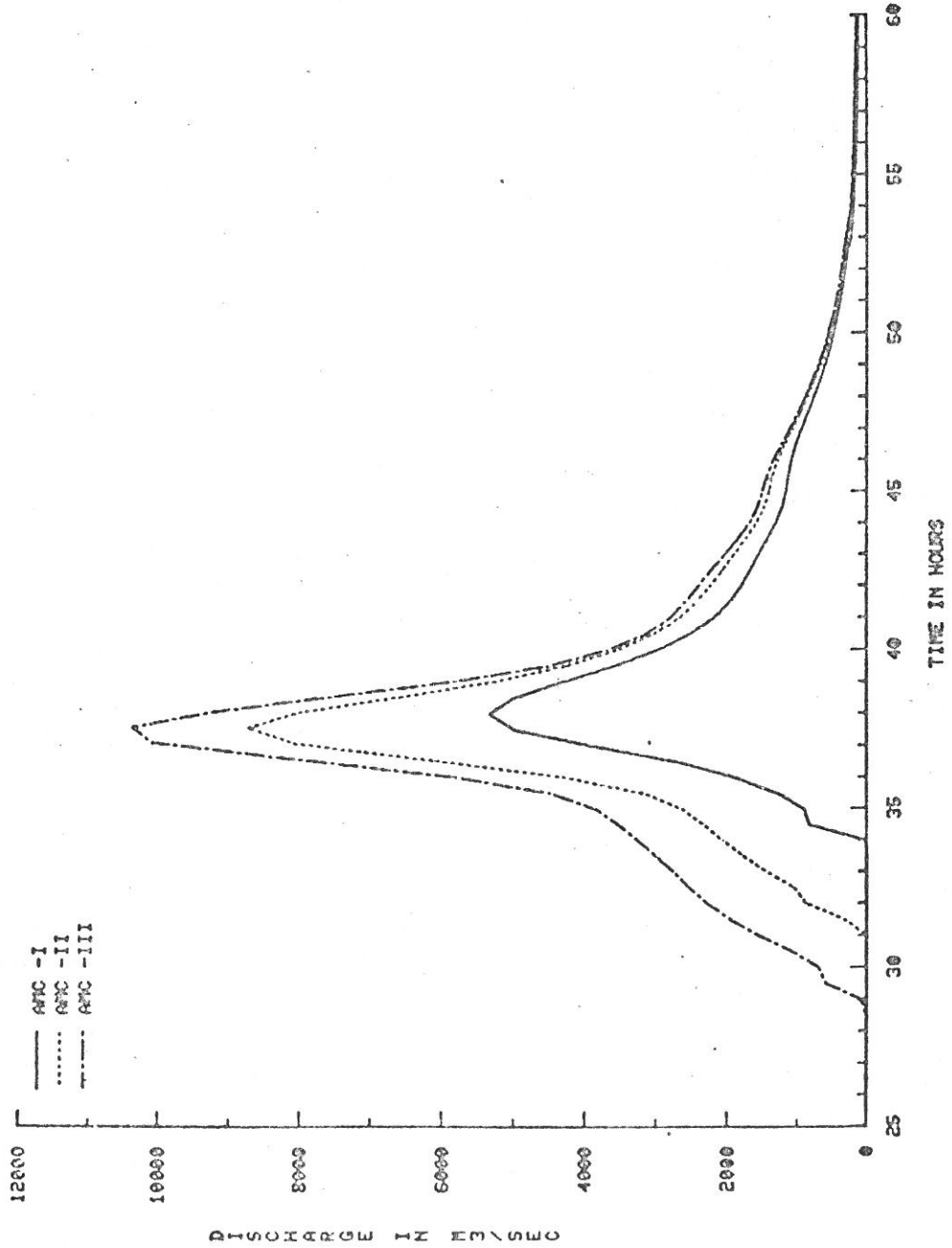


Figure 1.7.10. Simulated hydrograph for hurricane David (August, 1979).

the potential to substantially attenuate the flood hydrograph computed by the kinematic wave model.

Two possible storage areas were identified in the main channel of the Nizao River. The first one is located upstream of Rio Abajo and the second one upstream the junction of Banilejo River and the main channel.

Based on the new information made available by INDRHI (cross sections and map scale 1:20000) for the area upstream of Rio Abajo, it was concluded that the storage in the channel is not significant even for very high flows. Because a detailed map for the area upstream the junction of Banilejo and Nizao is not available a definite conclusion about the effect of storage in this location during high flows could not be made. However an approximate sensitivity analysis was conducted to see the effect of different storages on the peak discharge.

An estimate of elevation-discharge relation was obtained by using Manning's equation in the narrowest section downstream the junction. The values used for this estimation are:

$$\text{Slope} = = 0.0055$$

$$\text{Manning's } N = 0.04$$

$$\text{Base width} = 85 \text{ m}$$

$$z \text{ (side slope)} = \frac{11}{40} = 0.275$$

Using the parameters the following table was obtained.

<u>Depth</u>	<u>A</u>	<u>R</u>	<u>Q</u>	<u>V</u>
0.5	42.6	0.49	49.4	1.16
1.0	85.3	0.98	156	1.83
1.5	128.1	1.45	305	2.38
2.0	171.1	1.92	490	2.86
2.5	214.0	2.38	707	3.30
3.0	257.0	2.82	953	3.70
3.5	301.0	3.26	1227	4.08
4.0	344.0	3.69	1525	4.43
4.5	388.0	4.11	1847	4.76
5.0	432.0	4.53	2192	5.07
5.5	476.0	4.94	2557	5.37
6.0	520.0	5.34	2943	5.66

From the maps scale 1:50000 one can estimate the volume of storage for different depths. Two runs with different storage values were made to see their effect on the hydrograph at Paso del Ermitano.

(a) FIRST RUN: (Upper Limit)

For a maximum stage level of 6.0 m assume a maximum average flooded area of 3 km^2 . The storage for this stage is $\frac{1}{3} \times 3 \times 10^6 \times 6 = 6 \times 10^6 \text{ m}^3$. Assuming storage S is related to stage d , in the form, $S = ad^3$ the stage-storage data are obtained as given below:

<u>Depth (m)</u>	<u>STORAGE $\times 10^3 \text{ m}^3$</u>
0.5	3.5
1.0	27.8
1.5	93.8
2.0	222.2
2.5	434.0
3.0	750.0
3.5	1191.0
4.0	1777.8
4.5	2531.3
5.0	3472.2
5.5	4621.5
6.0	6000.0

(b) SECOND RUN

Assume that the maximum storage for a stage of 6 m is $2000 \times 10^3 \text{ m}^3$. Then the storage ordinate in the storage-elevation curve will be one third of the one used in the first run.

These two assumed reservoirs were incorporated into the Hurricane David HEC-file. The corresponding results for the larger reservoir are included in Appendix 1.7.B. It was seen that the effect of the storage on the peak discharge is insignificant in both cases and therefore it was not considered in subsequent analysis.

APPENDIX 1.7.A

SAMPLE OUTPUT OF RAINFALL WEIGHTING
USING THIESSEN POLYGON METHOD

S.059	54	27.70	77.70
S.060	55	26.60	75.70
S.061	56	27.20	74.80
S.062	57	29.00	74.40
S.063	58	31.00	72.70
S.064	59	31.20	71.00
S.065	60	30.90	69.80
S.066	61	31.80	68.70
S.067	62	32.20	67.50
S.068	63	33.00	66.80
S.069	64	33.80	67.10
S.070	65	35.20	67.40
S.071	66	37.10	65.70
S.072	67	39.30	64.10
S.073	68	40.40	60.20
S.074	69	43.10	59.70
S.075	70	43.20	58.60
S.076	71	43.80	58.50
S.077	72	44.80	56.00
S.078	73	44.30	54.70
S.079	74	45.40	54.00
S.080	75	45.80	53.00
S.081	76	47.30	52.80
S.082	77	47.70	51.20
S.083	78	47.00	50.20
S.084	79	46.80	49.00
S.085	80	47.70	45.80
S.086	81	49.70	44.60
S.087	82	49.80	43.80
S.088	83	51.20	42.60
S.089	84	54.30	42.50
S.090	85	54.30	41.50
S.091	86	55.00	40.80
S.092	87	57.60	41.20
S.093	88	58.80	40.80
S.094	89	59.80	39.80
S.095	90	61.20	38.70
S.096	91	60.70	36.80
S.097	92	61.20	34.80
S.098	93	62.50	34.90
S.099	94	64.70	34.30
S.100			

INPUT STATION	DATA
VALDESIA	64.80
EL RIO	28.70
QUEMADOS	46.40
VALLENUEVO	22.70
CONSTANZA	19.20
BANI	59.50
VILLA ALTAGRACIA	76.20
AZUA	16.60
PALO DE CAJA	52.20
MEDINA	79.20
JUMA-BONAO	54.00
ENCOMBE	94.20
AZUA HATILLO	37.40
END	

INPUT DATA	DATA
VALDESIA	313.7
EL RIO	18.8
QUEMADOS	47.3
VALLENUEVO	35.7

12.2
331.5
173.8
182.4
153.7
263.7
50.5
487.1
63.2

CONSTANZA
BANI
VILLA ALTAGRACIA
AZUA
PALO DE CAJA
MEDINA
JUAMA-BONAO
ENGOMBE
AZUA HATILLO
END

S.121
S.122
S.123
S.124
S.125
S.126
S.127
S.128
S.129
S.130

PROC THIESEN

STOP

S.131

S.132

DIAGNOSTIC CHECK OF COORDINATE SYSTEM : NIZAO BASIN - DOMINICAN REPUBLIC NON-RECORDING WEIGHTS SER.85

SPHERICAL RECTANGULAR

CONTROL POINTS	DEG.MIN. SEC LONGITUDE	DEG.MIN. SEC LATITUDE	X-COORDINATE	Y-COORDINATE	Z-ELEVATION	DIAGNOSTICS
MINIMUM			25.000	30.000		
MAXIMUM			75.000	85.000		
STA 1 VALDESIA			64.800	35.400	.000	INSIDE OF RANGE
STA 2 EL RIO			28.700	99.300	.000	*OUTSIDE OF RANGE*
STA 3 QUEMADOS			46.400	89.200	.000	*OUTSIDE OF RANGE*
STA 4 VALLENUE			22.700	80.100	.000	*OUTSIDE OF RANGE*
STA 5 CONSTANZ			19.200	91.800	.000	*OUTSIDE OF RANGE*
STA 6 BAHU			59.500	21.100	.000	*OUTSIDE OF RANGE*
STA 7 VILLA AL			76.200	64.900	.000	*OUTSIDE OF RANGE*
STA 8 AZUA			16.600	41.200	.000	*OUTSIDE OF RANGE*
STA 9 PALO DE			52.200	49.200	.000	INSIDE OF RANGE
STA 10 MEDINA			79.200	49.600	.000	*OUTSIDE OF RANGE*
STA 11 JUMA-BON			54.000	90.200	.000	*OUTSIDE OF RANGE*
STA 12 ENCOMBE			94.200	40.200	.000	*OUTSIDE OF RANGE*
STA 13 AZUA HAT			37.400	34.300	.000	INSIDE OF RANGE
BOUNDARY POINT 1			65.700	34.200		INSIDE OF RANGE
BOUNDARY POINT 2			68.500	32.900		INSIDE OF RANGE
BOUNDARY POINT 3			68.800	34.000		INSIDE OF RANGE
BOUNDARY POINT 4			70.000	35.100		INSIDE OF RANGE
BOUNDARY POINT 5			71.000	35.200		INSIDE OF RANGE
BOUNDARY POINT 6			72.500	35.000		INSIDE OF RANGE
BOUNDARY POINT 7			73.000	36.000		INSIDE OF RANGE
BOUNDARY POINT 8			72.500	37.000		INSIDE OF RANGE
BOUNDARY POINT 9			71.500	38.000		INSIDE OF RANGE
BOUNDARY POINT 10			70.800	39.000		INSIDE OF RANGE
BOUNDARY POINT 11			70.000	40.000		INSIDE OF RANGE
BOUNDARY POINT 12			69.300	41.800		INSIDE OF RANGE
BOUNDARY POINT 13			69.700	42.500		INSIDE OF RANGE
BOUNDARY POINT 14			69.500	43.300		INSIDE OF RANGE

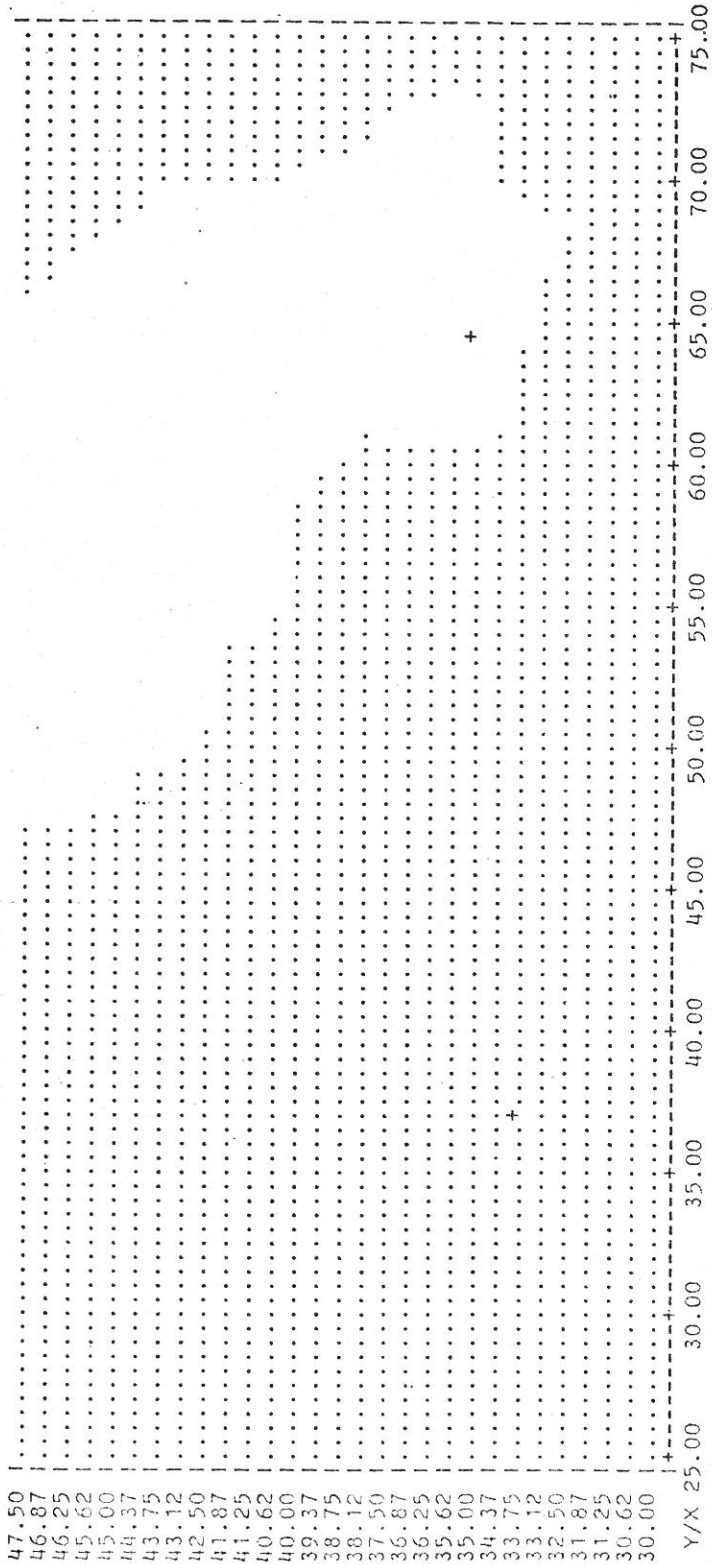
BOUNDARY POINT 15	69.700	44.200	INSIDE OF RANGE
BOUNDARY POINT 16	69.000	44.300	INSIDE OF RANGE
BOUNDARY POINT 17	65.300	48.000	INSIDE OF RANGE
BOUNDARY POINT 18	65.000	49.800	INSIDE OF RANGE
BOUNDARY POINT 19	64.200	51.600	INSIDE OF RANGE
BOUNDARY POINT 20	65.000	52.000	INSIDE OF RANGE
BOUNDARY POINT 21	65.200	53.000	INSIDE OF RANGE
BOUNDARY POINT 22	65.000	54.000	INSIDE OF RANGE
BOUNDARY POINT 23	64.000	55.000	INSIDE OF RANGE
BOUNDARY POINT 24	63.300	56.000	INSIDE OF RANGE
BOUNDARY POINT 25	63.200	58.000	INSIDE OF RANGE
BOUNDARY POINT 26	63.900	59.000	INSIDE OF RANGE
BOUNDARY POINT 27	65.200	61.000	INSIDE OF RANGE
BOUNDARY POINT 28	62.500	63.000	INSIDE OF RANGE
BOUNDARY POINT 29	62.300	64.600	INSIDE OF RANGE
BOUNDARY POINT 30	60.800	65.800	INSIDE OF RANGE
BOUNDARY POINT 31	60.200	67.000	INSIDE OF RANGE
BOUNDARY POINT 32	60.000	69.800	INSIDE OF RANGE
BOUNDARY POINT 33	58.300	70.900	INSIDE OF RANGE
BOUNDARY POINT 34	55.200	71.600	INSIDE OF RANGE
BOUNDARY POINT 35	53.600	72.000	INSIDE OF RANGE
BOUNDARY POINT 36	53.000	73.900	INSIDE OF RANGE
BOUNDARY POINT 37	51.500	74.200	INSIDE OF RANGE
BOUNDARY POINT 38	51.300	75.000	INSIDE OF RANGE
BOUNDARY POINT 39	49.300	74.000	INSIDE OF RANGE
BOUNDARY POINT 40	48.200	74.500	INSIDE OF RANGE
BOUNDARY POINT 41	47.600	74.500	INSIDE OF RANGE
BOUNDARY POINT 42	46.600	73.500	INSIDE OF RANGE
BOUNDARY POINT 43	46.300	72.800	INSIDE OF RANGE
BOUNDARY POINT 44	45.000	72.900	INSIDE OF RANGE
BOUNDARY POINT 45	43.700	73.800	INSIDE OF RANGE
BOUNDARY POINT 46	41.800	76.000	INSIDE OF RANGE

BOUNDARY POINT 47	40.000	78.800	INSIDE OF RANGE
BOUNDARY POINT 48	37.200	79.500	INSIDE OF RANGE
BOUNDARY POINT 49	35.500	82.100	INSIDE OF RANGE
BOUNDARY POINT 50	33.000	80.600	INSIDE OF RANGE
BOUNDARY POINT 51	29.000	82.200	INSIDE OF RANGE
BOUNDARY POINT 52	28.500	80.700	INSIDE OF RANGE
BOUNDARY POINT 53	27.500	80.400	INSIDE OF RANGE
BOUNDARY POINT 54	27.700	77.700	INSIDE OF RANGE
BOUNDARY POINT 55	26.600	75.700	INSIDE OF RANGE
BOUNDARY POINT 56	27.200	74.800	INSIDE OF RANGE
BOUNDARY POINT 57	29.000	74.400	INSIDE OF RANGE
BOUNDARY POINT 58	31.000	72.700	INSIDE OF RANGE
BOUNDARY POINT 59	31.200	71.000	INSIDE OF RANGE
BOUNDARY POINT 60	30.900	69.800	INSIDE OF RANGE
BOUNDARY POINT 61	31.800	68.700	INSIDE OF RANGE
BOUNDARY POINT 62	32.200	67.500	INSIDE OF RANGE
BOUNDARY POINT 63	33.000	66.800	INSIDE OF RANGE
BOUNDARY POINT 64	33.800	67.100	INSIDE OF RANGE
BOUNDARY POINT 65	35.200	67.400	INSIDE OF RANGE
BOUNDARY POINT 66	37.100	65.700	INSIDE OF RANGE
BOUNDARY POINT 67	39.300	64.100	INSIDE OF RANGE
BOUNDARY POINT 68	40.400	60.200	INSIDE OF RANGE
BOUNDARY POINT 69	43.100	59.700	INSIDE OF RANGE
BOUNDARY POINT 70	43.200	58.600	INSIDE OF RANGE
BOUNDARY POINT 71	43.800	58.500	INSIDE OF RANGE
BOUNDARY POINT 72	44.800	56.000	INSIDE OF RANGE
BOUNDARY POINT 73	44.300	54.700	INSIDE OF RANGE
BOUNDARY POINT 74	45.400	54.000	INSIDE OF RANGE
BOUNDARY POINT 75	45.800	53.000	INSIDE OF RANGE
BOUNDARY POINT 76	47.300	52.800	INSIDE OF RANGE
BOUNDARY POINT 77	47.700	51.200	INSIDE OF RANGE
BOUNDARY POINT 78	47.000	50.200	INSIDE OF RANGE

BOUNDARY POINT 79	46.800	49.000	INSIDE OF RANGE
BOUNDARY POINT 80	47.700	45.800	INSIDE OF RANGE
BOUNDARY POINT 81	49.700	44.600	INSIDE OF RANGE
BOUNDARY POINT 82	49.800	43.800	INSIDE OF RANGE
BOUNDARY POINT 83	51.200	42.600	INSIDE OF RANGE
BOUNDARY POINT 84	54.300	42.500	INSIDE OF RANGE
BOUNDARY POINT 85	54.300	41.500	INSIDE OF RANGE
BOUNDARY POINT 86	55.000	40.800	INSIDE OF RANGE
BOUNDARY POINT 87	57.600	41.200	INSIDE OF RANGE
BOUNDARY POINT 88	58.800	40.800	INSIDE OF RANGE
BOUNDARY POINT 89	59.800	39.800	INSIDE OF RANGE
BOUNDARY POINT 90	61.200	38.700	INSIDE OF RANGE
BOUNDARY POINT 91	60.700	36.800	INSIDE OF RANGE
BOUNDARY POINT 92	61.200	34.800	INSIDE OF RANGE
BOUNDARY POINT 93	62.500	34.900	INSIDE OF RANGE
BOUNDARY POINT 94	64.700	34.300	INSIDE OF RANGE

DELINEATION OF BASIN BOUNDARY: NIZAO BASIN - DOMINICAN REPUBLIC NON-RECORDING WEIGHTS SEP. 85

85.00
84.37
83.75
83.12
82.50
81.87
81.25
80.62
80.00
79.37
78.75
78.12
77.50
76.87
76.25
75.62
75.00
74.37
73.75
73.12
72.50
71.87
71.25
70.62
70.00
69.37
68.75
68.12
67.50
66.87
66.25
65.62
65.00
64.37
63.75
63.12
62.50
61.87
61.25
60.62
60.00
59.37
58.75
58.12
57.50
56.87
56.25
55.62
55.00
54.37
53.75
53.12
52.50
51.87
51.25
50.62
50.00
49.37
48.75
48.12



TOTAL AREA = 904.062
 UNIT AREA = .312

DELTA X = .500
 DELTA Y = .625

LEGEND:
 I I OUTSIDE OF BASIN
 I I INSIDE OF BASIN
 I+ I STATION

Y/X	25.00	30.00	35.00	40.00	45.00	50.00	55.00	60.00	65.00	70.00	75.00
47.50											
46.87											
46.25											
45.62											
45.00											
44.37											
43.75											
43.12											
42.50											
41.87											
41.25											
40.62											
40.00											
39.37											
38.75											
38.12											
37.50											
36.87											
36.25											
35.62											
35.00											
34.37											
33.75											
33.12											
32.50											
31.87											
31.25											
30.62											
30.00											

STA. NO.	STA. NAME	SYMBOL	THIESSEN WEIGHT	POLYGON AREA	STA. DATA
1	VALDESIA	0	.14967	135.31	313.700
2	EL RIO	1	.00000	.00	18.800
3	QUEMADOS	2	.09644	87.19	47.300
4	VALLENGUE	3	.15900	143.75	35.700
5	CONSTANZ	4	.00000	.00	12.200
6	BANI	5	.00000	.00	331.500
7	VILLA AL	6	.03837	34.69	173.800
8	AZUA	7	.00000	.00	182.400
9	PALO DE	8	.54511	492.81	153.700
10	MEDINA	9	.00000	.00	263.700
11	JUMA-BON	A	.01141	10.31	50.500
12	ENGOMBE	B	.00000	.00	487.100
13	AZUA HAT	C	.00000	.00	63.200

AREAL AVERAGE OF DATA = 148.2178

THIESSEN METHOD OPTION CODE = 1

- 1 MIN STRAIGHT DISTANCE
- 2 MIN SQUARE DISTANCE
- 3 MIN ORTHOGONAL DISTANCE
- 4 MIN MAX LEG DISTANCE
- 5 MIN ABS (DATA) X STRAIGHT DISTANCE
- 6 MAX ABS (DEVIATE) X SPATIAL CORRELATION COEFF

-----+
 ! HYDROLOGIC MAPPING, INTERPOLATION AND AREAL AVERAGING SYSTEM !
 +-----

S.001 INPUT TITLE NIZAO BASIN - DOMINICAN REPUBLIC RECORDING STATIONS WEIGHTS SEP.85
 S.002

S.003 INPUT RANGE 1 0 75.00 25.00 85.00 30.00
 S.004

S.005 INPUT BOUNDARY

S.006	1	65.70	34.20
S.007	2	68.50	32.90
S.008	3	68.80	34.00
S.009	4	70.00	35.10
S.010	5	71.00	35.20
S.011	6	72.50	35.00
S.012	7	73.00	36.00
S.013	8	72.50	37.00
S.014	9	71.50	38.00
S.015	10	70.80	39.00
S.016	11	70.00	40.00
S.017	12	69.30	41.80
S.018	13	69.70	42.50
S.019	14	69.50	43.30
S.020	15	69.70	44.20
S.021	16	69.00	44.30
S.022	17	65.30	48.00
S.023	18	65.00	49.80
S.024	19	64.20	51.60
S.025	20	65.00	52.00
S.026	21	65.20	53.00
S.027	22	65.00	54.00
S.028	23	64.00	55.00
S.029	24	63.30	56.00
S.030	25	63.20	58.00
S.031	26	63.90	59.00
S.032	27	65.20	61.00
S.033	28	62.50	63.00
S.034	29	62.30	64.60
S.035	30	60.80	65.80
S.036	31	60.20	67.00
S.037	32	60.00	69.80
S.038	33	58.30	70.90
S.039	34	55.20	71.60
S.040	35	53.60	72.00
S.041	36	53.00	73.90
S.042	37	51.50	74.20
S.043	38	51.30	75.00
S.044	39	49.30	74.00
S.045	40	48.20	74.50
S.046	41	47.60	74.50
S.047	42	46.60	73.50
S.048	43	46.30	72.80
S.049	44	45.00	72.90
S.050	45	43.70	73.80
S.051	46	41.80	76.00
S.052	47	40.00	78.80
S.053	48	37.20	79.50
S.054	49	35.50	82.10
S.055	50	33.00	80.60
S.056	51	29.00	82.20
S.057	52	28.50	80.70
S.058	53	27.50	80.40

S.059	54	27.70	77.70
S.060	55	26.60	75.70
S.061	56	27.20	74.80
S.062	57	29.00	74.40
S.063	58	31.00	72.70
S.064	59	31.20	71.00
S.065	60	30.90	69.80
S.066	61	31.80	68.70
S.067	62	32.20	67.50
S.068	63	33.00	66.80
S.069	64	33.80	67.10
S.070	65	35.20	67.40
S.071	66	37.10	65.70
S.072	67	39.30	64.10
S.073	68	40.40	60.20
S.074	69	43.10	59.70
S.075	70	43.20	58.60
S.076	71	43.80	58.50
S.077	72	44.80	56.00
S.078	73	44.30	54.70
S.079	74	45.40	54.00
S.080	75	45.80	53.00
S.081	76	47.30	52.80
S.082	77	47.70	51.20
S.083	78	47.00	50.20
S.084	79	46.80	49.00
S.085	80	47.70	45.80
S.086	81	49.70	44.60
S.087	82	49.80	43.80
S.088	83	51.20	42.60
S.089	84	54.30	42.50
S.090	85	54.30	41.50
S.091	86	55.00	40.80
S.092	87	57.60	41.20
S.093	88	58.80	40.80
S.094	89	59.80	39.80
S.095	90	61.20	38.70
S.096	91	60.70	36.80
S.097	92	61.20	34.80
S.098	93	62.50	34.90
S.099	94	64.70	34.30
S.100	END		
S.101	INPUT STATION		
S.102	VALDESIA	64.80	35.40
S.103	JUMA BONAO	54.00	90.20
S.104	VALLENUEVO	22.70	80.10
S.105	END		
S.106	INPUT DATA		
S.107	VALDESIA	30.00	
S.108	JUMA BONAO	80.00	
S.109	VALLENUEVO	10.00	
S.110	END		
S.111	PROC THIESEN		
S.112	STOP		

* STMT GENERATED

DIAGNOSTIC CHECK OF COORDINATE SYSTEM : NIZAO BASIN - DOMINICAN REPUBLIC RECORDING STATIONS WEIGHTS SEP.85

RECTANGULAR

CONTROL POINTS DEG. MIN. SEC LONGITUDE DEG. MIN. SEC LATITUDE X-COORDINATE Y-COORDINATE Z-ELEVATION DIAGNOSTICS

CONTROL POINTS	DEG. MIN. SEC LONGITUDE	DEG. MIN. SEC LATITUDE	X-COORDINATE	Y-COORDINATE	Z-ELEVATION	DIAGNOSTICS
MINIMUM			25.000	30.000		
MAXIMUM			75.000	85.000		
STA 1 VALDESIA			64.800	35.400	.000	INSIDE OF RANGE
STA 2 JUMA BON			54.000	90.200	.000	*OUTSIDE OF RANGE*
STA 3 VALLENUE			22.700	80.100	.000	*OUTSIDE OF RANGE*
BOUNDARY POINT 1			65.700	34.200		INSIDE OF RANGE
BOUNDARY POINT 2			68.500	32.900		INSIDE OF RANGE
BOUNDARY POINT 3			68.800	34.000		INSIDE OF RANGE
BOUNDARY POINT 4			70.000	35.100		INSIDE OF RANGE
BOUNDARY POINT 5			71.000	35.200		INSIDE OF RANGE
BOUNDARY POINT 6			72.500	35.000		INSIDE OF RANGE
BOUNDARY POINT 7			73.000	36.000		INSIDE OF RANGE
BOUNDARY POINT 8			72.500	37.000		INSIDE OF RANGE
BOUNDARY POINT 9			71.500	38.000		INSIDE OF RANGE
BOUNDARY POINT 10			70.800	39.000		INSIDE OF RANGE
BOUNDARY POINT 11			70.000	40.000		INSIDE OF RANGE
BOUNDARY POINT 12			69.300	41.800		INSIDE OF RANGE
BOUNDARY POINT 13			69.700	42.500		INSIDE OF RANGE
BOUNDARY POINT 14			69.500	43.300		INSIDE OF RANGE
BOUNDARY POINT 15			69.700	44.200		INSIDE OF RANGE
BOUNDARY POINT 16			69.000	44.300		INSIDE OF RANGE
BOUNDARY POINT 17			65.300	48.000		INSIDE OF RANGE
BOUNDARY POINT 18			65.000	49.800		INSIDE OF RANGE
BOUNDARY POINT 19			64.200	51.600		INSIDE OF RANGE
BOUNDARY POINT 20			65.000	52.000		INSIDE OF RANGE
BOUNDARY POINT 21			65.200	53.000		INSIDE OF RANGE
BOUNDARY POINT 22			65.000	54.000		INSIDE OF RANGE
BOUNDARY POINT 23			64.000	55.000		INSIDE OF RANGE
BOUNDARY POINT 24			63.300	56.000		INSIDE OF RANGE

SPHERICAL

BOUNDARY POINT 25	63.200	58.000	INSIDE OF RANGE
BOUNDARY POINT 26	63.900	59.000	INSIDE OF RANGE
BOUNDARY POINT 27	65.200	61.000	INSIDE OF RANGE
BOUNDARY POINT 28	62.500	63.000	INSIDE OF RANGE
BOUNDARY POINT 29	62.300	64.600	INSIDE OF RANGE
BOUNDARY POINT 30	60.800	65.800	INSIDE OF RANGE
BOUNDARY POINT 31	60.200	67.000	INSIDE OF RANGE
BOUNDARY POINT 32	60.000	69.800	INSIDE OF RANGE
BOUNDARY POINT 33	58.300	70.900	INSIDE OF RANGE
BOUNDARY POINT 34	55.200	71.600	INSIDE OF RANGE
BOUNDARY POINT 35	53.600	72.000	INSIDE OF RANGE
BOUNDARY POINT 36	53.000	73.900	INSIDE OF RANGE
BOUNDARY POINT 37	51.500	74.200	INSIDE OF RANGE
BOUNDARY POINT 38	51.300	75.000	INSIDE OF RANGE
BOUNDARY POINT 39	49.300	74.000	INSIDE OF RANGE
BOUNDARY POINT 40	48.200	74.500	INSIDE OF RANGE
BOUNDARY POINT 41	47.600	74.500	INSIDE OF RANGE
BOUNDARY POINT 42	46.600	73.500	INSIDE OF RANGE
BOUNDARY POINT 43	46.300	72.800	INSIDE OF RANGE
BOUNDARY POINT 44	45.000	72.900	INSIDE OF RANGE
BOUNDARY POINT 45	43.700	73.800	INSIDE OF RANGE
BOUNDARY POINT 46	41.800	76.000	INSIDE OF RANGE
BOUNDARY POINT 47	40.000	78.800	INSIDE OF RANGE
BOUNDARY POINT 48	37.200	79.500	INSIDE OF RANGE
BOUNDARY POINT 49	35.500	82.100	INSIDE OF RANGE
BOUNDARY POINT 50	33.000	80.600	INSIDE OF RANGE
BOUNDARY POINT 51	29.000	82.200	INSIDE OF RANGE
BOUNDARY POINT 52	28.500	80.700	INSIDE OF RANGE
BOUNDARY POINT 53	27.500	80.400	INSIDE OF RANGE
BOUNDARY POINT 54	27.700	77.700	INSIDE OF RANGE
BOUNDARY POINT 55	26.600	75.700	INSIDE OF RANGE
BOUNDARY POINT 56	27.200	74.800	INSIDE OF RANGE

BOUNDARY POINT 57	29.000	74.400	INSIDE OF RANGE
BOUNDARY POINT 58	31.000	72.700	INSIDE OF RANGE
BOUNDARY POINT 59	31.200	71.000	INSIDE OF RANGE
BOUNDARY POINT 60	30.900	69.800	INSIDE OF RANGE
BOUNDARY POINT 61	31.800	68.700	INSIDE OF RANGE
BOUNDARY POINT 62	32.200	67.500	INSIDE OF RANGE
BOUNDARY POINT 63	33.000	66.800	INSIDE OF RANGE
BOUNDARY POINT 64	33.800	67.100	INSIDE OF RANGE
BOUNDARY POINT 65	35.200	67.400	INSIDE OF RANGE
BOUNDARY POINT 66	37.100	65.700	INSIDE OF RANGE
BOUNDARY POINT 67	39.300	64.100	INSIDE OF RANGE
BOUNDARY POINT 68	40.400	60.200	INSIDE OF RANGE
BOUNDARY POINT 69	43.100	59.700	INSIDE OF RANGE
BOUNDARY POINT 70	43.200	58.600	INSIDE OF RANGE
BOUNDARY POINT 71	43.800	58.500	INSIDE OF RANGE
BOUNDARY POINT 72	44.800	56.000	INSIDE OF RANGE
BOUNDARY POINT 73	44.300	54.700	INSIDE OF RANGE
BOUNDARY POINT 74	45.400	54.000	INSIDE OF RANGE
BOUNDARY POINT 75	45.800	53.000	INSIDE OF RANGE
BOUNDARY POINT 76	47.300	52.800	INSIDE OF RANGE
BOUNDARY POINT 77	47.700	51.200	INSIDE OF RANGE
BOUNDARY POINT 78	47.000	50.200	INSIDE OF RANGE
BOUNDARY POINT 79	46.800	49.000	INSIDE OF RANGE
BOUNDARY POINT 80	47.700	45.800	INSIDE OF RANGE
BOUNDARY POINT 81	49.700	44.600	INSIDE OF RANGE
BOUNDARY POINT 82	49.800	43.800	INSIDE OF RANGE
BOUNDARY POINT 83	51.200	42.600	INSIDE OF RANGE
BOUNDARY POINT 84	54.300	42.500	INSIDE OF RANGE
BOUNDARY POINT 85	54.300	41.500	INSIDE OF RANGE
BOUNDARY POINT 86	55.000	40.800	INSIDE OF RANGE
BOUNDARY POINT 87	57.600	41.200	INSIDE OF RANGE
BOUNDARY POINT 88	58.800	40.800	INSIDE OF RANGE

INSIDE OF RANGE
INSIDE OF RANGE
INSIDE OF RANGE
INSIDE OF RANGE
INSIDE OF RANGE
INSIDE OF RANGE
INSIDE OF RANGE

39.800
38.700
36.800
34.800
34.900
34.300

59.800
61.200
60.700
61.200
62.500
64.700

BOUNDARY POINT 89
BOUNDARY POINT 90
BOUNDARY POINT 91
BOUNDARY POINT 92
BOUNDARY POINT 93
BOUNDARY POINT 94

BOUNDARY POINT 57	29.000	74.400	INSIDE OF RANGE
BOUNDARY POINT 58	31.000	72.700	INSIDE OF RANGE
BOUNDARY POINT 59	31.200	71.000	INSIDE OF RANGE
BOUNDARY POINT 60	30.900	69.800	INSIDE OF RANGE
BOUNDARY POINT 61	31.800	68.700	INSIDE OF RANGE
BOUNDARY POINT 62	32.200	67.500	INSIDE OF RANGE
BOUNDARY POINT 63	33.000	66.800	INSIDE OF RANGE
BOUNDARY POINT 64	33.800	67.100	INSIDE OF RANGE
BOUNDARY POINT 65	35.200	67.400	INSIDE OF RANGE
BOUNDARY POINT 66	37.100	65.700	INSIDE OF RANGE
BOUNDARY POINT 67	39.300	64.100	INSIDE OF RANGE
BOUNDARY POINT 68	40.400	60.200	INSIDE OF RANGE
BOUNDARY POINT 69	43.100	59.700	INSIDE OF RANGE
BOUNDARY POINT 70	43.200	58.600	INSIDE OF RANGE
BOUNDARY POINT 71	43.800	58.500	INSIDE OF RANGE
BOUNDARY POINT 72	44.800	56.000	INSIDE OF RANGE
BOUNDARY POINT 73	44.300	54.700	INSIDE OF RANGE
BOUNDARY POINT 74	45.400	54.000	INSIDE OF RANGE
BOUNDARY POINT 75	45.800	53.000	INSIDE OF RANGE
BOUNDARY POINT 76	47.300	52.800	INSIDE OF RANGE
BOUNDARY POINT 77	47.700	51.200	INSIDE OF RANGE
BOUNDARY POINT 78	47.000	50.200	INSIDE OF RANGE
BOUNDARY POINT 79	46.800	49.000	INSIDE OF RANGE
BOUNDARY POINT 80	47.700	45.800	INSIDE OF RANGE
BOUNDARY POINT 81	49.700	44.600	INSIDE OF RANGE
BOUNDARY POINT 82	49.800	43.800	INSIDE OF RANGE
BOUNDARY POINT 83	51.200	42.600	INSIDE OF RANGE
BOUNDARY POINT 84	54.300	42.500	INSIDE OF RANGE
BOUNDARY POINT 85	54.300	41.500	INSIDE OF RANGE
BOUNDARY POINT 86	55.000	40.800	INSIDE OF RANGE
BOUNDARY POINT 87	57.600	41.200	INSIDE OF RANGE
BOUNDARY POINT 88	58.800	40.800	INSIDE OF RANGE

INSIDE OF RANGE
INSIDE OF RANGE
INSIDE OF RANGE
INSIDE OF RANGE
INSIDE OF RANGE
INSIDE OF RANGE
INSIDE OF RANGE

39.800
38.700
36.800
34.800
34.900
34.300

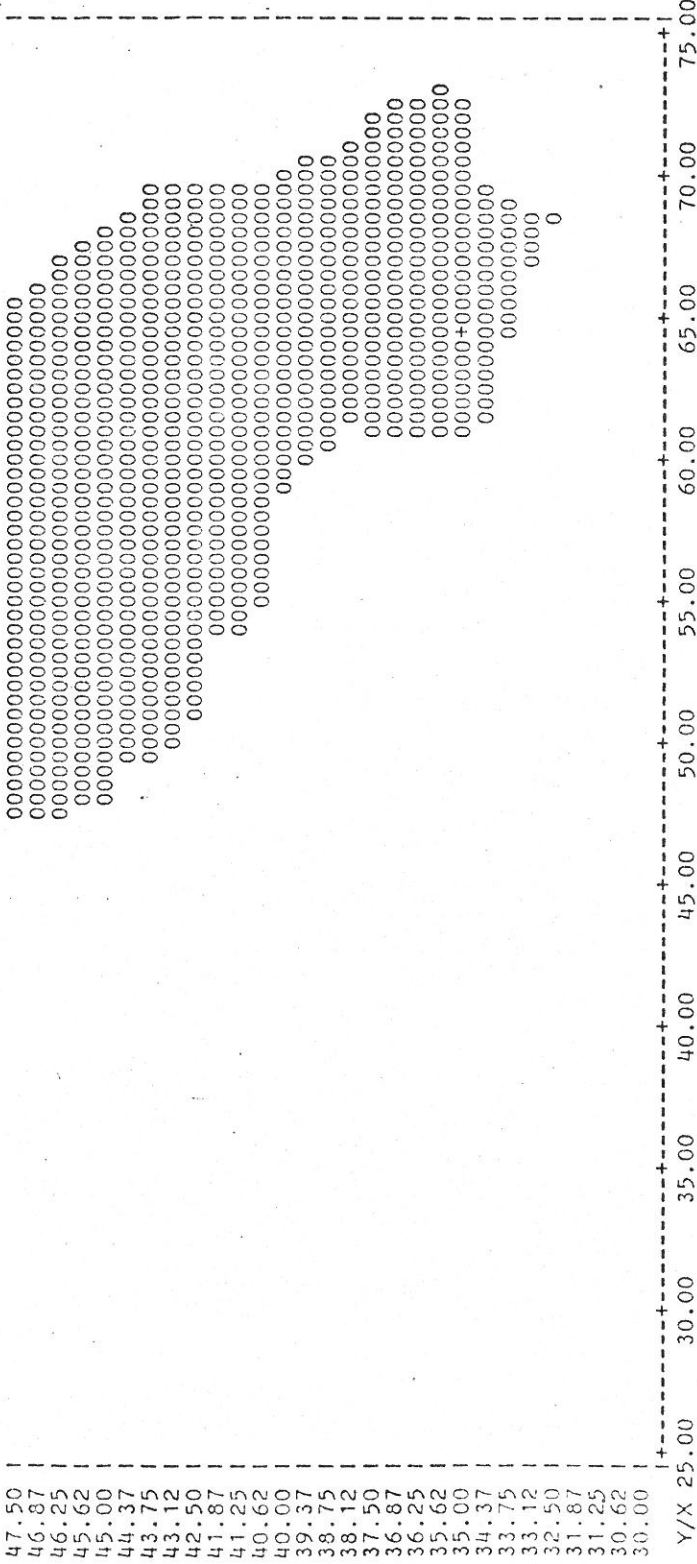
59.800
61.200
60.700
61.200
62.500
64.700

BOUNDARY POINT 89
BOUNDARY POINT 90
BOUNDARY POINT 91
BOUNDARY POINT 92
BOUNDARY POINT 93
BOUNDARY POINT 94

MAPPING BY THIESSEN POLYGON: NIZAO BASIN - DOMINICAN REPUBLIC RECORDING STATIONS WEIGHTS SEP. 85

85.00
84.37
83.75
83.12
82.50
81.87
81.25
80.62
80.00
79.37
78.75
78.12
77.50
76.87
76.25
75.62
75.00
74.37
73.75
73.12
72.50
71.87
71.25
70.62
70.00
69.37
68.75
68.12
67.50
66.87
66.25
65.62
65.00
64.37
63.75
63.12
62.50
61.87
61.25
60.62
60.00
59.37
58.75
58.12
57.50
56.87
56.25
55.62
55.00
54.37
53.75
53.12
52.50
51.87
51.25
50.62
50.00
49.37
48.75
48.12



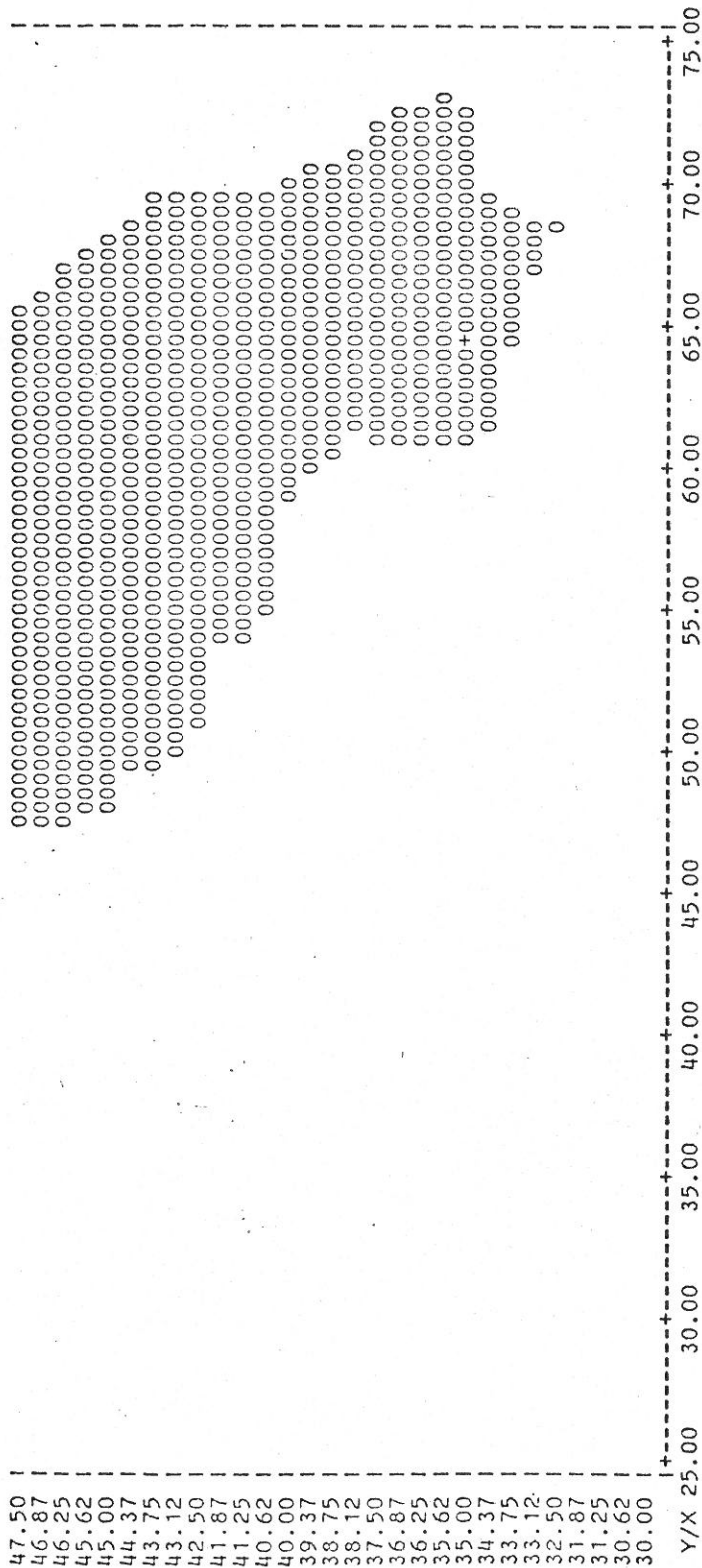


STA. NO.	STA. NAME	SYMBOL	THIESSEN WEIGHT	POLYGON AREA	STA. DATA
1	VALDESIA	0	.53128	480.31	30.000
2	JUMA BON	1	.20913	189.06	80.000
3	VALLENUE	2	.25959	234.69	10.000

AREAL AVERAGE OF DATA = 35.2644

THIESSEN METHOD OPTION CODE = 1

- 1 MIN STRAIGHT DISTANCE
- 2 MIN SQUARE DISTANCE
- 3 MIN ORTHOGONAL DISTANCE
- 4 MIN MAX LEG DISTANCE
- 5 MIN ABS (DATA) X STRAIGHT DISTANCE
- 6 MAX ABS (DEVIATE) X SPATIAL CORRELATION COEFF



AREAL AVERAGE OF DATA = 35.2644

THIESSEN METHOD OPTION CODE = 1

- 1 MIN STRAIGHT DISTANCE
- 2 MIN SQUARE DISTANCE
- 3 MIN ORTHOGONAL DISTANCE
- 4 MIN MAX LEG DISTANCE
- 5 MIN ABS (DATA) X STRAIGHT DISTANCE
- 6 MAX ABS (DEVIATE) X SPATIAL CORRELATION COEFF

APPENDIX 1.7.B

RESULTS OF HEC-1 MODEL RUN USING HURRICANE
DAVID DATA TO ASSESS EFFECTS OF UPSTREAM
NATURAL STORAGES

```

*****
* FLOOD HYDROGRAPH PACKAGE (HEC-1) *
* FEBRUARY 1981 *
* REVISED 31 JAN 85 *
* *
* RUN DATE 02/27/86 TIME 22.24.50 *
*****

```

```

*****
* U.S. ARMY CORPS OF ENGINEERS *
* THE HYDROLOGIC ENGINEERING CENTER *
* 609 SECOND STREET *
* DAVIS, CALIFORNIA 95616 *
* (916) 440-3285 OR (FTS) 448-3285 *
*****

```

```

X X XXXXXXXX XXXXX X
X X X X X X
X X X X X X
XXXXXXX XXXX XXXXX
X X X X X X
X X X XXXXXXXX XXXX
X X X XXX

```

THIS PROGRAM REPLACES ALL PREVIOUS VERSIONS OF HEC-1 KNOWN AS HEC1 (JAN 73), HEC1GS, HEC1DB, AND HEC1KW.

THE DEFINITIONS OF VARIABLES -RTIMP- AND -RTIOR- HAVE CHANGED FROM THOSE USED WITH THE 1973-STYLE INPUT STRUCTURE. THE DEFINITION OF -AMSKK- ON RM-CARD WAS CHANGED WITH REVISIONS DATED 28 SEP 81. THE VERSION RELEASED 31JAN85 CONTAINS NEW OPTIONS ON RL AND BA RECORDS, AND ADDS THE HL RECORD. SEE JANUARY 1985 INPUT DESCRIPTION FOR NEW DEFINITIONS.

HEC-1 INPUT

LINE	ID	1	2	3	4	5	6	7	8	9	10
55	PI	.00	.00	.00	.00	.00	.00	.00	.00	.00	.00
56	KK	SUB1A									
57	KM	RUNOFF FROM SUBBASIN 1A									
58	PR	4									
59	PW	1.0									
60	PT	50									
61	PW	1.00									
62	BA	70									
63	LS	80									
64	UK	2500	0.60	0.400	100						
65	RK	16000	0.1050	0.040	TRAP	45	10				
66	KK	SUB1B									
67	KM	RUNOFF FROM SUBBASIN 1B									
68	PR	3									
69	PW	0.30	0.70								
70	PT	40	50								
71	PW	0.30	0.70								
72	BA	70									
73	LS	80									
74	UK	2500	0.60	0.400	100						
75	RK	19000	0.0837	0.040	TRAP	45	10				
76	KK	SUB1C									
77	KM	COMBINE RUNOFF FROM SUB1A AND SUB1B									
78	HC	2									
79	KK	SUB1C									
80	KM	RUNOFF FROM SUBBASIN 1C									
81	PR	3									
82	PW	0.40	0.40	0.20	7						
83	PT	20	40	50							
84	PW	0.20	0.40	0.40							
85	BA	45									
86	LS	80									
87	UK	4500	0.50	0.300	100						
88	RK	7500	0.0133	0.026	TRAP	85	10	YES			
89	KK	SUB2A									
90	KM	RUNOFF FROM SUBBASIN 2A									
91	PR	3									
92	PW	0.80	0.20								
93	PT	20	40								
94	PW	0.2	0.8								
95	BA	24									
96	LS	80									
97	UK	2000	0.40	0.300	100						
98	RK	6500	0.0108	0.022	TRAP	85	10	YES			

SCHEMATIC DIAGRAM OF STREAM NETWORK

INPUT LINE NO.	(V) ROUTING	(--->) DIVERSION OR PUMP FLOW
56	SUB1A	
66		SUB1B
76	SUB1C.....	
79	SUB1C ***	
89	SUB2A ***	
99		SUB2B
109	SUB2C.....	
112	RESER	
121	SUB2C ***	
131		SUB3
141	SUB4.....	
144	SUB4 ***	
154		SUB5
164	SUB6.....	
167	SUB6 ***	

(***) RUNOFF ALSO COMPUTED AT THIS LOCATION

```

*****
* U.S. ARMY CORPS OF ENGINEERS
* THE HYDROLOGIC ENGINEERING CENTER
* 609 SECOND STREET
* DAVIS, CALIFORNIA 95616
* (916) 440-3285 OR (FTS) 448-3285
*****

```

```

*****
* FLOOD HYDROGRAPH PACKAGE (HEC-1)
* FEBRUARY 1981
* REVISED 31 JAN 85
*
* RUN DATE 02/27/86 TIME 22.24.50
*****

```

NIZAO BASIN - DOMINICAN REPUBLIC
KINEMATIC WAVE MODEL HURRICANE DAVID W/STORAGE

```

4 IO OUTPUT CONTROL VARIABLES
      IPRINT 1 PRINT CONTROL
      IPLOT 2 PLOT CONTROL
      QSCAL 0 HYDROGRAPH PLOT SCALE

IT HYDROGRAPH TIME DATA
      RMIN 60 MINUTES IN COMPUTATION INTERVAL
      IDATE 30AUG79 STARTING DATE
      ITIME 0900 STARTING TIME
      NQ 96 NUMBER OF HYDROGRAPH ORDINATES
      NDDATE 3SEP79 ENDING DATE
      NDTIME 0800 ENDING TIME

      COMPUTATION INTERVAL 1.00 HOURS
      TOTAL TIME BASE 95.00 HOURS

```

METRIC UNITS

DRAINAGE AREA	SQUARE KILOMETERS
PRECIPITATION DEPTH	MILLIMETERS
LENGTH, ELEVATION	METERS
FLOW	CUBIC METERS PER SECOND
STORAGE VOLUME	CUBIC METERS
SURFACE AREA	SQUARE METERS
TEMPERATURE	DEGREES CELSIUS

```

*****
*
* SUB1A
*
*****

```

```

56 KK

```

RUNOFF FROM SUBBASIN 1A

SUBBASIN RUNOFF DATA

```

62 BA SUBBASIN CHARACTERISTICS
      TAREA 70.00 SUBBASIN AREA

```

PRECIPITATION DATA

```

60 PT TOTAL STORM STATIONS 50
61 PW WEIGHTS 1.00

```

RECORDING STATIONS 4
WEIGHTS 1.00

SCS LOSS RATE
STRTL 12.70 INITIAL ABSTRACTION
CRVNR 80.00 CURVE NUMBER
RTIMP .00 PERCENT IMPERVIOUS AREA

KINEMATIC WAVE
OVERLAND-FLOW ELEMENT NO. 1
L 2500. OVERLAND FLOW LENGTH
S .6000 SLOPE
N .400 ROUGHNESS COEFFICIENT
PA 100.0 PERCENT OF SUBBASIN
MAIN CHANNEL
L 16000. CHANNEL LENGTH
S .1050 SLOPE
N .040 CHANNEL ROUGHNESS COEFFICIENT
CA 70.00 CONTRIBUTING AREA
SHAPE TRAP CHANNEL SHAPE
WD 45.00 BOTTOM WIDTH OR DIAMETER
Z 10.00 SIDE SLOPE
RUPSTQ NO ROUTE UPSTREAM HYDROGRAPH

PRECIPITATION STATION DATA

STATION	TOTAL	AVG. ANNUAL	WEIGHT
50	611.90	.00	1.00

TEMPORAL DISTRIBUTIONS

STATION	4, WEIGHT = 1.00	DT (MIN)	DX (FT)
.00	.00	60.00	4101.05
1.20	.80	60.00	26246.72
8.50	10.20		
31.20	52.10		
15.60	16.60		
4.40	2.00		

COMPUTED KINEMATIC PARAMETERS
ELEMENT ALPHA M
1 2.8854 1.667
3 .6898 1.549

HYDROGRAPH AT STATION SUB1A

DA MON HRMN	ORD	RAIN	LOSS	EXCESS	COMP Q	DA MON HRMN	ORD	RAIN	LOSS	EXCESS	COMP Q
30 AUG 0900	1	.00	.00	.00	0.	1 SEP 0900	49	.90	.01	.89	52.
30 AUG 1000	2	.00	.00	.00	0.	1 SEP 1000	50	1.30	.01	1.29	40.
30 AUG 1100	3	.00	.00	.00	0.	1 SEP 1100	51	1.30	.01	1.29	34.
30 AUG 1200	4	.00	.00	.00	0.	1 SEP 1200	52	4.40	.04	4.36	30.
30 AUG 1300	5	.00	.00	.00	0.	1 SEP 1300	53	2.70	.03	2.67	38.
30 AUG 1400	6	.10	.10	.00	0.	1 SEP 1400	54	2.00	.02	1.98	45.
30 AUG 1500	7	.10	.10	.00	0.	1 SEP 1500	55	.50	.00	.50	45.
30 AUG 1600	8	.10	.10	.00	0.	1 SEP 1600	56	.70	.01	.69	38.
30 AUG 1700	9	4.20	4.20	.00	0.	1 SEP 1700	57	1.90	.02	1.88	31.
30 AUG 1800	10	4.40	4.40	.00	0.	1 SEP 1800	58	.40	.00	.40	29.

30 AUG 1900	11	2.00	2.00	.00	0.	*	1 SEP 1900	59	.10	.00	.10	26.
30 AUG 2000	12	1.20	1.20	.00	0.	*	1 SEP 2000	60	.00	.00	.00	21.
30 AUG 2100	13	1.30	1.29	.01	0.	*	1 SEP 2100	61	.00	.00	.00	16.
30 AUG 2200	14	.80	4.77	.03	0.	*	1 SEP 2200	62	.00	.00	.00	13.
30 AUG 2300	15	5.40	4.76	.64	0.	*	1 SEP 2300	63	.00	.00	.00	10.
31 AUG 0000	16	1.60	1.27	.33	0.	*	2 SEP 0000	64	.00	.00	.00	9.
31 AUG 0100	17	.30	.23	.07	0.	*	2 SEP 0100	65	.00	.00	.00	7.
31 AUG 0200	18	1.90	1.43	.47	0.	*	2 SEP 0200	66	.00	.00	.00	6.
31 AUG 0300	19	1.90	1.36	.54	0.	*	2 SEP 0300	67	.00	.00	.00	5.
31 AUG 0400	20	3.00	2.01	.99	1.	*	2 SEP 0400	68	.00	.00	.00	4.
31 AUG 0500	21	6.50	3.87	2.63	2.	*	2 SEP 0500	69	.00	.00	.00	4.
31 AUG 0600	22	8.50	4.25	4.24	5.	*	2 SEP 0600	70	.00	.00	.00	3.
31 AUG 0700	23	11.59	4.70	6.89	18.	*	2 SEP 0700	71	.00	.00	.00	3.
31 AUG 0800	24	10.20	3.36	6.84	46.	*	2 SEP 0800	72	.00	.00	.00	3.
31 AUG 0900	25	12.39	3.36	9.03	81.	*	2 SEP 0900	73	.00	.00	.00	2.
31 AUG 1000	26	10.69	2.42	8.28	121.	*	2 SEP 1000	74	.00	.00	.00	2.
31 AUG 1100	27	13.39	2.55	10.84	146.	*	2 SEP 1100	75	.00	.00	.00	2.
31 AUG 1200	28	24.99	3.73	21.26	173.	*	2 SEP 1200	76	.00	.00	.00	2.
31 AUG 1300	29	24.99	2.81	22.18	267.	*	2 SEP 1300	77	.00	.00	.00	2.
31 AUG 1400	30	24.99	2.19	22.80	404.	*	2 SEP 1400	78	.00	.00	.00	1.
31 AUG 1500	31	34.38	2.33	32.05	436.	*	2 SEP 1500	79	.00	.00	.00	1.
31 AUG 1600	32	31.18	1.64	29.54	529.	*	2 SEP 1600	80	.00	.00	.00	1.
31 AUG 1700	33	38.58	1.60	36.98	589.	*	2 SEP 1700	81	.00	.00	.00	1.
31 AUG 1800	34	52.07	1.65	50.42	652.	*	2 SEP 1800	82	.00	.00	.00	1.
31 AUG 1900	35	60.47	1.43	59.04	860.	*	2 SEP 1900	83	.00	.00	.00	1.
31 AUG 2000	36	42.68	.80	41.88	1071.	*	2 SEP 2000	84	.00	.00	.00	1.
31 AUG 2100	37	23.39	.38	23.01	947.	*	2 SEP 2100	85	.00	.00	.00	1.
31 AUG 2200	38	19.29	.29	19.00	637.	*	2 SEP 2200	86	.00	.00	.00	1.
31 AUG 2300	39	23.89	.33	23.56	432.	*	2 SEP 2300	87	.00	.00	.00	1.
1 SEP 0000	40	11.49	.15	11.35	413.	*	3 SEP 0000	88	.00	.00	.00	1.
1 SEP 0100	41	9.40	.12	9.28	359.	*	3 SEP 0100	89	.00	.00	.00	1.
1 SEP 0200	42	15.59	.19	15.41	209.	*	3 SEP 0200	90	.00	.00	.00	1.
1 SEP 0300	43	30.59	.34	30.25	232.	*	3 SEP 0300	91	.00	.00	.00	1.
1 SEP 0400	44	16.59	.17	16.42	402.	*	3 SEP 0400	92	.00	.00	.00	1.
1 SEP 0500	45	5.00	.05	4.95	447.	*	3 SEP 0500	93	.00	.00	.00	0.
1 SEP 0600	46	2.60	.03	2.57	210.	*	3 SEP 0600	94	.00	.00	.00	0.
1 SEP 0700	47	.50	.00	.49	116.	*	3 SEP 0700	95	.00	.00	.00	0.
1 SEP 0800	48	1.50	.01	1.48	72.	*	3 SEP 0800	96	.00	.00	.00	0.

TOTAL RAINFALL = 611.90, TOTAL LOSS = 70.12, TOTAL EXCESS = 541.78

PEAK FLOW (CU M/S)	TIME (HR)	(CU M/S)	6-HR	MAXIMUM AVERAGE FLOW 24-HR	72-HR	95.00-HR
1071.	35.00	784.	408.	145.	110.	
		241.823	503.625	535.463	535.601	
		16928.	35254.	37482.	37492.	

CUMULATIVE AREA = 70.00 SQ KM

11800 58.0
11900 59.0
12000 60.0
12100 61.0
12200 62.0
12300 63.0
20000 64.0
20100 65.0
20200 66.0
20300 67.0
20400 68.0
20500 69.0
20600 70.0
20700 71.0
20800 72.0
20900 73.0
21000 74.0
21100 75.0
21200 76.0
21300 77.0
21400 78.0
21500 79.0
21600 80.0
21700 81.0
21800 82.0
21900 83.0
22000 84.0
22100 85.0
22200 86.0
22300 87.0
30000 88.0
30100 89.0
30200 90.0
30300 91.0
30400 92.0
30500 93.0
30600 94.0
30700 95.0
30800 96.0

*** **

*
* SUBTB *
*

RUNOFF FROM SUBBASIN 1B

SUBBASIN RUNOFF DATA

72 BA SUBBASIN CHARACTERISTICS
TAREA 70.00 SUBBASIN AREA

PRECIPITATION DATA

70 PT TOTAL STORM STATIONS 40 50
71 PW WEIGHTS .30 .70
68 PR RECORDING STATIONS 3 4
69 PW WEIGHTS .30 .70

73 LS SCS LOSS RATE 12.70 INITIAL ABSTRACTION
STRTL 80.00 CURVE NUMBER
GRVNR .00 PERCENT IMPERVIOUS AREA
RTIMP

74 UK KINEMATIC WAVE OVERLAND-FLOW ELEMENT NO. 1
L 2500. OVERLAND FLOW LENGTH
S .6000 SLOPE
N .4000 ROUGHNESS COEFFICIENT
PA 100.0 PERCENT OF SUBBASIN

75 RK MAIN CHANNEL CHANNEL LENGTH
L 19000.
S .0837 SLOPE
N .040 CHANNEL ROUGHNESS COEFFICIENT
CA 70.00 CONTRIBUTING AREA
SHAPE TRAP CHANNEL SHAPE
WD 45.00 BOTTOM WIDTH OR DIAMETER
Z 10.00 SIDE SLOPE
RUPSTQ NO ROUTE UPSTREAM HYDROGRAPH

PRECIPITATION STATION DATA

STATION TOTAL AVG. ANNUAL WEIGHT
40 450.10 .30
50 611.90 .70

TEMPORAL DISTRIBUTIONS

STATION 3, WEIGHT = .30
.00 .00 .00 .30 1.70
1.00 .00 .00 .00 .00
1.00 12.00 3.00 5.00 6.00
20.00 18.00 24.00 38.00 44.00
.00 .00 .00 .00 .00
.00 .00 .00 .00 .00

STATION 4, WEIGHT = .70

.00	.00	.10	4.20	4.40	2.00
1.20	1.30	1.60	1.90	3.00	6.50
8.50	11.60	10.70	25.00	34.40	34.40
31.20	38.60	42.70	23.90	11.50	9.40
15.60	30.60	2.60	.90	1.30	1.30
4.40	2.70	.70	.10		

COMPUTED KINEMATIC PARAMETERS

ELEMENT	ALPHA	M	DT (MIN)	DX (FT)
1	2.8854	1.667	60.00	4101.05
3	.6159	1.549	60.00	31167.98

HYDROGRAPH AT STATION SUB1B

DA	MON	HRMN	ORD	RAIN	LOSS	EXCESS	COMP Q	DA	MON	HRMN	ORD	RAIN	LOSS	EXCESS	COMP Q
30	AUG	0900	1	.00	.00	.00	0.	1	SEP	0900	49	.70	.01	.69	47.
30	AUG	1000	2	.00	.00	.00	0.	1	SEP	1000	50	1.01	.01	1.00	36.
30	AUG	1100	3	.00	.00	.00	0.	1	SEP	1100	51	1.01	.01	1.00	30.
30	AUG	1200	4	.00	.00	.00	0.	1	SEP	1200	52	3.41	.04	3.37	26.
30	AUG	1300	5	.10	.10	.00	0.	1	SEP	1300	53	2.09	.02	2.07	30.
30	AUG	1400	6	.64	.64	.00	0.	1	SEP	1400	54	1.55	.02	1.53	35.
30	AUG	1500	7	.41	.41	.00	0.	1	SEP	1500	55	.39	.00	.38	35.
30	AUG	1600	8	.18	.18	.00	0.	1	SEP	1600	56	.54	.01	.54	31.
30	AUG	1700	9	3.36	3.36	.00	0.	1	SEP	1700	57	1.47	.02	1.46	25.
30	AUG	1800	10	3.51	3.51	.00	0.	1	SEP	1800	58	.31	.00	.31	23.
30	AUG	1900	11	1.88	1.88	.00	0.	1	SEP	1900	59	.08	.00	.08	21.
30	AUG	2000	12	1.26	1.26	.00	0.	1	SEP	2000	60	.00	.00	.00	18.
30	AUG	2100	13	1.01	1.01	.00	0.	1	SEP	2100	61	.00	.00	.00	14.
30	AUG	2200	14	.62	.62	.00	0.	1	SEP	2200	62	.00	.00	.00	11.
30	AUG	2300	15	4.19	3.90	.29	0.	1	SEP	2300	63	.00	.00	.00	9.
31	AUG	0000	16	1.24	1.06	.18	0.	2	SEP	0000	64	.00	.00	.00	7.
31	AUG	0100	17	.23	.20	.04	0.	2	SEP	0100	65	.00	.00	.00	6.
31	AUG	0200	18	1.47	1.21	.27	0.	2	SEP	0200	66	.00	.00	.00	5.
31	AUG	0300	19	1.47	1.16	.32	0.	2	SEP	0300	67	.00	.00	.00	5.
31	AUG	0400	20	2.66	1.97	.68	0.	2	SEP	0400	68	.00	.00	.00	4.
31	AUG	0500	21	5.37	3.59	1.78	0.	2	SEP	0500	69	.00	.00	.00	4.
31	AUG	0600	22	6.92	3.98	2.95	1.	2	SEP	0600	70	.00	.00	.00	3.
31	AUG	0700	23	12.98	5.98	7.01	5.	2	SEP	0700	71	.00	.00	.00	3.
31	AUG	0800	24	8.91	3.28	5.63	25.	2	SEP	0800	72	.00	.00	.00	3.
31	AUG	0900	25	11.28	3.45	7.82	59.	2	SEP	0900	73	.00	.00	.00	2.
31	AUG	1000	26	10.29	2.63	7.66	94.	2	SEP	1000	74	.00	.00	.00	2.
31	AUG	1100	27	12.38	2.67	9.72	122.	2	SEP	1100	75	.00	.00	.00	2.
31	AUG	1200	28	24.70	4.14	20.56	151.	2	SEP	1200	76	.00	.00	.00	2.
31	AUG	1300	29	26.03	3.22	22.81	240.	2	SEP	1300	77	.00	.00	.00	2.
31	AUG	1400	30	24.04	2.29	21.74	394.	2	SEP	1400	78	.00	.00	.00	1.
31	AUG	1500	31	31.99	2.37	29.62	431.	2	SEP	1500	79	.00	.00	.00	1.
31	AUG	1600	32	30.84	1.77	29.06	494.	2	SEP	1600	80	.00	.00	.00	1.
31	AUG	1700	33	35.91	1.63	34.28	560.	2	SEP	1700	81	.00	.00	.00	1.
31	AUG	1800	34	48.37	1.69	46.68	616.	2	SEP	1800	82	.00	.00	.00	1.
31	AUG	1900	35	59.54	1.55	57.99	789.	2	SEP	1900	83	.00	.00	.00	1.
31	AUG	2000	36	47.73	.96	46.77	1022.	2	SEP	2000	84	.00	.00	.00	1.
31	AUG	2100	37	23.46	.40	23.06	993.	2	SEP	2100	85	.00	.00	.00	1.
31	AUG	2200	38	14.96	.24	14.73	684.	2	SEP	2200	86	.00	.00	.00	1.
31	AUG	2300	39	18.53	.28	18.25	412.	2	SEP	2300	87	.00	.00	.00	1.
1	SEP	0000	40	8.92	.13	8.79	327.	3	SEP	0000	88	.00	.00	.00	1.
1	SEP	0100	41	7.29	.10	7.19	267.	3	SEP	0100	89	.00	.00	.00	1.
1	SEP	0200	42	12.10	.16	11.94	182.	3	SEP	0200	90	.00	.00	.00	1.

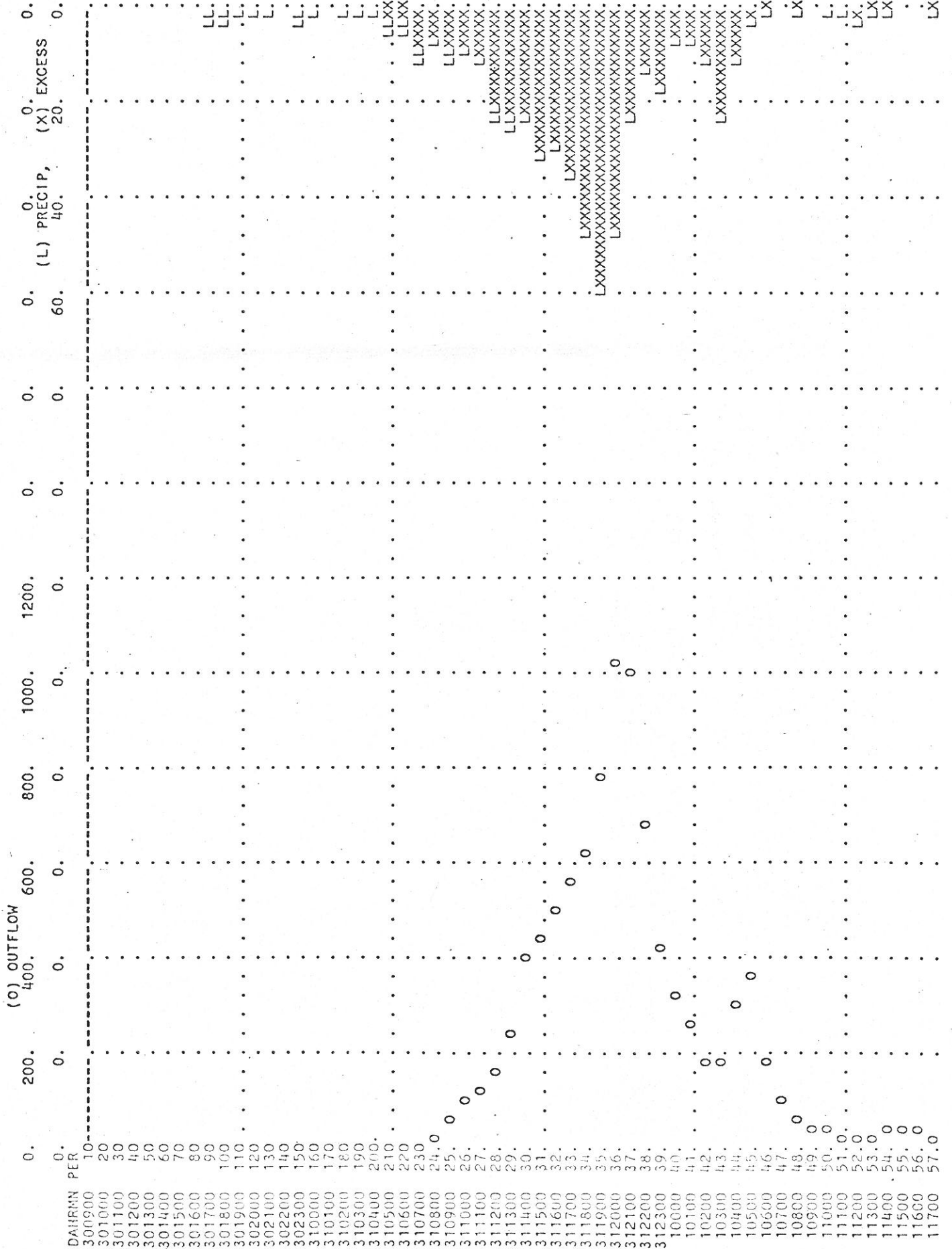
1 SEP 0300	43	23.73	.30	23.43	182.	*	3 SEP 0300	91	.00	.00	1.
1 SEP 0400	44	12.87	.15	12.72	290.	*	3 SEP 0400	92	.00	.00	1.
1 SEP 0500	45	3.88	.04	3.83	351.	*	3 SEP 0500	93	.00	.00	0.
1 SEP 0600	46	2.02	.02	1.99	188.	*	3 SEP 0600	94	.00	.00	0.
1 SEP 0700	47	.39	.00	.38	105.	*	3 SEP 0700	95	.00	.00	0.
1 SEP 0800	48	1.16	.01	1.15	67.	*	3 SEP 0800	96	.00	.00	0.

TOTAL RAINFALL = 563.36, TOTAL LOSS = 69.64, TOTAL EXCESS = 493.72

PEAK FLOW (CU M/S)	TIME (HR)	MAXIMUM AVERAGE FLOW 24-HR	72-HR
1022.	35.00	765.	132.
		(CU M/S)	(MM)
		236.041	488.875
		16523.	34221.
			100.
			488.959
			34227.

CUMULATIVE AREA = 70.00 SQ KM

STATION SUB1B



11800 58.0
11900 59.0
12000 60.0
12100 61.0
12200 62.0
12300 63.0
20000 64.0
20100 65.0
20200 66.0
20300 67.0
20400 68.0
20500 69.0
20600 70.0
20700 71.0
20800 72.0
20900 73.0
21000 74.0
21100 75.0
21200 76.0
21300 77.0
21400 78.0
21500 79.0
21600 80.0
21700 81.0
21800 82.0
21900 83.0
22000 84.0
22100 85.0
22200 86.0
22300 87.0
30000 88.0
30100 89.0
30200 90.0
30300 91.0
30400 92.0
30500 93.0
30600 94.0
30700 95.0
30800 96.0

*** ** ** ** ** ** ** ** ** ** ** ** ** ** ** ** ** ** ** ** ** ** ** ** ** ** ** ** ** ** ** ** ** ** ** ** ** ** ** ** ** ** ** ** ** ** ** ** ** ** ** ** ** ** ** ** ** ** ** ** ** **

*
* SUB1C *
*

COMBINE RUNOFF FROM SUB1A AND SUB1B

76 KK
78 HC
HYDROGRAPH COMBINATION 2 NUMBER OF HYDROGRAPHS TO COMBINE
ICOMP

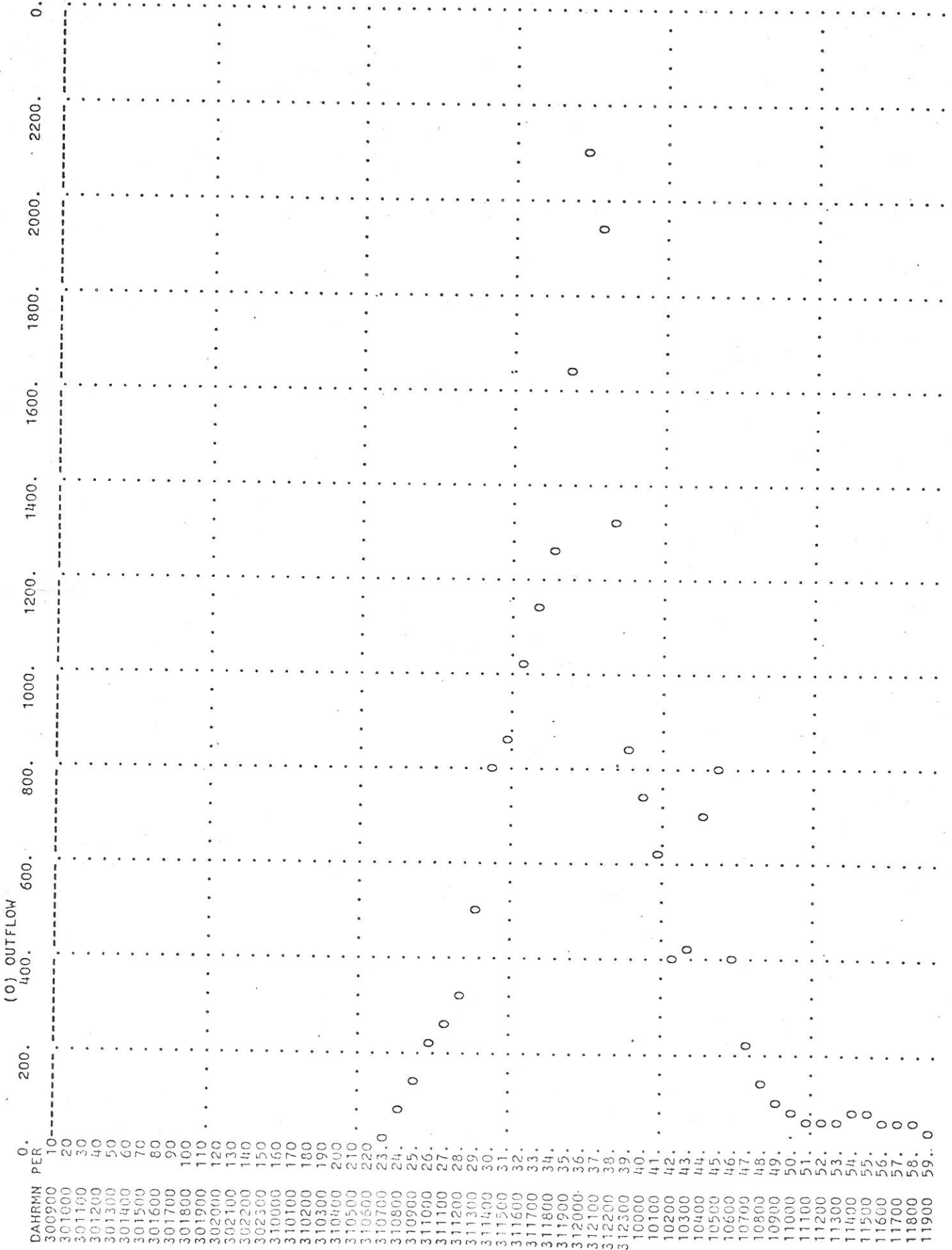
HYDROGRAPH AT STATION SUB1C
SUM OF 2 HYDROGRAPHS

DA	MON	HRMN	ORD	FLOW	DA	MON	HRMN	ORD	FLOW	DA	MON	HRMN	ORD	FLOW
30	AUG	0900	1	0.	31	AUG	0900	25	140.	1	SEP	0900	49	99.
30	AUG	1000	2	0.	31	AUG	1000	26	215.	1	SEP	1000	50	76.
30	AUG	1100	3	0.	31	AUG	1100	27	269.	1	SEP	1100	51	64.
30	AUG	1200	4	0.	31	AUG	1200	28	324.	1	SEP	1200	52	57.
30	AUG	1300	5	0.	31	AUG	1300	29	508.	1	SEP	1300	53	68.
30	AUG	1400	6	0.	31	AUG	1400	30	797.	1	SEP	1400	54	80.
30	AUG	1500	7	0.	31	AUG	1500	31	867.	1	SEP	1500	55	80.
30	AUG	1600	8	0.	31	AUG	1600	32	1023.	1	SEP	1600	56	69.
30	AUG	1700	9	0.	31	AUG	1700	33	1150.	1	SEP	1700	57	56.
30	AUG	1800	10	0.	31	AUG	1800	34	1267.	1	SEP	1800	58	53.
30	AUG	1900	11	0.	31	AUG	1900	35	1648.	1	SEP	1900	59	47.
30	AUG	2000	12	0.	31	AUG	2000	36	2093.	1	SEP	2000	60	39.
30	AUG	2100	13	0.	31	AUG	2100	37	1940.	1	SEP	2100	61	31.
30	AUG	2200	14	0.	31	AUG	2200	38	1321.	1	SEP	2200	62	24.
30	AUG	2300	15	0.	31	AUG	2300	39	844.	1	SEP	2300	63	19.
31	AUG	0000	16	0.	1	SEP	0000	40	740.	2	SEP	0000	64	16.
31	AUG	0100	17	0.	1	SEP	0100	41	623.	2	SEP	0100	65	13.
31	AUG	0200	18	0.	1	SEP	0200	42	391.	2	SEP	0200	66	11.
31	AUG	0300	19	0.	1	SEP	0300	43	414.	2	SEP	0300	67	10.
31	AUG	0400	20	1.	1	SEP	0400	44	692.	2	SEP	0400	68	8.
31	AUG	0500	21	2.	1	SEP	0500	45	798.	2	SEP	0500	69	7.
31	AUG	0600	22	6.	1	SEP	0600	46	398.	2	SEP	0600	70	6.
31	AUG	0700	23	23.	1	SEP	0700	47	221.	2	SEP	0700	71	6.
31	AUG	0800	24	71.	1	SEP	0800	48	139.	2	SEP	0800	72	5.

PEAK FLOW (CU M/S)	TIME (HR)	6-HR (CU M/S)	24-HR (CU M/S)	72-HR (CU M/S)	95.00-HR (CU M/S)
2093.	35.00	1545.	784.	277.	210.
		238.392	483.539	512.160	512.280
		33375.	67696.	71702.	71719.

CUMULATIVE AREA = 140.00 SQ KM

STATION SUB1C



**** ** ** ** ** ** ** ** ** ** ** ** ** ** ** ** ** ** ** ** ** ** ** ** ** ** ** ** ** ** ** ** ** ** ** ** ** **

* SUBIC *

RUNOFF FROM SUBBASIN 1C

SUBBASIN RUNOFF DATA

85 BA SUBBASIN CHARACTERISTICS
TAREA 45.00 SUBBASIN AREA

PRECIPITATION DATA

83 PT	TOTAL STORM STATIONS	20	40	50
84 PW	WEIGHTS	.20	.40	.40
81 PR	RECORDING STATIONS	3	4	7
82 PW	WEIGHTS	.40	.40	.20

86 LS SCS LOSS RATE 12.70 INITIAL ABSTRACTION
STR1L 80.00 CURVE NUMBER
CRVBR .00 PERCENT IMPERVIOUS AREA
RTIMP

87 UK KINEMATIC WAVE OVERLAND-FLOW ELEMENT NO. 1
L 4500. OVERLAND FLOW LENGTH
S .5000 SLOPE
N .300 ROUGHNESS COEFFICIENT
PA 100.0 PERCENT OF SUBBASIN

88 RK MAIN CHANNEL 7500. CHANNEL LENGTH
L .0133 SLOPE
S .026 CHANNEL ROUGHNESS COEFFICIENT
N 45.00 CONTRIBUTING AREA
CA TRAP CHANNEL SHAPE
SHAPE 85.00 BOTTOM WIDTH OR DIAMETER
WD 10.00 SIDE SLOPE
Z YES ROUTE UPSTREAM HYDROGRAPH
RUPSTQ

PRECIPITATION STATION DATA

STATION	TOTAL	AVG. ANNUAL	WEIGHT
20	300.00	.00	.20
40	450.10	.00	.40
50	611.90	.00	.40

TEMPORAL DISTRIBUTIONS

STATION	3,	WEIGHT = .40				
.00	.00	.00	.30	1.70	1.00	1.00
1.00	.00	.00	.00	.00	.00	1.00
1.00	12.00	3.00	5.00	6.00	6.00	16.00
20.00	18.00	24.00	38.00	44.00	16.00	16.00
.00	.00	.00	.00	.00	.00	.00
					.30	.30
					.00	.00
			20.00	20.00	1.00	1.00
			.00	.00	14.00	16.00
			.00	.00	.00	.00
			.00	.00	.00	.00

STATION 4, WEIGHT = .40

.00	.00	.00	.00	.00	.00	.00	.00	.00	.00	.00
1.20	1.30	.80	.00	1.10	1.60	1.10	1.10	1.90	4.20	2.00
8.50	11.60	10.20	5.40	10.70	12.40	13.40	13.40	1.90	1.90	6.50
31.20	38.60	52.10	60.50	42.70	60.50	23.40	23.40	25.00	25.00	34.40
15.60	30.60	16.60	5.00	2.60	5.00	.50	.50	11.50	11.50	9.40
4.40	2.70	2.00	.50	.70	.50	1.90	1.90	1.30	1.30	1.30

STATION 7, WEIGHT = .20

.00	.00	.00	.00	.00	.00	.00	.00	.00	.00	.00
8.00	3.00	3.50	3.50	5.00	5.00	5.00	5.00	10.00	9.00	12.00
12.00	12.00	16.00	12.00	64.00	64.00	56.00	56.00	10.00	10.00	12.00
10.00	.00	.00	.00	.00	.00	.00	.00	.00	.00	.00
.00	.00	.00	.00	.00	.00	.00	.00	.00	.00	.00

COMPUTED KINEMATIC PARAMETERS

ELEMENT	ALPHA	M	DT (MIN)	DX (FT)
1	3.5120	1.667	60.00	7381.89
3	.2183	1.593	60.00	12303.15

HYDROGRAPH AT STATION SUBIC

DA	MON	HRMN	ORD	RAIN	LOSS	EXCESS	COMP Q	DA	MON	HRMN	ORD	RAIN	LOSS	EXCESS	COMP Q
30	AUG	0900	1	.00	.00	.00	0.	1	SEP	0900	49	.42	.01	.42	151.
30	AUG	1000	2	.00	.00	.00	0.	1	SEP	1000	50	.61	.01	.60	113.
30	AUG	1100	3	.00	.00	.00	0.	1	SEP	1100	51	.61	.01	.60	91.
30	AUG	1200	4	.00	.00	.00	0.	1	SEP	1200	52	2.07	.03	2.04	78.
30	AUG	1300	5	.14	.14	.00	0.	1	SEP	1300	53	1.27	.02	1.25	80.
30	AUG	1400	6	.85	.85	.00	0.	1	SEP	1400	54	.94	.01	.93	90.
30	AUG	1500	7	.52	.52	.00	0.	1	SEP	1500	55	.24	.00	.23	95.
30	AUG	1600	8	.19	.19	.00	0.	1	SEP	1600	56	.33	.00	.33	88.
30	AUG	1700	9	2.12	2.12	.00	0.	1	SEP	1700	57	.90	.01	.88	76.
30	AUG	1800	10	2.21	2.21	.00	0.	1	SEP	1800	58	.19	.00	.19	67.
30	AUG	1900	11	1.41	1.41	.00	0.	1	SEP	1900	59	.05	.00	.05	61.
30	AUG	2000	12	1.04	1.04	.00	0.	1	SEP	2000	60	.00	.00	.00	53.
30	AUG	2100	13	.61	.61	.00	0.	1	SEP	2100	61	.00	.00	.00	44.
30	AUG	2200	14	.38	.38	.00	0.	1	SEP	2200	62	.00	.00	.00	36.
30	AUG	2300	15	2.54	2.54	.00	0.	1	SEP	2300	63	.00	.00	.00	29.
31	AUG	0000	16	.75	.75	.00	0.	2	SEP	0000	64	.00	.00	.00	24.
31	AUG	0100	17	.14	.14	.00	0.	2	SEP	0100	65	.00	.00	.00	20.
31	AUG	0200	18	.90	.88	.02	0.	2	SEP	0200	66	.00	.00	.00	17.
31	AUG	0300	19	1.84	1.73	.11	0.	2	SEP	0300	67	.00	.00	.00	15.
31	AUG	0400	20	2.12	1.88	.24	0.	2	SEP	0400	68	.00	.00	.00	13.
31	AUG	0500	21	3.77	3.06	.70	0.	2	SEP	0500	69	.00	.00	.00	11.
31	AUG	0600	22	6.36	4.51	1.85	1.	2	SEP	0600	70	.00	.00	.00	10.
31	AUG	0700	23	11.82	6.69	5.13	10.	2	SEP	0700	71	.00	.00	.00	9.
31	AUG	0800	24	7.04	3.22	3.83	44.	2	SEP	0800	72	.00	.00	.00	8.
31	AUG	0900	25	9.02	3.50	5.52	116.	2	SEP	0900	73	.00	.00	.00	7.
31	AUG	1000	26	9.04	2.96	6.09	207.	2	SEP	1000	74	.00	.00	.00	6.
31	AUG	1100	27	10.32	2.86	7.46	286.	2	SEP	1100	75	.00	.00	.00	6.
31	AUG	1200	28	21.67	4.70	16.97	360.	2	SEP	1200	76	.00	.00	.00	5.
31	AUG	1300	29	23.32	3.73	19.59	561.	2	SEP	1300	77	.00	.00	.00	5.
31	AUG	1400	30	21.43	2.63	18.81	898.	2	SEP	1400	78	.00	.00	.00	4.
31	AUG	1500	31	26.57	2.54	24.03	1062.	2	SEP	1500	79	.00	.00	.00	4.
31	AUG	1600	32	26.95	2.02	24.93	1246.	2	SEP	1600	80	.00	.00	.00	4.
31	AUG	1700	33	29.49	1.76	27.73	1414.	2	SEP	1700	81	.00	.00	.00	3.

31 AUG 1800	34	39.62	1.84	37.78	1557.	*	2 SEP 1800	82	.00	.00	3.
31 AUG 1900	35	49.23	1.73	47.50	1978.	*	2 SEP 1900	83	.00	.00	3.
31 AUG 2000	36	55.92	1.47	54.44	2545.	*	2 SEP 2000	84	.00	.00	3.
31 AUG 2100	37	31.75	.67	31.08	2599.	*	2 SEP 2100	85	.00	.00	2.
31 AUG 2200	38	11.45	.22	11.23	1964.	*	2 SEP 2200	86	.00	.00	2.
31 AUG 2300	39	13.61	.25	13.37	1258.	*	2 SEP 2300	87	.00	.00	2.
1 SEP 0000	40	7.30	.13	7.17	934.	*	3 SEP 0000	88	.00	.00	2.
1 SEP 0100	41	7.25	.12	7.13	784.	*	3 SEP 0100	89	.00	.00	2.
1 SEP 0200	42	9.70	.16	9.55	563.	*	3 SEP 0200	90	.00	.00	2.
1 SEP 0300	43	14.42	.22	14.19	520.	*	3 SEP 0300	91	.00	.00	2.
1 SEP 0400	44	7.82	.12	7.70	751.	*	3 SEP 0400	92	.00	.00	2.
1 SEP 0500	45	2.36	.03	2.32	892.	*	3 SEP 0500	93	.00	.00	1.
1 SEP 0600	46	1.22	.02	1.21	585.	*	3 SEP 0600	94	.00	.00	1.
1 SEP 0700	47	.24	.00	.23	345.	*	3 SEP 0700	95	.00	.00	1.
1 SEP 0800	48	.71	.01	.70	217.	*	3 SEP 0800	96	.00	.00	1.

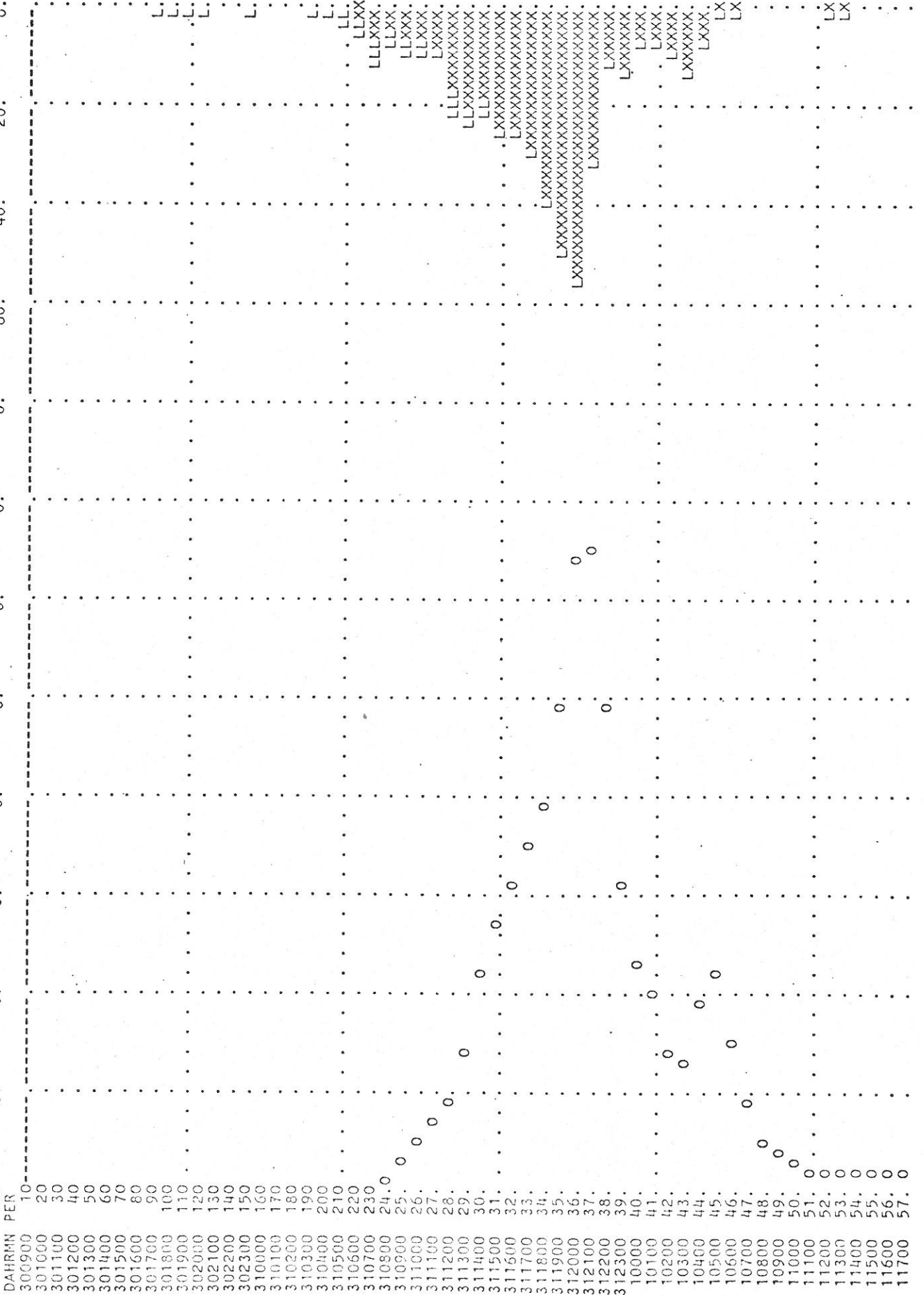
TOTAL RAINFALL = 484.80, TOTAL LOSS = 68.67, TOTAL EXCESS = 416.13

PEAK FLOW (CU M/S)	TIME (HR)	(CU M/S)	6-HR	24-HR	72-HR	95.00-HR
2599.	36.00	1996.	233.103	460.377	348.	264.
		(MM)	43124.	85170.	487.398	90180.
		(1000 CU M)			90169.	

CUMULATIVE AREA = 185.00 SQ KM

STATION SUBTC

(O) OUTFLOW 800. 0. 400. 0. 800. 0. 1200. 0. 1600. 0. 2000. 0. 2400. 0. 2800. 0. (L) PRECIP, 40. 0. 60. 0. 80. 0. (X) EXCESS 20. 0. 40. 0.



11800	58.0
11900	59.0
12000	60.0
12100	61.0
12200	62.0
12300	63.0
20000	64.0
20100	65.0
20200	66.0
20300	67.0
20400	68.0
20500	69.0
20600	70.0
20700	71.0
20800	72.0
20900	73.0
21000	74.0
21100	75.0
21200	76.0
21300	77.0
21400	78.0
21500	79.0
21600	80.0
21700	81.0
21800	82.0
21900	83.0
22000	84.0
22100	85.0
22200	86.0
22300	87.0
30000	88.0
30100	89.0
30200	90.0
30300	91.0
30400	92.0
30500	93.0
30600	94.0
30700	95.0
30800	96.0

**** **

*
* SUB2A *
*

RUNOFF FROM SUBBASIN 2A

SUBBASIN RUNOFF DATA

95 BA SUBBASIN CHARACTERISTICS
TAREA 24.00 SUBBASIN AREA

PRECIPITATION DATA

93 PT TOTAL STORM STATIONS 20 40
94 PW WEIGHTS .20 .80
91 PR RECORDING STATIONS 3 7
92 PW WEIGHTS .80 .20

96 LS SCS LOSS RATE 12.70 INITIAL ABSTRACTION
STRTL 80.00 CURVE NUMBER
GRVNR .00 PERCENT IMPERVIOUS AREA
RTIMP

KINEMATIC WAVE

OVERLAND-FLOW ELEMENT NO. 1
L 2000. OVERLAND FLOW LENGTH
S .4000 SLOPE
N .300 ROUGHNESS COEFFICIENT
PA 100.0 PERCENT OF SUBBASIN

98 RK MAIN CHANNEL

L 6500. CHANNEL LENGTH
S .0108 SLOPE
N .022 CHANNEL ROUGHNESS COEFFICIENT
CA 24.00 CONTRIBUTING AREA
SHAPE TRAP CHANNEL SHAPE
WD 85.00 BOTTOM WIDTH OR DIAMETER
Z 10.00 SIDE SLOPE
RUPSTQ YES ROUTE UPSTREAM HYDROGRAPH

PRECIPITATION STATION DATA

STATION TOTAL AVG. ANNUAL WEIGHT
20 300.00 .00 .20
40 450.10 .80

TEMPORAL DISTRIBUTIONS

STATION 3, WEIGHT = .80
.00 .00 .00 .30 1.70
1.00 .00 .00 .00 .00
1.00 12.00 3.00 5.00 6.00
20.00 18.00 24.00 38.00 44.00
.00 .00 .00 .00 .00
.00 .00 .00 .00 .00

STATION 7, WEIGHT = .20
 .00 .00 .00 .00 .00 .00 .00 .00 .00 .00
 .00 .00 .00 .00 .00 .00 .00 .00 .00 .00
 8.00 3.00 3.50 3.50 5.00 5.00 10.00 10.00 10.00 1.00
 12.00 12.00 16.00 12.00 64.00 64.00 56.00 10.00 10.00 13.00
 10.00 .00 .00 .00 .00 .00 .00 .00 .00 .00
 .00 .00 .00 .00 .00 .00 .00 .00 .00 .00

COMPUTED KINEMATIC PARAMETERS
 ELEMENT ALPHA M DT (MIN) DX (FT)
 1 3.1412 1.667 60.00 3280.84
 3 .2325 1.593 60.00 10662.73

HYDROGRAPH AT STATION SUB2A

DA	MON	HRMN	ORD	RAIN	LOSS	EXCESS	COMP Q	DA	MON	HRMN	ORD	RAIN	LOSS	EXCESS	COMP Q
30	AUG	0900	1	.00	.00	.00	0.	1	SEP	0900	49	.00	.00	.00	178.
30	AUG	1000	2	.00	.00	.00	0.	1	SEP	1000	50	.00	.00	.00	131.
30	AUG	1100	3	.00	.00	.00	0.	1	SEP	1100	51	.00	.00	.00	103.
30	AUG	1200	4	.00	.00	.00	0.	1	SEP	1200	52	.00	.00	.00	86.
30	AUG	1300	5	.37	.37	.00	0.	1	SEP	1300	53	.00	.00	.00	81.
30	AUG	1400	6	2.09	2.09	.00	0.	1	SEP	1400	54	.00	.00	.00	86.
30	AUG	1500	7	1.23	1.23	.00	0.	1	SEP	1500	55	.00	.00	.00	93.
30	AUG	1600	8	.37	.37	.00	0.	1	SEP	1600	56	.00	.00	.00	91.
30	AUG	1700	9	.37	.37	.00	0.	1	SEP	1700	57	.00	.00	.00	82.
30	AUG	1800	10	.37	.37	.00	0.	1	SEP	1800	58	.00	.00	.00	72.
30	AUG	1900	11	1.23	1.23	.00	0.	1	SEP	1900	59	.00	.00	.00	65.
30	AUG	2000	12	1.23	1.23	.00	0.	1	SEP	2000	60	.00	.00	.00	57.
30	AUG	2100	13	.00	.00	.00	0.	1	SEP	2100	61	.00	.00	.00	49.
30	AUG	2200	14	.00	.00	.00	0.	1	SEP	2200	62	.00	.00	.00	41.
30	AUG	2300	15	.00	.00	.00	0.	1	SEP	2300	63	.00	.00	.00	34.
31	AUG	0000	16	.00	.00	.00	0.	2	SEP	0000	64	.00	.00	.00	28.
31	AUG	0100	17	.00	.00	.00	0.	2	SEP	0100	65	.00	.00	.00	24.
31	AUG	0200	18	.00	.00	.00	0.	2	SEP	0200	66	.00	.00	.00	17.
31	AUG	0300	19	1.23	1.23	.00	0.	2	SEP	0300	67	.00	.00	.00	11.
31	AUG	0400	20	1.54	1.54	.00	0.	2	SEP	0400	68	.00	.00	.00	15.
31	AUG	0500	21	1.54	1.54	.00	0.	2	SEP	0500	69	.00	.00	.00	13.
31	AUG	0600	22	3.69	3.59	.10	0.	2	SEP	0600	70	.00	.00	.00	10.
31	AUG	0700	23	15.67	11.71	3.96	0.	2	SEP	0700	71	.00	.00	.00	9.
31	AUG	0800	24	4.76	2.72	2.04	28.	2	SEP	0800	72	.00	.00	.00	8.
31	AUG	0900	25	7.22	3.59	3.62	85.	2	SEP	0900	73	.00	.00	.00	7.
31	AUG	1000	26	8.91	3.74	5.17	182.	2	SEP	1000	74	.00	.00	.00	7.
31	AUG	1100	27	8.91	3.14	5.77	279.	2	SEP	1100	75	.00	.00	.00	6.
31	AUG	1200	28	22.73	6.12	16.61	367.	2	SEP	1200	76	.00	.00	.00	6.
31	AUG	1300	29	27.34	5.08	22.25	577.	2	SEP	1300	77	.00	.00	.00	6.
31	AUG	1400	30	21.19	2.89	18.30	956.	2	SEP	1400	78	.00	.00	.00	5.
31	AUG	1500	31	23.34	2.50	20.85	1162.	2	SEP	1500	79	.00	.00	.00	5.
31	AUG	1600	32	28.26	2.36	25.90	1345.	2	SEP	1600	80	.00	.00	.00	4.
31	AUG	1700	33	25.80	1.71	24.10	1544.	2	SEP	1700	81	.00	.00	.00	4.
31	AUG	1800	34	34.40	1.81	32.59	1699.	2	SEP	1800	82	.00	.00	.00	4.
31	AUG	1900	35	50.38	2.00	48.38	2111.	2	SEP	1900	83	.00	.00	.00	3.
31	AUG	2000	36	73.72	2.06	71.66	2752.	2	SEP	2000	84	.00	.00	.00	3.
31	AUG	2100	37	36.86	.78	36.08	3020.	2	SEP	2100	85	.00	.00	.00	3.
31	AUG	2200	38	3.07	.06	3.01	2360.	2	SEP	2200	86	.00	.00	.00	3.
31	AUG	2300	39	3.07	.06	3.01	1498.	2	SEP	2300	87	.00	.00	.00	2.
1	SEP	0000	40	2.46	.05	2.41	1010.	3	SEP	0000	88	.00	.00	.00	2.
1	SEP	0100	41	3.69	.07	3.62	829.	3	SEP	0100	89	.00	.00	.00	2.
1	SEP	0200	42	3.07	.06	3.02	628.	3	SEP	0200	90	.00	.00	.00	2.

1 SEP 0300	43	.00	.00	.00	552.	*	3 SEP 0300	91	.00	.00	2.
1 SEP 0400	44	.00	.00	.00	716.	*	3 SEP 0400	92	.00	.00	2.
1 SEP 0500	45	.00	.00	.00	871.	*	3 SEP 0500	93	.00	.00	2.
1 SEP 0600	46	.00	.00	.00	651.	*	3 SEP 0600	94	.00	.00	1.
1 SEP 0700	47	.00	.00	.00	409.	*	3 SEP 0700	95	.00	.00	1.
1 SEP 0800	48	.00	.00	.00	260.	*	3 SEP 0800	96	.00	.00	1.

TOTAL RAINFALL = 420.08, TOTAL LOSS = 67.64, TOTAL EXCESS = 352.44

PEAK FLOW	TIME	6-HR	24-HR	72-HR	95.00-HR
(CU M/S)	(HR)				
3020.	36.00	2244.	1080.	380.	288.
		(MM)	446.642	471.527	471.552
		(1000 CU M)	93348.	98549.	98554.

CUMULATIVE AREA = 209.00 SQ KM

11800 58.0
11900 59.0
12000 60.0
12100 61.0
12200 62.0
12300 63.0
20000 64.0
20100 65.0
20200 65.0
20300 67.0
20400 68.0
20500 69.0
20600 70.0
20700 71.0
20800 72.0
20900 73.0
21000 74.0
21100 75.0
21200 76.0
21300 77.0
21400 78.0
21500 79.0
21600 80.0
21700 81.0
21800 82.0
21900 83.0
22000 84.0
22100 85.0
22200 86.0
22300 87.0
30000 88.0
30100 89.0
30200 90.0
30300 91.0
30400 92.0
30500 93.0
30600 94.0
30700 95.0
30800 96.0

STATION 7. WEIGHT = .35
 .00 .00 .00 .00 .00 .00 .00 .00 .00 .00
 .00 .00 .00 .00 .00 .00 .00 .00 .00 .00
 8.00 3.00 3.50 5.00 5.00 5.00 5.00 5.00 5.00 5.00
 12.00 12.00 16.00 12.00 64.00 64.00 64.00 64.00 64.00 64.00
 10.00

COMPUTED KINEMATIC PARAMETERS
 ELEMENT ALPHA M DT (MIN) DX (FT)
 1 2.2212 1.667 60.00 5741.47
 3 .6034 1.549 60.00 18864.83

HYDROGRAPH AT STATION SUB2B

DA	MON	HRMN	ORD	RAIN	LOSS	EXCESS	COMP Q	#	DA	MON	HRMN	ORD	RAIN	LOSS	EXCESS	COMP Q	#
30	AUG	0900	1	.00	.00	.00	0.	*	1	SEP	0900	49	.00	.00	.00	18.	*
30	AUG	1000	2	.00	.00	.00	0.	*	1	SEP	1000	50	.00	.00	.00	15.	*
30	AUG	1100	3	.00	.00	.00	0.	*	1	SEP	1100	51	.00	.00	.00	12.	*
30	AUG	1200	4	.00	.00	.00	0.	*	1	SEP	1200	52	.00	.00	.00	10.	*
30	AUG	1300	5	.26	.26	.00	0.	*	1	SEP	1300	53	.00	.00	.00	9.	*
30	AUG	1400	6	1.50	1.50	.00	0.	*	1	SEP	1400	54	.00	.00	.00	8.	*
30	AUG	1500	7	.88	.88	.00	0.	*	1	SEP	1500	55	.00	.00	.00	7.	*
30	AUG	1600	8	.26	.26	.00	0.	*	1	SEP	1600	56	.00	.00	.00	6.	*
30	AUG	1700	9	.26	.26	.00	0.	*	1	SEP	1700	57	.00	.00	.00	5.	*
30	AUG	1800	10	.26	.26	.00	0.	*	1	SEP	1800	58	.00	.00	.00	5.	*
30	AUG	1900	11	.88	.88	.00	0.	*	1	SEP	1900	59	.00	.00	.00	4.	*
30	AUG	2000	12	.88	.88	.00	0.	*	1	SEP	2000	60	.00	.00	.00	4.	*
30	AUG	2100	13	.00	.00	.00	0.	*	1	SEP	2100	61	.00	.00	.00	3.	*
30	AUG	2200	14	.00	.00	.00	0.	*	1	SEP	2200	62	.00	.00	.00	3.	*
30	AUG	2300	15	.00	.00	.00	0.	*	1	SEP	2300	63	.00	.00	.00	3.	*
31	AUG	0000	16	.00	.00	.00	0.	*	2	SEP	0000	64	.00	.00	.00	2.	*
31	AUG	0100	17	.00	.00	.00	0.	*	2	SEP	0100	65	.00	.00	.00	2.	*
31	AUG	0200	18	.00	.00	.00	0.	*	2	SEP	0200	66	.00	.00	.00	2.	*
31	AUG	0300	19	1.90	1.90	.00	0.	*	2	SEP	0300	67	.00	.00	.00	2.	*
31	AUG	0400	20	1.36	1.36	.00	0.	*	2	SEP	0400	68	.00	.00	.00	2.	*
31	AUG	0500	21	1.36	1.36	.00	0.	*	2	SEP	0500	69	.00	.00	.00	2.	*
31	AUG	0600	22	4.68	4.63	.05	0.	*	2	SEP	0600	70	.00	.00	.00	1.	*
31	AUG	0700	23	12.00	9.59	2.41	0.	*	2	SEP	0700	71	.00	.00	.00	1.	*
31	AUG	0800	24	4.31	2.75	1.55	0.	*	2	SEP	0800	72	.00	.00	.00	1.	*
31	AUG	0900	25	6.07	3.42	2.65	1.	*	2	SEP	0900	73	.00	.00	.00	1.	*
31	AUG	1000	26	7.66	3.70	3.97	3.	*	2	SEP	1000	74	.00	.00	.00	1.	*
31	AUG	1100	27	7.66	3.15	4.52	10.	*	2	SEP	1100	75	.00	.00	.00	1.	*
31	AUG	1200	28	18.85	6.06	12.79	18.	*	2	SEP	1200	76	.00	.00	.00	1.	*
31	AUG	1300	29	21.90	5.04	16.86	49.	*	2	SEP	1300	77	.00	.00	.00	1.	*
31	AUG	1400	30	18.51	3.20	15.31	111.	*	2	SEP	1400	78	.00	.00	.00	1.	*
31	AUG	1500	31	19.80	2.70	17.10	169.	*	2	SEP	1500	79	.00	.00	.00	1.	*
31	AUG	1600	32	23.33	2.52	20.81	215.	*	2	SEP	1600	80	.00	.00	.00	1.	*
31	AUG	1700	33	21.56	1.87	19.70	263.	*	2	SEP	1700	81	.00	.00	.00	1.	*
31	AUG	1800	34	28.75	2.00	26.75	289.	*	2	SEP	1800	82	.00	.00	.00	1.	*
31	AUG	1900	35	39.19	2.10	37.10	341.	*	2	SEP	1900	83	.00	.00	.00	1.	*
31	AUG	2000	36	69.17	2.60	66.57	460.	*	2	SEP	2000	84	.00	.00	.00	1.	*
31	AUG	2100	37	40.69	1.11	39.57	744.	*	2	SEP	2100	85	.00	.00	.00	1.	*
31	AUG	2200	38	4.75	.12	4.63	820.	*	2	SEP	2200	86	.00	.00	.00	1.	*
31	AUG	2300	39	4.75	.11	4.63	338.	*	2	SEP	2300	87	.00	.00	.00	1.	*
1	SEP	0000	40	3.80	.09	3.71	180.	*	3	SEP	0000	88	.00	.00	.00	1.	*
1	SEP	0100	41	5.70	.13	5.57	123.	*	3	SEP	0100	89	.00	.00	.00	1.	*
1	SEP	0200	42	4.75	.11	4.64	103.	*	3	SEP	0200	90	.00	.00	.00	0.	*
1	SEP	0300	43	.00	.00	.00	91.	*	3	SEP	0300	91	.00	.00	.00	0.	*

1 SEP 0400	44	.00	.00	.00	68.	*	3 SEP 0400	92	.00	.00	0.
1 SEP 0500	45	.00	.00	.00	50.	*	3 SEP 0500	93	.00	.00	0.
1 SEP 0600	46	.00	.00	.00	37.	*	3 SEP 0600	94	.00	.00	0.
1 SEP 0700	47	.00	.00	.00	28.	*	3 SEP 0700	95	.00	.00	0.
1 SEP 0800	48	.00	.00	.00	22.	*	3 SEP 0800	96	.00	.00	0.

TOTAL RAINFALL = 377.68, TOTAL LOSS = 66.79, TOTAL EXCESS = 310.89

PEAK FLOW TIME

(CU M/S) (HR)

820. 37.00

(CU M/S)

6-HR

24-HR

72-HR

95.00-HR

.00

.00

0.

(MM)

492.

190.

65.

49.

.00

.00

0.

(1000 CU M)

189.936

293.428

300.679

300.683

.00

.00

0.

CUMULATIVE AREA = 56.00 SQ KM

16838.

.00

.00

0.

STATION SUB2B

DAHRMN PER	(O) OUTFLOW	PRECIP.	EXCESS
	0. 100. 200. 300. 400. 500. 600. 700. 800. 900.	(L) 40.	(X) 20.
300900	10		
301000	20		
301100	30		
301200	40		
301300	50		
301400	60		
301500	70		
301600	80		
301700	90		
301800	100		
301900	110		
302000	120		
302100	130		
302200	140		
302300	150		
310000	160		
310100	170		
310200	180		
310300	190		
310400	200		
310500	210		
310600	220		
310700	230		
310800	240		
310900	250		
311000	260		
311100	27.0		
311200	28.0		
311300	29.0		
311400	30.0		
311500	31.0		
311600	32.0		
311700	33.0		
311800	34.0		
311900	35.0		
312000	36.0		
312100	37.0		
312200	38.0		
312300	39.0		
10000	40.0		
10100	41.0		
10200	42.0		
10300	43.0		
10400	44.0		
10500	45.0		
10600	46.0		
10700	47.0		
10800	48.0		
10900	49.0		
11000	50.0		
11100	51.0		
11200	52.0		
11300	53.0		
11400	54.0		
11500	55.0		
11600	56.0		
11700	57.0		

11800
11900
12000
12100
12200
12300
20000
20100
20200
20300
20400
20500
20600
20700
20800
20900
21000
21100
21200
21300
21400
21500
21600
21700
21800
21900
22000
22100
22200
22300
30000
30100
30200
30300
30400
30500
30600
30700
30800

*** ** ** ** ** ** ** ** ** ** ** ** ** ** ** ** ** ** ** ** ** ** ** ** ** ** ** ** ** ** ** ** ** ** ** ** ** ** ** ** ** ** ** ** ** ** ** ** ** ** ** ** ** ** ** ** **

 # SUB2C #
 # *****

COMBINE RUNOFF FROM SUB2A AND SUB2B

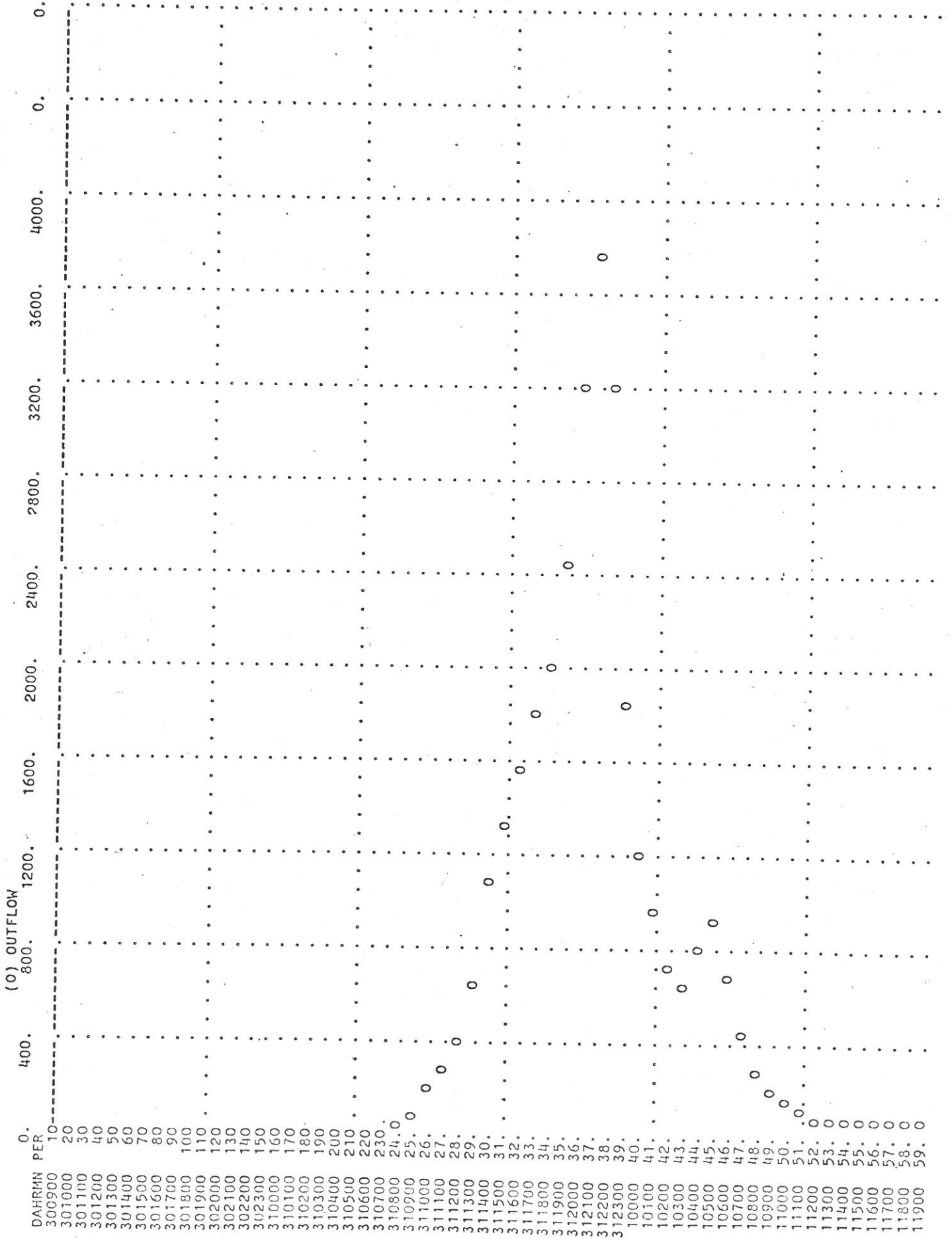
111 HC HYDROGRAPH COMBINATION 2 NUMBER OF HYDROGRAPHS TO COMBINE
 ICOMP

HYDROGRAPH AT STATION SUB2C
 SUM OF 2 HYDROGRAPHS

DA	MON	HRMN	ORD	FLOW	DA	MON	HRMN	ORD	FLOW	DA	MON	HRMN	ORD	FLOW	DA	MON	HRMN	ORD	FLOW
30	AUG	0900	1	0.	31	AUG	0900	25	86.	1	SEP	0900	49	196.	2	SEP	0900	73	9.
30	AUG	1000	2	0.	31	AUG	1000	26	185.	1	SEP	1000	50	146.	2	SEP	1000	74	8.
30	AUG	1100	3	0.	31	AUG	1100	27	288.	1	SEP	1100	51	115.	2	SEP	1100	75	8.
30	AUG	1200	4	0.	31	AUG	1200	28	386.	1	SEP	1200	52	96.	2	SEP	1200	76	7.
30	AUG	1300	5	0.	31	AUG	1300	29	626.	1	SEP	1300	53	90.	2	SEP	1300	77	6.
30	AUG	1400	6	0.	31	AUG	1400	30	1067.	1	SEP	1400	54	94.	2	SEP	1400	78	6.
30	AUG	1500	7	0.	31	AUG	1500	31	1332.	1	SEP	1500	55	100.	2	SEP	1500	79	5.
30	AUG	1600	8	0.	31	AUG	1600	32	1560.	1	SEP	1600	56	97.	2	SEP	1600	80	5.
30	AUG	1700	9	0.	31	AUG	1700	33	1806.	1	SEP	1700	57	87.	2	SEP	1700	81	5.
30	AUG	1800	10	0.	31	AUG	1800	34	1987.	1	SEP	1800	58	77.	2	SEP	1800	82	4.
30	AUG	1900	11	0.	31	AUG	1900	35	2452.	1	SEP	1900	59	69.	2	SEP	1900	83	4.
30	AUG	2000	12	0.	31	AUG	2000	36	3212.	1	SEP	2000	60	61.	2	SEP	2000	84	4.
30	AUG	2100	13	0.	31	AUG	2100	37	3764.	1	SEP	2100	61	53.	2	SEP	2100	85	4.
30	AUG	2200	14	0.	31	AUG	2200	38	3181.	1	SEP	2200	62	44.	2	SEP	2200	86	3.
30	AUG	2300	15	0.	31	AUG	2300	39	1836.	1	SEP	2300	63	37.	2	SEP	2300	87	3.
31	AUG	0000	16	0.	1	SEP	0000	40	1190.	2	SEP	0000	64	31.	3	SEP	0000	88	3.
31	AUG	0100	17	0.	1	SEP	0100	41	953.	2	SEP	0100	65	26.	3	SEP	0100	89	3.
31	AUG	0200	18	0.	1	SEP	0200	42	731.	2	SEP	0200	66	22.	3	SEP	0200	90	3.
31	AUG	0300	19	0.	1	SEP	0300	43	643.	2	SEP	0300	67	19.	3	SEP	0300	91	2.
31	AUG	0400	20	0.	1	SEP	0400	44	784.	2	SEP	0400	68	17.	3	SEP	0400	92	2.
31	AUG	0500	21	0.	1	SEP	0500	45	920.	2	SEP	0500	69	15.	3	SEP	0500	93	2.
31	AUG	0600	22	0.	1	SEP	0600	46	688.	2	SEP	0600	70	13.	3	SEP	0600	94	2.
31	AUG	0700	23	0.	1	SEP	0700	47	438.	2	SEP	0700	71	12.	3	SEP	0700	95	2.
31	AUG	0800	24	28.	1	SEP	0800	48	282.	2	SEP	0800	72	10.	3	SEP	0800	96	2.

PEAK FLOW (CU M/S)	TIME (HR)	6-HR (CU M/S)	MAXIMUM AVERAGE FLOW 24-HR	72-HR	95.00-HR
3764.	36.00	2736.	1270.	445.	337.
		223.022	414.168	435.419	435.444
		59101.	109754.	115386.	115393.
CUMULATIVE AREA = 265.00 SQ KM					

STATION SUB2C



12000	60.0
12100	61.0
12200	62.0
12300	63.0
20000	64.0
20100	65.0
20200	66.0
20300	67.0
20400	68.0
20500	69.0
20600	70.0
20700	71.0
20800	72.0
20900	73.0
21000	74.0
21100	75.0
21200	76.0
21300	77.0
21400	78.0
21500	79.0
21600	80.0
21700	81.0
21800	82.0
21900	83.0
22000	84.0
22100	85.0
22200	86.0
22300	87.0
30000	88.0
30100	89.0
30200	90.0
30300	91.0
30400	92.0
30500	93.0
30600	94.0
30700	95.0
30800	96.0

*** ** ** ** ** ** ** ** ** ** ** ** ** ** ** ** ** ** ** ** ** ** ** ** **

 * RESER *

ROUTE FLOW THROUGH RESERVOIR

HYDROGRAPH ROUTING DATA

STATION	STORAGE ROUTING	NSIPS	IJYP	RSVRIC	X	1	NUMBER OF SUBREACHES	0.0	TYPE OF INITIAL CONDITION	.00	INITIAL CONDITION	.00	WORKING R	AND D	COEFFICIENT
112 KK	STORAGE	3472.2	3.5	27.8	93.8	222.2	434.0	750.0	1191.0	1777.8	2531.3				
114 RS	ELEVATION	5.00	1.00	6.00	1.50	2.00	2.50	3.00	3.50	4.00	4.50				
115 SV	DISCHARGE	2192.	49.	156.	305.	490.	707.	953.	1227.	1525.	1847.				

*** WARNING *** MODIFIED PULS ROUTING MAY BE NUMERICALLY UNSTABLE FOR OUTFLOWS BETWEEN 0. TO 1227.
 THE ROUTED HYDROGRAPH SHOULD BE EXAMINED FOR OSCILLATIONS OR OUTFLOWS GREATER THAN PEAK INFLOWS.
 THIS CAN BE CORRECTED BY DECREASING THE TIME INTERVAL OR INCREASING STORAGE (USE A LONGER REACH.)

HYDROGRAPH AT STATION RESER

DA	MON	HRMN	ORD	OUTFLOW	STORAGE	STAGE	DA	MON	HRMN	ORD	OUTFLOW	STORAGE	STAGE
30	AUG	0900	1	0.	0.	4.2	2	SEP	0100	65	26.	1.8	3
30	AUG	1000	2	0.	2091.1	4.5	2	SEP	0200	66	22.	1.6	.2
30	AUG	1100	3	0.	2577.9	4.9	2	SEP	0300	67	19.	1.4	.2
30	AUG	1200	4	0.	3348.3	5.6	2	SEP	0400	68	17.	1.2	.2
30	AUG	1300	5	0.	4923.2	6.3	2	SEP	0500	69	15.	1.0	.1
30	AUG	1400	6	0.	6948.3	6.6	2	SEP	0600	70	13.	.9	.1
30	AUG	1500	7	0.	7579.4	5.8	2	SEP	0700	71	12.	.8	.1
30	AUG	1600	8	0.	5481.7	4.5	2	SEP	0800	72	10.	.7	.1
30	AUG	1700	9	0.	2554.5	3.3	2	SEP	0900	73	9.	.6	.1
30	AUG	1800	10	0.	1036.6	2.7	2	SEP	1000	74	9.	.6	.1
30	AUG	1900	11	0.	569.5	2.4	2	SEP	1100	75	8.	.5	.1
30	AUG	2000	12	0.	389.4	2.5	2	SEP	1200	76	7.	.5	.1
30	AUG	2100	13	0.	459.2	2.8	2	SEP	1300	77	6.	.5	.1
30	AUG	2200	14	0.	647.5	2.7	2	SEP	1400	78	6.	.4	.1
30	AUG	2300	15	0.	544.0	2.0	2	SEP	1500	79	5.	.4	.1
31	AUG	0000	16	0.	236.2	1.5	2	SEP	1600	80	5.	.4	.1
31	AUG	0100	17	0.	90.3	1.2	2	SEP	1700	81	5.	.3	.0
31	AUG	0200	18	0.	49.1	1.0	2	SEP	1800	82	4.	.3	.0
31	AUG	0300	19	0.	26.5	.8	2	SEP	1900	83	4.	.3	.0
31	AUG	0400	20	0.	18.6	.7	2	SEP	2000	84	4.	.3	.0

```

31 AUG 0500 21 0.0 0.0 0.0 0.0 1 SEP 1300 53 89.0 12.5 0.7 * 2 SEP 2100 85 4.0 0.3
31 AUG 0600 22 0.0 0.0 0.0 0.0 1 SEP 1400 54 94.0 13.7 0.7 * 2 SEP 2200 86 3.0 0.2
31 AUG 0700 23 0.0 0.0 0.0 0.0 1 SEP 1500 55 99.0 14.8 0.7 * 2 SEP 2300 87 3.0 0.2
31 AUG 0800 24 27.0 1.9 0.3 * 1 SEP 1600 56 98.0 14.6 0.7 * 3 SEP 0000 88 3.0 0.2
31 AUG 0900 25 82.0 11.0 0.7 * 1 SEP 1700 57 87.0 12.2 0.7 * 3 SEP 0100 89 3.0 0.2
31 AUG 1000 26 175.0 36.3 1.1 * 1 SEP 1800 58 78.0 9.9 0.6 * 3 SEP 0200 90 3.0 0.2
31 AUG 1100 27 274.0 80.1 1.4 * 1 SEP 1900 59 69.0 8.0 0.6 * 3 SEP 0300 91 2.0 0.2
31 AUG 1200 28 368.0 137.5 1.7 * 1 SEP 2000 60 62.0 6.3 0.5 * 3 SEP 0400 92 2.0 0.2
31 AUG 1300 29 559.0 289.7 2.2 * 1 SEP 2100 61 53.0 4.3 0.5 * 3 SEP 0500 93 2.0 0.2
31 AUG 1400 30 909.0 693.7 2.9 * 1 SEP 2200 62 45.0 3.2 0.4 * 3 SEP 0600 94 2.0 0.1
31 AUG 1500 31 1220.0 1179.1 3.5 * 1 SEP 2300 63 37.0 2.6 0.4 * 3 SEP 0700 95 2.0 0.1
31 AUG 1600 32 1436.0 1603.4 3.9 * 2 SEP 0000 64 31.0 2.2 0.3 * 3 SEP 0800 96 2.0 0.1

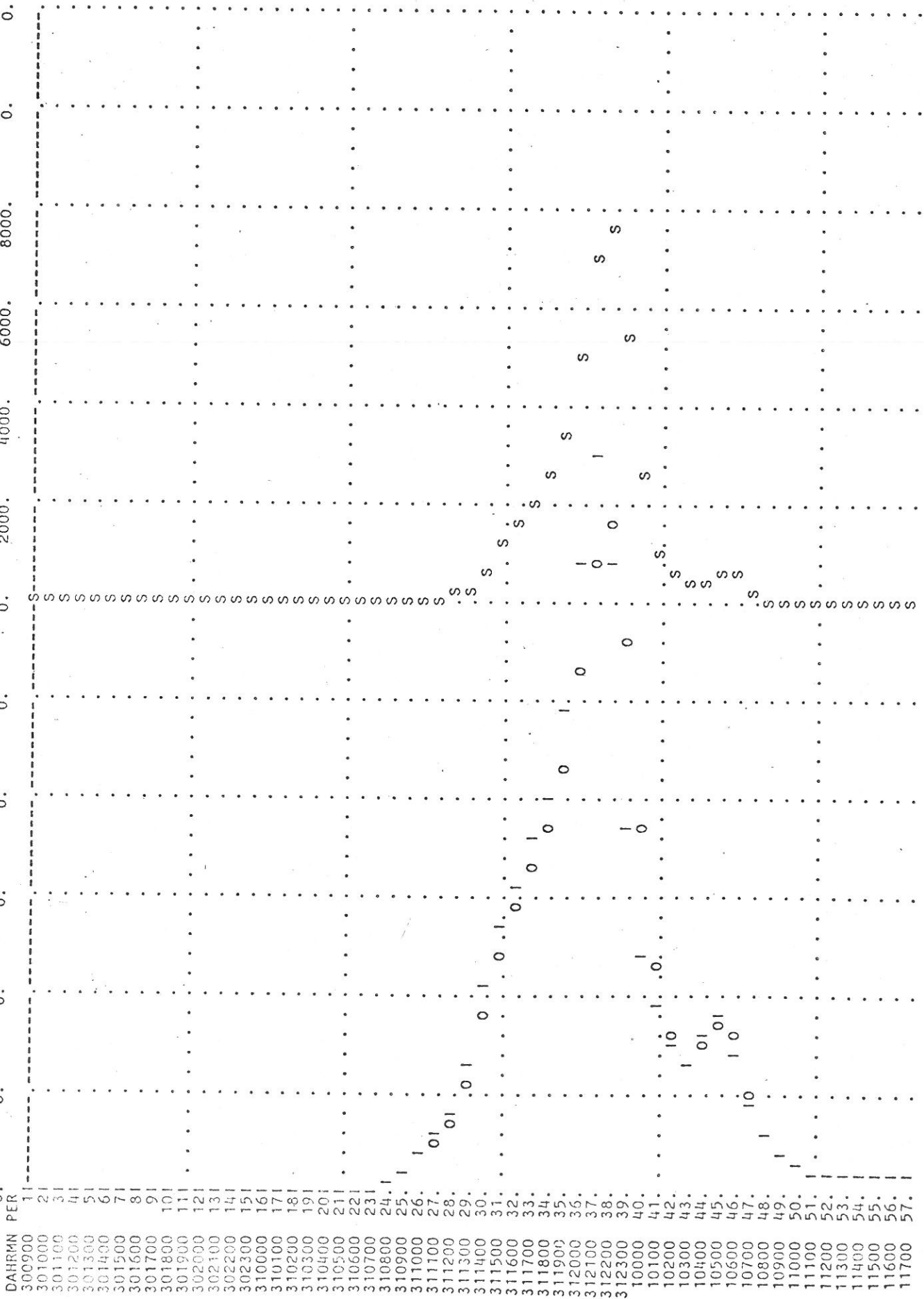
```

PEAK FLOW (CU M/S)	TIME (HR)	6-HR	24-HR	MAXIMUM AVERAGE FLOW 72-HR	95.00-HR
3385.	37.00	2673.	1270.	445.	337.
		217.895	414.205	435.419	435.444
		57742.	109764.	115386.	115393.
PEAK STORAGE (1000 CU M)	TIME (HR)	6-HR	24-HR	MAXIMUM AVERAGE STORAGE 72-HR	95.00-HR
7579.	37.00	5141.	1814.	607.	460.
PEAK STAGE (METERS)	TIME (HR)	6-HR	24-HR	MAXIMUM AVERAGE STAGE 72-HR	95.00-HR
6.57	37.00	5.63	3.34	1.29	.98

CUMULATIVE AREA = 265.00 SQ KM

STATION RESER

(I) INFLOW, (O) OUTFLOW
 1000. 1500. 2000. 3500. 4000. STORAGE
 0. 500. 1000. 2000. 3000. 4000. 6000. 8000. 0. 0. 0. 0.



11800	58.
11900	59.
12000	60.
12100	61.
12200	62.
12300	63.
20000	64.
20100	65.
20200	66.
20300	67.
20400	68.
20500	69.
20600	70.
20700	71.
20800	72.
20900	73.
21000	74.
21100	75.
21200	76.
21300	77.
21400	78.
21500	79.
21600	80.
21700	81.
21800	82.
21900	83.
22000	84.
22100	85.
22200	86.
22300	87.
30000	88.
30100	89.
30200	90.
30300	91.
30400	92.
30500	93.
30600	94.
30700	95.
30800	96.

*** **

 * SUB2C *

RUNOFF FROM SUBBASIN 2C

SUBBASIN RUNOFF DATA

127 BA SUBBASIN CHARACTERISTICS
 TAREA 164.00 SUBBASIN AREA

PRECIPITATION DATA

125 PT TOTAL STORM STATIONS 20
 126 PW WEIGHTS 1.00
 123 PR RECORDING STATIONS 7
 124 PW WEIGHTS 1.00

128 LS SCS LOSS RATE 12.70 INITIAL ABSTRACTION
 STRL 80.00 CURVE NUMBER
 CRVNR .00 PERCENT IMPERVIOUS AREA
 RTIMP

129 UK KINEMATIC WAVE
 OVERLAND-FLOW ELEMENT NO. 1
 L 4500. OVERLAND FLOW LENGTH
 S .4000 SLOPE
 N .300 ROUGHNESS COEFFICIENT
 PA 100.0 PERCENT OF SUBBASIN

130 RK MAIN CHANNEL
 L 21000. CHANNEL LENGTH
 S .0086 SLOPE
 N .022 CHANNEL ROUGHNESS COEFFICIENT
 CA 164.00 CONTRIBUTING AREA
 SHAPE TRAP CHANNEL SHAPE
 WD .70.00 BOTTOM WIDTH OR DIAMETER
 Z 10.00 SIDE SLOPE
 RUPSTQ YES ROUTE UPSTREAM HYDROGRAPH

PRECIPITATION STATION DATA

STATION	TOTAL	AVG. ANNUAL	WEIGHT
20	300.00	.00	1.00

TEMPORAL DISTRIBUTIONS

STATION	7,	WEIGHT =	1.00				
.00	.00	.00	.00	.00	.00	.00	.00
.00	.00	.00	.00	.00	.00	.00	.00
8.00	3.00	3.50	5.00	5.00	5.00	5.00	5.00
12.00	12.00	16.00	12.00	64.00	12.00	64.00	56.00
10.00							

COMPUTED KINEMATIC PARAMETERS

ELEMENT ALPHA M DT (MIN) DX (FT)
 1 3.1412 1.667 60.00 7381.89
 3 .2454 1.581 60.00 34448.82

HYDROGRAPH AT STATION SUB2C

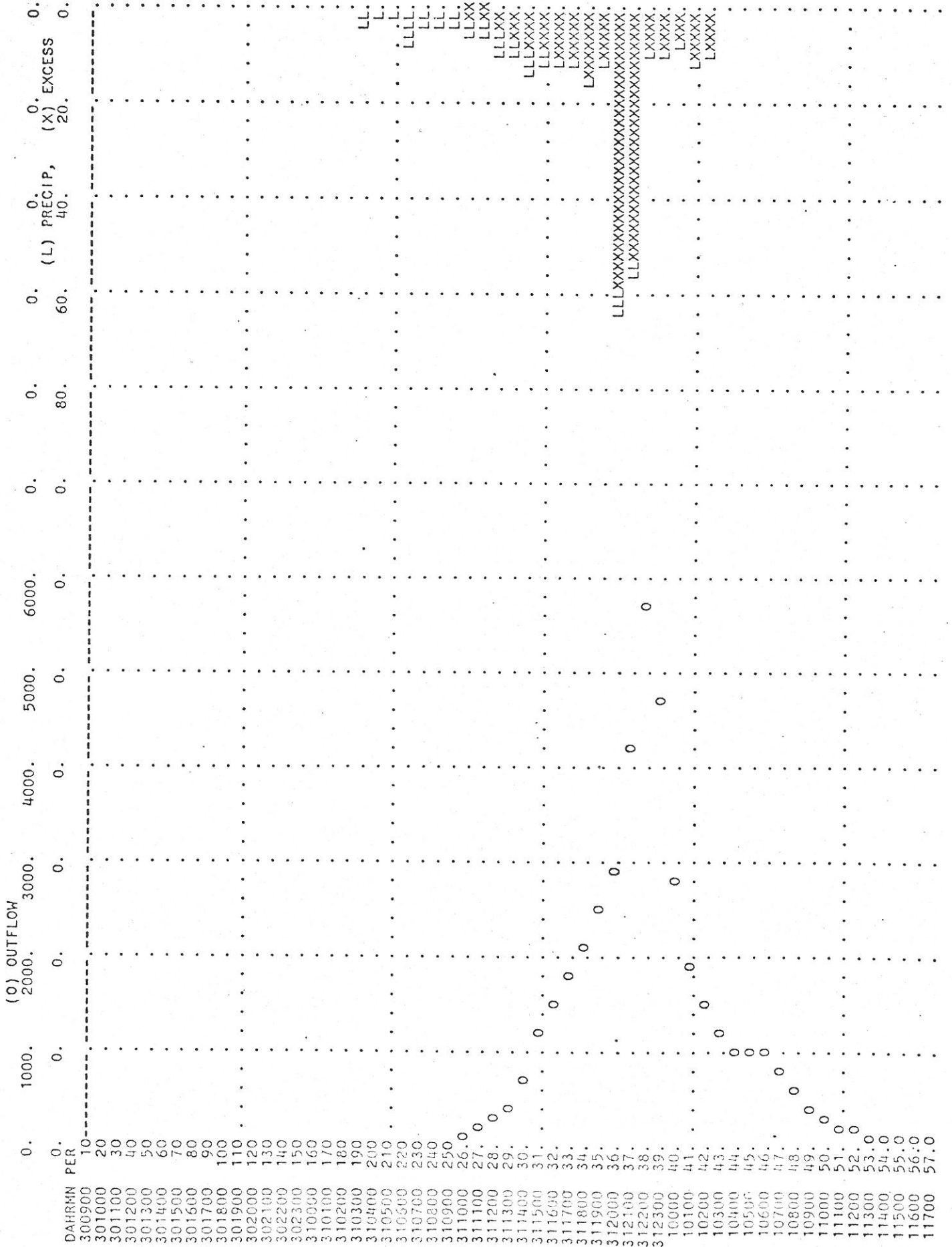
DA	MON	HRMN	ORD	RAIN	LOSS	EXCESS	COMP Q	#	DA	MON	HRMN	ORD	RAIN	LOSS	EXCESS	COMP Q	#
30	AUG	0900	1	.00	.00	.00	0.	*	1	SEP	0900	49	.00	.00	.00	375.	*
30	AUG	1000	2	.00	.00	.00	0.	*	1	SEP	1000	50	.00	.00	.00	267.	*
30	AUG	1100	3	.00	.00	.00	0.	*	1	SEP	1100	51	.00	.00	.00	203.	*
30	AUG	1200	4	.00	.00	.00	0.	*	1	SEP	1200	52	.00	.00	.00	161.	*
30	AUG	1300	5	.00	.00	.00	0.	*	1	SEP	1300	53	.00	.00	.00	133.	*
30	AUG	1400	6	.00	.00	.00	0.	*	1	SEP	1400	54	.00	.00	.00	117.	*
30	AUG	1500	7	.00	.00	.00	0.	*	1	SEP	1500	55	.00	.00	.00	111.	*
30	AUG	1600	8	.00	.00	.00	0.	*	1	SEP	1600	56	.00	.00	.00	113.	*
30	AUG	1700	9	.00	.00	.00	0.	*	1	SEP	1700	57	.00	.00	.00	113.	*
30	AUG	1800	10	.00	.00	.00	0.	*	1	SEP	1800	58	.00	.00	.00	106.	*
30	AUG	1900	11	.00	.00	.00	0.	*	1	SEP	1900	59	.00	.00	.00	95.	*
30	AUG	2000	12	.00	.00	.00	0.	*	1	SEP	2000	60	.00	.00	.00	85.	*
30	AUG	2100	13	.00	.00	.00	0.	*	1	SEP	2100	61	.00	.00	.00	76.	*
30	AUG	2200	14	.00	.00	.00	0.	*	1	SEP	2200	62	.00	.00	.00	68.	*
30	AUG	2300	15	.00	.00	.00	0.	*	1	SEP	2300	63	.00	.00	.00	60.	*
31	AUG	0000	16	.00	.00	.00	0.	*	2	SEP	0000	64	.00	.00	.00	51.	*
31	AUG	0100	17	.00	.00	.00	0.	*	2	SEP	0100	65	.00	.00	.00	44.	*
31	AUG	0200	18	.00	.00	.00	0.	*	2	SEP	0200	66	.00	.00	.00	38.	*
31	AUG	0300	19	4.00	4.00	.00	0.	*	2	SEP	0300	67	.00	.00	.00	29.	*
31	AUG	0400	20	1.00	1.00	.00	0.	*	2	SEP	0400	68	.00	.00	.00	26.	*
31	AUG	0500	21	1.00	1.00	.00	0.	*	2	SEP	0500	69	.00	.00	.00	23.	*
31	AUG	0600	22	8.00	7.97	.03	0.	*	2	SEP	0600	70	.00	.00	.00	20.	*
31	AUG	0700	23	3.00	2.75	.25	0.	*	2	SEP	0700	71	.00	.00	.00	18.	*
31	AUG	0800	24	3.50	2.92	.58	0.	*	2	SEP	0800	72	.00	.00	.00	17.	*
31	AUG	0900	25	3.50	2.65	.85	0.	*	2	SEP	0900	73	.00	.00	.00	15.	*
31	AUG	1000	26	5.00	3.38	1.62	90.	*	2	SEP	1000	74	.00	.00	.00	14.	*
31	AUG	1100	27	5.00	2.98	2.02	159.	*	2	SEP	1100	75	.00	.00	.00	13.	*
31	AUG	1200	28	10.00	5.02	4.98	282.	*	2	SEP	1200	76	.00	.00	.00	12.	*
31	AUG	1300	29	9.00	3.69	5.31	443.	*	2	SEP	1300	77	.00	.00	.00	11.	*
31	AUG	1400	30	13.00	4.32	8.68	739.	*	2	SEP	1400	78	.00	.00	.00	10.	*
31	AUG	1500	31	12.00	3.22	8.78	1153.	*	2	SEP	1500	79	.00	.00	.00	9.	*
31	AUG	1600	32	12.00	2.67	9.33	1520.	*	2	SEP	1600	80	.00	.00	.00	9.	*
31	AUG	1700	33	12.00	2.25	9.75	1833.	*	2	SEP	1700	81	.00	.00	.00	8.	*
31	AUG	1800	34	16.00	2.50	13.50	2113.	*	2	SEP	1800	82	.00	.00	.00	8.	*
31	AUG	1900	35	12.00	1.59	10.41	2461.	*	2	SEP	1900	83	.00	.00	.00	7.	*
31	AUG	2000	36	64.00	5.83	58.17	2904.	*	2	SEP	2000	84	.00	.00	.00	7.	*
31	AUG	2100	37	56.00	3.07	52.93	4250.	*	2	SEP	2100	85	.00	.00	.00	6.	*
31	AUG	2200	38	10.00	.43	9.57	5714.	*	2	SEP	2200	86	.00	.00	.00	6.	*
31	AUG	2300	39	10.00	.40	9.60	4724.	*	2	SEP	2300	87	.00	.00	.00	6.	*
1	SEP	0000	40	8.00	.31	7.69	2827.	*	3	SEP	0000	88	.00	.00	.00	5.	*
1	SEP	0100	41	12.00	.43	11.57	1925.	*	3	SEP	0100	89	.00	.00	.00	5.	*
1	SEP	0200	42	10.00	.34	9.66	1464.	*	3	SEP	0200	90	.00	.00	.00	5.	*
1	SEP	0300	43	.00	.00	.00	1217.	*	3	SEP	0300	91	.00	.00	.00	4.	*
1	SEP	0400	44	.00	.00	.00	1023.	*	3	SEP	0400	92	.00	.00	.00	4.	*
1	SEP	0500	45	.00	.00	.00	989.	*	3	SEP	0500	93	.00	.00	.00	4.	*
1	SEP	0600	46	.00	.00	.00	961.	*	3	SEP	0600	94	.00	.00	.00	4.	*
1	SEP	0700	47	.00	.00	.00	784.	*	3	SEP	0700	95	.00	.00	.00	4.	*
1	SEP	0800	48	.00	.00	.00	593.	*	3	SEP	0800	96	.00	.00	.00	3.	*

TOTAL RAINFALL = 300.00, TOTAL LOSS = 64.71, TOTAL EXCESS = 235.29

PEAK FLOW (CU M/S)	TIME (HR)	6-HR (CU M/S)	24-HR MAXIMUM AVERAGE FLOW	72-HR MAXIMUM AVERAGE FLOW	95.00-HR
5714.	37.00	3769. 189.753 81404.	1696. 341.562 146530.	592. 357.951 153561.	1449. 357.951 153561.

CUMULATIVE AREA = 429.00 SQ KM

STATION SUB2C



11800 58.0
11900 59.0
12000 60.0
12100 61.0
12200 62.0
12300 63.0
20000 64.0
20100 65.0
20200 66.0
20300 67.0
20400 68.0
20500 69.0
20600 70.0
20700 71.0
20800 72.0
20900 73.0
21000 74.0
21100 75.0
21200 76.0
21300 77.0
21400 78.0
21500 79.0
21600 80.0
21700 81.0
21800 82.0
21900 83.0
22000 84.0
22100 85.0
22200 86.0
22300 87.0
30000 88.0
30100 89.0
30200 90.0
30300 91.0
30400 92.0
30500 93.0
30600 94.0
30700 95.0
30800 96.0

*** ** ** ** ** ** ** ** ** ** ** ** ** ** ** ** ** ** ** ** ** ** ** ** ** ** ** ** ** ** ** ** ** ** ** ** ** ** ** ** ** ** ** ** ** **

*
* SUB3 *
*

RUNOFF FROM SUBBASIN 3

SUBBASIN RUNOFF DATA

137 BA SUBBASIN CHARACTERISTICS TAREA 106.00 SUBBASIN AREA

PRECIPITATION DATA

135 PT TOTAL STORM STATIONS 20 30 60
136 PW WEIGHTS .75 .05 .20

133 PR RECORDING STATIONS 3 7
134 PW WEIGHTS .05 .95

138 LS SCS LOSS RATE 12.70 INITIAL ABSTRACTION
STRFL 80.00 CURVE NUMBER
CRYHBR .00 PERCENT IMPERVIOUS AREA
RTIMP

139 UK KINEMATIC WAVE OVERLAND-FLOW ELEMENT NO. 1
OVERLAND-FLOW LENGTH 2000.
L SLOPE .4000
S ROUGHNESS COEFFICIENT .300
N PERCENT OF SUBBASIN 100.0
PA MAIN CHANNEL

140 RK MAIN CHANNEL L 21500. CHANNEL LENGTH
S .0293 SLOPE
N .030 CHANNEL ROUGHNESS COEFFICIENT
CA 106.00 CONTRIBUTING AREA
SHAPE TRAP CHANNEL SHAPE
WD .45.00 BOTTOM WIDTH OR DIAMETER
Z 10.00 SIDE SLOPE
RUPSTQ NO ROUTE UPSTREAM HYDROGRAPH

PRECIPITATION STATION DATA

STATION TOTAL AVG. ANNUAL WEIGHT
20 300.00 .75
30 251.30 .05
60 359.50 .20

TEMPORAL DISTRIBUTIONS

STATION 3, WEIGHT = .05
1.00 .00 .30 1.70 1.00 .30 .30 1.00
1.00 .00 .00 .00 .00 .00 .00 1.00
1.00 12.00 3.00 5.00 6.00 16.00 20.00 14.00 16.00
20.00 18.00 24.00 38.00 44.00 44.00 20.00 14.00 16.00
.00

STATION 7, WEIGHT = .95
 .00 .00 .00 .00 .00 .00 .00 .00
 .00 .00 .00 .00 .00 .00 .00 .00
 8.00 3.00 3.50 3.50 5.00 5.00 5.00 5.00
 12.00 12.00 16.00 12.00 64.00 64.00 64.00 64.00
 10.00

COMPUTED KINEMATIC PARAMETERS

ELEMENT ALPHA M DT (MIN) DX (FT)
 1 3.1412 1.667 60.00 3280.84
 3 .4858 1.549 60.00 35269.03

HYDROGRAPH AT STATION SUB3

DA	MON	HRMN	ORD	RAIN	LOSS	EXCESS	COMP Q	*	DA	MON	HRMN	ORD	RAIN	LOSS	EXCESS	COMP Q	*
30	AUG	0900	1	.00	.00	.00	0.	*	1	SEP	0900	49	.00	.00	.00	0.	*
30	AUG	1000	2	.00	.00	.00	0.	*	1	SEP	1000	50	.00	.00	.00	0.	*
30	AUG	1100	3	.00	.00	.00	0.	*	1	SEP	1100	51	.00	.00	.00	0.	*
30	AUG	1200	4	.00	.00	.00	0.	*	1	SEP	1200	52	.00	.00	.00	0.	*
30	AUG	1300	5	.02	.02	.00	0.	*	1	SEP	1300	53	.00	.00	.00	0.	*
30	AUG	1400	6	.09	.09	.00	0.	*	1	SEP	1400	54	.00	.00	.00	0.	*
30	AUG	1500	7	.05	.05	.00	0.	*	1	SEP	1500	55	.00	.00	.00	0.	*
30	AUG	1600	8	.02	.02	.00	0.	*	1	SEP	1600	56	.00	.00	.00	0.	*
30	AUG	1700	9	.02	.02	.00	0.	*	1	SEP	1700	57	.00	.00	.00	0.	*
30	AUG	1800	10	.02	.02	.00	0.	*	1	SEP	1800	58	.00	.00	.00	0.	*
30	AUG	1900	11	.05	.05	.00	0.	*	1	SEP	1900	59	.00	.00	.00	0.	*
30	AUG	2000	12	.05	.05	.00	0.	*	1	SEP	2000	60	.00	.00	.00	0.	*
30	AUG	2100	13	.00	.00	.00	0.	*	1	SEP	2100	61	.00	.00	.00	0.	*
30	AUG	2200	14	.00	.00	.00	0.	*	1	SEP	2200	62	.00	.00	.00	0.	*
30	AUG	2300	15	.00	.00	.00	0.	*	2	SEP	2300	63	.00	.00	.00	0.	*
31	AUG	0000	16	.00	.00	.00	0.	*	2	SEP	0000	64	.00	.00	.00	0.	*
31	AUG	0100	17	.00	.00	.00	0.	*	2	SEP	0100	65	.00	.00	.00	0.	*
31	AUG	0200	18	.00	.00	.00	0.	*	2	SEP	0200	66	.00	.00	.00	0.	*
31	AUG	0300	19	3.94	3.94	.00	0.	*	2	SEP	0300	67	.00	.00	.00	0.	*
31	AUG	0400	20	1.04	1.04	.00	0.	*	2	SEP	0400	68	.00	.00	.00	0.	*
31	AUG	0500	21	1.04	1.04	.00	0.	*	2	SEP	0500	69	.00	.00	.00	0.	*
31	AUG	0600	22	7.94	7.94	.04	0.	*	2	SEP	0600	70	.00	.00	.00	0.	*
31	AUG	0700	23	3.58	3.23	.35	0.	*	2	SEP	0700	71	.00	.00	.00	0.	*
31	AUG	0800	24	3.60	2.93	.67	0.	*	2	SEP	0800	72	.00	.00	.00	0.	*
31	AUG	0900	25	3.71	2.73	.98	0.	*	2	SEP	0900	73	.00	.00	.00	0.	*
31	AUG	1000	26	5.24	3.43	1.81	0.	*	2	SEP	1000	74	.00	.00	.00	0.	*
31	AUG	1100	27	5.24	3.01	2.23	2.	*	2	SEP	1100	75	.00	.00	.00	0.	*
31	AUG	1200	28	10.68	5.13	5.55	9.	*	2	SEP	1200	76	.00	.00	.00	0.	*
31	AUG	1300	29	9.91	3.84	6.06	38.	*	2	SEP	1300	77	.00	.00	.00	0.	*
31	AUG	1400	30	13.54	4.23	9.31	96.	*	2	SEP	1400	78	.00	.00	.00	0.	*
31	AUG	1500	31	12.65	3.18	9.48	175.	*	2	SEP	1500	79	.00	.00	.00	0.	*
31	AUG	1600	32	12.86	2.67	10.20	239.	*	2	SEP	1600	80	.00	.00	.00	0.	*
31	AUG	1700	33	12.76	2.22	10.54	279.	*	2	SEP	1700	81	.00	.00	.00	0.	*
31	AUG	1800	34	17.01	2.46	14.56	301.	*	2	SEP	1800	82	.00	.00	.00	0.	*
31	AUG	1900	35	13.80	1.67	12.13	353.	*	2	SEP	1900	83	.00	.00	.00	0.	*
31	AUG	2000	36	65.35	5.45	59.90	387.	*	2	SEP	2000	84	.00	.00	.00	0.	*
31	AUG	2100	37	56.01	2.85	53.16	1117.	*	2	SEP	2100	85	.00	.00	.00	0.	*
31	AUG	2200	38	9.85	.40	9.46	1572.	*	2	SEP	2200	86	.00	.00	.00	0.	*
31	AUG	2300	39	9.85	.37	9.48	890.	*	2	SEP	2300	87	.00	.00	.00	0.	*
1	SEP	0000	40	7.88	.28	7.60	377.	*	3	SEP	0000	88	.00	.00	.00	0.	*
1	SEP	0100	41	11.82	.40	11.42	262.	*	3	SEP	0100	89	.00	.00	.00	0.	*
1	SEP	0200	42	9.85	.31	9.54	268.	*	3	SEP	0200	90	.00	.00	.00	0.	*
1	SEP	0300	43	.00	.00	.00	302.	*	3	SEP	0300	91	.00	.00	.00	0.	*

1 SEP 0400	44	.00	.00	.00	182.	*	3 SEP 0400	92	.00	.00	0.
1 SEP 0500	45	.00	.00	.00	83.	*	3 SEP 0500	93	.00	.00	0.
1 SEP 0600	46	.00	.00	.00	48.	*	3 SEP 0600	94	.00	.00	0.
1 SEP 0700	47	.00	.00	.00	32.	*	3 SEP 0700	95	.00	.00	0.
1 SEP 0800	48	.00	.00	.00	22.	*	3 SEP 0800	96	.00	.00	0.

TOTAL RAINFALL = 309.46, TOTAL LOSS = 65.01, TOTAL EXCESS = 244.46

PEAK FLOW (CU M/S)	TIME (HR)	6-HR	24-HR	72-HR	95.00-HR
1572.	37.00	776.	295.	99.	75.
		158.204	240.297	242.919	242.919
		16770.	25471.	25749.	25749.

CUMULATIVE AREA = 106.00 SQ KM

11800	580
11900	590
12000	600
12100	610
12200	620
12300	630
20000	640
20100	650
20200	660
20300	670
20400	680
20500	690
20600	700
20700	710
20800	720
20900	730
21000	740
21100	750
21200	760
21300	770
21400	780
21500	790
21600	800
21700	810
21800	820
21900	830
22000	840
22100	850
22200	860
22300	870
30000	880
30100	890
30200	900
30300	910
30400	920
30500	930
30600	940
30700	950
30800	960

*** ** ** ** **

* SUB4 *

141 KK

COMBINE RUNOFF FROM SUB2C AND SUB3

143 HC

HYDROGRAPH COMBINATION 2 NUMBER OF HYDROGRAPHS TO COMBINE
ICOMP

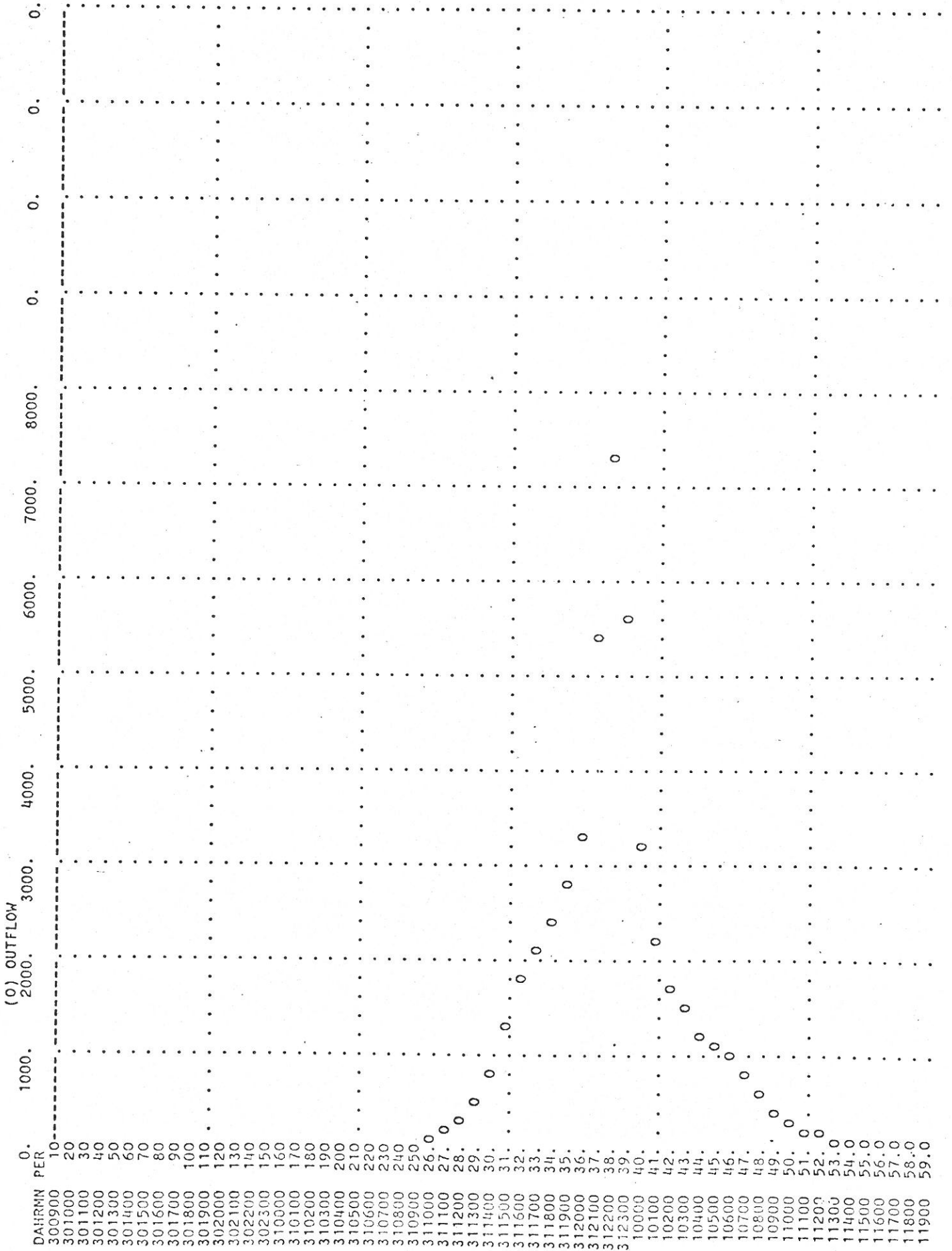
HYDROGRAPH AT STATION SUB4
SUM OF 2 HYDROGRAPHS

DA	MON	HRMN	ORD	FLOW	DA	MON	HRMN	ORD	FLOW	DA	MON	HRMN	ORD	FLOW
30	AUG	0900	1	0.	31	AUG	0900	25	3.	1	SEP	0900	49	391.
30	AUG	1000	2	0.	31	AUG	1000	26	90.	1	SEP	1000	50	280.
30	AUG	1100	3	0.	31	AUG	1100	27	161.	1	SEP	1100	51	213.
30	AUG	1200	4	0.	31	AUG	1200	28	291.	1	SEP	1200	52	169.
30	AUG	1300	5	0.	31	AUG	1300	29	481.	1	SEP	1300	53	140.
30	AUG	1400	6	0.	31	AUG	1400	30	835.	1	SEP	1400	54	123.
30	AUG	1500	7	0.	31	AUG	1500	31	1328.	1	SEP	1500	55	117.
30	AUG	1600	8	0.	31	AUG	1600	32	1759.	1	SEP	1600	56	118.
30	AUG	1700	9	0.	31	AUG	1700	33	2113.	1	SEP	1700	57	117.
30	AUG	1800	10	0.	31	AUG	1800	34	2414.	1	SEP	1800	58	109.
30	AUG	1900	11	0.	31	AUG	1900	35	2815.	1	SEP	1900	59	98.
30	AUG	2000	12	0.	31	AUG	2000	36	3292.	1	SEP	2000	60	88.
30	AUG	2100	13	0.	31	AUG	2100	37	5366.	1	SEP	2100	61	79.
30	AUG	2200	14	0.	31	AUG	2200	38	7286.	1	SEP	2200	62	70.
30	AUG	2300	15	0.	31	AUG	2300	39	5614.	1	SEP	2300	63	62.
31	AUG	0000	16	0.	1	SEP	0000	40	3204.	2	SEP	0000	64	53.
31	AUG	0100	17	0.	1	SEP	0100	41	2187.	2	SEP	0100	65	46.
31	AUG	0200	18	0.	1	SEP	0200	42	1732.	2	SEP	0200	66	39.
31	AUG	0300	19	0.	1	SEP	0300	43	1518.	2	SEP	0300	67	31.
31	AUG	0400	20	0.	1	SEP	0400	44	1205.	2	SEP	0400	68	27.
31	AUG	0500	21	0.	1	SEP	0500	45	1077.	2	SEP	0500	69	24.
31	AUG	0600	22	0.	1	SEP	0600	46	1009.	2	SEP	0600	70	21.
31	AUG	0700	23	0.	1	SEP	0700	47	816.	2	SEP	0700	71	19.
31	AUG	0800	24	0.	1	SEP	0800	48	575.	2	SEP	0800	72	17.

PEAK FLOW (CU M/S)	TIME (HR)	6-HR (CU M/S)	24-HR (CU M/S)	MAXIMUM AVERAGE FLOW 72-HR (MM)	95.00-HR (1000 CU M)
7286.	37.00	4544.	1991.	692.	524.
		183.451	321.474	335.160	335.160
		98146.	171989.	179310.	179310.

CUMULATIVE AREA = 535.00 SQ KM

STATION SUB4



12000 60.0
12100 61.0
12200 62.0
12300 63.0
20000 64.0
20100 65.0
20200 66.0
20300 67.0
20400 68.0
20500 69.0
20600 70.0
20700 71.0
20800 72.0
20900 73.0
21000 74.0
21100 75.0
21200 76.0
21300 77.0
21400 78.0
21500 79.0
21600 80.0
21700 81.0
21800 82.0
21900 83.0
22000 84.0
22100 85.0
22200 86.0
22300 87.0
30000 88.0
30100 89.0
30200 90.0
30300 91.0
30400 92.0
30500 93.0
30600 94.0
30700 95.0
30800 96.0

*** ** ** ** ** ** ** ** ** ** ** ** ** ** ** ** ** ** ** ** ** ** ** ** ** ** ** ** ** ** ** ** ** ** ** ** ** ** ** ** ** ** ** ** ** ** ** ** ** ** ** ** ** ** ** ** ** **

* SUB4 *

RUNOFF FROM SUBBASIN 4

SUBBASIN RUNOFF DATA

150 BA SUBBASIN CHARACTERISTICS
TAREA 103.00 SUBBASIN AREA

PRECIPITATION DATA

148 PT TOTAL STORM STATIONS 10 20
149 PW WEIGHTS .10 .90
146 PR RECORDING STATIONS 7 9
147 PW WEIGHTS .90 .10

151 LS SCS LOSS RATE 12.70 INITIAL ABSTRACTION
STRTL 80.00 CURVE NUMBER
CRVBR .00 PERCENT IMPERVIOUS AREA
RTIMP

152 UK KINEMATIC WAVE OVERLAND-FLOW ELEMENT NO. 1
L 5000. OVERLAND FLOW LENGTH
S .5000 SLOPE
N .300 ROUGHNESS COEFFICIENT
PA 100.0 PERCENT OF SUBBASIN

153 RK MAIN CHANNEL 11500. CHANNEL LENGTH
L .0148 SLOPE
S .030 CHANNEL ROUGHNESS COEFFICIENT
N 103.00 CONTRIBUTING AREA
CA TRAP CHANNEL SHAPE
WD 80.00 BOTTOM WIDTH OR DIAMETER
Z 10.00 SIDE SLOPE
RUPSTQ YES ROUTE UPSTREAM HYDROGRAPH

PRECIPITATION STATION DATA

STATION	TOTAL	AVG. ANNUAL	WEIGHT
10	300.10	.00	.10
20	300.00	.00	.90

TEMPORAL DISTRIBUTIONS

STATION	7, WEIGHT = .90	.00	.00	.00	.00	.00	.00	.00	.00
.00	.00	.00	.00	.00	.00	.00	.00	.00	.00
8.00	.00	.00	.00	.00	.00	.00	.00	.00	.00
12.00	3.50	3.50	5.00	5.00	10.00	10.00	10.00	13.00	12.00
10.00	12.00	16.00	64.00	56.00	10.00	10.00	8.00	8.00	12.00

STATION 9, WEIGHT = .10

.00	.00	.00	.00	.00	.00	.00	.00	.00	.00
.00	.00	.00	.00	.00	.00	.00	.00	.00	.00
.00	.00	.00	.00	.00	.00	.00	.00	.00	.00
14.00	16.00	20.00	20.00	20.00	5.50	6.00	6.00	6.00	11.10
1.70					16.00	84.00	14.00	14.00	5.50

COMPUTED KINEMATIC PARAMETERS

ELEMENT	ALPHA	M	DT (MIN)	DX (FT)
1	3.5120	1.667	60.00	8202.10
3	.2103	1.589	60.00	18864.83

HYDROGRAPH AT STATION SUB4

DA	MON	HRMN	ORD	RAIN	LOSS	EXCESS	COMP Q	#	DA	MON	HRMN	ORD	RAIN	LOSS	EXCESS	COMP Q	#
30	AUG	0900	1	.00	.00	.00	0.	*	1	SEP	0900	49	.08	.00	.08	522.	*
30	AUG	1000	2	.00	.00	.00	0.	*	1	SEP	1000	50	.10	.00	.10	378.	*
30	AUG	1100	3	.00	.00	.00	0.	*	1	SEP	1100	51	.12	.00	.12	285.	*
30	AUG	1200	4	.00	.00	.00	0.	*	1	SEP	1200	52	.12	.00	.12	224.	*
30	AUG	1300	5	.00	.00	.00	0.	*	1	SEP	1300	53	.12	.00	.12	183.	*
30	AUG	1400	6	.00	.00	.00	0.	*	1	SEP	1400	54	.12	.00	.12	156.	*
30	AUG	1500	7	.00	.00	.00	0.	*	1	SEP	1500	55	.12	.00	.12	139.	*
30	AUG	1600	8	.00	.00	.00	0.	*	1	SEP	1600	56	.10	.00	.10	132.	*
30	AUG	1700	9	.00	.00	.00	0.	*	1	SEP	1700	57	.10	.00	.10	130.	*
30	AUG	1800	10	.00	.00	.00	0.	*	1	SEP	1800	58	.10	.00	.10	125.	*
30	AUG	1900	11	.00	.00	.00	0.	*	1	SEP	1900	59	.15	.00	.15	117.	*
30	AUG	2000	12	.00	.00	.00	0.	*	1	SEP	2000	60	.15	.00	.15	106.	*
30	AUG	2100	13	.00	.00	.00	0.	*	1	SEP	2100	61	.00	.00	.00	95.	*
30	AUG	2200	14	.00	.00	.00	0.	*	1	SEP	2200	62	.00	.00	.00	86.	*
30	AUG	2300	15	.00	.00	.00	0.	*	1	SEP	2300	63	.00	.00	.00	77.	*
31	AUG	0000	16	.00	.00	.00	0.	*	2	SEP	0000	64	.00	.00	.00	68.	*
31	AUG	0100	17	.00	.00	.00	0.	*	2	SEP	0100	65	.00	.00	.00	60.	*
31	AUG	0200	18	.00	.00	.00	0.	*	2	SEP	0200	66	.00	.00	.00	52.	*
31	AUG	0300	19	3.60	3.60	.00	0.	*	2	SEP	0300	67	.00	.00	.00	45.	*
31	AUG	0400	20	.90	.90	.00	0.	*	2	SEP	0400	68	.00	.00	.00	38.	*
31	AUG	0500	21	.90	.90	.00	0.	*	2	SEP	0500	69	.00	.00	.00	32.	*
31	AUG	0600	22	7.20	7.20	.00	0.	*	2	SEP	0600	70	.00	.00	.00	29.	*
31	AUG	0700	23	2.70	2.70	.10	0.	*	2	SEP	0700	71	.00	.00	.00	26.	*
31	AUG	0800	24	3.15	2.77	.38	0.	*	2	SEP	0800	72	.00	.00	.00	23.	*
31	AUG	0900	25	3.20	2.57	.63	0.	*	2	SEP	0900	73	.00	.00	.00	21.	*
31	AUG	1000	26	5.05	3.63	1.42	0.	*	2	SEP	1000	74	.00	.00	.00	20.	*
31	AUG	1100	27	5.10	3.21	1.89	151.	*	2	SEP	1100	75	.00	.00	.00	18.	*
31	AUG	1200	28	9.40	4.99	4.41	216.	*	2	SEP	1200	76	.00	.00	.00	17.	*
31	AUG	1300	29	8.70	3.79	4.91	399.	*	2	SEP	1300	77	.00	.00	.00	15.	*
31	AUG	1400	30	13.39	4.70	8.69	723.	*	2	SEP	1400	78	.00	.00	.00	14.	*
31	AUG	1500	31	11.91	3.34	8.57	1232.	*	2	SEP	1500	79	.00	.00	.00	13.	*
31	AUG	1600	32	12.20	2.83	9.37	1751.	*	2	SEP	1600	80	.00	.00	.00	12.	*
31	AUG	1700	33	12.40	2.40	10.00	2189.	*	2	SEP	1700	81	.00	.00	.00	11.	*
31	AUG	1800	34	16.40	2.63	13.77	2555.	*	2	SEP	1800	82	.00	.00	.00	10.	*
31	AUG	1900	35	12.80	1.72	11.08	3001.	*	2	SEP	1900	83	.00	.00	.00	9.	*
31	AUG	2000	36	59.20	5.56	53.64	3488.	*	2	SEP	2000	84	.00	.00	.00	8.	*
31	AUG	2100	37	58.80	3.33	55.47	5722.	*	2	SEP	2100	85	.00	.00	.00	8.	*
31	AUG	2200	38	10.40	.43	9.97	8462.	*	2	SEP	2200	86	.00	.00	.00	8.	*
31	AUG	2300	39	10.40	.43	9.97	6969.	*	2	SEP	2300	87	.00	.00	.00	7.	*
1	SEP	0000	40	17.75	.30	10.93	4014.	*	3	SEP	0000	88	.00	.00	.00	7.	*
1	SEP	0100	41	11.35	.42	10.93	2711.	*	3	SEP	0100	89	.00	.00	.00	7.	*
1	SEP	0200	42	9.17	.32	8.85	2135.	*	3	SEP	0200	90	.00	.00	.00	6.	*
1	SEP	0300	43	.17	.01	.16	1851.	*	3	SEP	0300	91	.00	.00	.00	6.	*
1	SEP	0400	44	.17	.01	.16	1484.	*	3	SEP	0400	92	.00	.00	.00	6.	*

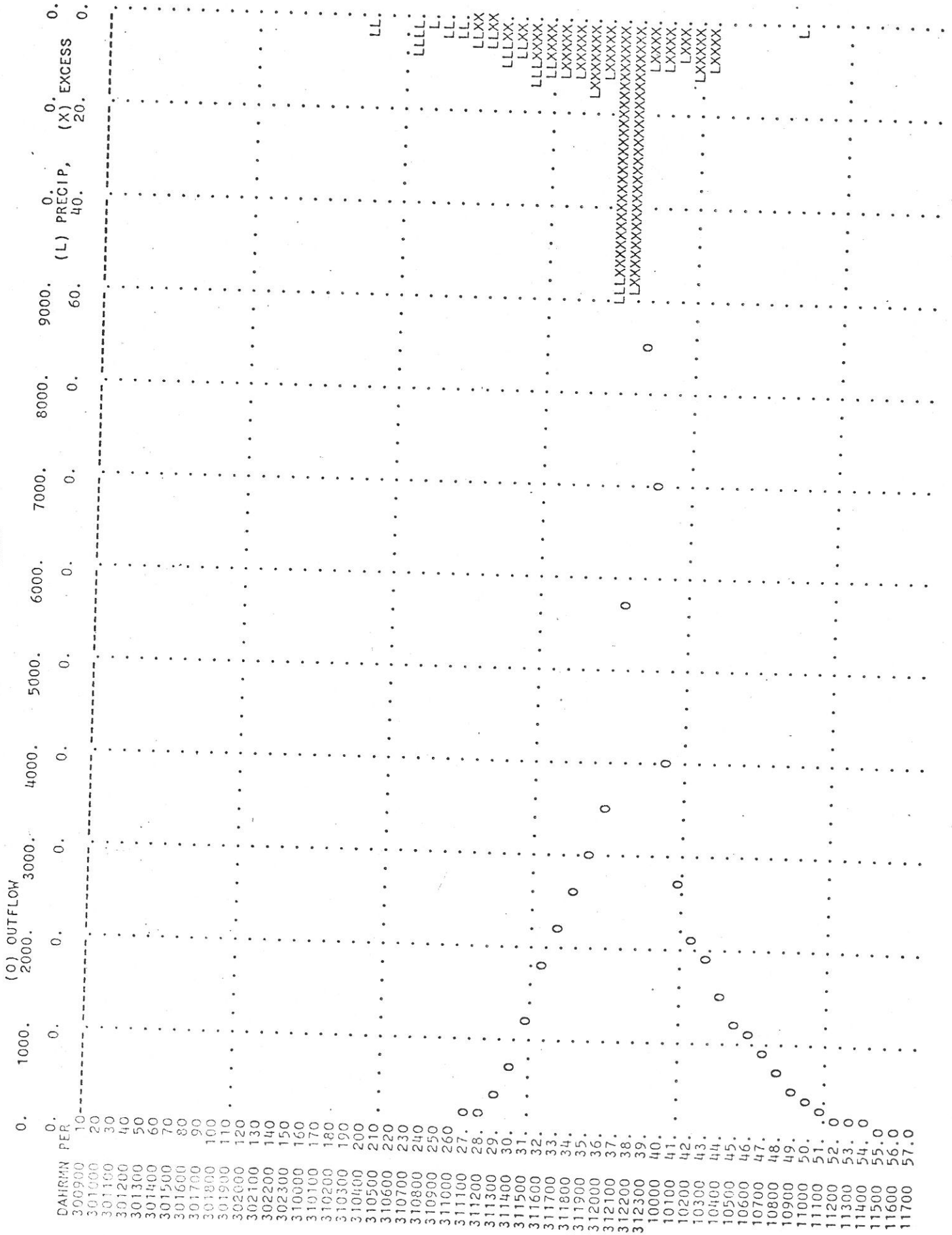
1 SEP 0500	45	.80	.03	.77	1246.	*	3 SEP 0500	93	.00	.00	5.
1 SEP 0600	46	.29	.01	.28	1123.	*	3 SEP 0600	94	.00	.00	5.
1 SEP 0700	47	.11	.00	.11	949.	*	3 SEP 0700	95	.00	.00	5.
1 SEP 0800	48	1.02	.03	.99	721.	*	3 SEP 0800	96	.00	.00	5.

TOTAL RAINFALL = 300.01, TOTAL LOSS = 64.71, TOTAL EXCESS = 235.30

PEAK FLOW (CU M/S)	TIME (HR)	6-HR	24-HR	72-HR	95.00-HR
8462.	37.00	5252.	2255.	784.	594.
		177.801	305.423	318.590	318.590
		(1000 CU M)	194860.	203260.	203260.

CUMULATIVE AREA = 638.00 SQ KM

STATION SUB4



11800 58.0
11900 59.0
12000 60.0
12100 61.0
12200 62.0
12300 63.0
20000 64.0
20100 65.0
20200 66.0
20300 67.0
20400 68.0
20500 69.0
20600 70.0
20700 71.0
20800 72.0
20900 73.0
21000 74.0
21100 75.0
21200 76.0
21300 77.0
21400 78.0
21500 79.0
21600 80.0
21700 81.0
21800 82.0
21900 83.0
22000 84.0
22100 85.0
22200 86.0
22300 87.0
30000 88.0
30100 89.0
30200 90.0
30300 91.0
30400 92.0
30500 93.0
30600 94.0
30700 95.0
30800 96.0

STATION	.00	.00	.00	.00	.00	.00	.00	.00	.00	.00	.00	.00	.00	.00	.00	.00	.00	.00
.00	.00	.00	.00	.00	.00	.00	.00	.00	.00	.00	.00	.00	.00	.00	.00	.00	.00	.00
14.00	1.70	1.70	1.70	20.00	20.00	20.00	84.00	4.00	4.00	6.00	6.00	14.00	14.00	14.00	14.00	16.90	16.90	16.90
1.20	1.20	1.20	1.20	1.20	1.20	1.20	1.00	1.00	1.00	1.00	1.00	1.50	1.50	1.50	1.50	1.50	1.50	1.50

9, WEIGHT = .10

COMPUTED KINEMATIC PARAMETERS

ELEMENT	ALPHA	M	DT (MIN)	DX (FT)
1	3.1412	1.667	60.00	4921.26
3	.3826	1.549	60.00	50853.02

HYDROGRAPH AT STATION SUB5

DA	MON	HRMN	ORD	RAIN	LOSS	EXCESS	COMP Q	#	DA	MON	HRMN	ORD	RAIN	LOSS	EXCESS	COMP Q
30	AUG	0900	1	.00	.00	.00	0.	*	1	SEP	0900	49	.08	.00	.08	39.
30	AUG	1000	2	.00	.00	.00	0.	*	1	SEP	1000	50	.10	.00	.10	33.
30	AUG	1100	3	.00	.00	.00	0.	*	1	SEP	1100	51	.12	.00	.12	27.
30	AUG	1200	4	.00	.00	.00	0.	*	1	SEP	1200	52	.12	.00	.12	23.
30	AUG	1300	5	.00	.00	.00	0.	*	1	SEP	1300	53	.12	.00	.12	20.
30	AUG	1400	6	.00	.00	.00	0.	*	1	SEP	1400	54	.12	.00	.12	17.
30	AUG	1500	7	.00	.00	.00	0.	*	1	SEP	1500	55	.12	.00	.12	15.
30	AUG	1600	8	.00	.00	.00	0.	*	1	SEP	1600	56	.10	.00	.10	13.
30	AUG	1700	9	.00	.00	.00	0.	*	1	SEP	1700	57	.10	.00	.10	12.
30	AUG	1800	10	.00	.00	.00	0.	*	1	SEP	1800	58	.10	.00	.10	10.
30	AUG	1900	11	.00	.00	.00	0.	*	1	SEP	1900	59	.15	.00	.15	9.
30	AUG	2000	12	.00	.00	.00	0.	*	1	SEP	2000	60	.15	.00	.15	9.
30	AUG	2100	13	.00	.00	.00	0.	*	1	SEP	2100	61	.00	.00	.00	8.
30	AUG	2200	14	.00	.00	.00	0.	*	1	SEP	2200	62	.00	.00	.00	7.
30	AUG	2300	15	.00	.00	.00	0.	*	1	SEP	2300	63	.00	.00	.00	6.
31	AUG	0000	16	.00	.00	.00	0.	*	2	SEP	0000	64	.00	.00	.00	6.
31	AUG	0100	17	.00	.00	.00	0.	*	2	SEP	0100	65	.00	.00	.00	5.
31	AUG	0200	18	.00	.00	.00	0.	*	2	SEP	0200	66	.00	.00	.00	5.
31	AUG	0300	19	3.67	3.67	.00	0.	*	2	SEP	0300	67	.00	.00	.00	4.
31	AUG	0400	20	.92	.92	.00	0.	*	2	SEP	0400	68	.00	.00	.00	4.
31	AUG	0500	21	.92	.92	.00	0.	*	2	SEP	0500	69	.00	.00	.00	4.
31	AUG	0600	22	7.34	7.34	.13	0.	*	2	SEP	0600	70	.00	.00	.00	4.
31	AUG	0700	23	2.75	2.63	.41	0.	*	2	SEP	0700	71	.00	.00	.00	3.
31	AUG	0800	24	3.21	2.80	.67	0.	*	2	SEP	0800	72	.00	.00	.00	3.
31	AUG	0900	25	3.26	2.59	1.50	0.	*	2	SEP	0900	73	.00	.00	.00	3.
31	AUG	1000	26	5.15	3.65	1.50	0.	*	2	SEP	1000	74	.00	.00	.00	3.
31	AUG	1100	27	5.20	3.23	1.97	0.	*	2	SEP	1100	75	.00	.00	.00	2.
31	AUG	1200	28	9.59	5.00	4.58	1.	*	2	SEP	1200	76	.00	.00	.00	2.
31	AUG	1300	29	8.87	3.79	5.08	6.	*	2	SEP	1300	77	.00	.00	.00	2.
31	AUG	1400	30	13.66	4.69	8.96	21.	*	2	SEP	1400	78	.00	.00	.00	2.
31	AUG	1500	31	12.15	3.33	8.82	70.	*	2	SEP	1500	79	.00	.00	.00	2.
31	AUG	1600	32	12.44	2.81	9.63	147.	*	2	SEP	1600	80	.00	.00	.00	2.
31	AUG	1700	33	12.65	2.39	10.26	225.	*	2	SEP	1700	81	.00	.00	.00	2.
31	AUG	1800	34	16.73	2.61	14.11	267.	*	2	SEP	1800	82	.00	.00	.00	2.
31	AUG	1900	35	13.05	1.70	11.35	323.	*	2	SEP	1900	83	.00	.00	.00	1.
31	AUG	2000	36	60.37	5.50	54.87	357.	*	2	SEP	2000	84	.00	.00	.00	1.
31	AUG	2100	37	59.97	3.29	56.68	835.	*	2	SEP	2100	85	.00	.00	.00	1.
31	AUG	2200	38	10.61	.45	10.16	1541.	*	2	SEP	2200	86	.00	.00	.00	1.
31	AUG	2300	39	10.61	.42	10.18	1141.	*	2	SEP	2300	87	.00	.00	.00	1.
1	SEP	0000	40	7.90	.30	7.61	517.	*	3	SEP	0000	88	.00	.00	.00	1.
1	SEP	0100	41	11.58	.41	11.17	287.	*	3	SEP	0100	89	.00	.00	.00	1.

1 SEP 0200	42	9.35	.31	9.04	282.	*	3 SEP 0200	90	.00	.00	1.
1 SEP 0300	43	.17	.01	.17	288.	*	3 SEP 0300	91	.00	.00	1.
1 SEP 0400	44	.17	.01	.17	220.	*	3 SEP 0400	92	.00	.00	1.
1 SEP 0500	45	.82	.03	.79	134.	*	3 SEP 0500	93	.00	.00	1.
1 SEP 0600	46	.30	.01	.29	77.	*	3 SEP 0600	94	.00	.00	1.
1 SEP 0700	47	.11	.00	.11	60.	*	3 SEP 0700	95	.00	.00	1.
1 SEP 0800	48	1.04	.03	1.01	46.	*	3 SEP 0800	96	.00	.00	1.

TOTAL RAINFALL = 305.96, TOTAL LOSS = 64.90, TOTAL EXCESS = 241.06

PEAK FLOW (CU M/S)	1541.	37.00	(CU M/S)	291.	100.	75.
			(MM)	230.461	236.756	25806.
			(1000 CU.M)	155.128	25806.	25806.
				16909.		

CUMULATIVE AREA = 109.00 SQ KM

MAXIMUM AVERAGE FLOW
24-HR 291. 100. 75.
72-HR 230.461 236.756
95.00-HR 25806. 25806.

11800 58.0
11900 590
12000 600
12100 610
12200 620
12300 630
20000 640
20100 650
20200 660
20300 670
20400 680
20500 690
20600 700
20700 710
20800 720
20900 730
21000 740
21100 750
21200 760
21300 770
21400 780
21500 790
21600 800
21700 810
21800 820
21900 830
22000 840
22100 850
22200 860
22300 870
30000 880
30100 890
30200 900
30300 910
30400 920
30500 930
30600 940
30700 950
30800 960

 * SUB6 *
 *

 164 KK

 * SUB6 *
 *

 166 HC

 * SUB6 *
 *

COMBINE RUNOFF FROM SUB4 AND SUB5
 HYDROGRAPH COMBINATION 2 NUMBER OF HYDROGRAPHS TO COMBINE
 ICOMP

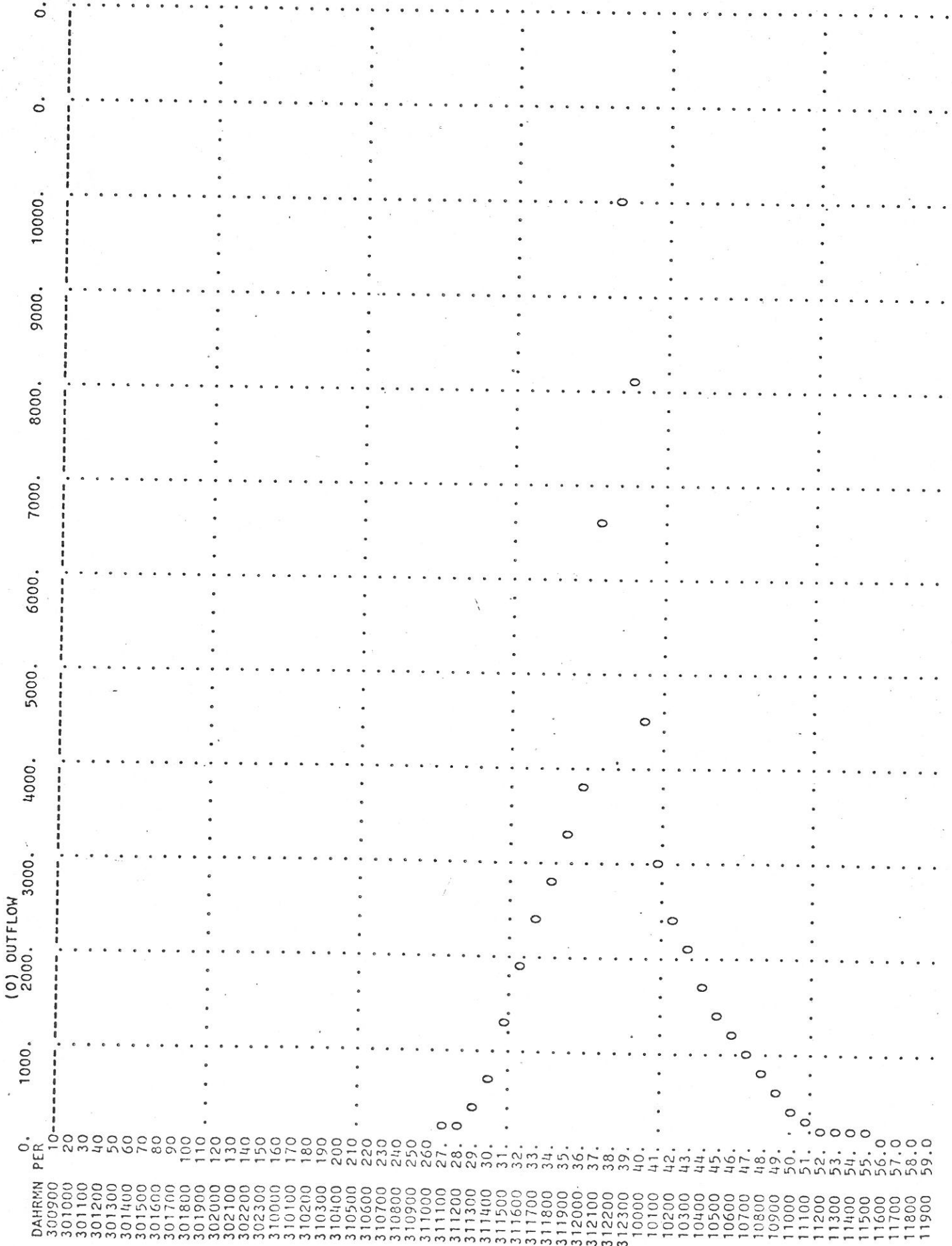
HYDROGRAPH AT STATION SUB6
 SUM OF 2 HYDROGRAPHS

DA	MON	HRMN	ORD	FLOW	DA	MON	HRMN	ORD	FLOW	DA	MON	HRMN	ORD	FLOW
30	AUG	0900	1	0.	31	AUG	0900	25	0.	1	SEP	0900	49	560.
30	AUG	1000	2	0.	31	AUG	1000	26	0.	1	SEP	1000	50	410.
30	AUG	1100	3	0.	31	AUG	1100	27	151.	1	SEP	1100	51	312.
30	AUG	1200	4	0.	31	AUG	1200	28	217.	1	SEP	1200	52	248.
30	AUG	1300	5	0.	31	AUG	1300	29	404.	1	SEP	1300	53	203.
30	AUG	1400	6	0.	31	AUG	1400	30	744.	1	SEP	1400	54	173.
30	AUG	1500	7	0.	31	AUG	1500	31	1302.	1	SEP	1500	55	154.
30	AUG	1600	8	0.	31	AUG	1600	32	1898.	1	SEP	1600	56	145.
30	AUG	1700	9	0.	31	AUG	1700	33	2414.	1	SEP	1700	57	141.
30	AUG	1800	10	0.	31	AUG	1800	34	2921.	1	SEP	1800	58	136.
30	AUG	1900	11	0.	31	AUG	1900	35	3324.	1	SEP	1900	59	126.
30	AUG	2000	12	0.	31	AUG	2000	36	3845.	1	SEP	2000	60	114.
30	AUG	2100	13	0.	31	AUG	2100	37	6558.	1	SEP	2100	61	103.
30	AUG	2200	14	0.	31	AUG	2200	38	10003.	1	SEP	2200	62	93.
30	AUG	2300	15	0.	31	AUG	2300	39	8110.	1	SEP	2300	63	84.
31	AUG	0000	16	0.	1	SEP	0000	40	4531.	2	SEP	0000	64	74.
31	AUG	0100	17	0.	1	SEP	0100	41	2998.	2	SEP	0100	65	65.
31	AUG	0200	18	0.	1	SEP	0200	42	2417.	2	SEP	0200	66	57.
31	AUG	0300	19	0.	1	SEP	0300	43	2138.	2	SEP	0300	67	49.
31	AUG	0400	20	0.	1	SEP	0400	44	1705.	2	SEP	0400	68	42.
31	AUG	0500	21	0.	1	SEP	0500	45	1380.	2	SEP	0500	69	36.
31	AUG	0600	22	0.	1	SEP	0600	46	1199.	2	SEP	0600	70	32.
31	AUG	0700	23	0.	1	SEP	0700	47	1009.	2	SEP	0700	71	29.
31	AUG	0800	24	0.	1	SEP	0800	48	768.	2	SEP	0800	72	26.

PEAK FLOW (CU M/S)	TIME (HR)	6-HR (CU M/S)	24-HR (CU M/S)	72-HR (CU M/S)	95.00-HR (CU M/S)
10003.	37.00	6035.	2545.	884.	670.
		174,493	294,374	306,649	306,649
		130346.	219897.	229067.	229067.

CUMULATIVE AREA = 747.00 SQ KM

STATION SUB6



COMPUTED KINEMATIC PARAMETERS
 ELEMENT 1 3
 ALPHA 3.1412 .1886
 M
 DT (MIN) 60.00 60.00
 DX (FT) 8202.10 22145.67

HYDROGRAPH AT STATION SUB6

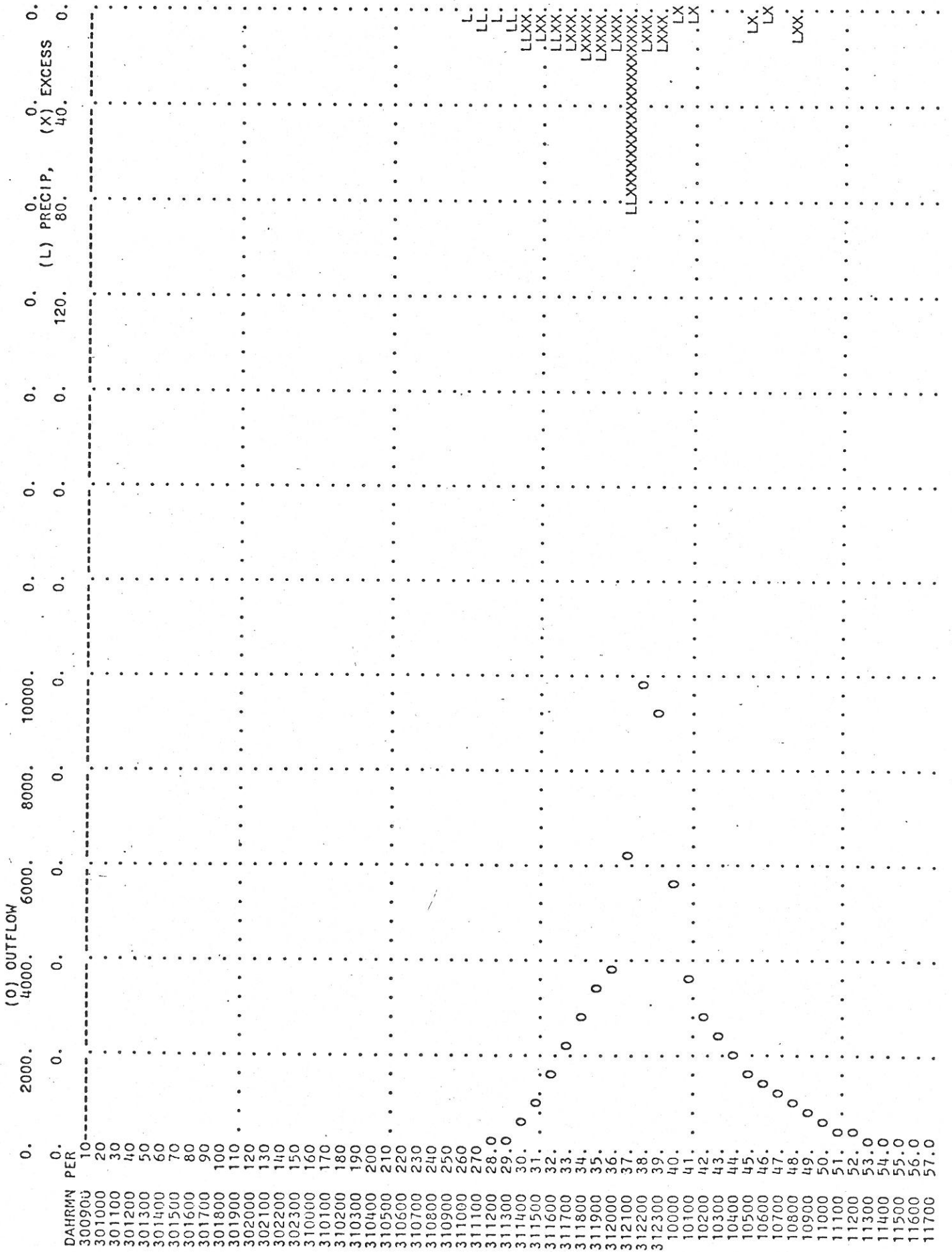
DA	MON	HRMN	ORD	RAIN	LOSS	EXCESS	COMP Q	DA	MON	HRMN	ORD	RAIN	LOSS	EXCESS	COMP Q
30	AUG	0900	1	.00	.00	.00	0.	1	SEP	0900	49	.80	.03	.77	735.
30	AUG	1000	2	.00	.00	.00	0.	1	SEP	1000	50	1.00	.04	.96	559.
30	AUG	1100	3	.00	.00	.00	0.	1	SEP	1100	51	1.20	.04	1.16	427.
30	AUG	1200	4	.00	.00	.00	0.	1	SEP	1200	52	1.20	.04	1.16	337.
30	AUG	1300	5	.00	.00	.00	0.	1	SEP	1300	53	1.20	.04	1.16	274.
30	AUG	1400	6	.00	.00	.00	0.	1	SEP	1400	54	1.20	.04	1.16	229.
30	AUG	1500	7	.00	.00	.00	0.	1	SEP	1500	55	1.20	.04	1.16	199.
30	AUG	1600	8	.00	.00	.00	0.	1	SEP	1600	56	1.00	.03	.97	179.
30	AUG	1700	9	.00	.00	.00	0.	1	SEP	1700	57	1.00	.03	.97	168.
30	AUG	1800	10	.00	.00	.00	0.	1	SEP	1800	58	1.00	.03	.97	161.
30	AUG	1900	11	.00	.00	.00	0.	1	SEP	1900	59	1.50	.05	1.45	154.
30	AUG	2000	12	.00	.00	.00	0.	1	SEP	2000	60	1.50	.05	1.45	144.
30	AUG	2100	13	.00	.00	.00	0.	1	SEP	2100	61	.00	.00	.00	134.
30	AUG	2200	14	.00	.00	.00	0.	1	SEP	2200	62	.00	.00	.00	121.
30	AUG	2300	15	.00	.00	.00	0.	2	SEP	0000	64	.00	.00	.00	97.
31	AUG	0000	16	.00	.00	.00	0.	2	SEP	0100	65	.00	.00	.00	86.
31	AUG	0100	17	.00	.00	.00	0.	2	SEP	0200	66	.00	.00	.00	76.
31	AUG	0200	18	.00	.00	.00	0.	2	SEP	0300	67	.00	.00	.00	67.
31	AUG	0300	19	.00	.00	.00	0.	2	SEP	0400	68	.00	.00	.00	59.
31	AUG	0400	20	.00	.00	.00	0.	2	SEP	0500	69	.00	.00	.00	51.
31	AUG	0500	21	.00	.00	.00	0.	2	SEP	0600	70	.00	.00	.00	45.
31	AUG	0600	22	.00	.00	.00	0.	2	SEP	0700	71	.00	.00	.00	39.
31	AUG	0700	23	.00	.00	.00	0.	2	SEP	0800	72	.00	.00	.00	35.
31	AUG	0800	24	.00	.00	.00	0.	2	SEP	0900	73	.00	.00	.00	32.
31	AUG	0900	25	.50	.50	.00	0.	2	SEP	1000	74	.00	.00	.00	29.
31	AUG	1000	26	5.50	5.50	.00	0.	2	SEP	1100	75	.00	.00	.00	27.
31	AUG	1100	27	6.00	6.00	.00	0.	2	SEP	1200	76	.00	.00	.00	25.
31	AUG	1200	28	4.00	3.84	1.16	221.	2	SEP	1300	77	.00	.00	.00	23.
31	AUG	1300	29	6.00	4.97	1.03	282.	2	SEP	1400	78	.00	.00	.00	21.
31	AUG	1400	30	16.90	10.44	6.46	556.	2	SEP	1500	79	.00	.00	.00	20.
31	AUG	1500	31	11.10	4.95	6.15	1055.	2	SEP	1600	80	.00	.00	.00	18.
31	AUG	1600	32	14.00	4.88	9.12	1679.	2	SEP	1700	81	.00	.00	.00	16.
31	AUG	1700	33	16.00	4.30	11.70	2269.	2	SEP	1800	82	.00	.00	.00	15.
31	AUG	1800	34	20.00	4.09	15.91	2760.	2	SEP	1900	83	.00	.00	.00	14.
31	AUG	1900	35	20.00	3.13	16.87	3302.	2	SEP	2000	84	.00	.00	.00	13.
31	AUG	2000	36	16.00	2.02	13.98	3878.	2	SEP	2100	85	.00	.00	.00	12.
31	AUG	2100	37	84.00	6.70	77.30	6113.	2	SEP	2200	86	.00	.00	.00	11.
31	AUG	2200	38	14.00	.73	13.27	9872.	2	SEP	2300	87	.00	.00	.00	10.
31	AUG	2300	39	14.00	.66	13.34	9153.	3	SEP	0000	88	.00	.00	.00	9.
1	SEP	0000	40	5.50	.24	5.26	5530.	3	SEP	0100	89	.00	.00	.00	9.
1	SEP	0100	41	5.50	.07	5.43	3595.	3	SEP	0200	90	.00	.00	.00	9.
1	SEP	0200	42	1.70	.07	1.63	2736.	3	SEP	0300	91	.00	.00	.00	8.
1	SEP	0300	43	1.70	.07	1.63	2326.	3	SEP	0400	92	.00	.00	.00	8.
1	SEP	0400	44	8.00	.32	7.68	1560.	3	SEP	0500	93	.00	.00	.00	8.
1	SEP	0500	45	8.00	.11	7.89	1341.	3	SEP	0600	94	.00	.00	.00	7.
1	SEP	0600	46	2.90	.11	2.79	1150.	3	SEP	0700	95	.00	.00	.00	7.
1	SEP	0700	47	1.10	.04	1.06	927.	3	SEP	0800	96	.00	.00	.00	7.
1	SEP	0800	48	10.20	.37	9.83									

TOTAL RAINFALL = 300.10, TOTAL LOSS = 64.71, TOTAL EXCESS = 235.39

PEAK FLOW (CU M/S)	TIME (HR)	(CU M/S)	6-HR	24-HR	72-HR	95.00-HR
9872.	37.00	6332.	170.976	2669.	931.	706.
		(MM)	136781.	288.257	301.803	301.803
		(1000 CU M)		230606.	241442.	241442.

CUMULATIVE AREA = 800.00 SQ KM

STATION SUB6



11800 58.0
11900 59.0
12000 60.0
12100 61.0
12200 62.0
12300 63.0
20000 64.0
20100 65.0
20200 66.0
20300 67.0
20400 68.0
20500 69.0
20600 70.0
20700 71.0
20800 72.0
20900 73.0
21000 74.0
21100 75.0
21200 76.0
21300 77.0
21400 78.0
21500 79.0
21600 80.0
21700 81.0
21800 82.0
21900 83.0
22000 84.0
22100 85.0
22200 86.0
22300 87.0
30000 88.0
30100 89.0
30200 90.0
30300 91.0
30400 92.0
30500 93.0
30600 94.0
30700 95.0
30800 96.0

RUNOFF SUMMARY, AVERAGE FLOW IN CUBIC METERS PER SECOND
AREA IN SQUARE KILOMETERS

OPERATION	STATION	PEAK FLOW	TIME OF PEAK	AVERAGE FLOW FOR MAXIMUM PERIOD			BASIN AREA	MAXIMUM STAGE	TIME OF MAX STAGE
				6-HOUR	24-HOUR	72-HOUR			
HYDROGRAPH AT	SUB1A	1071.14	35.00	783.68	408.03	144.61	70.00		
HYDROGRAPH AT	SUB1B	1022.03	35.00	764.95	375.54	132.03	70.00		
2 COMBINED AT	SUB1C	2093.18	35.00	1545.13	783.51	276.63	140.00		
HYDROGRAPH AT	SUB1C	2599.15	36.00	1996.48	985.76	347.87	185.00		
HYDROGRAPH AT	SUB2A	3019.69	36.00	2243.72	1080.42	380.21	209.00		
HYDROGRAPH AT	SUB2B	820.34	37.00	492.43	190.18	64.96	56.00		
2 COMBINED AT	SUB2C	3763.89	36.00	2736.15	1270.31	445.16	265.00		
ROUTED TO	RESER	3385.25	37.00	2673.25	1270.42	445.16	265.00	6.57 37.00	
HYDROGRAPH AT	SUB2C	5713.95	37.00	3768.70	1695.95	592.44	429.00		
HYDROGRAPH AT	SUB3	1571.98	37.00	776.37	294.81	99.34	106.00		
2 COMBINED AT	SUB4	7285.94	37.00	4543.80	1990.61	691.78	535.00		
HYDROGRAPH AT	SUB4	8461.53	37.00	5251.71	2255.32	784.18	638.00		
HYDROGRAPH AT	SUB5	1541.23	37.00	782.82	290.74	99.56	109.00		
2 COMBINED AT	SUB6	10002.77	37.00	6034.53	2545.11	883.74	747.00		
HYDROGRAPH AT	SUB6	9871.84	37.00	6332.44	2669.05	931.49	800.00		

*** NORMAL END OF HEC-1 ***

1.8 STREAMFLOW FORECASTING MODEL

1.8.1 Selection of Models

Several existing simulation models were considered and reviewed to constitute the flood forecasting model for this project. The selection process is basically in the light of the state of the art of the model equations used, both requirement and their availability, the computer facility capabilities at INDHRI, and that the model provides suitable and acceptable results at reasonable cost within the forecast lead time frame. In view of these, two models were particularly examined namely the National Weather Service River Forecast System (National Weather Service, 1984) and the Flood Hydrograph Package (HEC-1) of the U.S. Army Corps of Engineers (1985). The NWSRFS is an operational model for continuous time streamflow simulation and real-time river forecasting. It is composed of several models developed independently including Sacramento soil moisture accounting model (SAC), snow accumulation and ablation model, routing models such as layered coefficient routing technique, Muskingum routing and unit hydrograph, precipitation and temperature models and extended streamflow prediction model. Also, this model contains procedures for data processing and analyses for calibration, testing and forecasting and auxiliary programs for data preprocessing and data file manipulation prior to and after model runs. On the other hand, the HEC-1 model simulates single event rainfall/snowmelt runoff processes. In this model, the surface runoff response of the river basin to precipitation is represented as a network system of hydrologic and hydraulic components such as overland flow plane, stream channel, pump station, diversion channel or a reservoir. For a given precipitation hydrograph the rainfall excess is derived

using loss rate equations and routed via a unit hydrograph or kinematic wave method to obtain the surface runoff hydrograph. A baseflow component can also be added to the surface runoff at a basin or subbasin outlet using empirical methods. Sophisticated hydrologic analysis of basin wide flow-frequencies and analysis of expected annual flood damages may also be accomplished.

So far, the NWSRFS and HEC-1 models had found applications in several countries especially in the United States. Both models, however, have advantages and disadvantages which are unique but complementary to each other. On one hand, an attractive component of the NWSRFS model is the Sacramento soil moisture accounting model which has conceptually sound rainfall to channel inflow components transformation as compared to using precipitation loss equations in the HEC-1 model. On the other hand, the use of kinematic wave routing model of HEC-1 is more favored to the other routing techniques but is not available in the NWSRFS model. Thus, combining the Sacramento soil moisture accounting model of the NWSRFS and the kinematic wave routing procedure of HEC-1 is believed to have more promising applications.

1.8.2 Development of SACKW Model

The latest version of the NWSRFS model obtained for this project is too large and requires a tremendous amount of core memory. Besides a PC-version of the NWSRFS model is not yet available and its use is definitely precluded in the IBM-34 computer at INDRHI. In view of this, steps were taken to develop a small version of NWSRFS model which specifically involved adapting the Sacramento soil-moisture accounting model (SAC) of the huge NWSRFS model. Similarly, the kinematic wave (KW) routing model of the HEC-1 model has been incorporated in the SAC

model. In essence therefore, the streamflow forecasting model finally developed is the combination of Sacramento soil moisture accounting and kinematic wave routing referred to as SACKW model here.

For purposes of model calibration, the constrained Rosenbrock optimization routine presented by Kuester and Mize (1973) has also been adapted as an option to automatically calibrate the parameters of the SAC component of the model. This method is a sequential search technique which has been proven effective in finding the maximum or minimum of a multivariable, nonlinear objective function subject to nonlinear inequality constraints. This optimization technique is readily adaptable to SAC model since no derivatives are required.

1.8.3 Description of Model Components

Primarily, the SACKW model can be partitioned into two major components, namely: The Sacramento soil moisture accounting, and the kinematic wave routing. Given below are descriptions of each component. A subsection is also included to describe the watershed partitioning and timing considerations of the model.

1.8.3.1 Sacramento Soil Moisture Accounting Model

Referring to Figure 1.8.1, the Sacramento model computes various runoff components which are added together as total channel inflow and subsurface discharge through a soil moisture accounting procedure from a linkage of five basic soil moisture storages. Another function of the Sacramento model deals with the evapotranspiration process which has significant role in moisture movement in the hydrologic cycle.

The use of five basic storages and their linking mechanism is intended to provide a simple but effective representation of the vertical and horizontal movement of water through and over the soil. As

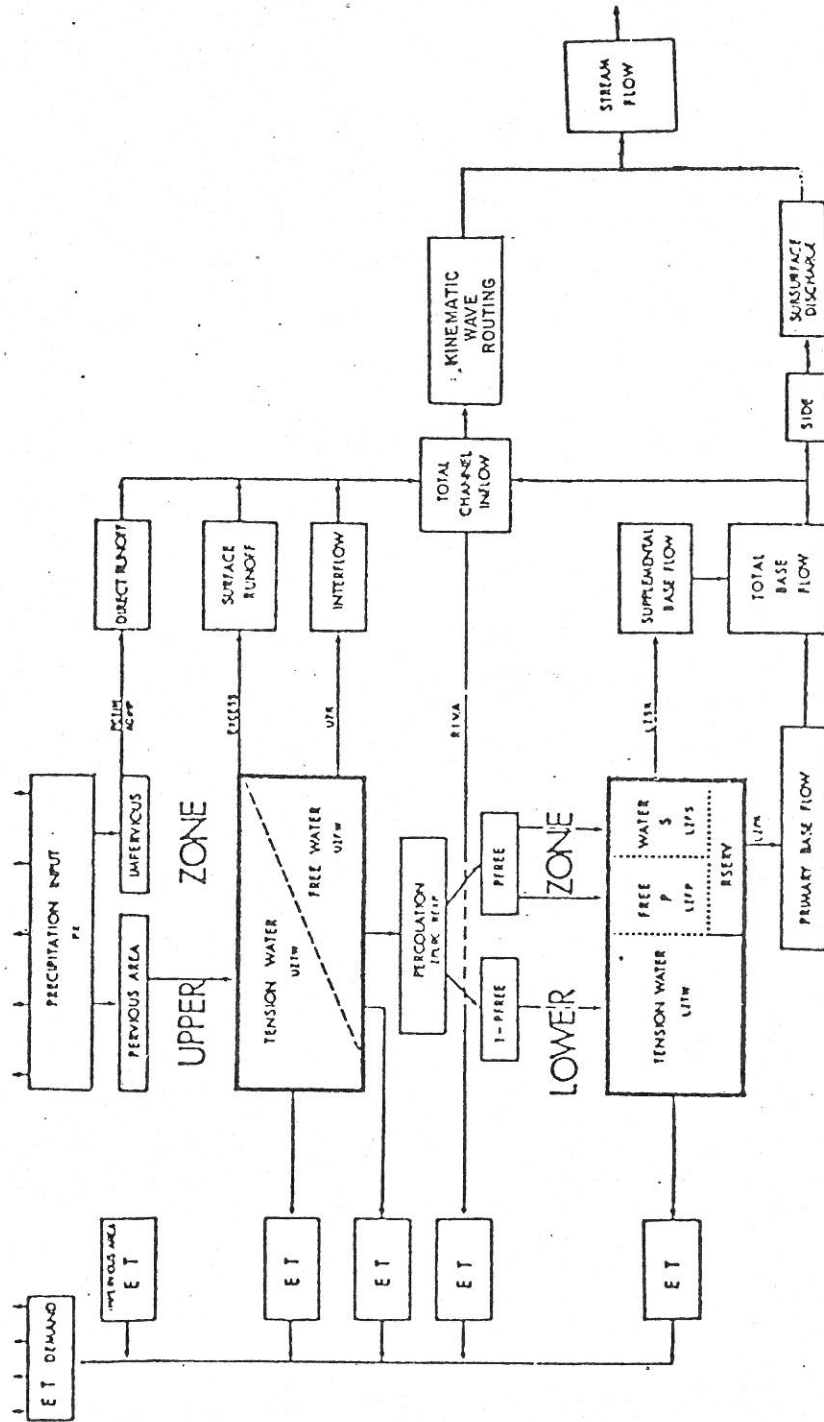


Figure 1.8.1. Flow chart of Sacramento soil-moisture accounting model.

shown in Figure 1.8.1, the five storages are: 1) the upper zone tension water which is the volume of water tightly bound to soil molecules but can be removed by evapotranspiration but occurs within such a shallow layer of soil that it is rapidly replaced by rainfall before sufficient moisture can accumulate to initiate the runoff process, 2) the upper zone free water is that water needed to produce fully effective wetting front which is a key factor to the percolation process and provides the source for rapid drainage in the form of interflow, 3) the lower zone tension water is that volume of water utilized by plants for evapotranspiration but not readily transferred from roots to leaf systems as in the shallower upper zone tension water, 4) the lower zone supplemental free water represents the source of rapidly draining component of subsurface runoff known as supplemental baseflow, and, 5) the lower zone primary water provides the source of slowly draining runoff component referred to as primary baseflow.

Three subprocesses in the Sacramento model worthwhile mentioning are the percolation process, evapotranspiration process and runoff process. The percolation process essentially centers on computing the water that percolates to deeper soil through vertical drainage prior to interflow calculation. The percolation rate is controlled by the amount of water available for percolation in the upper zone free water and the deficiency of lower zone moisture volume translated into the lower zone percolation demand as shown in Figure 1.8.2. The evapotranspiration process consists of evaporation from the area covered by surface water or phreatophyte vegetation and evapotranspiration from upper zone and lower zone water storages. Evaporation is computed at a potential rate while evapotranspiration from the soil moisture storages varies with the

PBASE = The continuing percolation rate under saturated conditions.

ZPERC = The number of PBASE units which must be added to the continuing saturated percolation rate to define the maximum percolation condition.

REXP = The exponent which defines the curvature in the percolation curve with changes in the lower zone soil moisture deficiency.

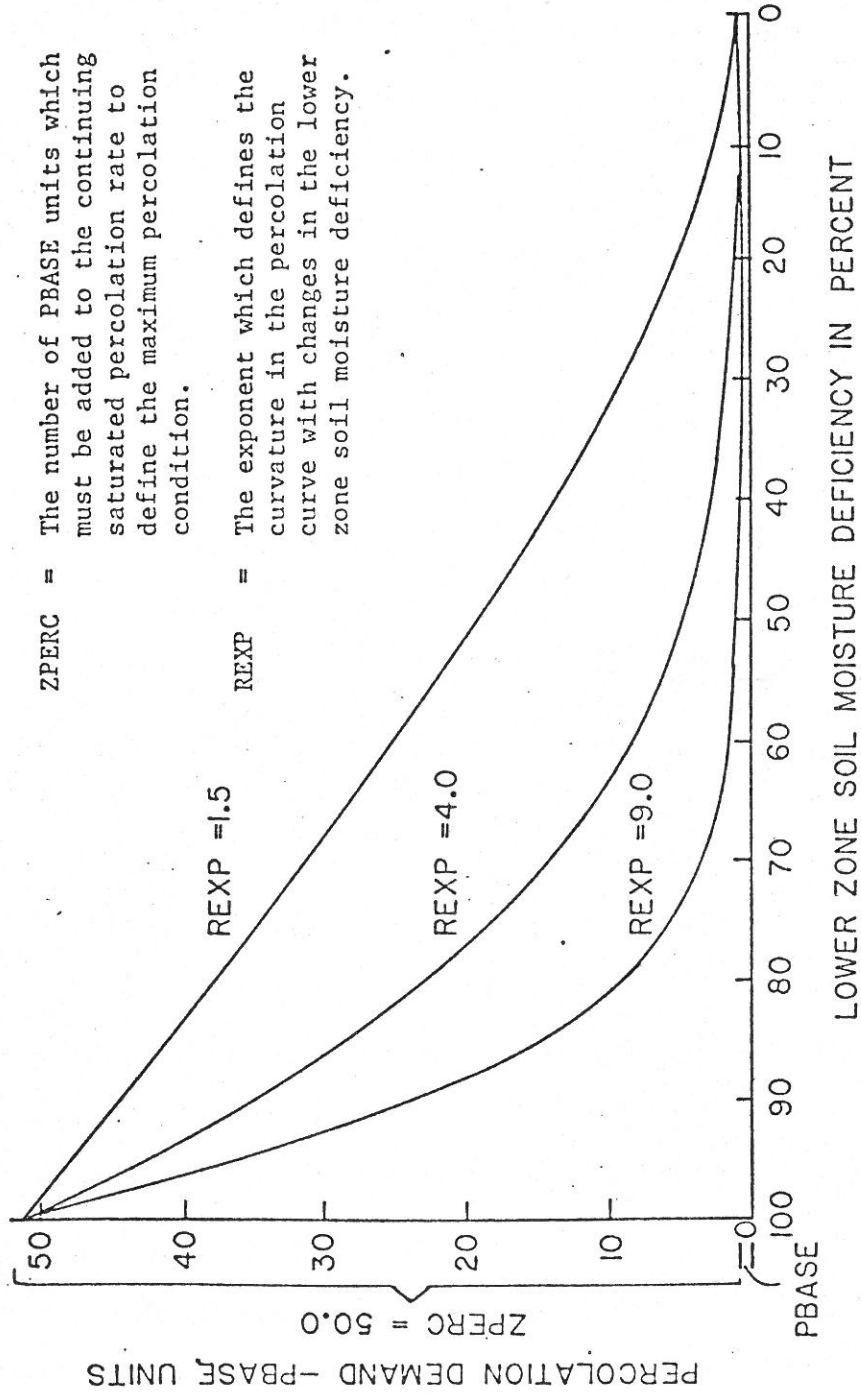


Figure 1.8.2. Relationship of percolation demand and lower zone moisture deficiency.

volume and distribution of tension water storage and evapotranspiration demand. Starting with a saturated soil, and exposing it to a constant evapotranspiration demand would produce an effective evapotranspiration use curve of the type illustrated in Figure 1.8.3.

The runoff resulting from soil-moisture accounting are given in five basic forms. These are: 1) the impervious runoff from impervious areas, and direct runoff from temporary impervious area, 2) surface runoff which occurs when the upper zone free water storage is full and the precipitation intensity exceeds the rate of percolation and interflow, 3) interflow resulting from lateral drainage of the upper zone free water storage, 4) supplemental baseflow, and 5) primary baseflow. The first three runoff components represents the total channel inflow while the latter two is the total baseflow. In the SACKW model, the so called total channel inflow constitute the surface runoff contribution to the stream flow hydrograph routed via the kinematic wave routing methodology and the total baseflow is the subsurface runoff contribution to streamflow. This baseflow component is added to the routed streamflow at the basin or subbasin outlet using a linear, decay weighting function of current and some specified previous time total baseflows.

1.8.3.2 Kinematic Wave Routing Model

The kinematic wave model provides the mechanism of water movement over the land surface and in stream channels towards the basin or subbasin outlet. The input to this model is the total channel inflow or hydrograph which is assumed to be uniform over the subbasin. In determining the subbasin runoff by the kinematic wave method, three conceptual elements are used: flow planes, collector channels and a

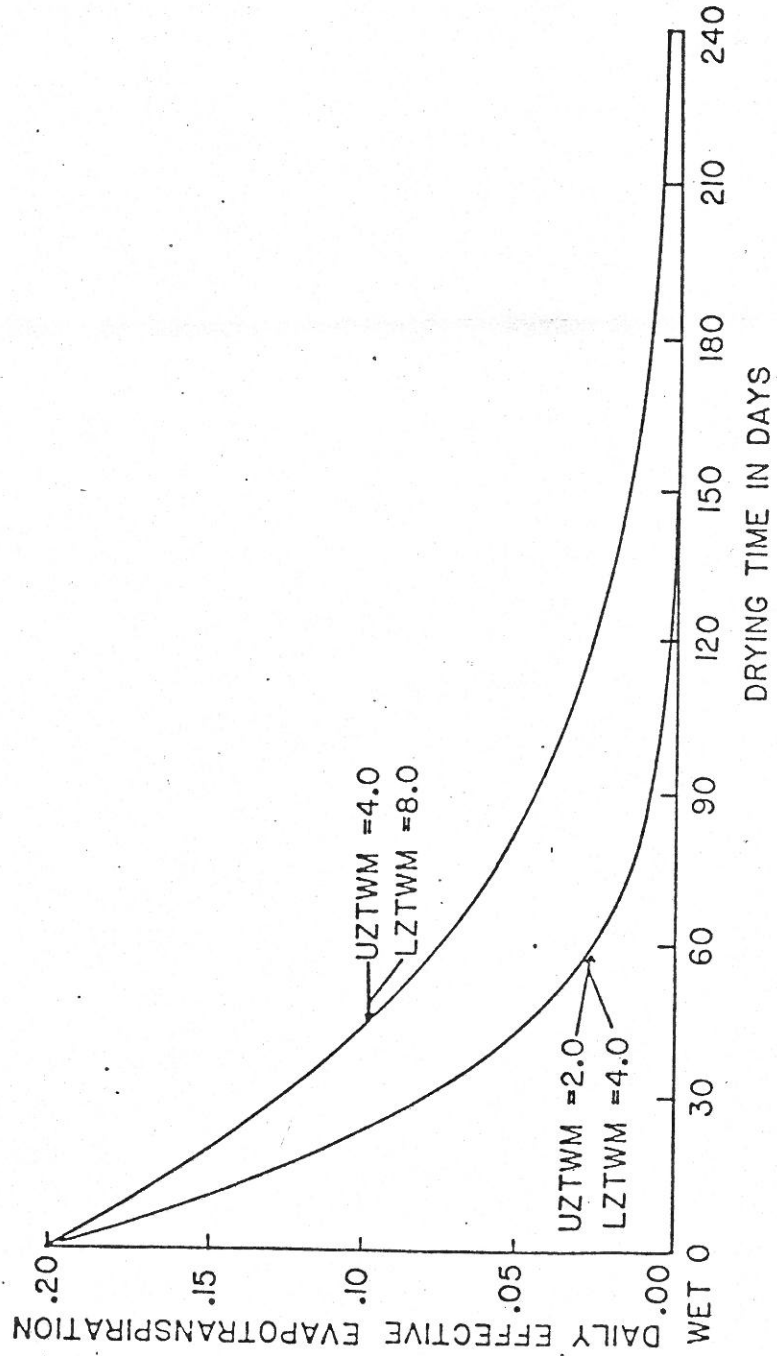


Figure 1.8.3. Daily effective evapotranspiration from initially wet soils exposed to a constant evapotranspiration demand.

main channel as shown in Figure 1.8.4. The kinematic wave method assumes that the bed slope and water surface slope are equal and acceleration effects are negligible. In this manner, the overland or channel flow can be represented as a power function of cross-sectional area with power coefficients related to flow geometry and surface roughness. The movement of flood wave is described solely by the continuity equation in partial differential form. Through combining the flow and continuity equations, a finite difference approximation can be developed and likewise solved by finite difference methods. For further details of the kinematic wave method and its solution, the HEC-1 User's Manual (U.S. Corps of Engineers, 1985) may be consulted.

1.8.3.3 Watershed Partitioning and Timing Considerations

Partitioning of the watershed provides the distributed parameter capability of the model. This is done to account for the spatial and temporal variabilities of the physical and hydrological characteristics of the basing, the climatic variables, and basin-wide response characteristics. In the SACKW model, the watershed can be partitioned into two levels. The first level, partitions the watershed into subwatersheds where each subwatershed is a homogeneous unit of the SAC model in terms of the SAC model parameters. Rainfall and evapotranspiration are assumed homogeneous or uniform over one subwatershed. The second level of partitioning divides further a subwatershed into smaller homogeneous units representing individual flow planes. Each flow plane is assumed to have homogeneous kinematic wave parameters.

The timing considerations for the model refers to the time basis of model operation. The model is set up to simulate basin hydrology on an

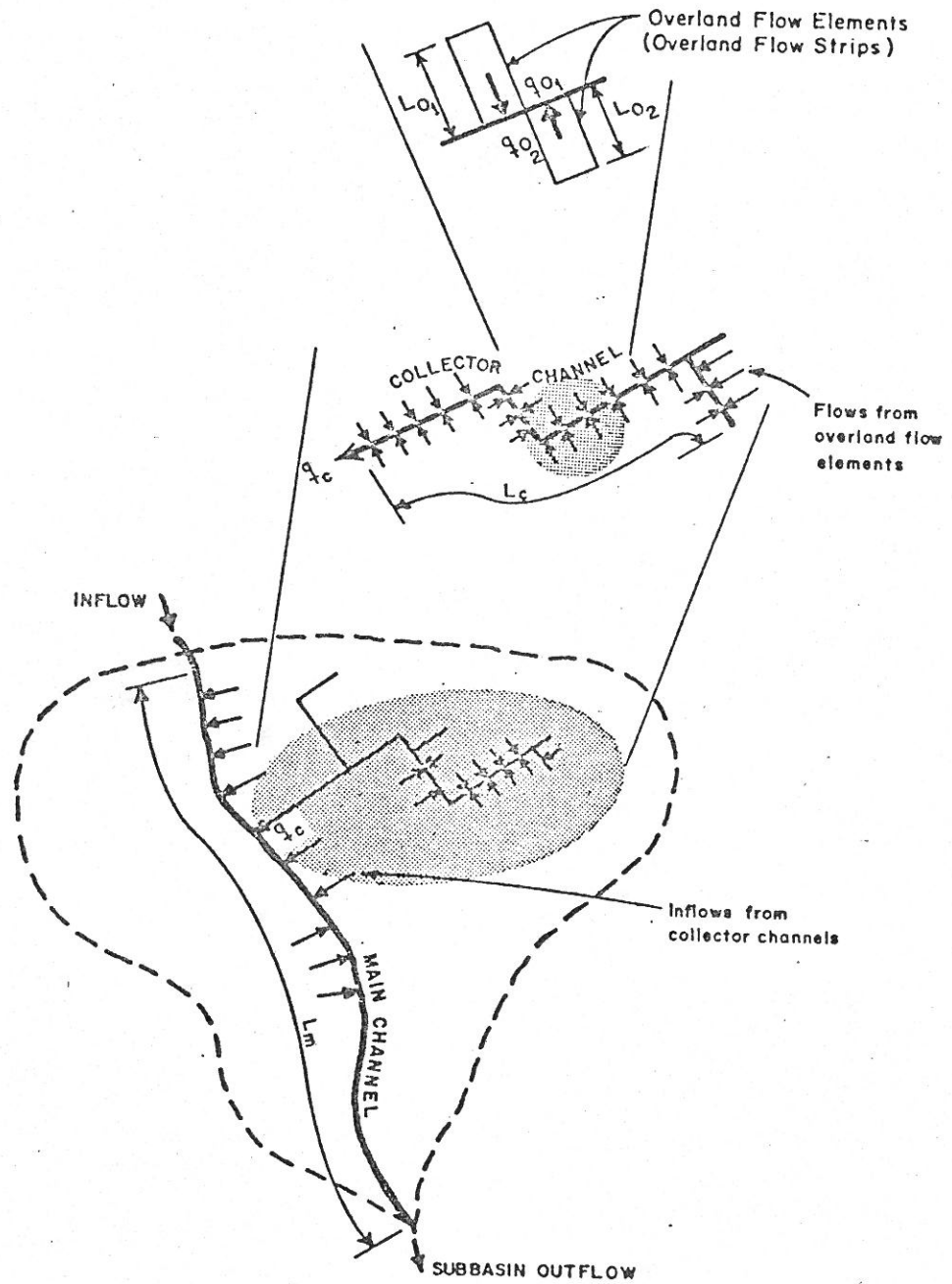


Figure 1.8.4. Relationship between flow elements in kinematic wave routing.

hourly basis and on longer time intervals which is a multiple integer of 1 hour. In the SAC model, the computational time interval is always done in an hourly basis so that rainfall, evapotranspiration and streamflow if given in longer time intervals are uniformly transformed into hourly data. Model outputs however are given on time intervals equal to those specified in the input data. In the kinematic wave routing computations, the time interval may vary depending on the stability criteria requirements of the finite-difference numerical sequence.

1.8.4 Model Calibration

This section reports the SACKW model calibration for the Nizao basin. In the ensuing text, the SACKW user's manual given in Appendix 1.8.A may be referred to which describes the model usage and capabilities, input requirements, program description, output information and some guidelines for model calibration.

Shown in Figure 1.8.5 is the Nizao basin and its watershed partitioning. It is decided that the basin be partitioned into three subwatersheds: La Estrechura, Palo de Caja and Paso del Ermitano, for purposes of the SAC model (first level partitioning). For each subwatershed, further partitioning is made to constitute the homogeneous flowplanes in the kinematic wave routing (second level partitioning).

A total of four years (1972-1975) of data is used for model calibration. The rainfall data used which are available hourly are areal averages from nine stations using optimal interpolation. Three sets of areal averaged time series were obtained corresponding to the three first level subwatersheds. Since not all of the nine rainfall stations are recording at one time or another, the areal averaging was

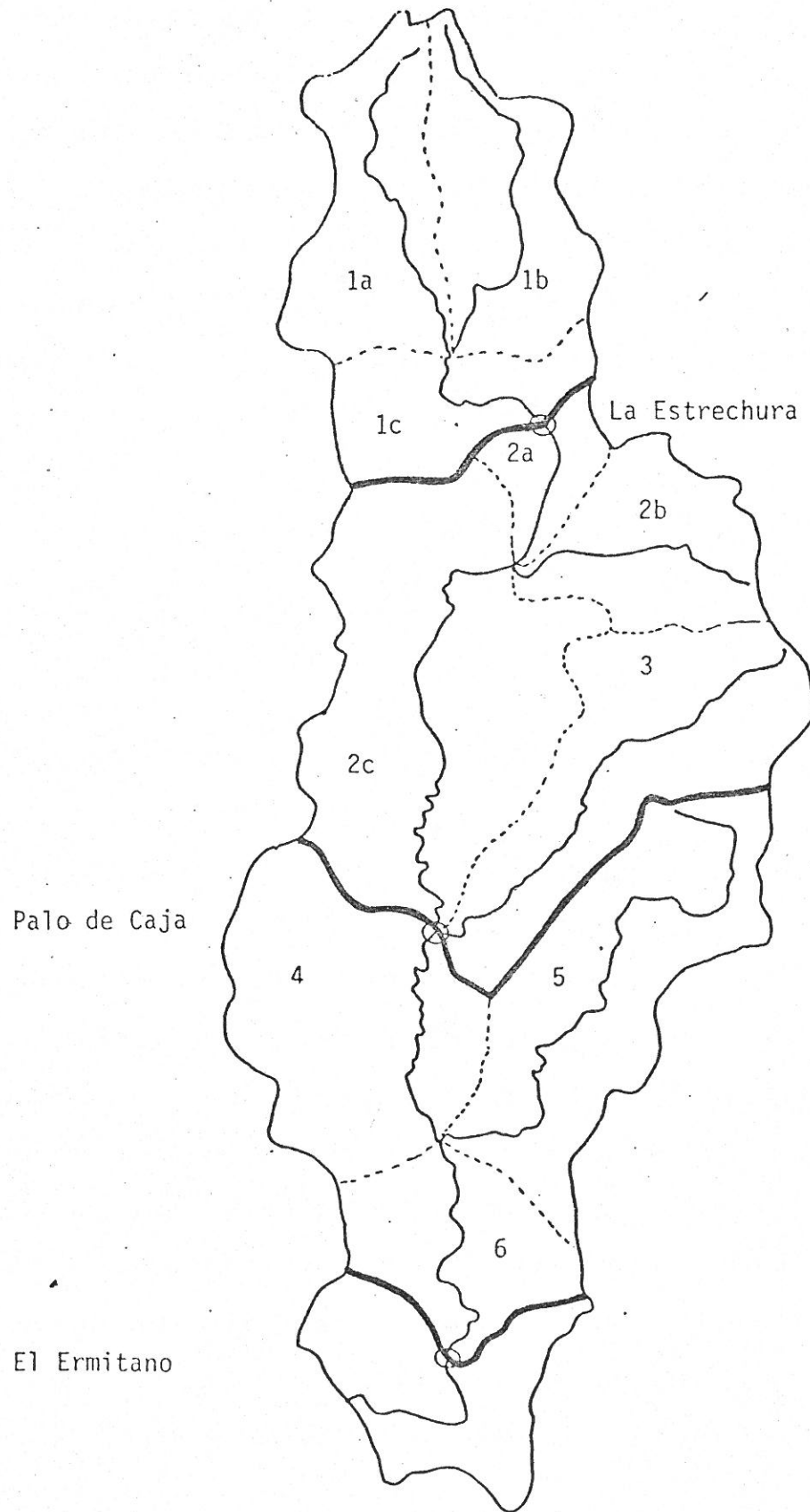
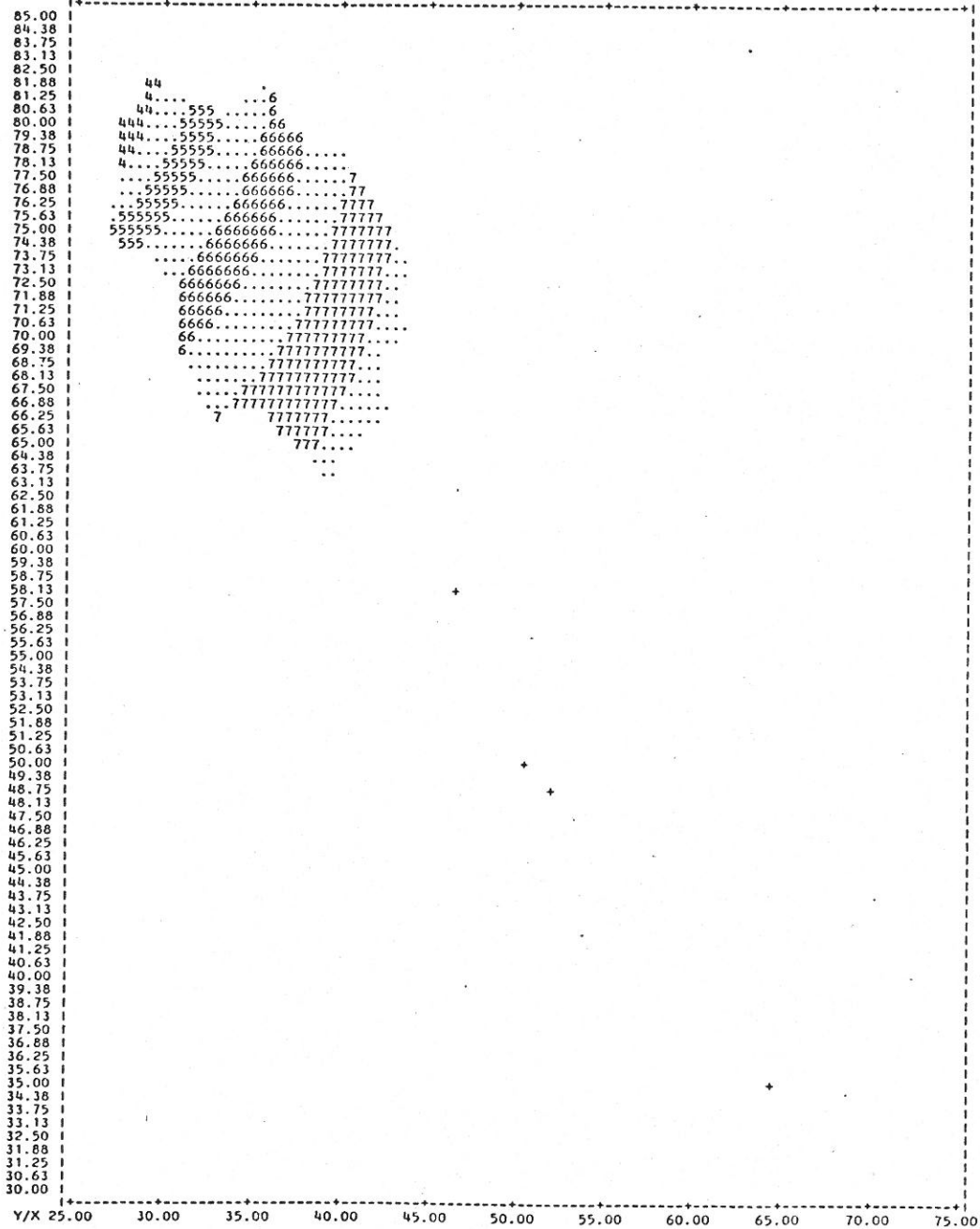


Figure 1.8.5. First and second level watershed partitions of Nizao basin for SACKW model.

done on a case to case basis. As an example, in case of nine stations recording at the same time, Figure 1.8.6. to 1.8.8 show the optimal interpolation areal averages and the weights of each station using a hypothetical data. Nevertheless, Figures 1.8.9 through 1.8.20 show the areal averaged rainfall series plotted on a daily basis for each subwatershed and for each year. Daily streamflow at the outlet of each subwatershed were used for purposes of calibration. Figures 1.8.9 to 1.8.20 also show the daily streamflow for each subwatershed and for each year. The daily evapotranspiration demand data required in the model is obtained from the monthly pan evaporation of Valdesia station after converting them to daily values and multiplying by an adjustment coefficient of 0.7. This demand data is assumed to be uniform all over the basin.

Based on some guidelines suggested for model calibration in Appendix 1.8.A and references herein, the model input parameters were set up for the first year of data to be calibrated. The input file for this first year is given in Figure 1.8.21. The kinematic wave routing model parameters used were those obtained from the HEC-1 model calibration which is presented in Section 1.7.2 of the report. Beginning with year 1972, the best parameter estimates of the SAC model are obtained where some refinements were made using the optimization routine. The least squares objective function was used in the optimization runs between observed and computed daily stream flows at each subwatershed. Based on these 1972 model parameters, the calibration proceeds to 1973, then to 1974, and finally 1975. It is assumed that for four years, the kinematic wave parameters are the same such that only the SAC model parameters change. Given in Figure 1.8.22

OPTIMAL INTERPOLATION BASED ON STRAIGHT DATA LAESTRECHURA SUBBASIN



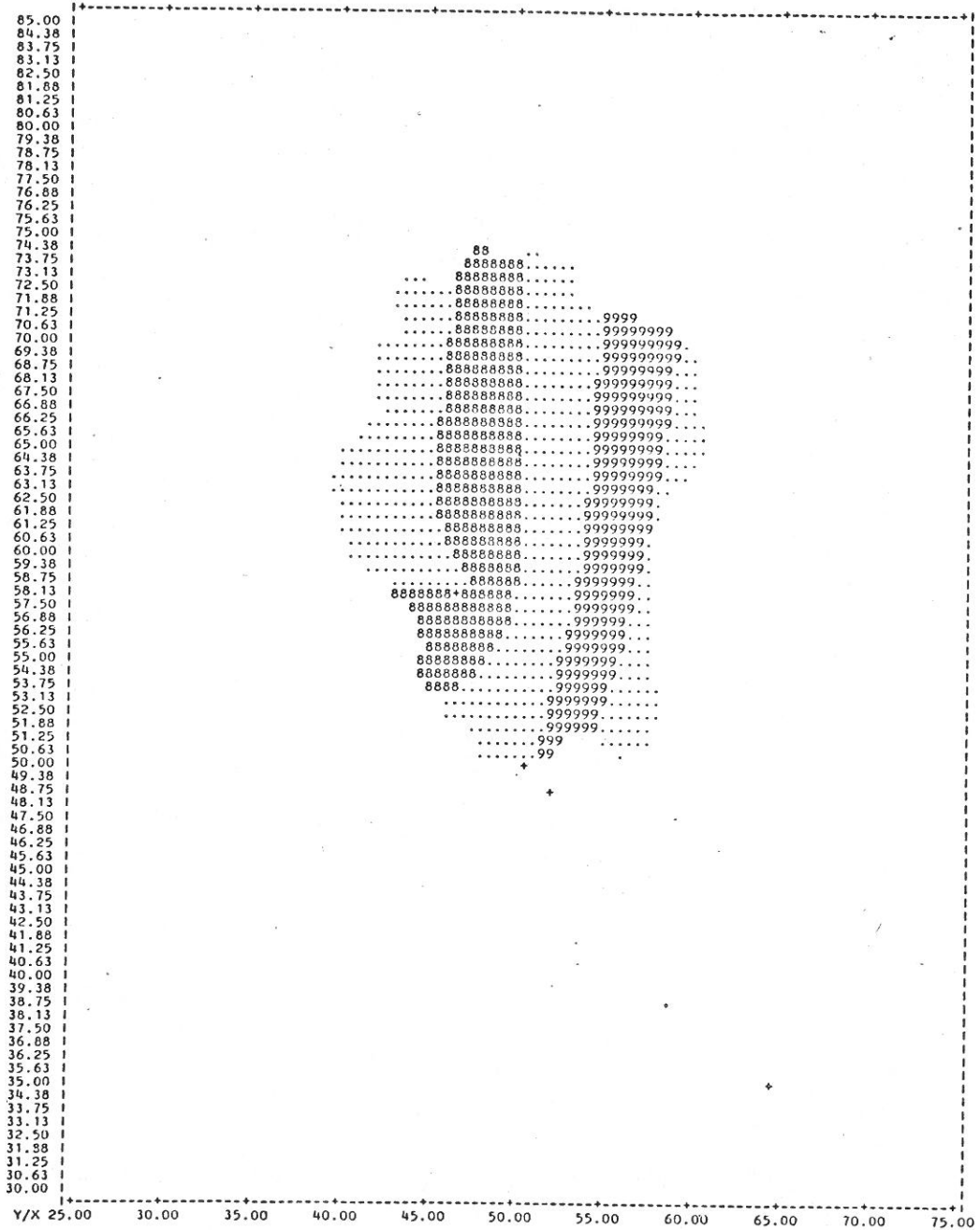
OPTIMAL AREAL AVERAGING

STA.NO.	STA. NAME	INPUT DATA	OPTIMAL WEIGHT
1	EL RIO	28.700	.0784
2	FRICOMBE	94.200	.0374
3	JUNA BON	54.000	.0664
4	LALAGUNA	50.800	.0334
5	QUEMADOS	46.400	.1696
6	NIZAO	46.900	.2404
7	PAI OUECA	52.200	.0240
8	VALDESIA	64.800	.0364
9	VALLENUE	22.640	.3141

SUM OF OPTIMAL WEIGHT = 1.00000
 OPTIMAL AREAL MEAN = 40.917
 STD. ERROR OF AREAL MEAN = 12.7183

Figure 1.8.6. Sample output of rainfall areal averaging for La Estrechura subbasin.

OPTIMAL INTERPOLATION BASED ON STRAIGHT DATA PALO DE CAJA SUBBASIN



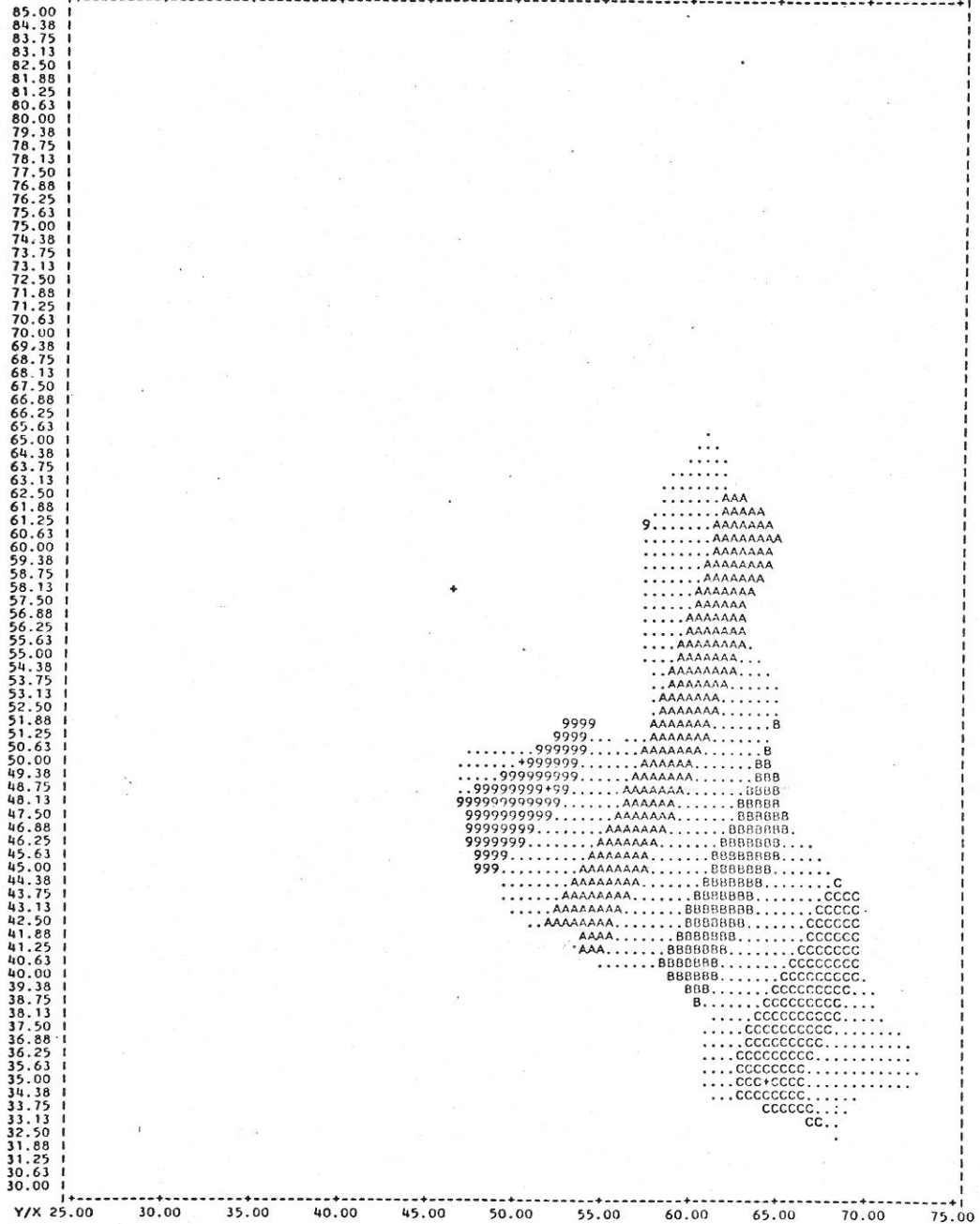
OPTIMAL AREAL AVERAGING

STA. NO.	STA. NAME	INPUT DATA	OPTIMAL WEIGHT
1	EL RIO	28.700	.0269
2	ENCOMBE	94.200	.0456
3	JUMA BON	54.000	.0839
4	LALAGUNA	50.800	.1298
5	QUEMADOS	46.400	.0765
6	NIZAO	46.900	.0423
7	PALODECA	52.200	.0872
8	VALDESLA	64.800	.0482
9	VALLENUE	22.640	.0596

SUM OF OPTIMAL WEIGHT = 1.00000
 OPTIMAL AREAL MEAN = 49.510
 STD. ERROR OF AREAL MEAN = 9.9004

Figure 1.8.7. Sample output of rainfall areal averaging for Palo de Caja subbasin.

OPTIMAL INTERPOLATION BASED ON STRAIGHT DATA PASO DEL ERMITANO SUBBASIN



OPTIMAL AREAL AVERAGING

STA. NO.	STA. NAME	INPUT DATA	OPTIMAL WEIGHT
1	EL RIO	28.700	.0191
2	ENCOMHE	94.200	.0807
3	JUMA RON	54.000	.0404
4	LALAGUNA	50.800	.0310
5	QUEMADOS	46.400	.0204
6	NIZAO	46.900	.0216
7	PAI ODECA	52.200	.2892
8	VALDESTA	64.800	.3430
9	VALLINUE	22.640	.0226

SUM OF OPTIMAL WEIGHT = 1.00000
 OPTIMAL AREAL MEAN = 58.132
 STD. ERROR OF AREAL MEAN = 10.9390

Figure 1.8.8. Sample output of rainfall areal averaging for Paso del Ermitano subbasin.

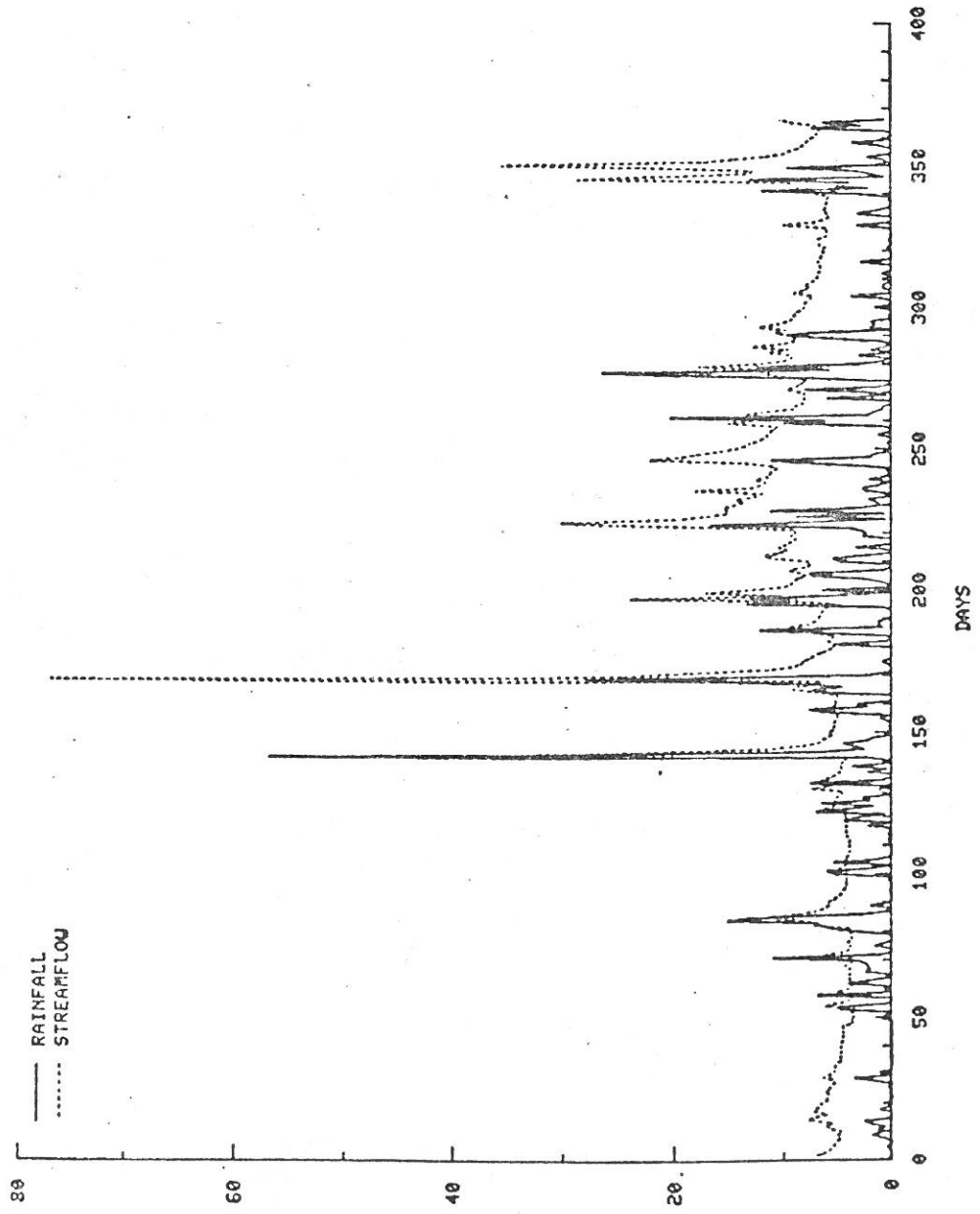


Figure 1.8.9. LA ESTRECHURA - 1972

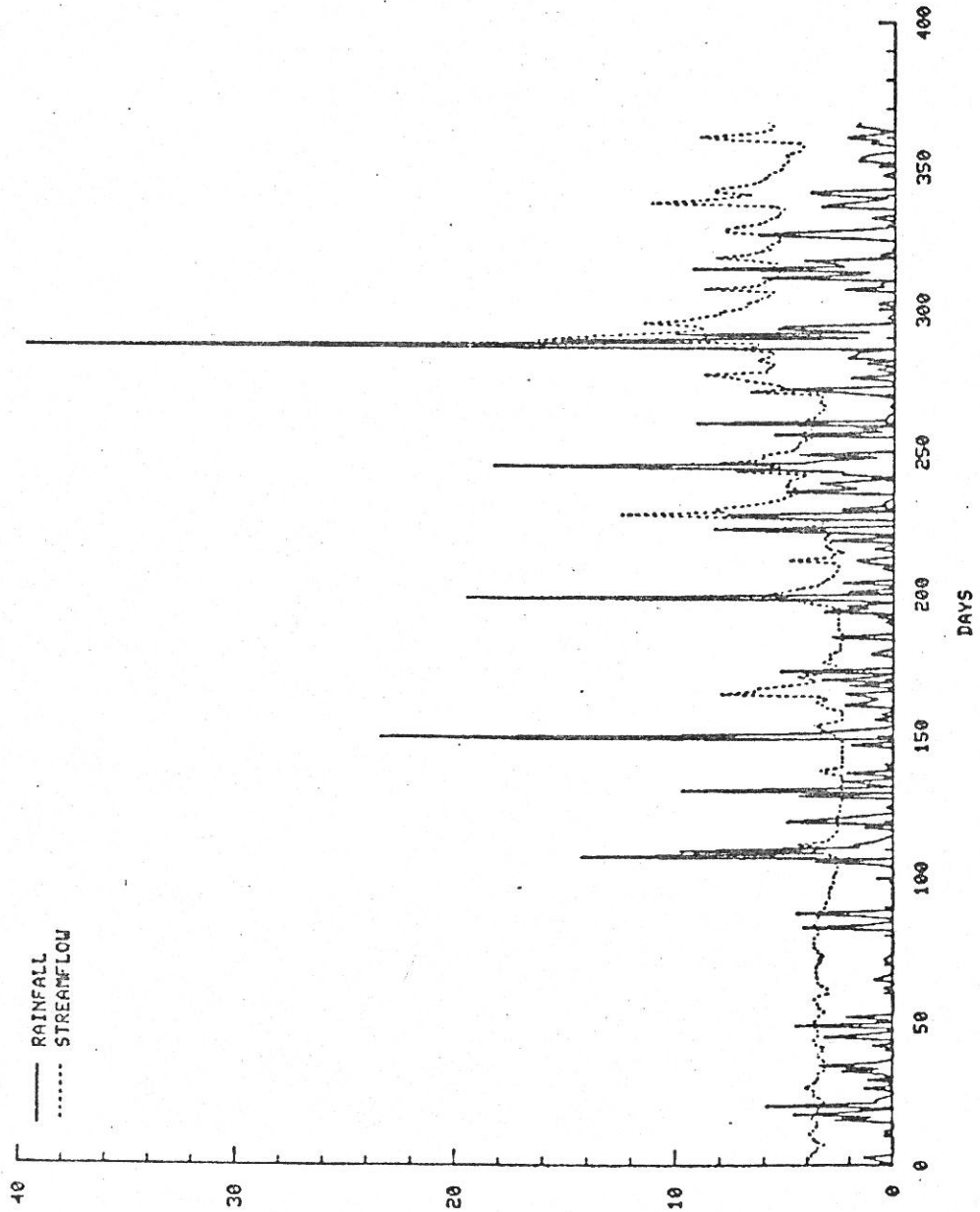


Figure 1.8.10. LA ESTRECHURA - 1973

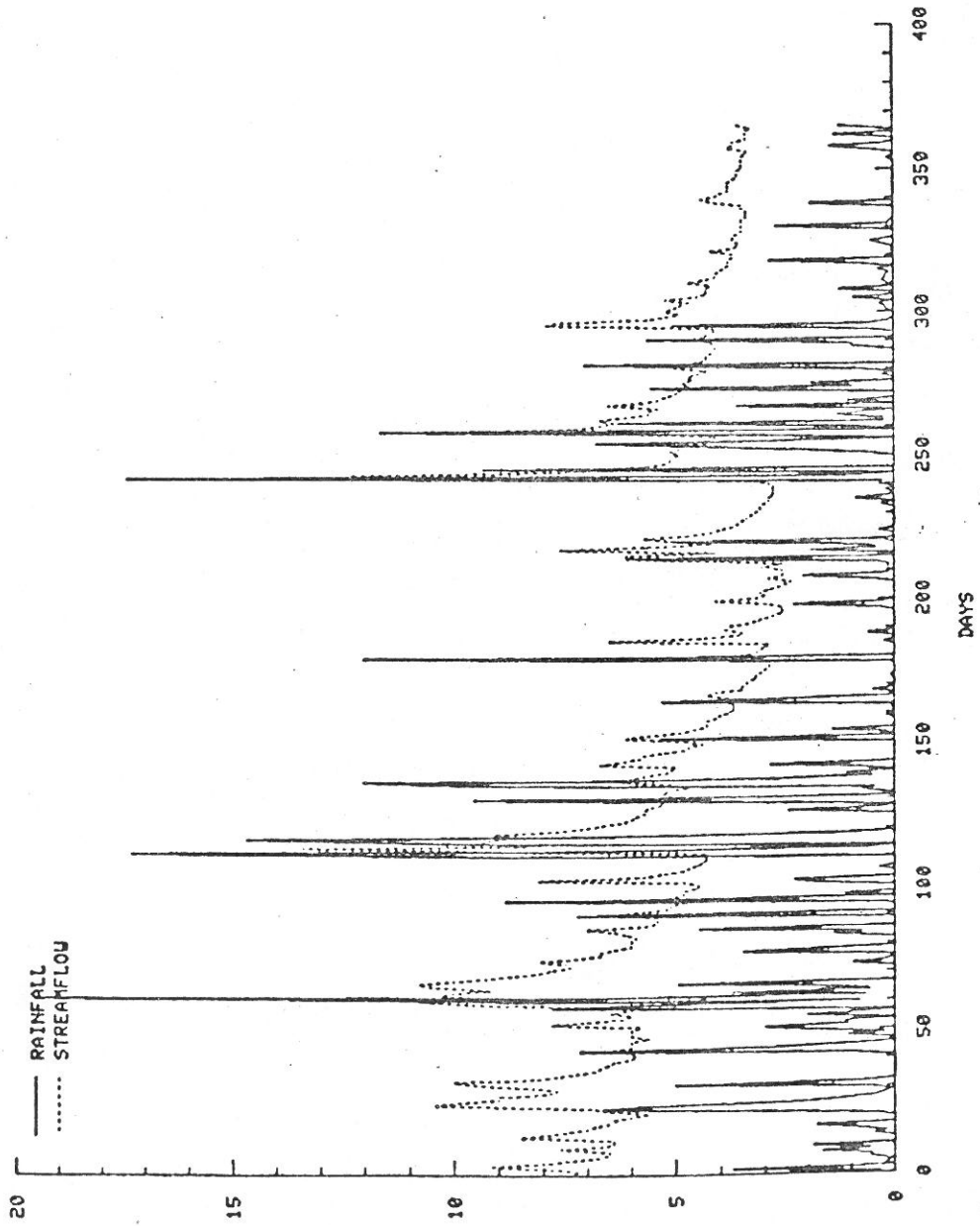


Figure 1.8.11. LA ESTRECHURA - 1974

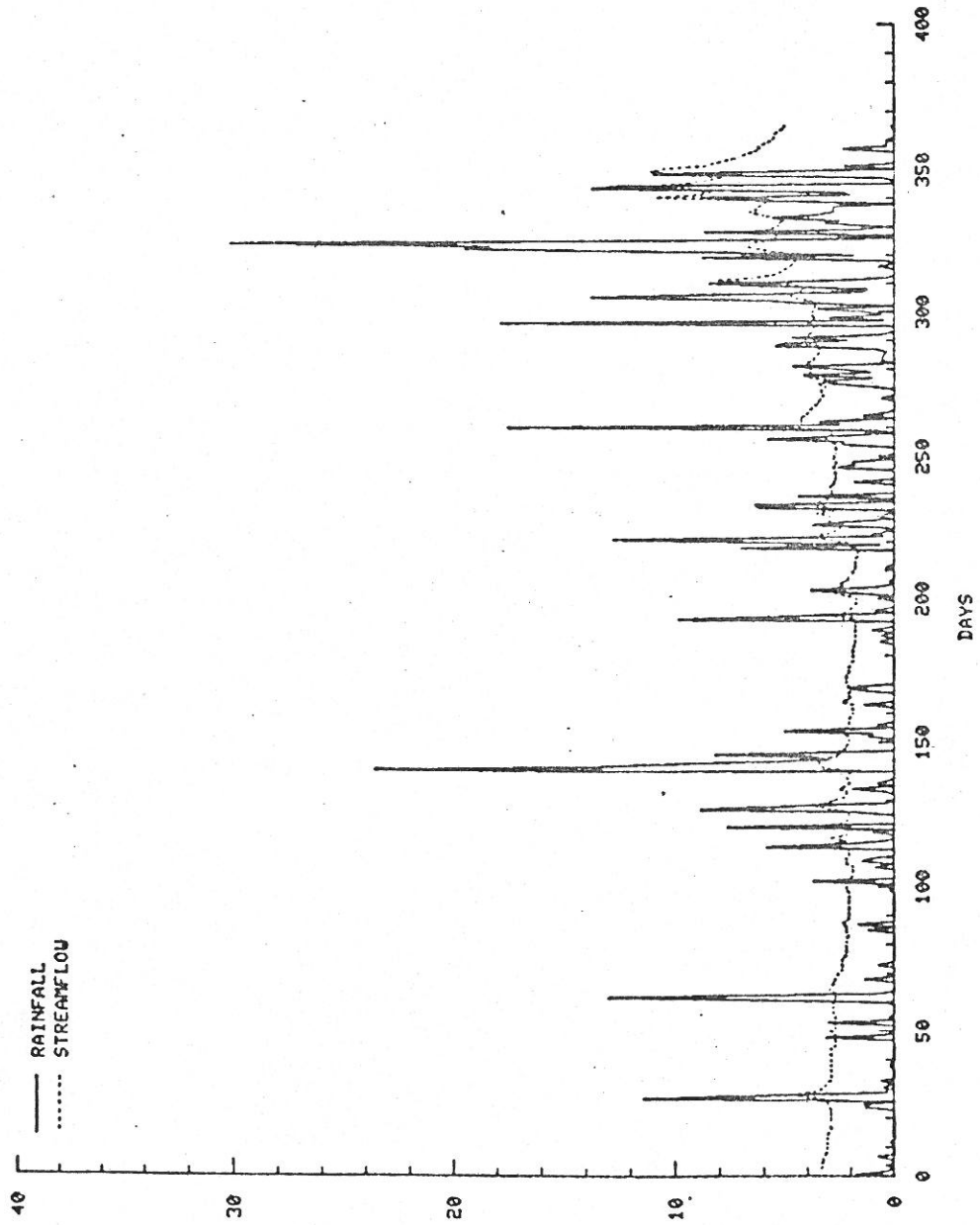


Figure 1.8.12. LA ESTRECHURA - 1975

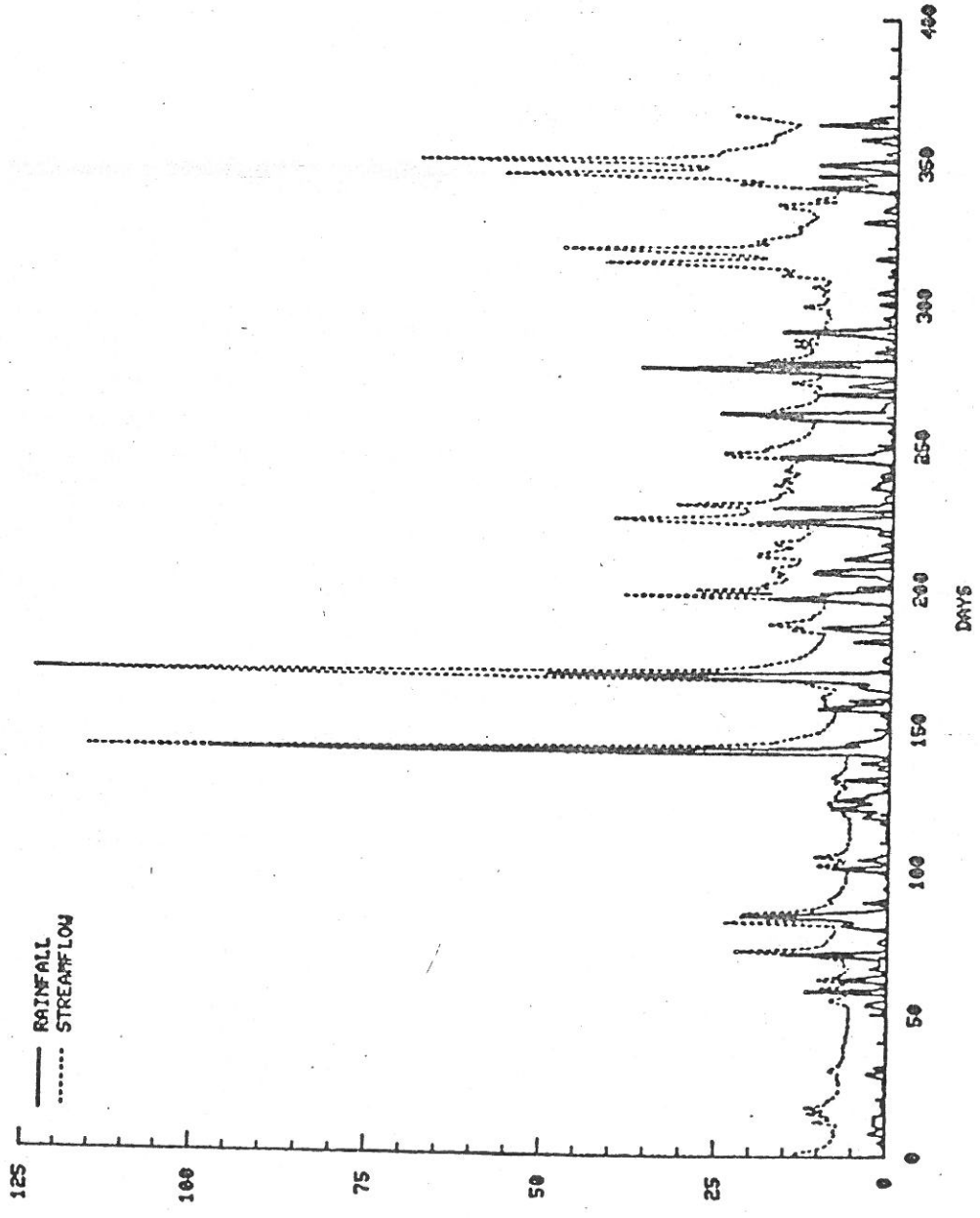


Figure 1.8.13. PALO DE CAJA - 1972

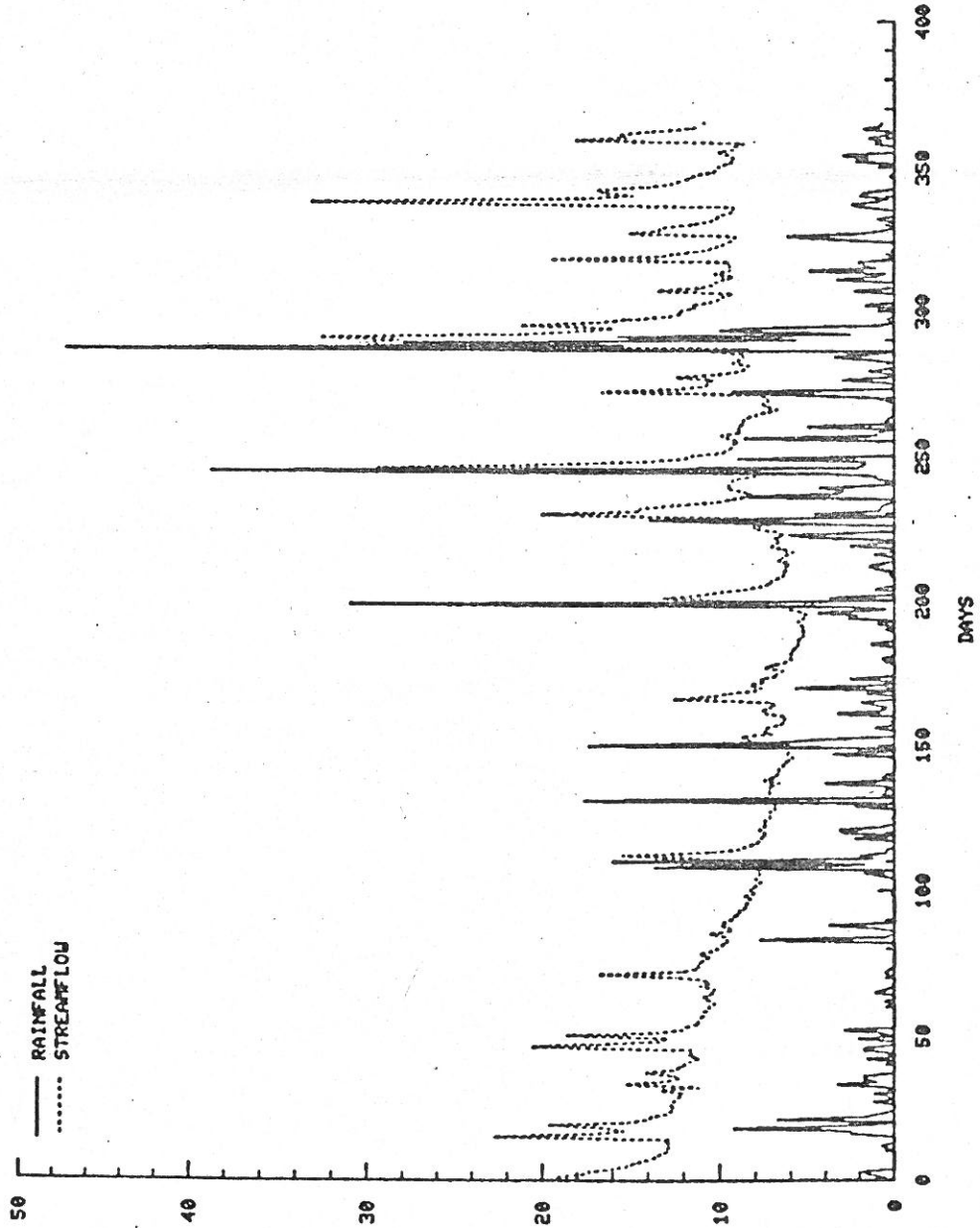


Figure 1.8.14. PALO DE CAJA - 1973

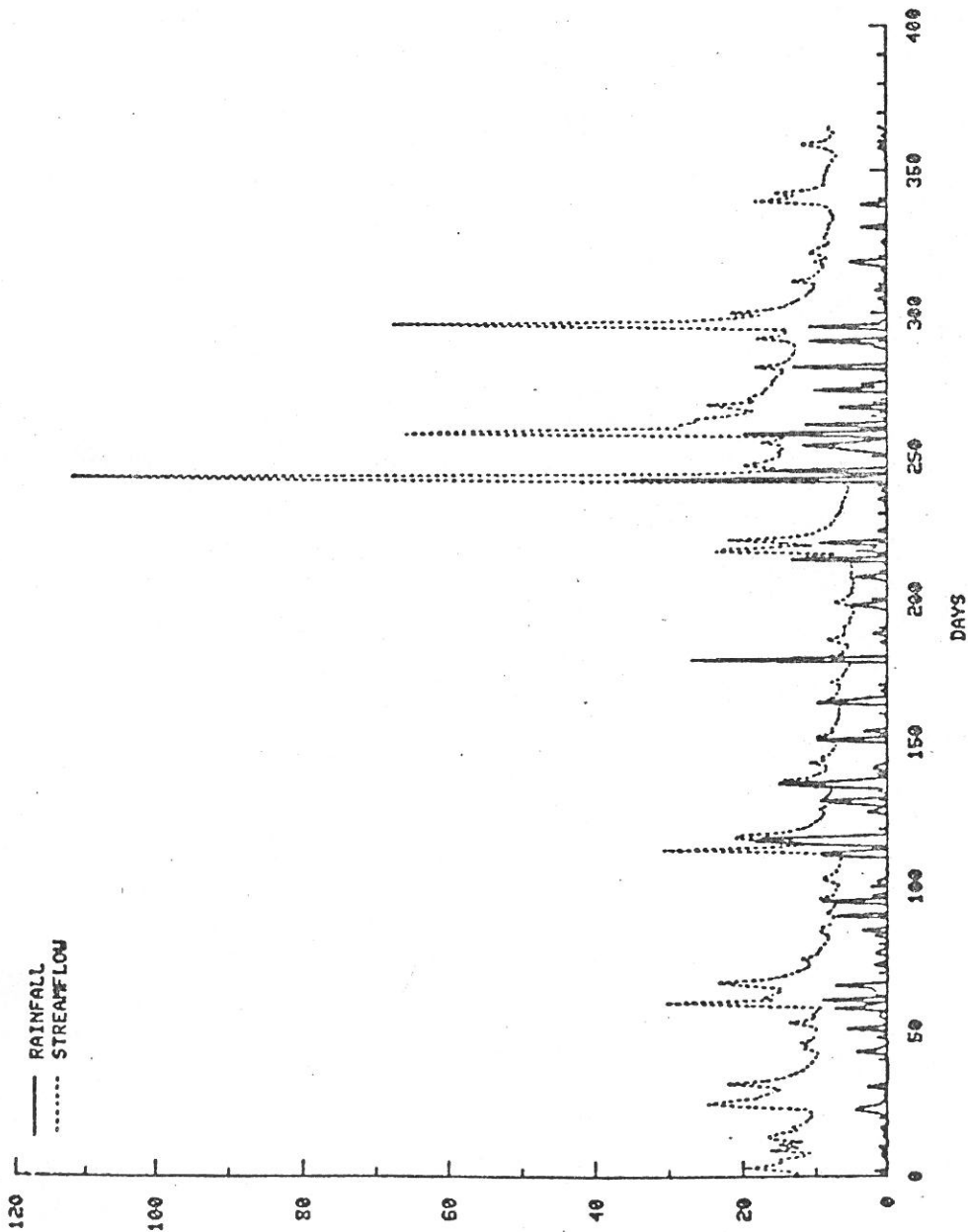


Figure 1.8.15. PALO DE CAJA - 1974

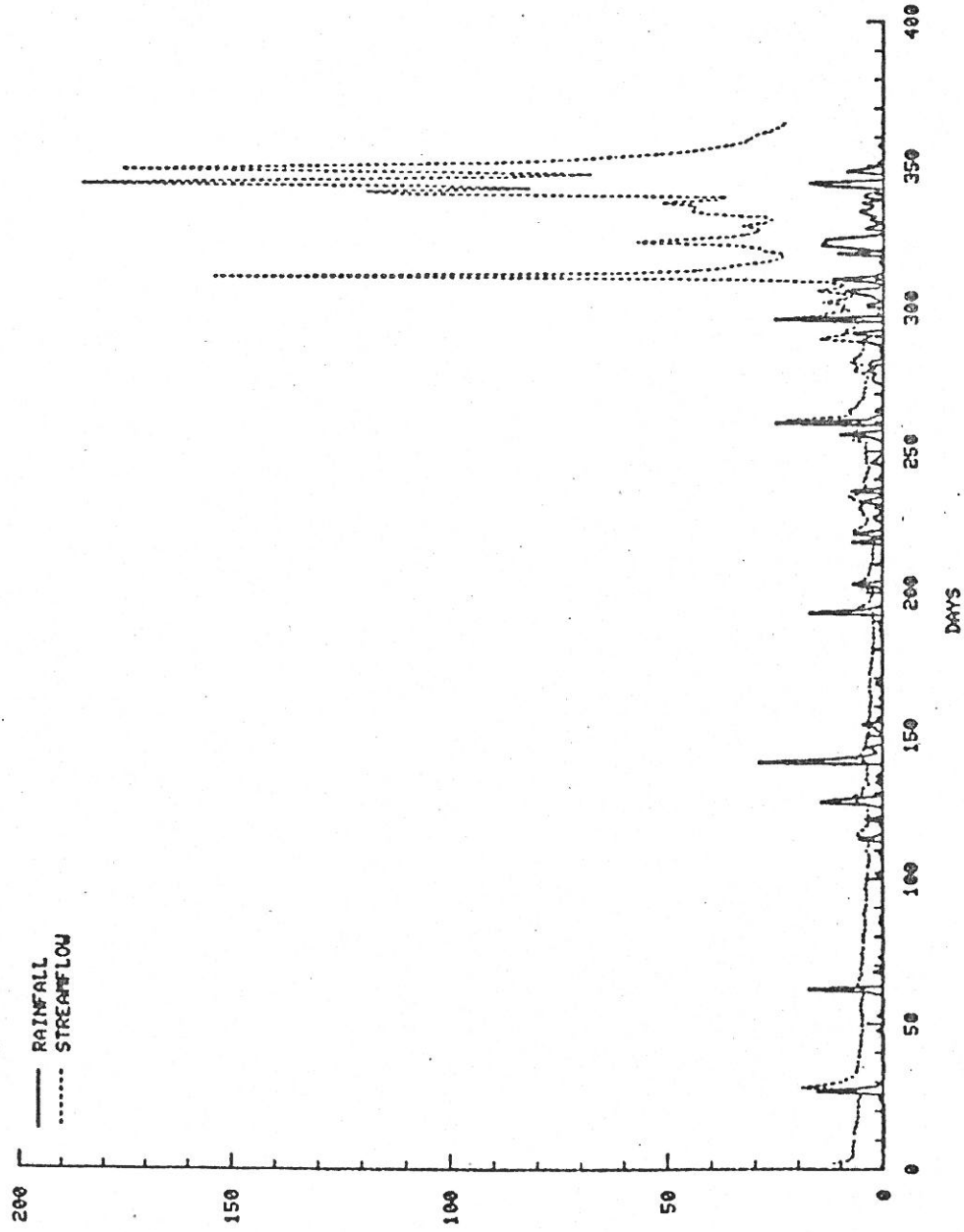


Figure 1.8.16. PALO DE CAJA - 1975

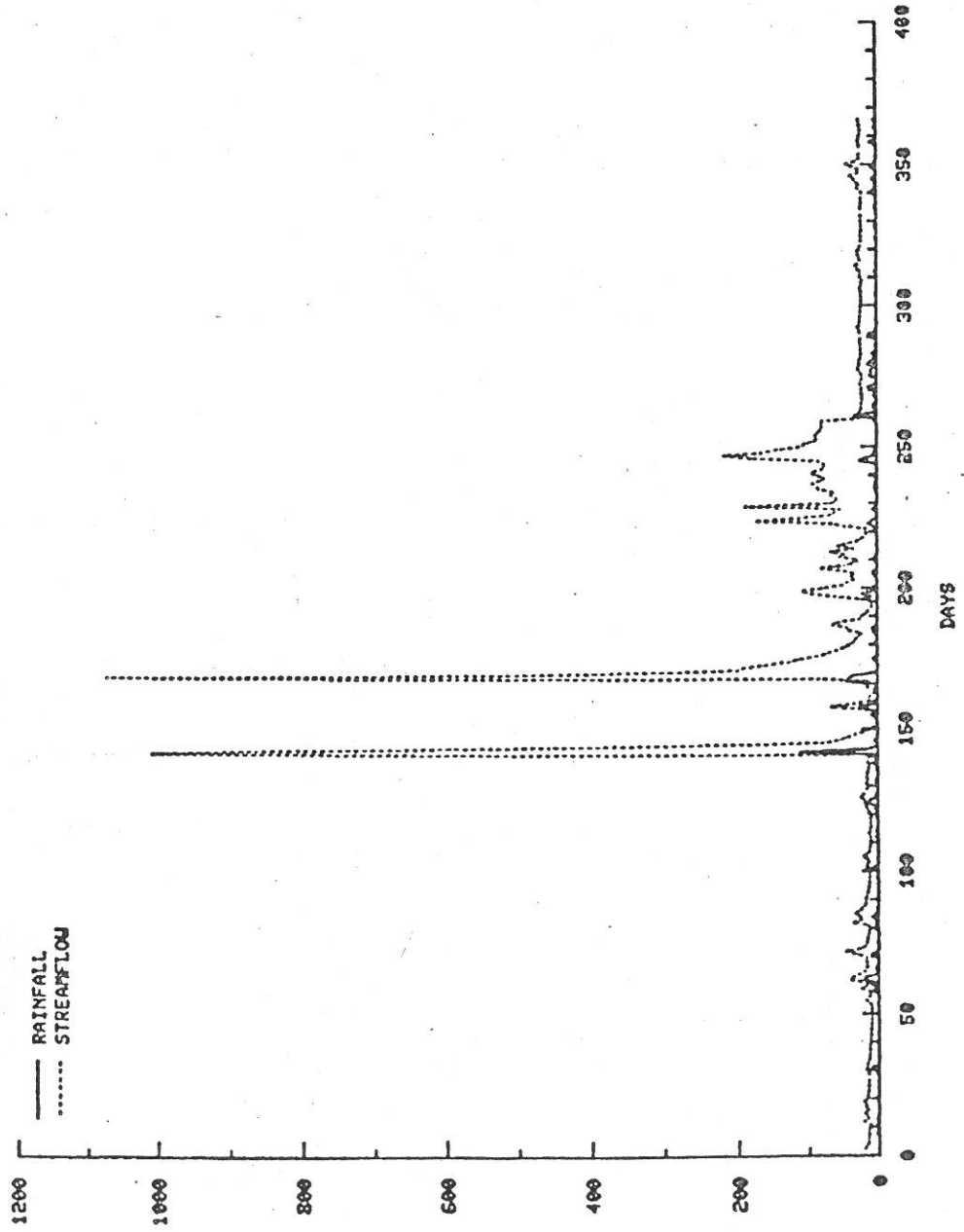


Figure 1.8.17. PASO DEL ERMITANO - 1972

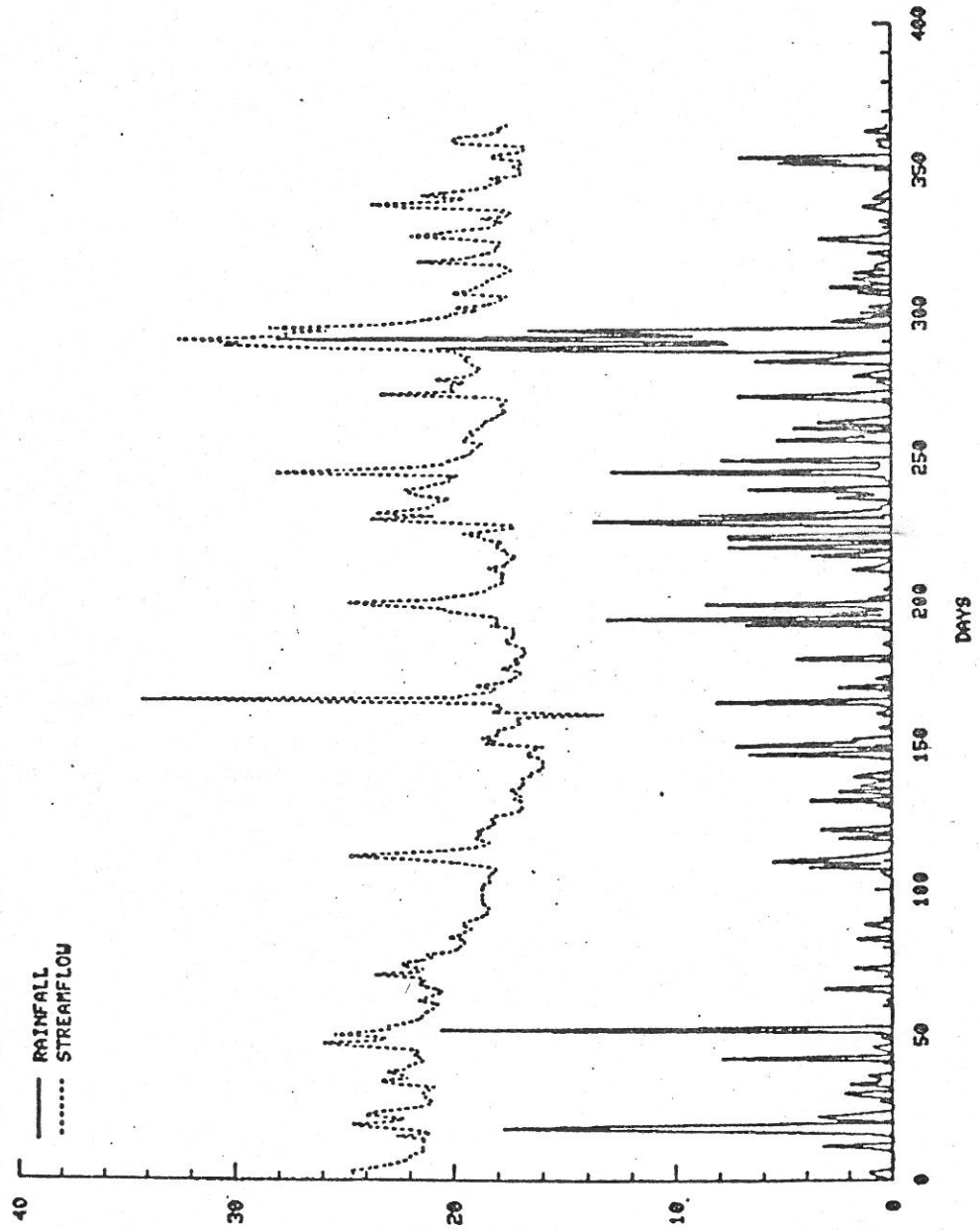


Figure 1.8.18. PASO DEL ERMITANO - 1973

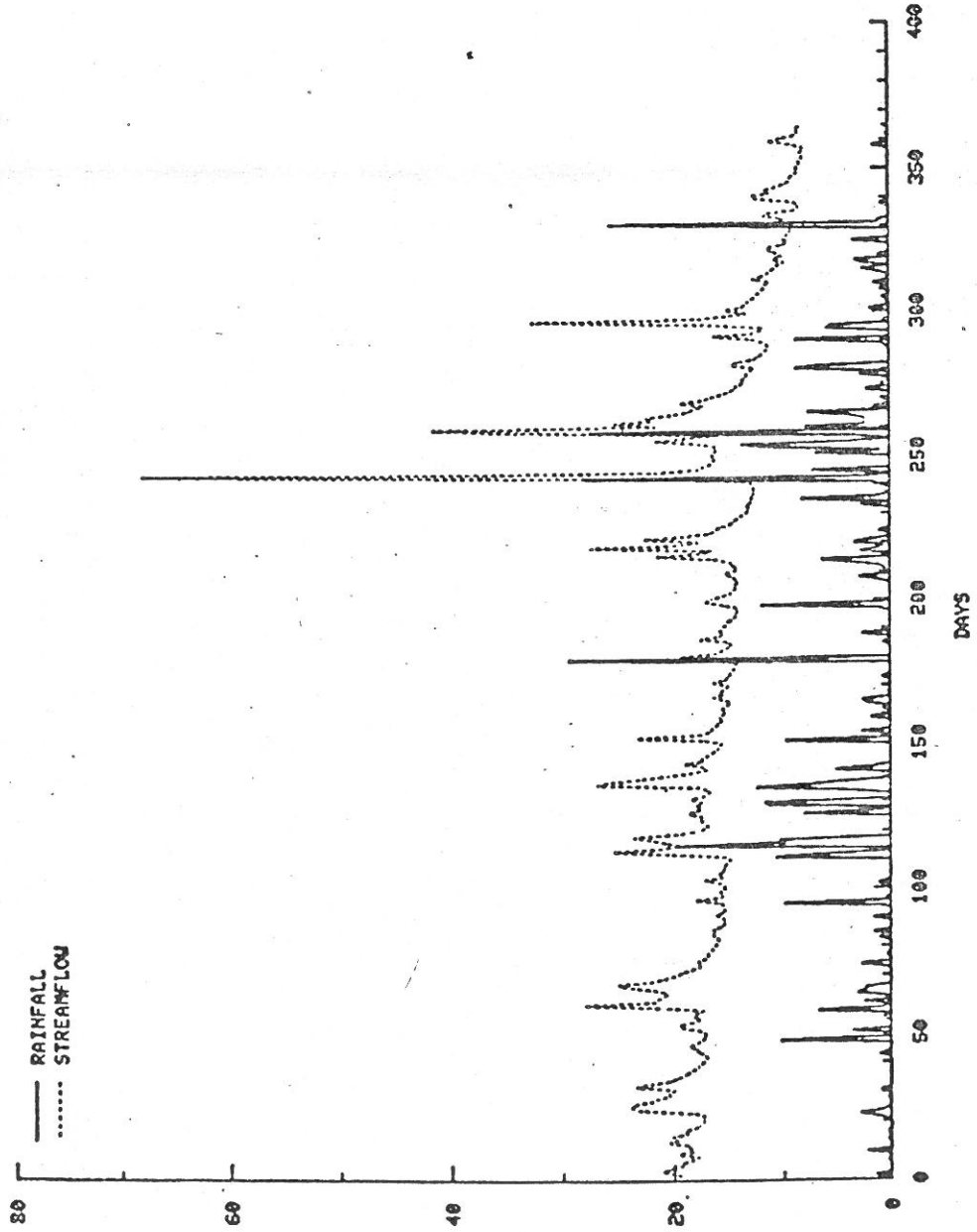


Figure 1.8.19. PASO DEL ERMITANO - 1974

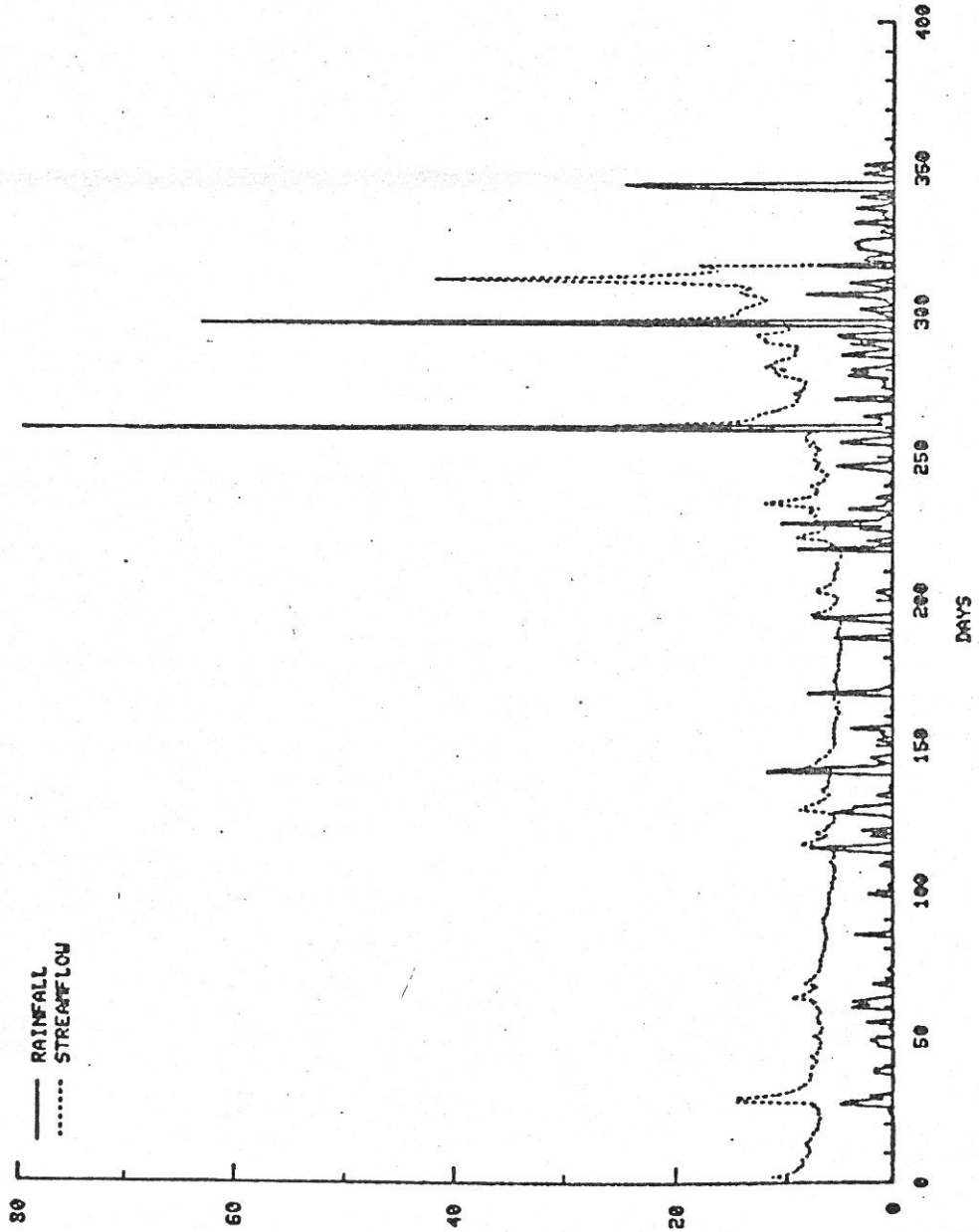


Figure 1.8.20. PASO DEL ERMITANO - 1975

0.3	0.5	0.2		
ROUTE				
103.0				
2000.	0.4	0.3	100.	
33000.	0.0293	0.05	00.0	
5.0	10.0	1.0	14.0	0.0
BASEF	2			
0.3	0.5	0.2		
ADD	2			
END				
RAIN				
1 8784	7	1	1	
(20X, F10.0)				
ETDATA				
24	9	1	1	
(10F8.0)				
FLOW				
24 366	8	1	1	
(20X, F10.0)				
END				
BASIN C - PASO DEL ERMITANO (1972)				
17	0 730 8784	1		
1.0	25.0	0.0001	1.0	1.0
ZPERC	33.70	15.0	150.	5.0
REXP	4.83	0.5	5.0	0.1
SIDE	0.006	0.001	10.0	0.001
UZK	0.900	0.15	0.95	0.01
ADIMP	0.092	0.001	0.5	0.001
RSERV	0.440	0.01	.50	0.01
RIVA	0.829	0.25	0.90	0.01
PCTIM	0.908	0.10	.90	0.01
LZPK	0.012	0.01	.80	0.01
LZSK	0.476	0.01	.50	0.01
PEREE	0.994	0.001	.999	0.0001
UZTWC	98.0	50.0	150.0	2.0
UZFWC	57.0	40.0	80.0	2.0
LZTWC	195.0	40.0	200.0	.5.0
LZFSC	74.65	60.0	150.0	2.0
LZFPC	501.0	400.0	800.0	5.0
ADIMC	350.0	100.0	400.0	10.0
UZTWM	150.0			
UZFWM	100.0			
LZTWM	300.0			
LZF5M	200.0			
LZFPM	800.0			
PXADJ	1.0			
PEADJ	0.7			
END				
ROUTE				
103.0				
5000.0	0.5	0.3	100.0	
18000.	0.0148	0.04	0.0	
5.0	80.0	1.0	23.6	1.0
BASEF	2			
0.3	0.5	0.2		

Figure 1.8.21 (continuation)

ROUTE				
109.0				
3000.	0.4	0.3	100.	
34000.	0.0323	0.05	00.0	
5.0	10.0	1.0	23.6	0.0
BASEF	2			
0.3	0.5	0.2		
ADD	2			
ROUTE				
53.0				
5000.0	0.4	0.3	100.0	
18000.	0.0119	0.040	00.0	
5.0	80.0	1.0	23.6	1.0
BASEF	2			
0.3	0.5	0.2		
END				
RAIN				
1 8784	7	1	1	
(30X, F10.0)				
ETDATA				
24 9	1	1		
(10F8.0)				
FLOW				
24 366	8	1	1	
(30X, F10.0)				

Figure 1.8.21 (continuation)

LA ESTRECHURA

	1972	1973	1974	1975	AVE
ZPERC	87.000	97.000	17.000	147.000	87.000
REXP	.709	.659	1.909	1.819	1.274
SIDE	.176	.207	.175	.219	.194
UZK	.010	.006	.005	.005	.007
ADIMP	.115	.115	.085	.073	.097
RSERV	.350	.500	.593	.621	.516
RIVA	.570	.880	.880	.990	.830
PCTIM	.089	.014	.003	.001	.027
LZPK	.010	.009	.013	.025	.014
LZSK	.006	.005	.015	.029	.014
PEREE	.793	.793	.800	.801	.797
UZTWC	96.000	100.000	148.300	84.600	
UZFWC	34.000	2.320	.103	.010	
LZTWC	197.000	300.000	295.100	274.000	
LZFSC	34.000	99.000	84.680	19.790	
LZFPC	405.000	320.000	267.200	92.020	
ADIMC	115.000	390.000	439.800	349.000	
UZTWM	150.000	150.000	150.000	150.000	150.000
UZFWM	100.000	100.000	100.000	100.000	100.000
LZTWM	300.000	301.000	301.000	301.000	300.750
LZFMS	200.000	200.000	200.000	200.000	200.000
LZFPM	800.000	800.000	800.000	800.000	800.000
PXADJ	1.000	1.000	1.000	1.000	1.000
PEADJ	.700	.700	.700	.700	.700

PALO DE CAJA

	1972	1973	1974	1975	AVE
ZPERC	178.000	198.000	140.000	145.000	165.250
REXP	.400	.275	.440	.285	.350
SIDE	1.050	1.057	1.049	1.000	1.039
UZK	.030	.020	.050	.002	.026
ADIMP	.145	.141	.171	.121	.145
RSERV	.225	.365	.585	.585	.440
RIVA	.339	.439	.739	.939	.614
PCTIM	.015	.001	.031	.001	.012
LZPK	.003	.008	.011	.041	.016
LZSK	.002	.002	.004	.015	.006
PEREE	.987	.987	.989	.999	.991
UZTWC	43.000	100.000	148.300	85.400	
UZFWC	32.500	4.160	.008	.006	
LZTWC	42.500	278.000	214.800	160.800	
LZFSC	44.000	183.000	118.200	65.360	
LZFPC	324.500	721.700	322.900	137.200	
ADIMC	217.500	385.500	437.500	348.900	
UZTWM	150.000	150.000	150.000	150.000	150.000
UZFWM	100.000	100.000	100.000	100.000	100.000
LZTWM	300.000	300.000	300.000	300.000	300.000
LZFMS	200.000	200.000	200.000	200.000	200.000
LZFPM	800.000	800.000	800.000	800.000	800.000
PXADJ	1.000	1.000	1.000	1.000	1.000
PEADJ	.700	.700	.700	.700	.700

PASO DEL ERMITANO

	1972	1973	1974	1975	AVE
ZPERC	33.700	199.300	218.700	199.600	162.825

Figure 1.8.22 Summary of SAC model parameters calibrated for years 1972 to 1975.

REXP	4.830	.502	.130	.100	1.391
SIDE	.006	.372	.360	.236	.244
UZK	.900	.125	.135	.152	.328
ADIMP	.092	.004	.003	.002	.025
RSERV	.440	.470	.750	.190	.463
RIVA	.829	.879	.880	.320	.727
PCTIM	.908	.155	.040	.051	.289
LZPK	.012	.010	.010	.004	.009
LZSK	.476	.075	.016	.003	.143
PFREE	.994	.975	.948	.742	.915
UZWTC	98.000	100.000	148.300	84.800	
UZFWC	57.000	.000	.000	.010	
LZWTC	195.000	160.000	137.400	126.500	
LZFSC	74.650	4.600	3.824	20.230	
LZFPC	501.000	165.000	187.400	127.900	
ADIMC	350.000	386.000	438.800	348.400	
UZWTC	150.000	150.000	150.000	150.000	150.000
UZFWC	100.000	100.000	100.000	100.000	100.000
LZWTC	300.000	300.000	300.000	300.000	300.000
LZFSC	200.000	200.000	200.000	200.000	200.000
LZFPC	800.000	800.000	800.000	800.000	800.000
PXADJ	1.000	1.000	1.000	1.000	1.000
PEADJ	.700	.700	.700	.700	.700

Figure 1.8.22 (continuation)

is the summary SAC model parameters for each year and for each subwatershed. Figures 1.8.23 through 1.8.34 show the observed and computed streamflow for each year and for each subwatershed.

It is seen from these figures that some years in some subwatersheds are not fairly satisfactory. Generally, the fit between observed and computed streamflows worsens in going from La Estrechura to Paso del Ermitano (i.e. upstream to downstream). This can be expected since any calibration inadequacy at the upstream flow points are carried to the downstream flow points. It appears that the best fit is in year 1972.

One major problem encountered in the model calibration is the rainfall data where inconsistencies with respect to streamflow are found. For example, the worst fit between observed and computed streamflow experienced in La Estrechura in 1974 can well be attributed to the rainfall data as seen in the plots of rainfall with streamflow in Figure 1.8.12. Admittedly, this is difficult to resolve since the inconsistency can be caused by inadequacies in areal averaging, representativeness of rainfall stations recording at this time, or sampling errors. It may be noted that rainfall is the most critical input data in the model as shown in Figure 1.8.A. , Appendix 1.8.A.2.

1.8.5 Model Testing by Forecasting

This section exemplifies the application of the SACKW model in a forecasting mode. The interest here is to arrive at the streamflow estimate of Paso del Ermitano during the hurricanes David and Frederic that hit the Dominican Republic around late August, 1979 and early September, 1979, respectively. For purposes of this exercise, the model parameters calibrated for year 1972 are used since it is found to have the best fit as well as this year experienced a high flood flow regime.

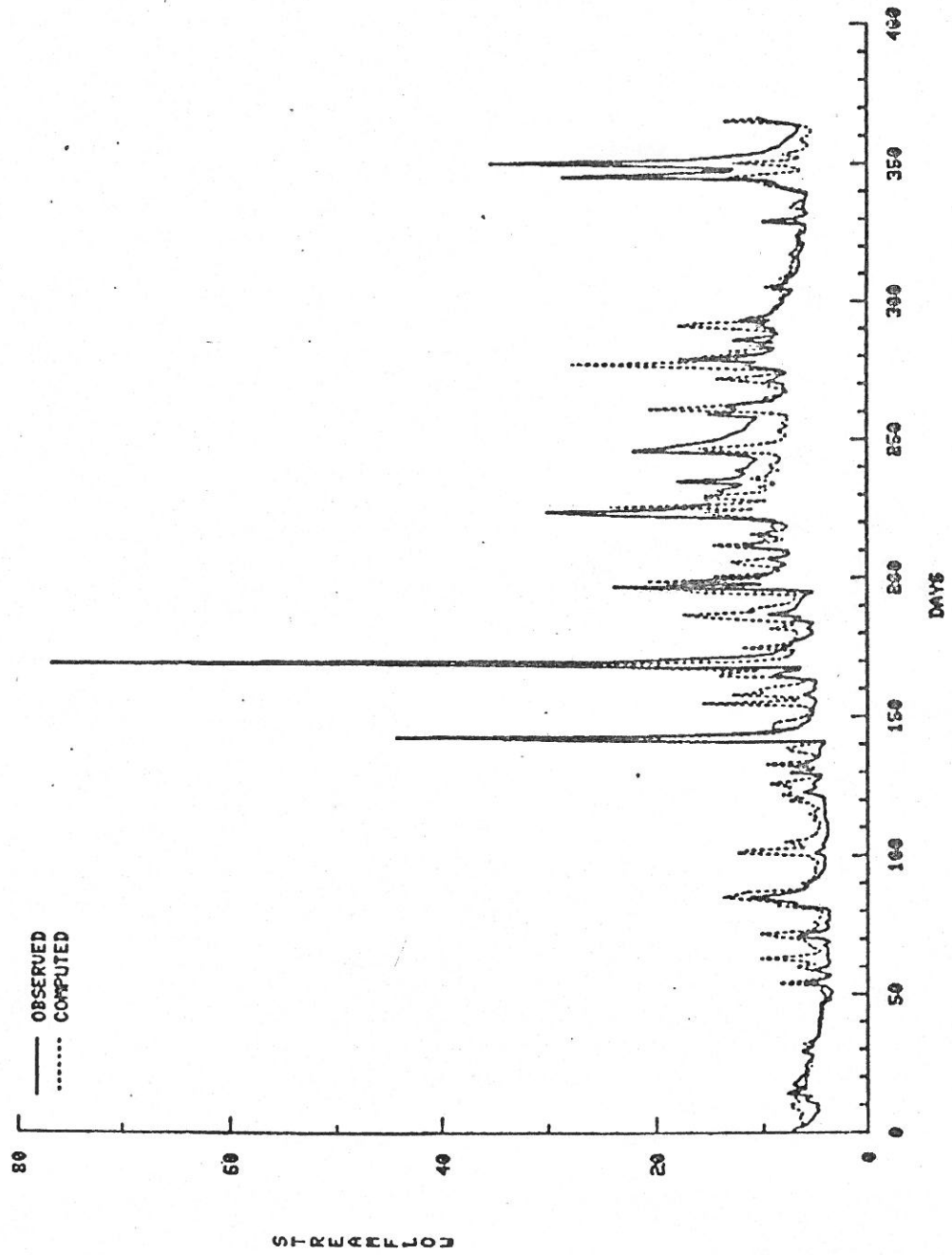


Figure 1.8.23. LA ESTRECHURA - 1972

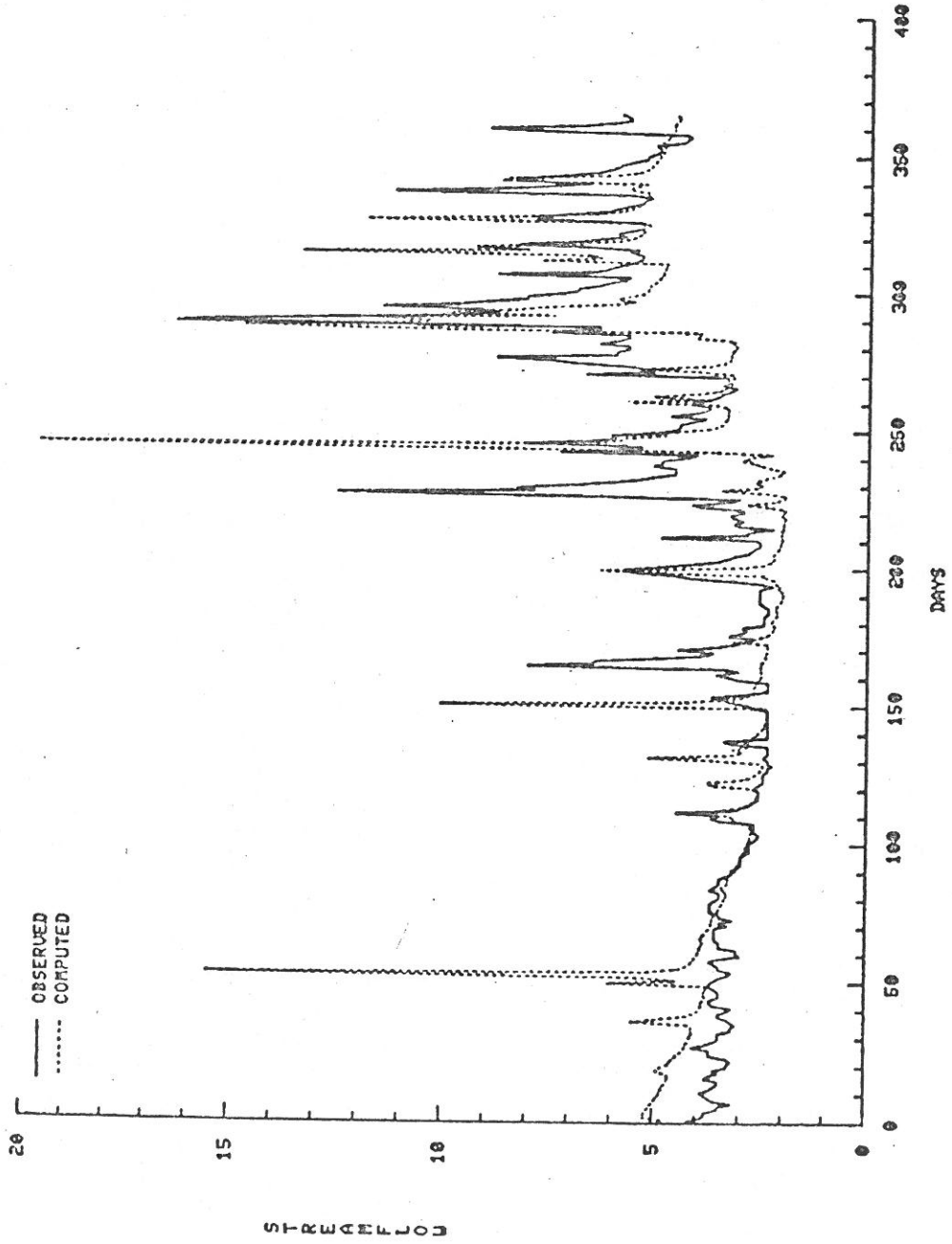


Figure 1.8.24. LA ESTRECHURA - 1973

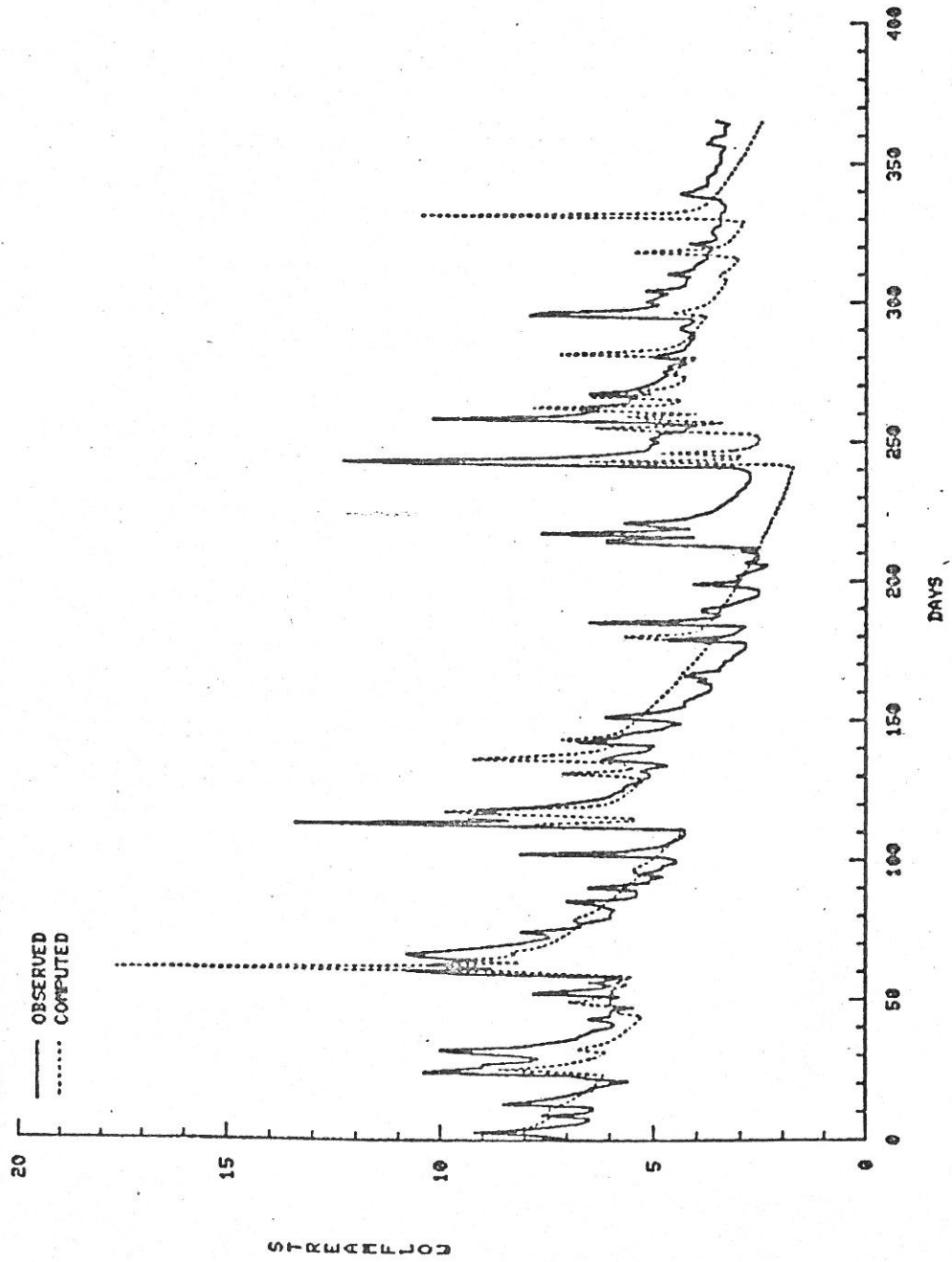


Figure 1.8.25. LA ESTRECHURA - 1974

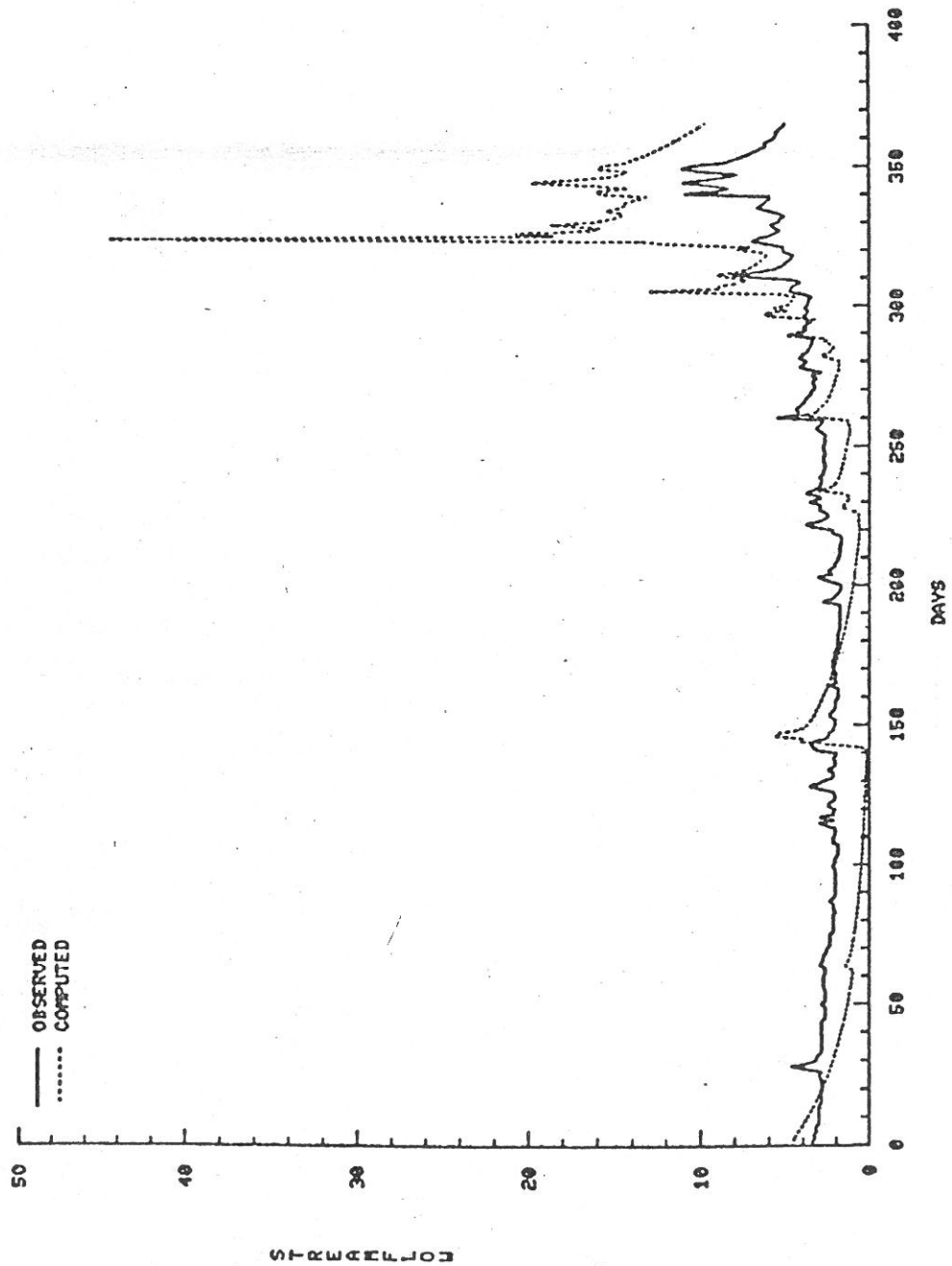


Figure 1.8.26. LA ESTRECHURA - 1975

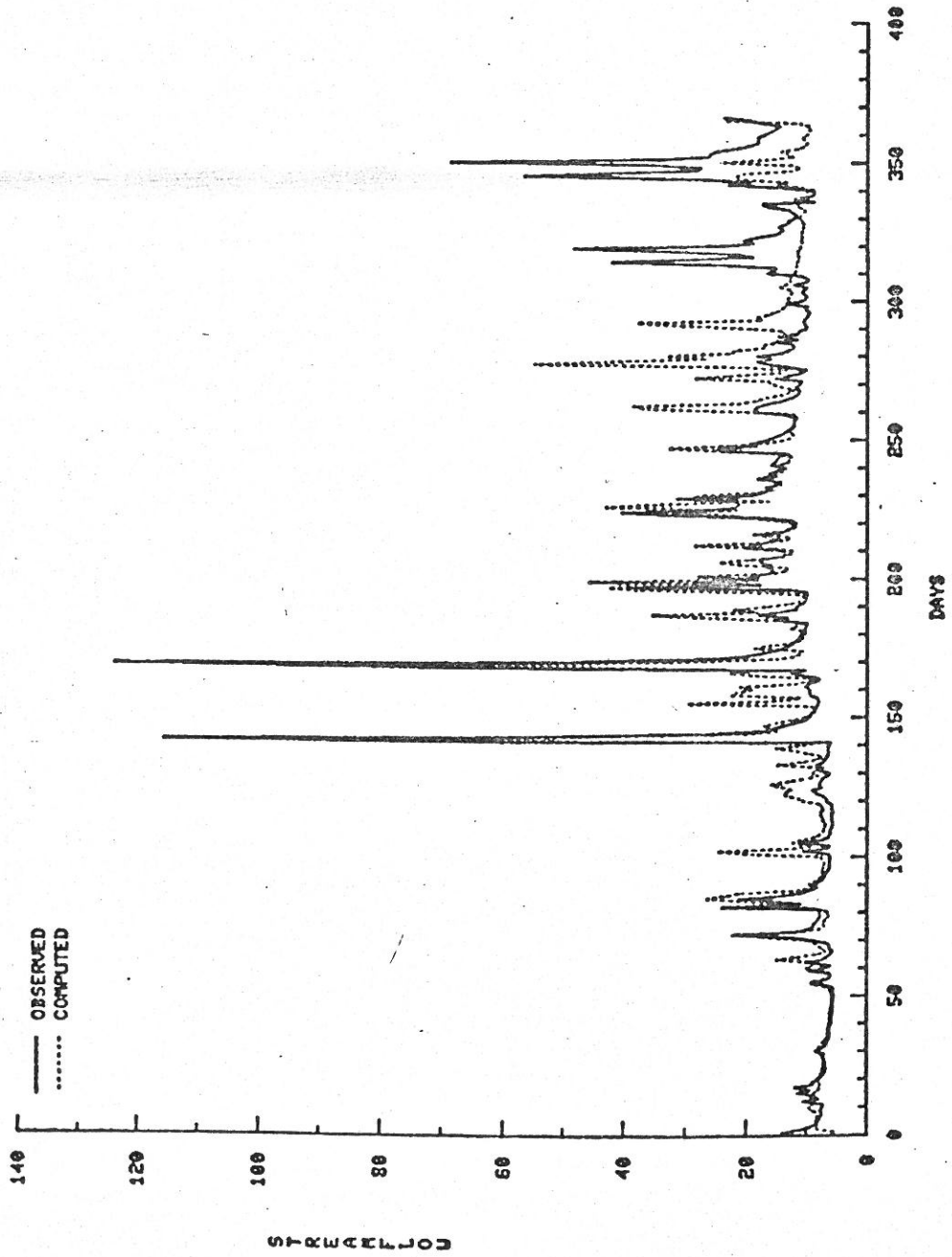


Figure 1.8.27. PALO DE CAJA - 1972

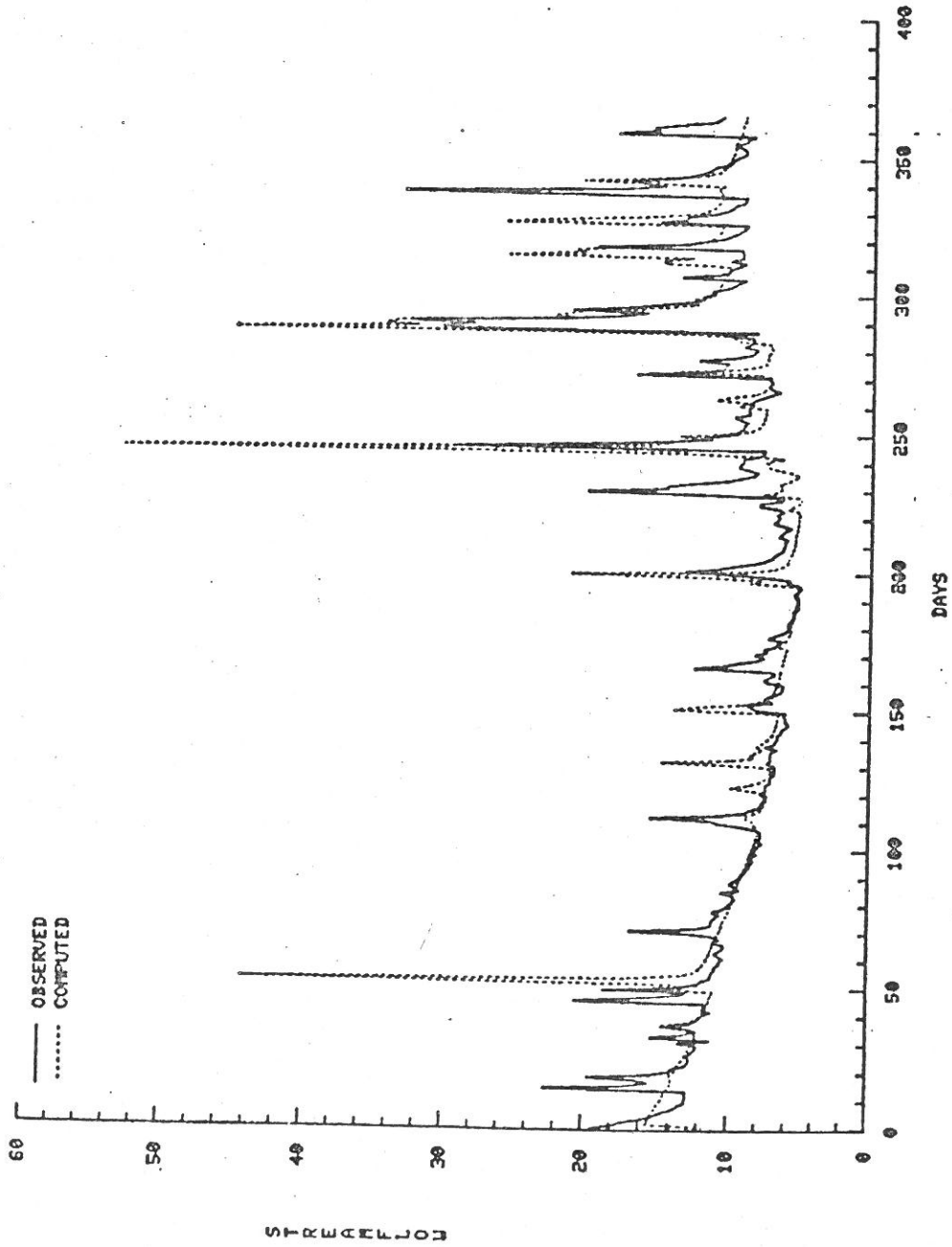


Figure 1.8.28. PALO DE CAJA - 1973

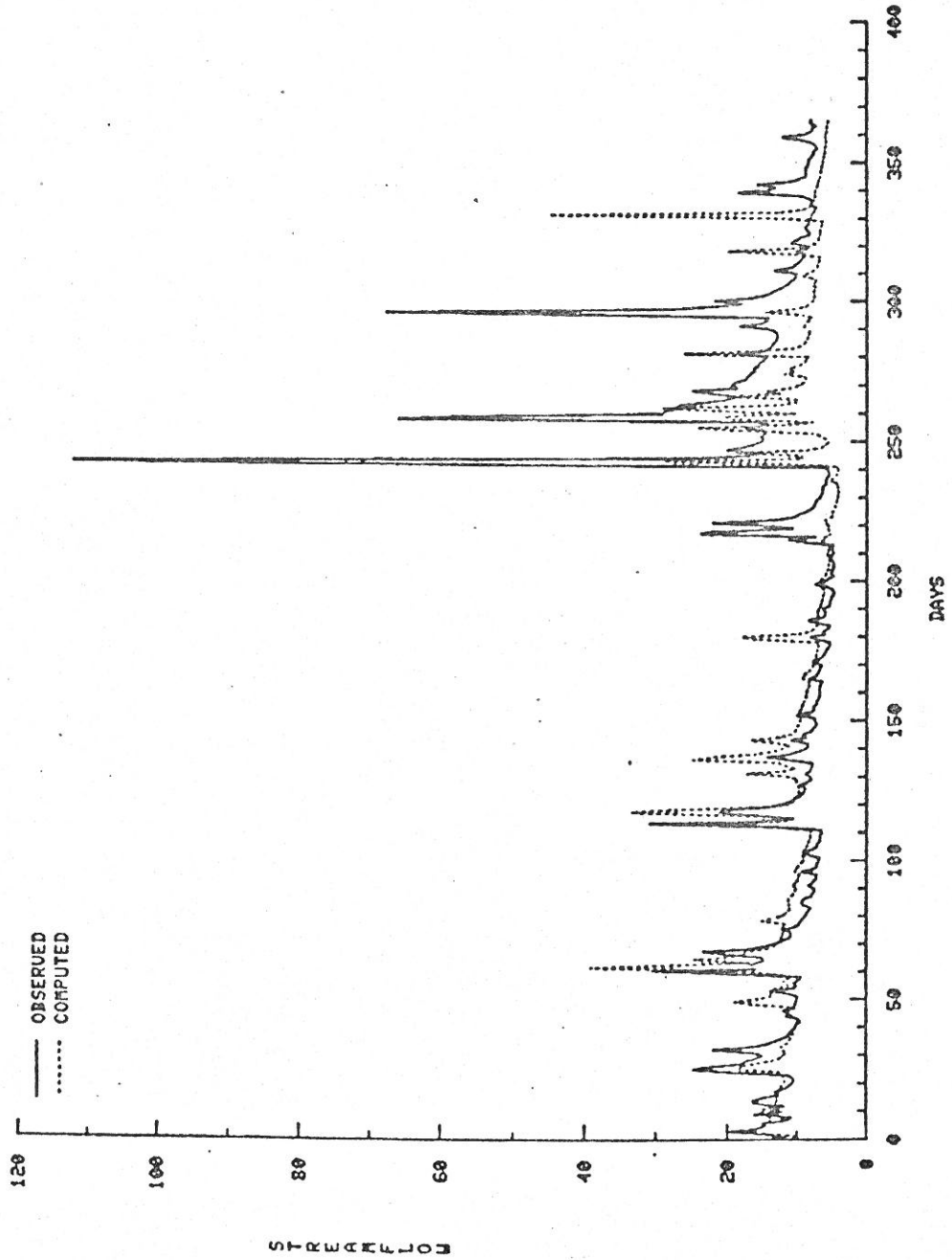


Figure 1.8.29. PALO DE CAJA - 1974

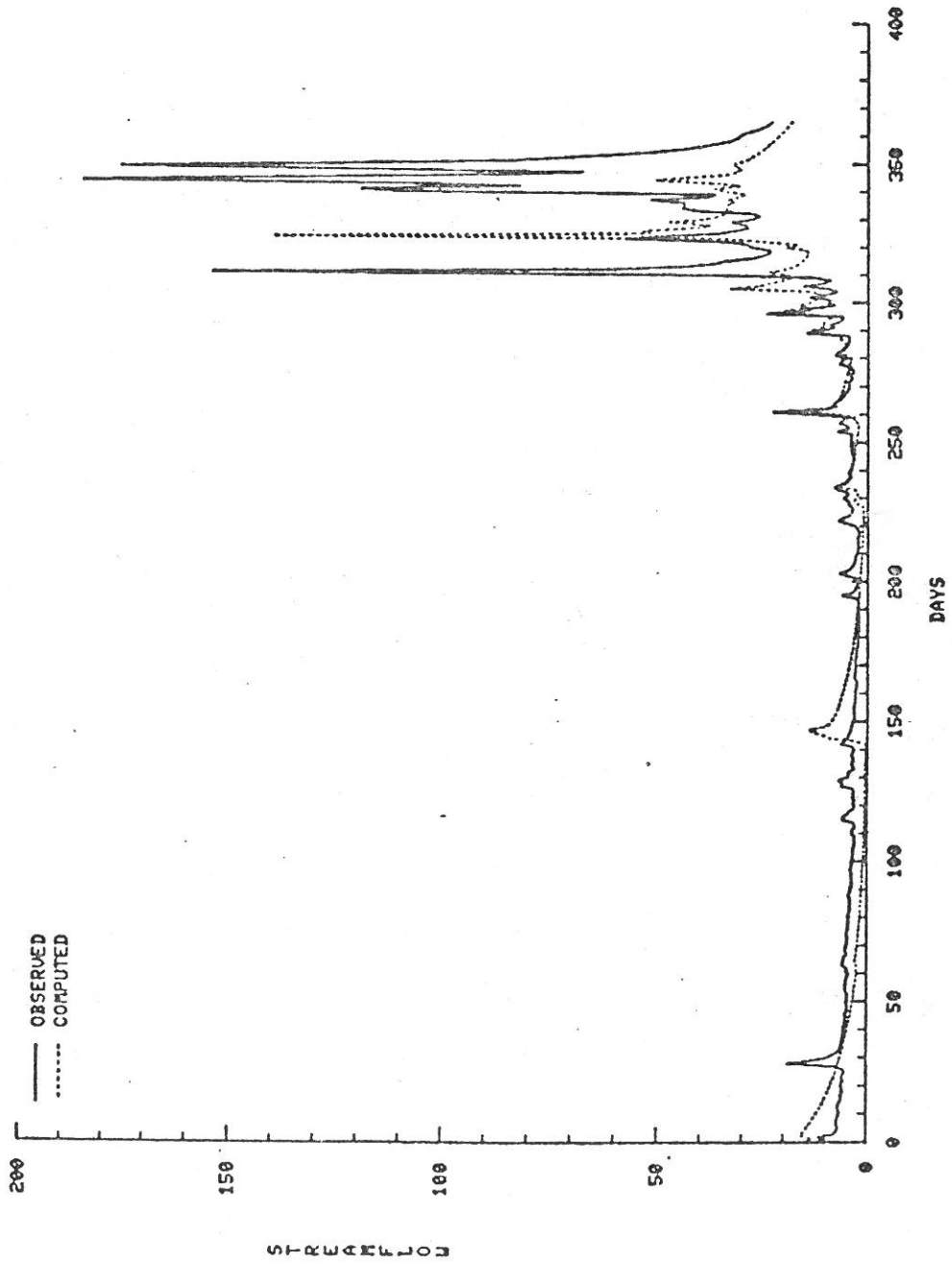


Figure 1.8.30. PALO DE CAJA - 1975

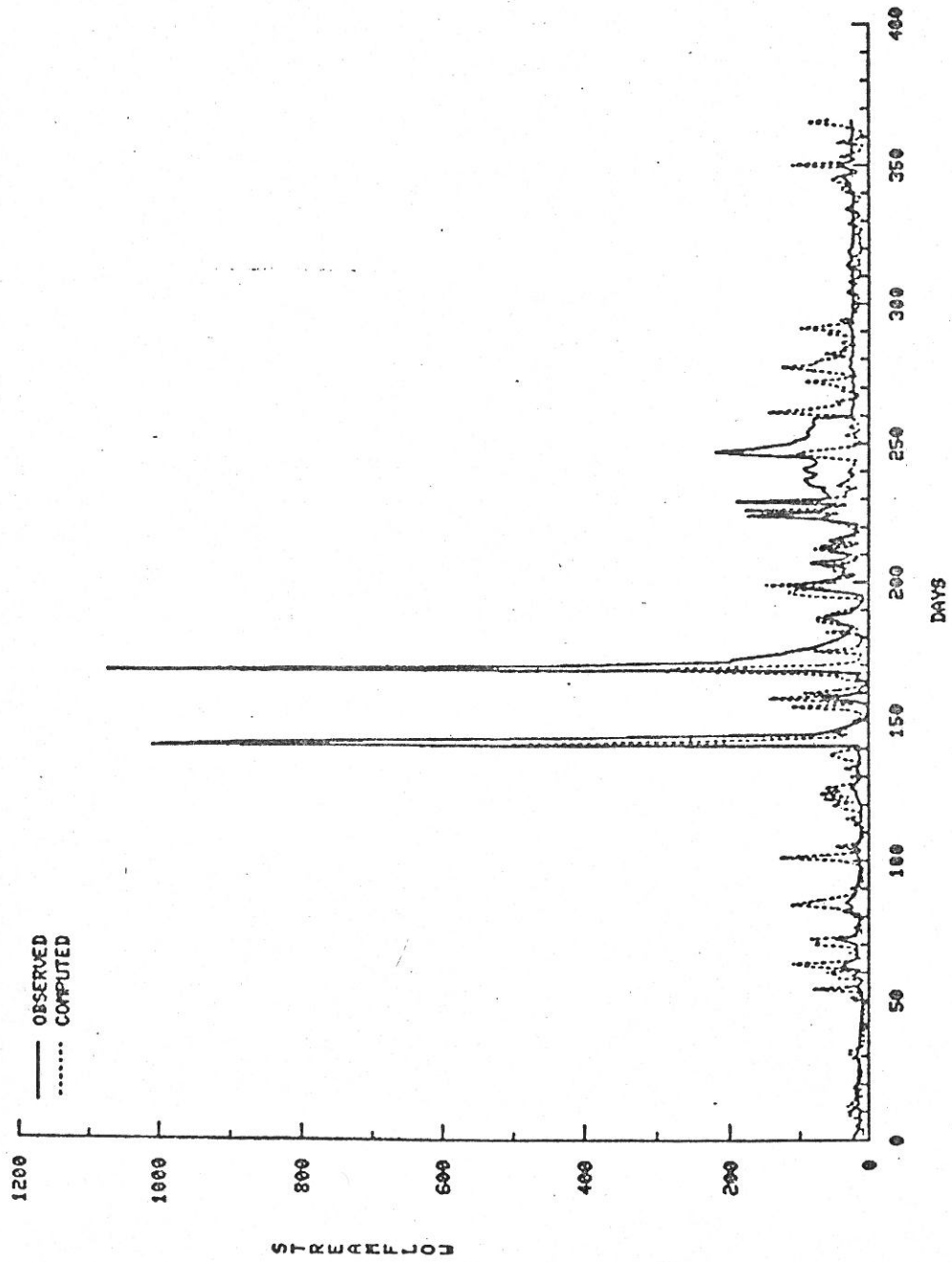


Figure 1.8.31. PASO DEL ERMITANO - 1972

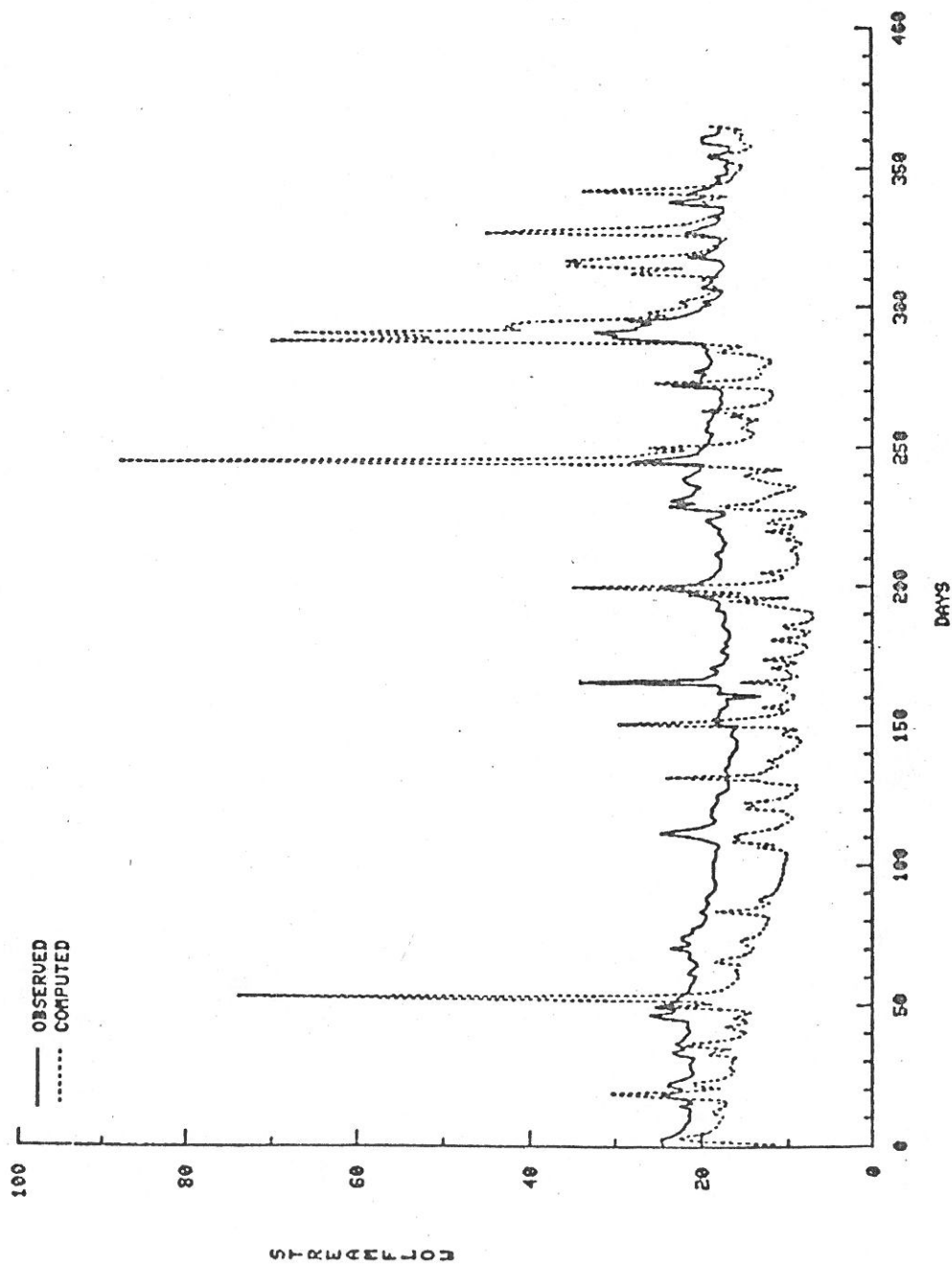


Figure 1.8.32. PASO DEL ERMITANO - 1973

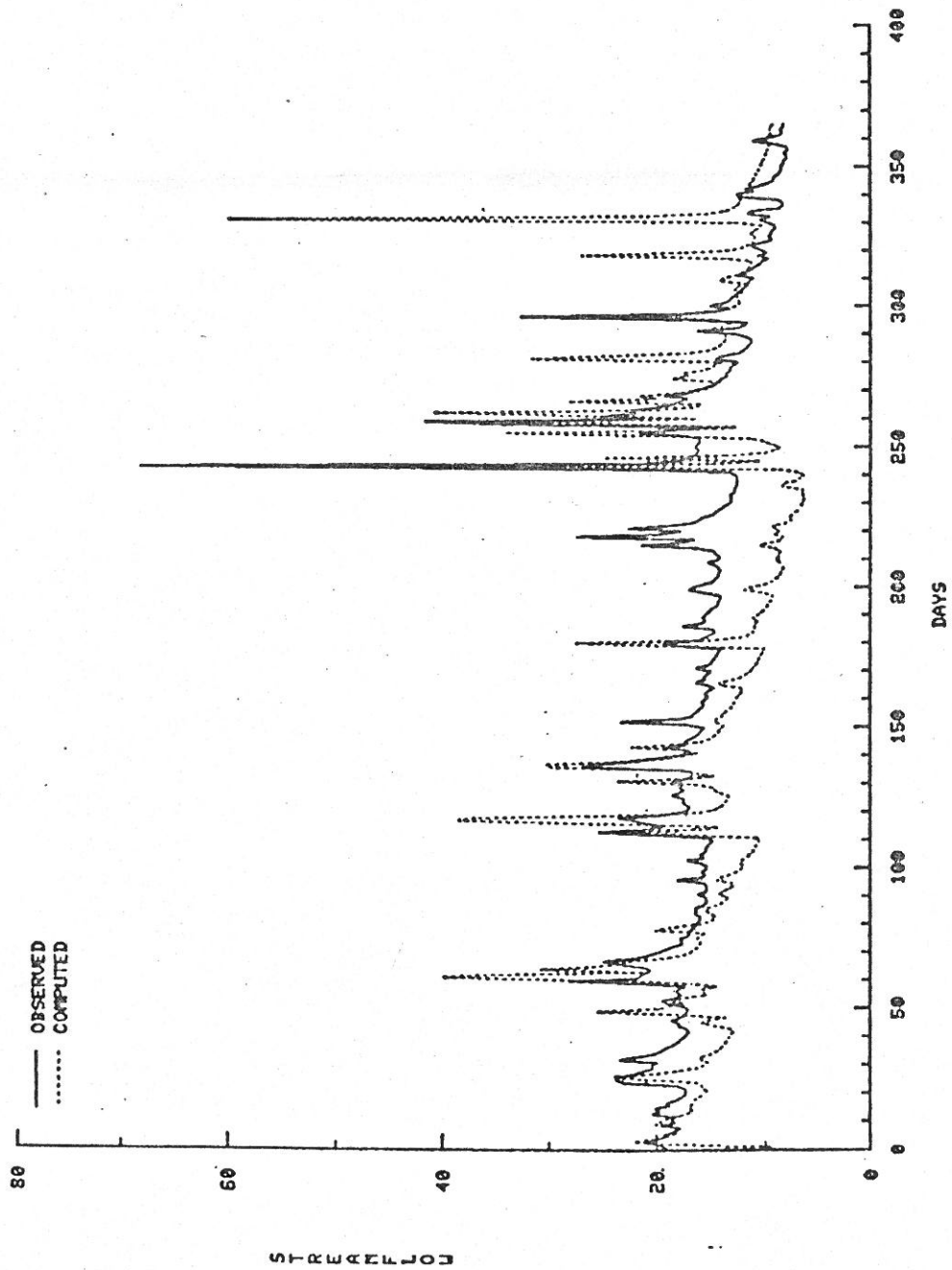


Figure 1.8.33. PASO DEL ERMITANO - 1974

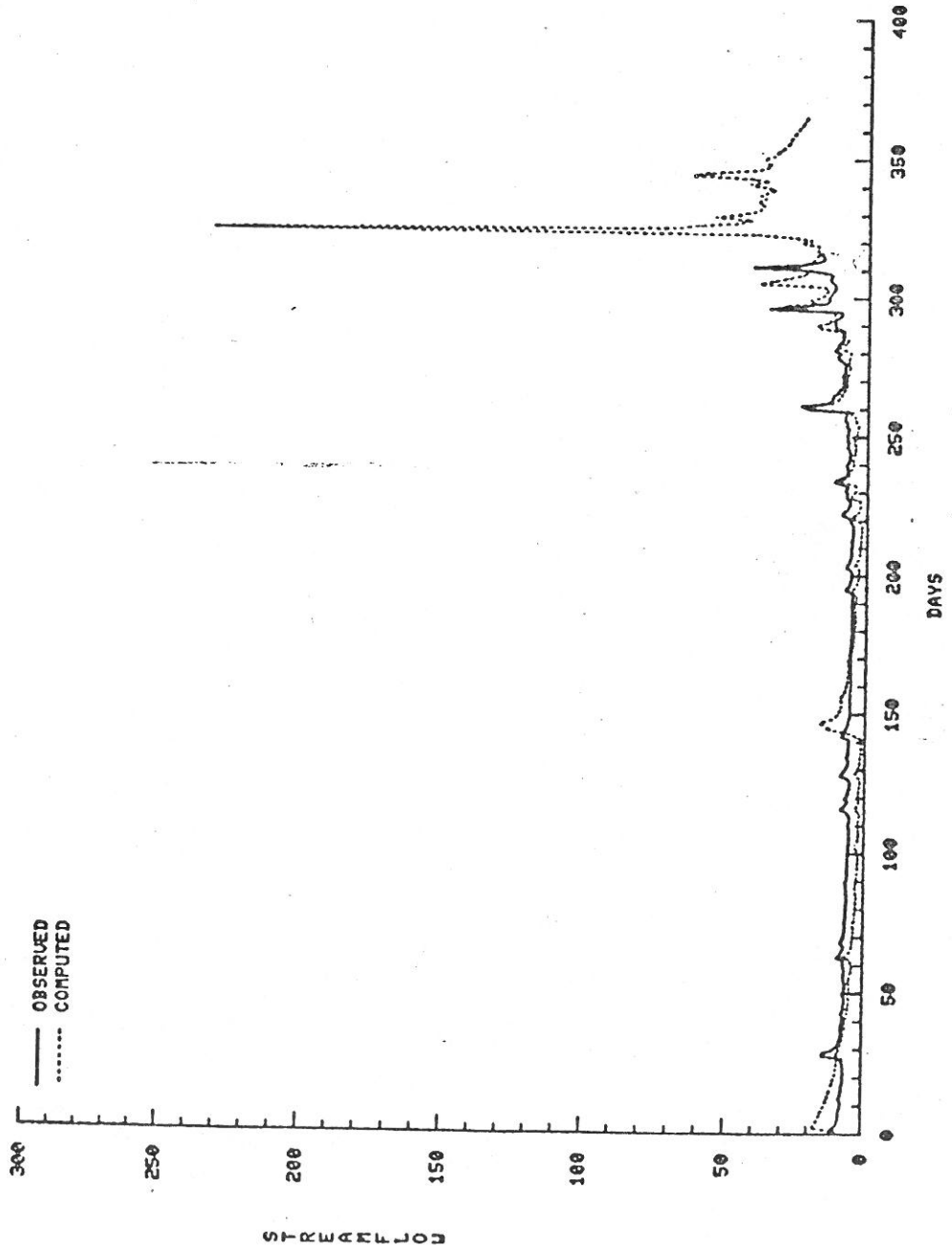


Figure 1.8.34. PASO DEL ERMITANO - 1975

which occurred around the month of June. Note that there is no streamflow record for the hurricane period of interest. However, hourly rainfall data is available.

The input data file for the model is shown in Figure 1.8.35. To obtain the forecast streamflow during the hurricane period of interest, the model has to be ran from year 1976 up to 1979. Results from this run are shown in Figures 1.8.36 and 1.8.37 assuming that there is an evapotranspiration (ET) demand and without ET demand, respectively. For hurricane David, the highest streamflows estimated are 6983 cms for the case where there is ET demand and 7074 cms for the case with no ET demand. Both highest flows occurred on the 21st hour of August 31, 1979. These flows are rather lower than those obtained using the HEC-1 model but they can be considered and admissibly within the streamflow regime in that hurricane David period.

1.8.6 References

- Burnash, R.J.C., 1985. Real-time forecasting with the Sacramento watershed model. In Proc. 14th Annual Hydrology Day, Colorado State University, Ft. Collins, CO, pp. 103-113.
- Burnash, R.J.C., R.L. Ferral and R.A. McGuire, 1979. A generalized streamflow simulation system conceptual modeling for digital computers. National Weather Service, California Dept. of Water Resources, March, 1979 (second printing).
- Kuester, J.L. and J.H. Mize, 1973. Optimization techniques with FORTRAN. McGraw-Hill Book Co., New York.
- National Weather Service, 1984. NWS-river forecast system manual calibration version 3.0. National Weather Service, Silver Spring, Md.
- U.S. Army Corps of Engineers, 1985. HEC-1 flood hydrograph package user's manual. The Hydrologic Engineering Center, Dan's, CA, January, 1985 (revised edition).

```

BASIN A - LA ESTRECHURA (1976-1979) 1972 PARAMETERS
  0      0 8766 8766      1
ZPERC      87.00
REXP       0.709
SIDE       0.176
UZK        0.010
ADIMP      0.115
RSERV      0.35
RIVA       0.57
PCTIM      0.089
LZPK       0.010
LZSK       0.006
PFREE      0.793
UZTWM      150.0
UZFWM      100.0
LZTWM      300.0
LZFSM      200.0
LZFPM      800.0
UZTWC      120.4
UZFWC      0.005
LZTWC      297.27
LZFSC      41.06
LZFPC      185.43
ADIMC      416.5
PXADJ      1.0
PEADJ      0.7
END
ROUTE
70.0
2500.0     0.6      0.4      100.0
16000.     0.105    0.06    0.0
5.0        15.0     1.0     5.10     0.0
BASEF      2
0.3        0.5      0.2
ROUTE
70.0
2500.     0.6      0.4      100.
20000.    0.0837   0.06    70.0
5.0        15.0     1.0     5.10     0.0
BASEF      2
0.3        0.5      0.2
ADD        2
ROUTE
45.0
4500.0    0.5      0.3      100.0
8500.0    0.0133   0.040   45.0
5.0        85.0     1.0     5.10     1.0
BASEF      2
0.3        0.5      0.2
END
RAIN
  135064    7      1      1
(10X,F10.0)
ETDATA
  24      9      1      1

```

Figure 1.8.35 Model input file for model testing by forecasting hurricanes David and Frederic flood flow regime.

(10F8.0)

FLOW

1 0

END

BASIN B - PALO DE CAJA (1976-1979) 1972 PARAMETERS

0 0 8766 8766 1

ZPERC	178.0			
REXP	0.400			
SIDE	1.050			
UZK	0.030			
ADIMP	0.145			
RSERV	0.225			
RIVA	0.339			
PCTIM	0.015			
LZPK	0.003			
LZSK	0.002			
PFREE	0.987			
UZTWM	150.0			
UZFWM	100.0			
LZTWM	300.0			
LZF5M	200.0			
LZFPM	800.0			
UZTWC	120.4	40.0	150.0	2.0
UZFWC	.000	00.0	80.0	2.0
LZTWC	75.0	40.0	300.0	5.0
LZF5C	56.05	40.0	200.0	2.0
LZFPC	92.07	100.0	500.0	5.0
ADIMC	414.4	100.0	450.0	10.0
PXADJ	1.0			
PEADJ	0.7			

END

ROUTE

24.0

2000.0	0.4	0.3	100.0	
7500.0	0.108	0.040	0.0	
5.0	85.0	1.0	23.1	1.0

BASEF 2

0.3 0.5 0.2

ROUTE

56.0

3500.	0.2	0.3	100.	
13500.	0.0452	0.05	0.0	
5.0	10.0	1.0	23.1	0.0

BASEF 2

0.3 0.5 0.2

ADD 2

ROUTE

164.0

4500.0	0.4	0.3	100.0	
33500.0	0.0086	0.040	00.0	
5.0	70.0	1.0	23.1	1.0

BASEF 2

0.3 0.5 0.2

ROUTE

103.0

Figure 1.8.35 (continuation)

2000.	0.4	0.3	100.	
33000.	0.0293	0.05	00.0	
5.0	10.0	1.0	23.1	0.0
BASEF	2			
0.3	0.5	0.2		
ADD	2			
END				
RAIN				
135064	7	1	1	
(20X,F10.0)				
ETDATA				
24	9	1	1	
(10F8.0)				
FLOW				
1	0			
END				
BASIN C - PASO DEL ERMITANO (1976-1979)	1972	PARAMETERS		
0	0 8766 8766	1		
ZPERC	33.70			
REXP	4.83			
SIDE	0.006			
UZK	0.900			
ADIMP	0.092			
RSERV	0.440			
RIVA	0.829			
PCTIM	0.908			
LZPK	0.012			
LZSK	0.476			
PFREE	0.994			
UZWWM	150.0			
UZFWM	100.0			
LZTWM	300.0			
LZFWM	200.0			
LZFPM	800.0			
UZTWC	84.8	50.0	150.0	2.0
UZFWC	0.01	00.0	80.0	2.0
LZTWC	126.5	40.0	200.0	5.0
LZFSC	20.23	60.0	150.0	2.0
LZFPC	127.9	0.0	800.0	5.0
ADIMC	348.4	100.0	400.0	10.0
PXADJ	1.0			
PEADJ	0.7			
END				
ROUTE				
103.0				
5000.0	0.5	0.3	100.0	
18000.	0.0148	0.04	0.0	
5.0	80.0	1.0	29.18	1.0
BASEF	2			
0.3	0.5	0.2		
ROUTE				
109.0				
3000.	0.4	0.3	100.	
34000.	0.0323	0.05	00.0	
5.0	10.0	1.0	29.18	0.0

Figure 1.8.35 (continuation)

BASEF	2				
0.3	0.5	0.2			
ADD	2				
ROUTE					
53.0					
5000.0	0.4	0.3	100.0		
18000.	0.0119	0.040	00.0		
5.0	80.0	1.0	29.18	1.0	
BASEF	2				
0.3	0.5	0.2			
END					
RAIN					
135064	7	1	1		
(30X, F10.0)					
ETDATA					
24	9	1	1		
(10F8.0)					
FLOW					
1	0				

Figure 1.8.35 (continuation)

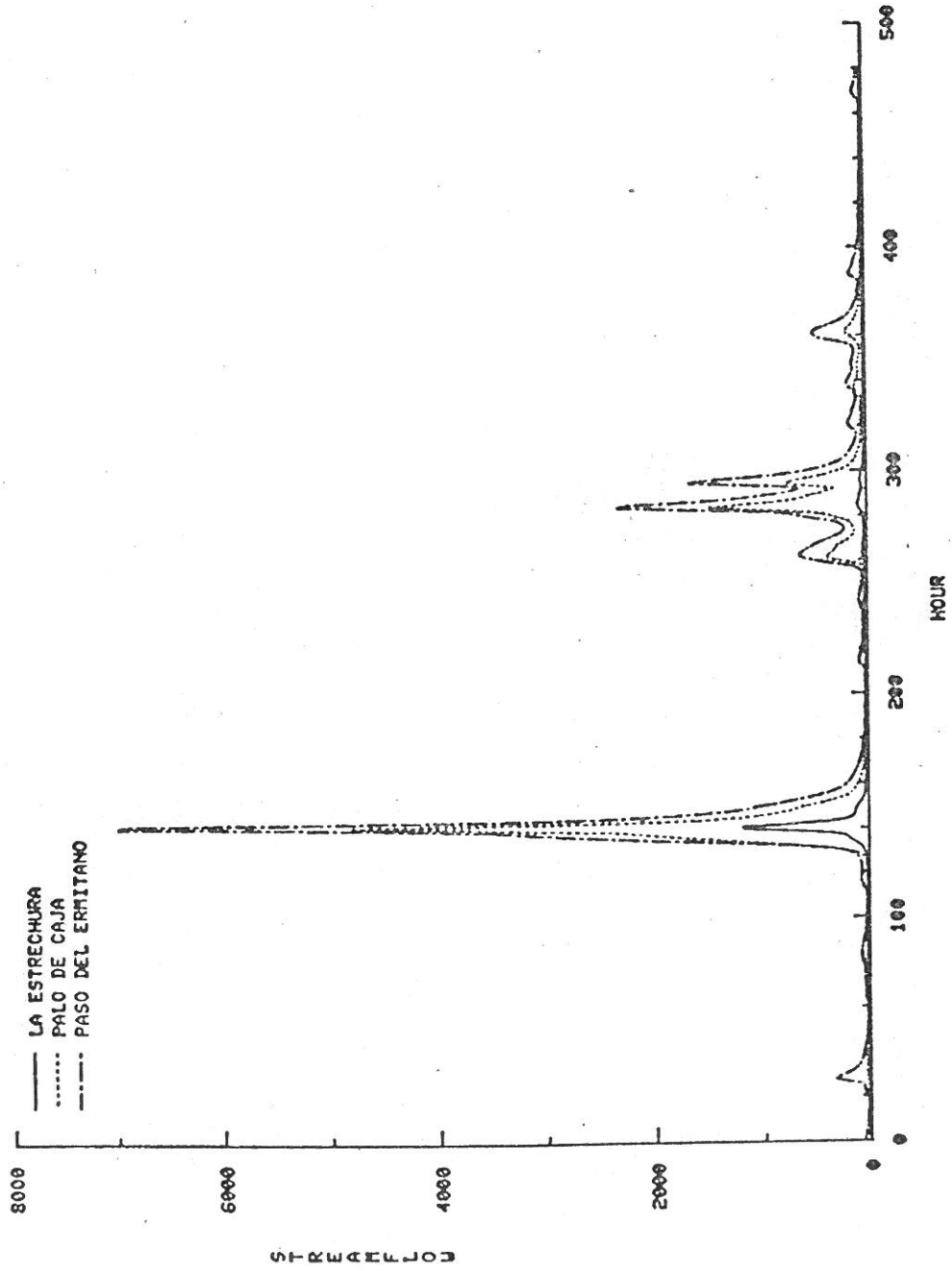


Figure 1.8.36. Results of model run with evapotranspiration demand for hurricanes David and Frederic (Aug 26-Sep 14, 1979).

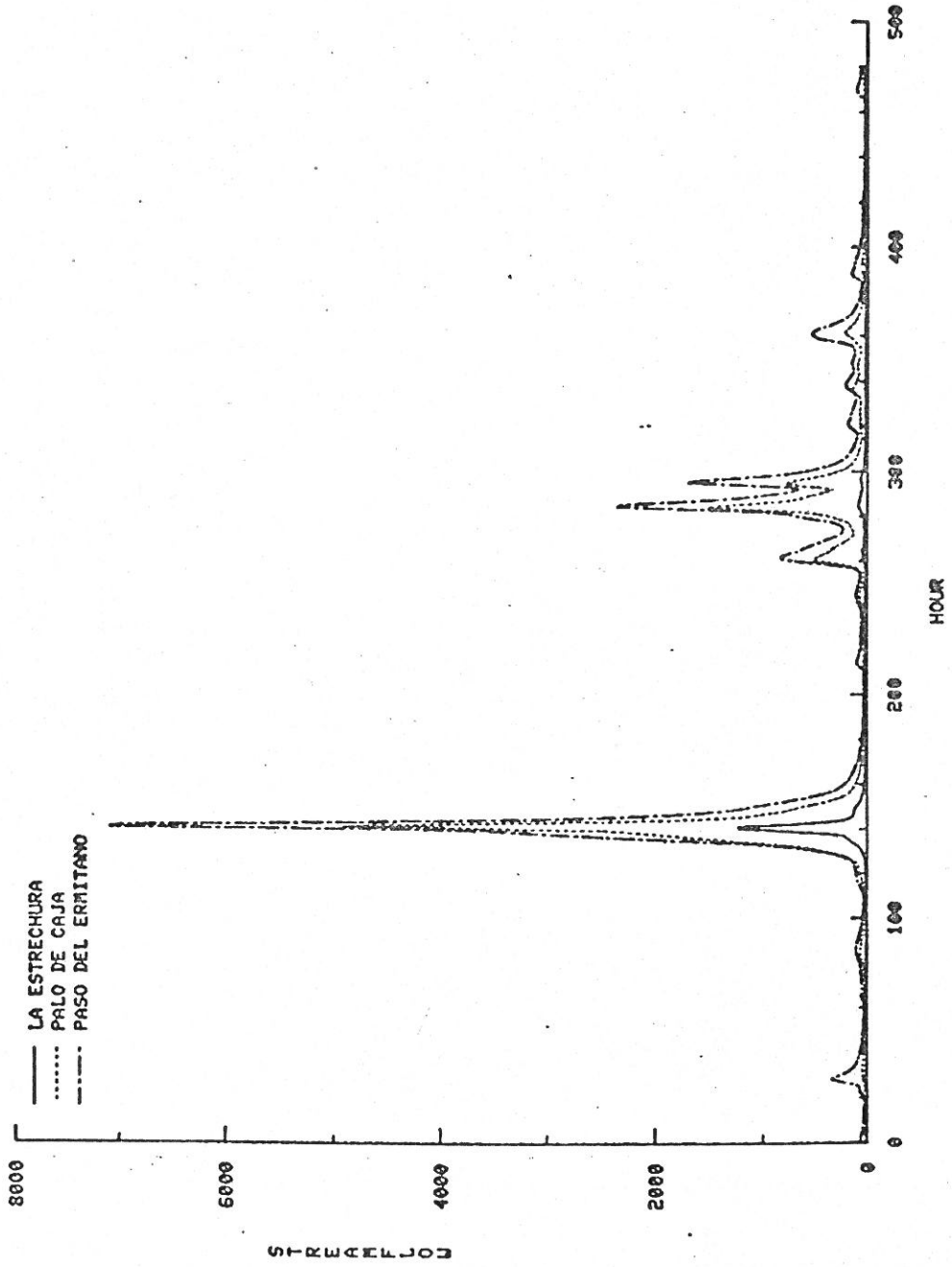


Figure 1.8.37. Results of model run without evapotranspiration demand for hurricanes David and Frederic (Aug 26-Sep 14, 1979).

APPENDIX 1.8.A

SACKW MODEL USER'S MANUAL

1.8.A.1 Introduction

The program SACKW is a conceptual-hydraulic model of watershed which simulates the various elements of the hydrologic cycle. Basically, the model starts with simulating the basin hydrology through the conceptual based Sacramento soil moisture accounting (SAC) model to derive the different runoff components in the basin accruing from input precipitation. Then, the pertinent runoff components are hydraulically routed through overland-flow planes and channels to arrive at the streamflows at the basin outlet using the kinematic wave (KW) routing methodology. The model has two operational modes: the calibration mode for model parameter estimation, and the forecasting mode for simulating basin hydrology under some specified or known set of model parameters. In the calibration mode, the parameters of the SAC model may be calibrated manually or automatically using the constrained Rosenbrock optimization technique. The time scale of model simulation is at the least on an hourly basis and can be at longer time intervals but as multiples of one hour.

The ensuing sections present the description of Program SACKW, input requirements, output information, some guidelines for model usage and parameter calibration, and sample model application.

1.8.A.2 Program Description

Program SACKW is written in FORTRAN 77 which can be ran in mainframe computers or desktop computers. The program listing is given in Appendix 1.8.B. It is composed of 14 subprograms and a main program.

A descriptive flowchart of the program operation and sequence is given in Figure 1.8.A.1. Included in the flowchart are the pertinent program and subprograms used in the various operations. Given below are brief descriptions of the main program and subprograms.

MAIN PROGRAM

The MAIN program reads all input data and controls the overall sequence of program operations.

Subroutine SETPAR

This subprogram checks and sets up default values of model parameters for the Sacramento soil moisture accounting model.

Subroutine SACROUT

Subroutine SACROUT resets model parameters, sequences the soil moisture accounting and routing operations, and output some execution results.

Subroutine SACSMA

Subroutine SACSMA specifically controls the time loop of the Sacramento soil moisture accounting, performs water balance and prints results of soil moisture accounting.

Subroutine SMAONE

This subroutine performs all soil moisture accounting computations for one time step.

Subroutine KINWAVE

This subroutine controls the timing and sequencing of kinematic wave routing operations. Additionally, the routed streamflows and computed baseflows from different flow planes and subbasins are combined in this subroutine.

Subroutine KINOFF

This subroutine determines the runoff hydrograph for each flow plane using the kinematic wave method.

Subroutine FDKRUT

Subroutine FDKRUT generates overland flow runoff hydrograph or stream discharge hydrographs.

Subroutine ROFGRD

Subroutine ROFGRD computes the incremental length and time required in finite-difference grid computations for overland flow.

Subroutine FLOGRO

This subroutine computes the size and number incremental lengths required in finite-difference solution in the stream discharge routing computation.

Subroutine FRMMTC

Converts input data from metric units to English units in the kinematic wave routing computations.

Subroutine TOMTRC

Converts computed results in English units to metric system in the kinematic wave routing computation.

Subroutine OPTIM

Subroutine OPTIM performs the constrained Rosenrock optimization algorithm for the SAC model parameters.

Function FPOBS

This subprogram computes the value of the objective function for the optimization routine based on observed and computed streamflows.

Subroutine UPDATE

This subprogram generates an updated model input data file if derived to incorporate adjusted model parameters when optimized and sending volumes of soil-moisture contents for eventual use in the future.

1.8.A.3 Input and Output Information

Generally, the program input can be summarized in the following sequence:

1. Model run information
2. Control parameters for input, output and optimization options
3. Soil-moisture accounting model parameters
4. Kinematic wave model parameters
5. Hydrologic input data control parameters

A detailed description and sequence of the program input is given in Appendix 1.8.C. To facilitate in model inputting some data or parameter input (data sets) require a five-letter word identifier as a leading record or contained in the beginning of an input record. For further clarification on the input requirements, Section 1.8.A.5. presents a sample program application.

The program output is primarily in the form of tabular and graphical displays on a line printer. The tabular outputs include summaries of soil-moisture contents, runoff components, evapotranspiration and rainfall. Tabular and graphical outputs of observed and simulated streamflows at subbasin outlets are also given. There are printing frequency control options provided in the program. A sample program output is given in Section 1.8.A.5.

1.8.A.4. Some Guidelines For Model Usage and Parameter Calibration

For details of setting-up the parameter values of the SAC model and kinematic wave routing model, the manuals prepared by Burnash, et al., 1979 and the U.S. Corps of Engineers (1985), respectively can be consulted. The ensuing text presents only some guidelines on model parameter calibration with emphasis on the SAC model parameters.

For the SAC model, Burnash (1985) have shown that from several tests conducted on the sensitivity of the model, the rainfall input data practically accounts for all variations in the computed streamflows as opposed to the rest of the model parameters. This result is shown in Figure 1.8.A.2. in which a particular runoff hydrograph is ten times as sensitive to a shift in the rainfall input as it is to a similar change in the most sensitive parameters. In view of this, an important aspect in using the model is to resolve the question of handling the rainfall data. As done in the model calibration for Nizao basin, the rainfall input data has been defined for each subwatershed based on areally averaging point rainfall time series data from several stations. An areal averaging technique such as Thiessen method or optimal interpolation technique can be used for this purpose.

As mentioned earlier, the initial model parameters of SAC model may be obtained from guidelines given by Burnash, et al., (1979). Once this is set-up, some parameters may be refined by manual calibration or automatically through the optimization algorithm. Generally, the SAC model parameters to be calibrated are: UZK, REXP, ZPERC, SIDE, UZTWM, UZFWM, LZTWM, LZFSM and LZFPM. In the case of manually calibrating the model parameters, the following guidelines may be useful.

1. If surface runoff is excessive and baseflow is too low, the percolation could be inadequate. A possible action is to increase LZFSM and LZFPF which increases percolation and potential baseflow.

2. If initial runoff is inadequate, decreasing UZTWM allows runoff to take place sooner.

3. If surface runoff is generally excessive, the following action may be taken: i) raise ZPERC and thus increase percolation, ii) enlarge LZFSM and LZFPF which results in higher potential baseflow, and iii) lower REXP to increase continuing percolation and also alter shape of percolation curve.

4. If streamflow rising limbs are underforecast, too much water may be required to fill LZTWM. Action: reduce LZTWM by as much as water balance residual.

5. If streamflow rising limbs are overforecast but recession limbs underforecast, interflow could be inadequate which can be corrected by increasing UZFWM.

6. If the streamflow hydrograph baseflow time is too wide and rising peaks are flat the impervious area parameter ADIMP may be small. Decreasing ADIMP sharpens rising peaks and diminishes baseflow area.

7. To increase the general level of surface runoff and slightly decrease the trailing baseflow, ZPERC may be reduced.

For the kinematic wave model, most of the parameters can be obtained from basin topographic maps, such as basin areas and overland-flow and channel slopes, widths and lengths. The channel geometry may be obtained from actual photographs with scales or river cross-sectional data. The parameter that may require some calibration is the roughness

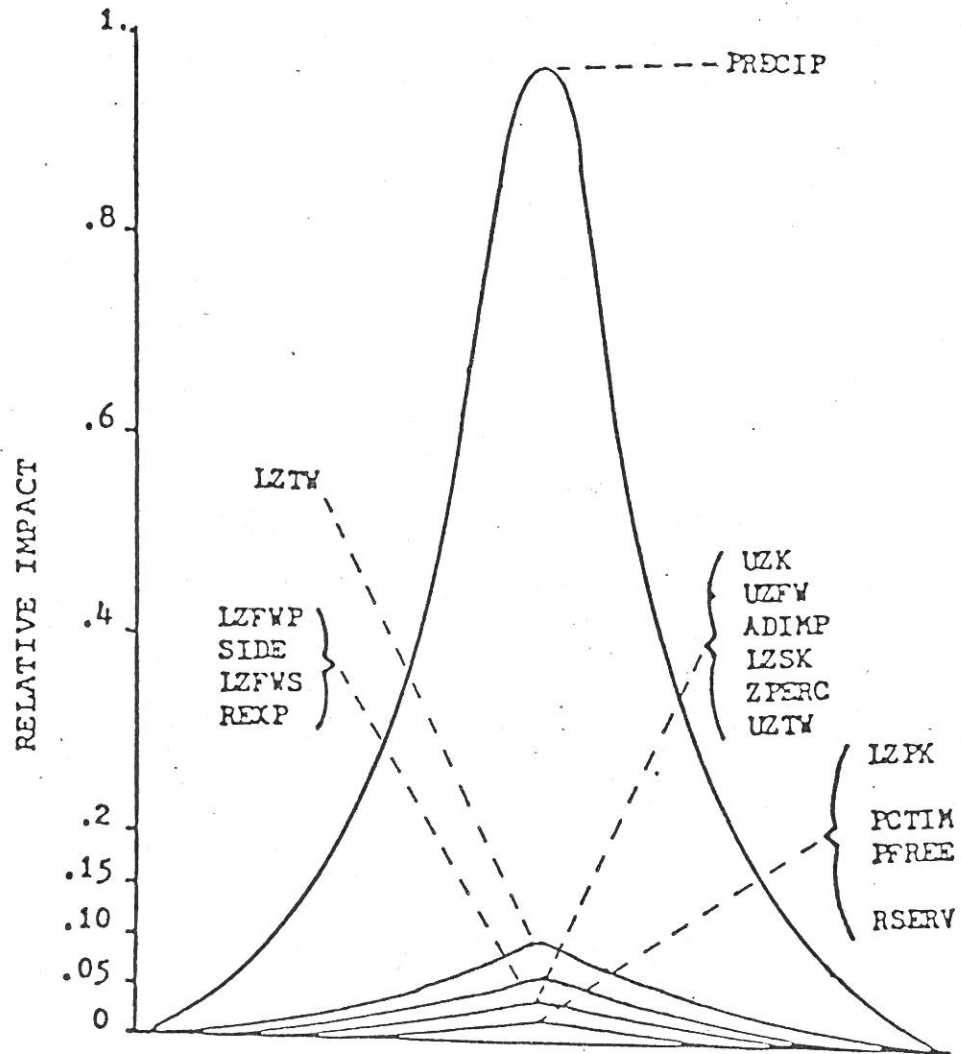
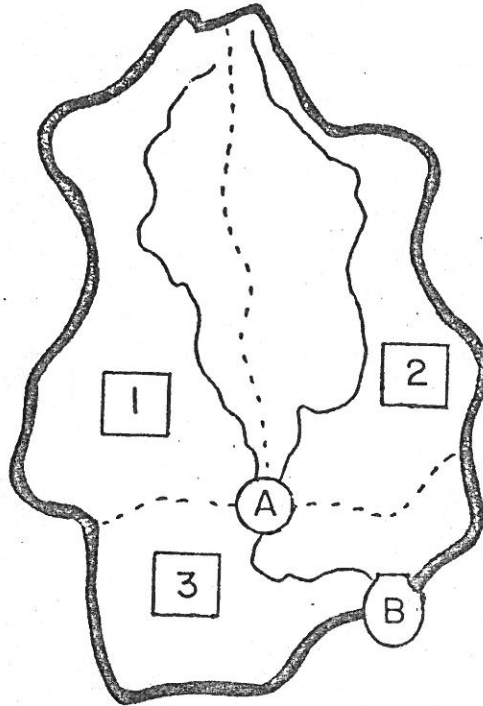
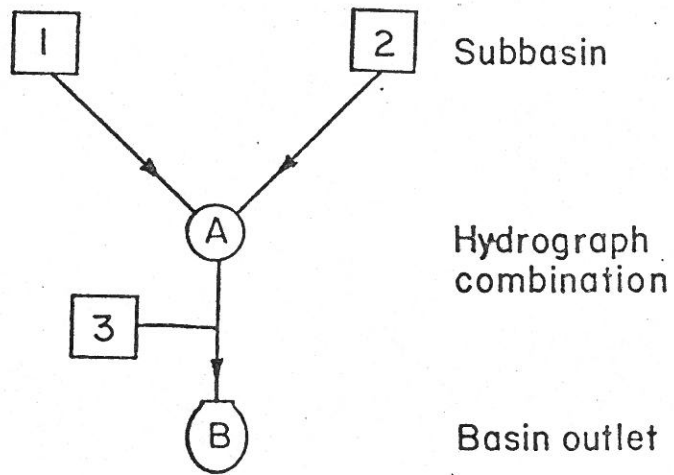


Figure 1.8.A.2. Incremental effect of 10% changes in basic input or parameter values as evaluated with the Sacramento model.



Example River Basin



Example River Basin Schematic

Figure 1.8.A.3. Sample river basin and schematic representation.

coefficient since flow-plane or channel heterogeneity effects may be difficult to fully parameterize into some lumped or average values.

1.8.A.5. Sample Model Application

This section presents a sample model application for illustration purposes. One watershed is used in the example with three overland-flow elements as shown in Figure 1.8.A.3. The input hydrologic data are rainfall, streamflow and evapotranspiration demand which are on an hourly basis. Given in Figure 1.8.A.4. is the input data file for the model. In this case, the hydrologic data are read as part of the overall model input file.

It is worthwhile to mention the manner in which the kinematic wave routing parameters are inputted with respect to the river basin configuration. Referring to Figures 1.8.A.3. and 1.8.A.4., the overland-flow element 1 runoff hydrograph is computed first which corresponds to the first "ROUTE" operation in the data file. Then the hydrograph of overland-flow element 2 is computed in the second "ROUTE" operation. In both cases, the "ROUTE" operation are followed by "BASEF" operations so that baseflow components are added at their outlet. At point A, the two hydrographs are combined using the "ADD" operation. The hydrograph for overland-flow element 3 is computed in the third "ROUTE" operation plus the contribution of baseflow upon issuing the last "BASEF" operation. Note that in this third "ROUTE" operation, the variable ARUPF(.) is set equal to 1.0 which indicates that the upstream hydrograph (at point A) is also routed together with the flows in subbasin 3.

The program output for this run is given in Figure 1.8.A.5. In this output all soil-moisture accounting results are printed on an hourly

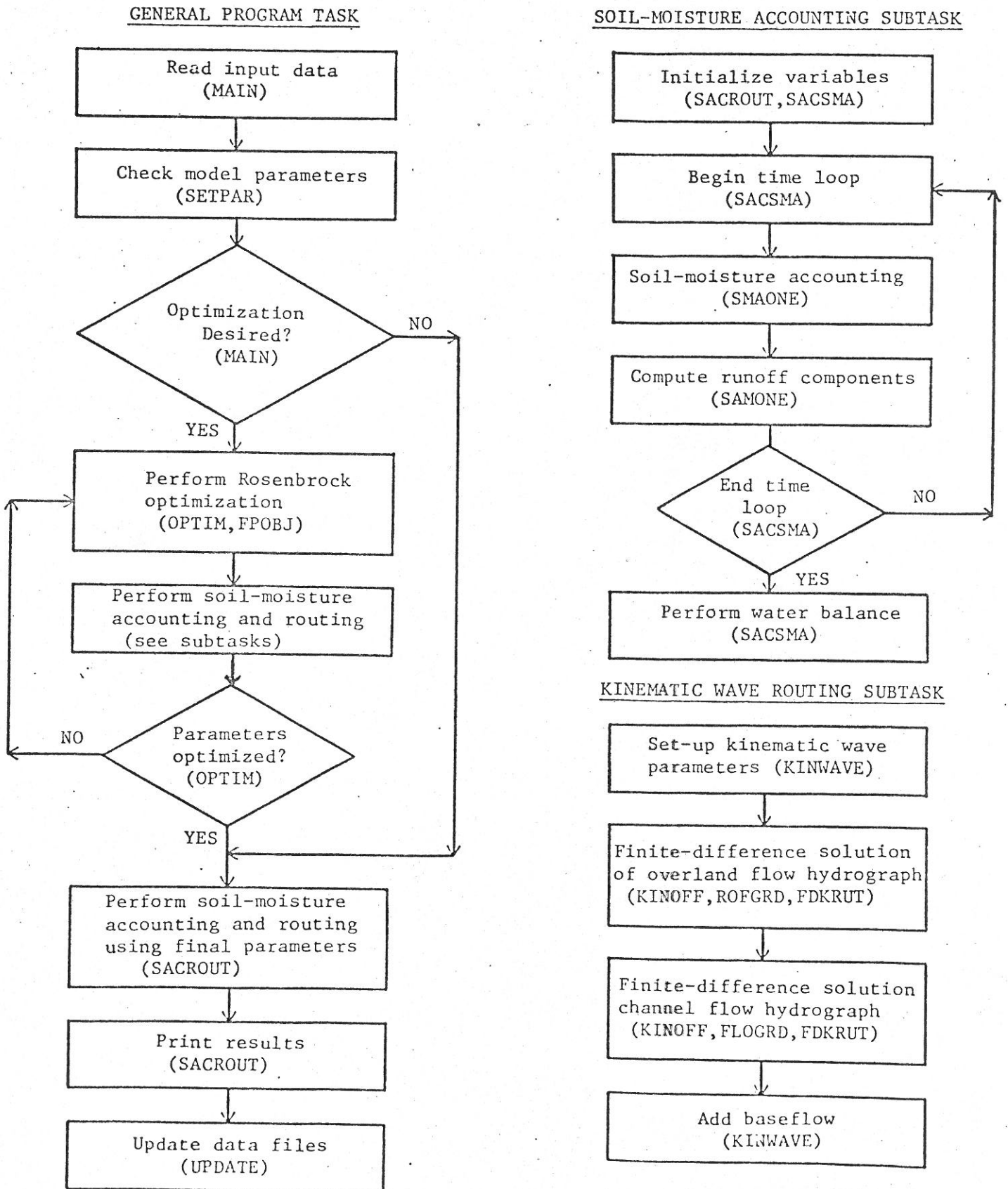


Figure 1.8.A.1. Descriptive flow chart of PROGRAM SACKW.

SAMPLE MODEL APPLICATION

	5	0	72	1	1				
0.01		1	25	1	1				
ZPERC	147.0			15.0		160.	5.0		
REXP	0.507			0.4		5.0	0.1		
SIDE	0.183			0.0001		10.0	0.001		
UZK	0.0096			0.001		0.90	0.001		
ADIMP	0.1175			0.001		0.5	0.001		
RSERV	0.42			0.01		.50	0.01		
RIVA	0.84			0.25		0.90	0.01		
PCTIM	0.089			0.05		.80	0.01		
LZPK	0.010			0.005		.80	0.001		
LZSK	0.0055			0.003		.50	0.001		
PFREE	0.790			0.01		.999	0.0001		
UZTWC	89.0			50.0		150.0	2.0		
UZFWC	33.0			30.0		80.0	2.0		
LZTWC	120.0			40.0		250.0	5.0		
LZFSC	33.00			30.0		150.0	2.0		
LZFPC	405.0			400.0		800.0	5.0		
ADIMC	102.0			90.0		400.0	10.0		
UZTWM	150.0								
UZFWM	100.0								
LZTWM	300.0								
LZFMS	200.0								
LZFPM	800.0								
PXADJ	1.0								
PEADJ	0.7								
END									
ROUTE									
70.0									
2500.0	0.6			0.4		100.0			
16000.	0.105			0.06		0.0			
5.0	15.0			1.0		4.92	0.0		
BASEF	2								
ROUTE									
70.0									
2500.	0.6			0.4		100.			
20000.	0.0837			0.06		70.0			
5.0	15.0			1.0		4.92	0.0		
BASEF	2								
ADD	2								
ROUTE									
45.0									
4500.0	0.5			0.3		100.0			
8500.0	0.0133			0.040		45.0			
5.0	85.0			1.0		4.92	1.0		
BASEF	2								
END									
RAIN									
	1	72	0	1	0				
(8F10.0)									
0.6	0.2			0.2		2.70	0.7	1.10	0.5
2.50	57.6			15.1		0.90	0.4	0.0	0.0
0.00	0.0			0.0		0.0	0.0	0.0	0.0
0.0	0.0			0.0		0.0	0.0	0.0	0.0
0.0	0.0			0.0		0.0	0.0	0.0	0.0

Figure 1.8.A.4 Input data file for sample model application.

0.0	0.0	0.0	0.0	0.0	0.0	0.0	0.0
0.0	0.0	0.0	0.0	0.0	0.0	0.0	0.0
0.0	0.0	0.0	0.0	0.0	0.0	0.0	0.0
0.0	0.0	0.0	0.0	0.0	0.0	0.0	0.0
ETDATA							
1	0	1	0				
(8F10.0)							
0.438	0.438	0.438	0.438	0.438	0.438	0.438	0.438
0.438	0.438	0.438	0.438	0.179	0.179	0.179	0.179
0.179	0.179	0.179	0.179	0.179	0.179	0.179	0.179
0.179	0.179	0.179	0.179	0.179	0.179	0.179	0.179
0.179	0.179	0.179	0.179	0.217	0.217	0.217	0.217
0.271	0.271	0.271	0.271	0.271	0.271	0.271	0.271
0.271	0.271	0.271	0.271	0.271	0.271	0.271	0.271
0.271	0.271	0.271	0.271	0.271	0.271	0.271	0.271
0.271	0.271	0.271	0.271	0.271	0.271	0.271	0.271
FLOW							
1	1	0	1	0			
(8F10.0)							
4.92	4.92	5.57	5.57	5.57	5.57	5.57	5.57
6.27	6.27	6.64	7.84	17.4	39.46	50.1	39.46
30.44	24.64	20.41	17.40	15.34	10.65	12.86	12.28
11.72	11.18	10.65	10.14	9.65	9.65	9.18	8.71
8.71	8.27	7.84	7.84	7.84	7.84	7.84	7.42
7.42	7.02	7.02	7.02	7.02	7.02	7.02	7.02
7.02	6.64	6.27	6.27	6.27	6.27	6.27	6.27
6.57	6.27	6.27	6.27	6.27	6.27	6.27	6.27
7.02	6.64	6.27	6.27	6.27	6.27	6.27	6.27
END							

Figure 1.8.A.4 (continuation)

BASIN 1 RUN: SAMPLE MODEL APPLICATION

LIST OF PARAMETERS TO BE OPTIMIZED

NAME	VALUE	MINIMUM	MAXIMUM	STEP SIZES
ZPERC	147.00000	15.00000	160.00000	5.00000
REXP	0.50700	0.40000	5.00000	0.10000
SIDE	0.18300	0.00010	10.00000	0.00100
UZK	0.00960	0.00100	0.90000	0.00100
ADIMP	0.11750	0.00100	0.50000	0.00100

LIST OF PARAMETERS MANUALLY CALIBRATED

NAME	VALUE
RSERV	0.42000
RIVA	0.84000
PCTIM	0.08900
LZPK	0.01000
LZSK	0.00550
PFREE	0.79000
UZTWC	89.00000
UZFWC	33.00000
LZT-C	120.00000
LZFSC	33.00000
LZFPC	405.00000
ADIMC	102.00000
UZTWM	150.00000
UZFYM	100.00000
LZTWM	300.00000
LZFSM	200.00000
LZFFM	800.00000
PXADJ	1.00000
PEADJ	0.70000

OPTIMIZATION PARAMETERS :
 OBJECTIVE FUNCTION TYPE 1
 MAXIMUM NUMBER OF ITERATIONS 25
 CONVERGENCE CRITERION 0.10000E-01
 STEP SIZE UPDATE OPTION 1
 PRINTING FREQUENCY 1

NUMBER OF PARAMETERS TO BE OPTIMIZED = 5
 NUMBER OF PARAMETERS MANUALLY CALIBRATED = 19

PARAMETERS TO BE OPTIMIZED AFTER CHECKING

ITEM NAME	VALUE	MINIMUM	MAXIMUM	STEP SIZES
1 ZPERC	147.00000	15.00000	160.00000	5.00000
2 REXP	0.50700	0.40000	5.00000	0.10000
3 SIDE	0.18300	0.00010	10.00000	0.00100
4 UZK	0.00960	0.00100	0.90000	0.00100
5 ADIMP	0.11750	0.00100	0.50000	0.00100

PARAMETERS MANUALLY CALIBRATED AFTER CHECKING

ITEM NAME	VALUE
6 RSERV	0.42000
7 RIVA	0.84000
8 PCTIM	0.08900

9	LZPK	0.01000
10	LZSK	0.00550
11	PFREE	0.79000
12	UZTWC	89.00000
13	UZFWC	33.00000
14	LZTWC	120.00000
15	LZFSC	33.00000
16	LZFPC	405.00000
17	ADIMC	102.00000
18	UZTWM	150.00000
19	UZFWM	100.00000
20	LZTWM	300.00000
21	LZFWM	200.00000
22	LZFFM	800.00000
23	PXADJ	1.00000
24	PEADJ	0.70000

OPTIMIZATION BY ROSENBROCK HILLCLIMB PROCEDURE

STAGE 1
 FUNCTION -0.90737742E+01
 LATERAL PROGRESS 0.24414062E-06
 PROGRESS 0.16250006E+02
 NUMBER OF FUNCTION EVALUATIONS = 70
 VALUES OF X(.) AT THIS STAGE
 X(1) = 0.158250E+03
 X(2) = 0.492937E+00
 X(3) = 0.183000E+00
 X(4) = 0.959803E-02
 X(5) = 0.117500E+00

STAGE 1
 FUNCTION -0.90737672E+01
 LATERAL PROGRESS 0.24414062E-06
 PROGRESS 0.16250006E+02
 NUMBER OF FUNCTION EVALUATIONS = 76
 VALUES OF X(.) AT THIS STAGE
 X(1) = 0.158260E+03
 X(2) = 0.492831E+00
 X(3) = 0.183000E+00
 X(4) = 0.959803E-02
 X(5) = 0.117500E+00

FINAL DIRECTION VECTOR MATRIX
 V(1, 1) = 0.99999903
 V(1, 2) = -0.00086538
 V(1, 3) = 0.00000000
 V(1, 4) = -0.00000012
 V(1, 5) = 0.00000002
 V(2, 1) = -0.00086538
 V(2, 2) = -0.99999962
 V(2, 3) = 0.00001736
 V(2, 4) = -0.00013889
 V(2, 5) = 0.00001736
 V(3, 1) = -0.00000012
 V(3, 2) = 0.00000000
 V(3, 3) = 0.12309149
 V(3, 4) = -0.98473193
 V(3, 5) = 0.12309149
 V(4, 1) = -0.00000012
 V(4, 2) = 0.00000000
 V(4, 3) = 0.00000000
 V(4, 4) = -0.99227788
 V(4, 5) = 0.12403473
 V(5, 1) = -0.00000002
 V(5, 2) = 0.00000000
 V(5, 3) = 0.00000000
 V(5, 4) = 0.00000000
 V(5, 5) = 1.00000000

FINAL STEP SIZES
 S(1) = 0.976563E-02
 S(2) = 0.195312E-03
 S(3) = 0.122070E-06
 S(4) = 0.122070E-06
 S(5) = 0.122070E-06

Figure 1.8.A.5 (continuation)

SACRAMENTO SOIL-MOISTURE ACCOUNTING OPERATION

PARAMETER VALUES - CAPACITIES ARE IN MM.

PX-ADJ	PE-ADJ	UZTWM	UZFWM	UZK	PCTIM	ADIMP	RIVA
1.000	0.700	150.000	100.000	0.010	0.089	0.118	0.840

DAILY ET DIST. ASSUMED UNIFORM

PBASE	ZPERC	REXP	LZTWM	LZFSM	LZFPM	LZSK	LZPK	RSERV	SIDE
9.100	158.260	0.493	300.000	200.000	800.000	0.006	0.010	0.420	0.183
								0.790	

SOIL-MOISTURE CONTENTS (MM)

UZTWC	UZFWC	LZTWC	LZFSC	LZFPC	ADIMC
89.000	33.000	120.000	33.000	405.000	102.000

DETAILED SOIL-MOISTURE ACCOUNTING OUTPUT

UNITS ARE IN MM.

DAY HR	UZTWC	UZFWC	LZTWC	LZFSC	LZFPC	ADIMC	PERC	IMP	DIR	SUR	INT	SUP	PRI	TOT-RO	ET-DMD	ACT-ET	RAIN
1 1	89.4	20.511	122.59	36.974	410.71	102.41	12.479	0.053	0.000	0.000	0.008	0.005	0.115	0.105	0.307	0.269	0.60
1 2	89.4	12.705	124.19	39.436	414.22	102.43	7.800	0.018	0.000	0.000	0.005	0.006	0.116	0.069	0.307	0.269	0.20
1 3	89.5	7.758	125.20	40.985	416.40	102.44	4.943	0.018	0.000	0.000	0.003	0.006	0.117	0.068	0.307	0.269	0.20
1 4	92.0	4.637	125.82	41.955	417.71	104.96	3.118	0.240	0.000	0.000	0.000	0.006	0.117	0.290	0.307	0.269	2.70
1 5	92.5	2.537	126.22	42.604	418.53	105.46	2.099	0.062	0.000	0.000	0.001	0.006	0.117	0.115	0.307	0.270	0.70
1 6	93.4	1.389	126.43	42.953	418.90	106.37	1.147	0.098	0.000	0.000	0.000	0.007	0.118	0.151	0.307	0.270	1.10
1 7	93.7	0.761	126.53	43.140	419.03	106.68	0.628	0.045	0.000	0.000	0.000	0.007	0.118	0.099	0.307	0.270	0.50
1 8	93.5	0.417	126.57	43.237	419.02	106.48	0.344	0.000	0.000	0.000	0.000	0.007	0.118	0.055	0.307	0.270	0.00
1 9	95.8	0.229	126.58	43.286	418.93	108.79	0.188	0.223	0.000	0.000	0.000	0.007	0.118	0.277	0.307	0.270	2.50
1 10	150.0	3.349	126.57	43.308	418.81	166.18	0.103	5.126	0.001	0.000	0.000	0.007	0.118	5.185	0.307	0.270	57.60
1 11	150.0	14.545	127.32	44.418	420.35	180.90	3.595	1.344	0.009	0.000	0.002	0.007	0.118	1.480	0.307	0.279	15.10
1 12	150.0	5.970	129.30	47.323	424.52	181.75	5.613	0.080	0.001	0.000	0.004	0.007	0.119	0.210	0.307	0.279	0.90
1 13	150.0	3.595	129.80	47.323	424.52	181.75	3.823	0.036	0.000	0.000	0.002	0.007	0.119	0.164	0.125	0.114	0.40
1 14	149.9	3.595	129.80	48.044	425.49	181.63	2.373	0.000	0.000	0.000	0.001	0.007	0.119	0.128	0.125	0.114	0.00
1 15	149.7	1.987	130.14	48.528	426.09	181.50	1.608	0.000	0.000	0.000	0.001	0.007	0.119	0.127	0.125	0.114	0.00
1 16	149.6	1.099	130.33	48.790	426.34	181.38	0.888	0.000	0.000	0.000	0.000	0.007	0.120	0.127	0.125	0.114	0.00
1 17	149.5	0.608	130.43	48.930	426.39	181.25	0.491	0.000	0.000	0.000	0.000	0.008	0.120	0.127	0.125	0.114	0.00
1 18	149.4	0.337	130.49	49.002	426.35	181.13	0.271	0.000	0.000	0.000	0.000	0.008	0.120	0.127	0.125	0.114	0.00
1 19	149.2	0.186	130.52	49.037	426.24	181.00	0.150	0.000	0.000	0.000	0.000	0.008	0.120	0.127	0.125	0.114	0.00
1 20	149.1	0.103	130.54	49.051	426.10	180.88	0.083	0.000	0.000	0.000	0.000	0.008	0.120	0.127	0.125	0.114	0.00
1 21	149.0	0.057	130.54	49.054	425.95	180.76	0.046	0.000	0.000	0.000	0.000	0.008	0.120	0.127	0.125	0.114	0.00
1 22	148.9	0.032	130.55	49.050	425.78	180.63	0.025	0.000	0.000	0.000	0.000	0.008	0.120	0.127	0.125	0.114	0.00
1 23	148.8	0.017	130.55	49.043	425.61	180.51	0.014	0.000	0.000	0.000	0.000	0.008	0.120	0.127	0.125	0.114	0.00
1 24	148.6	0.010	130.55	49.034	425.43	180.38	0.008	0.000	0.000	0.000	0.000	0.008	0.120	0.126	0.125	0.114	0.00
2 25	148.5	0.010	130.55	49.023	425.26	180.26	0.000	0.000	0.000	0.000	0.000	0.008	0.119	0.126	0.125	0.114	0.00
2 26	148.4	0.010	130.55	49.012	425.08	180.13	0.000	0.000	0.000	0.000	0.000	0.008	0.119	0.126	0.125	0.114	0.00
2 27	148.3	0.010	130.55	49.001	424.90	180.01	0.000	0.000	0.000	0.000	0.000	0.008	0.119	0.126	0.125	0.114	0.00
2 28	148.1	0.010	130.55	48.989	424.72	179.89	0.000	0.000	0.000	0.000	0.000	0.008	0.119	0.126	0.125	0.114	0.00
2 29	148.0	0.010	130.55	48.978	424.54	179.76	0.000	0.000	0.000	0.000	0.000	0.008	0.119	0.126	0.125	0.114	0.00
2 30	147.9	0.010	130.55	48.967	424.37	179.64	0.000	0.000	0.000	0.000	0.000	0.008	0.119	0.126	0.125	0.114	0.00
2 31	147.8	0.010	130.55	48.956	424.19	179.51	0.000	0.000	0.000	0.000	0.000	0.008	0.119	0.126	0.125	0.114	0.00
2 32	147.6	0.010	130.55	48.944	424.01	179.39	0.000	0.000	0.000	0.000	0.000	0.008	0.119	0.126	0.125	0.114	0.00
2 33	147.5	0.010	130.55	48.933	423.83	179.27	0.000	0.000	0.000	0.000	0.000	0.008	0.119	0.125	0.125	0.114	0.00
2 34	147.4	0.010	130.55	48.922	423.66	179.14	0.000	0.000	0.000	0.000	0.000	0.008	0.119	0.125	0.125	0.114	0.00
2 35	147.3	0.010	130.55	48.911	423.48	179.02	0.000	0.000	0.000	0.000	0.000	0.008	0.119	0.125	0.125	0.114	0.00
2 36	147.1	0.010	130.55	48.899	423.30	178.90	0.000	0.000	0.000	0.000	0.000	0.008	0.119	0.125	0.125	0.114	0.00
2 37	147.0	0.010	130.55	48.888	423.12	178.75	0.000	0.000	0.000	0.000	0.000	0.008	0.119	0.125	0.152	0.138	0.00
2 38	146.8	0.010	130.55	48.877	422.95	178.60	0.000	0.000	0.000	0.000	0.000	0.008	0.119	0.125	0.152	0.138	0.00
2 39	146.7	0.010	130.55	48.866	422.77	178.45	0.000	0.000	0.000	0.000	0.000	0.008	0.119	0.124	0.152	0.138	0.00
2 40	146.5	0.010	130.54	48.854	422.59	178.30	0.000	0.000	0.000	0.000	0.000	0.008	0.119	0.124	0.152	0.138	0.00
2 41	146.4	0.010	130.54	48.843	422.42	178.12	0.000	0.000	0.000	0.000	0.000	0.008	0.119	0.124	0.190	0.172	0.00

Figure 1.8.A.5 (continuation)

KINEMATIC WAVE ROUTING FOR SUBBASIN 1

TOTAL AREA = 70.00

OVERLAND FLOW ELEMENT 1

CHLNG = 2500.00
 SLOPE = 0.60000
 RCMAN = 0.40000
 PAREA = 100.000

MAIN CHANNEL

CHLNG = 16000.00
 SLOPE = 0.10500
 RCMAN = 0.06000
 SAREA = 70.000

ISHAPE= 5
 CHWDT = 15.000
 ZLNG = 1.000

FLOIC = 4.920

COMPUTED KINEMATIC PARAMETERS

ELEMENT	ALPHA	M	DT (MIN)	DX (MT)
1	2.8854	1.667	60.00	833.55
2	0.7189	1.604	60.00	8002.05

ADD BASEFLOW TO SUBBASIN 1 USING A LINEAR DECAY FUNCTION OF THE FORM:

ADDED BF AT TIME T = SUM OF W(L) * BF(T-L) FOR L = 0 TO 2 ; WHERE

W(L): 0.3000 0.5000 0.2000

KINEMATIC WAVE ROUTING FOR SUBBASIN 2

TOTAL AREA = 70.00

OVERLAND FLOW ELEMENT 1

CHLNG = 2500.00
 SLOPE = 0.60000
 RCMAN = 0.40000
 PAREA = 100.000

MAIN CHANNEL

CHLNG = 20000.00
 SLOPE = 0.08370
 RCMAN = 0.06000
 SAREA = 70.000

ISHAPE= 5
 CHWDT = 15.000
 ZLNG = 1.000

FLOIC = 4.920

Figure 1.8.A.5 (continuation)

COMPUTED KINEMATIC PARAMETERS

ELEMENT	ALPHA	M	DT (MIN)	DX (MT)
1	2.8854	1.667	60.00	833.55
2	0.6419	1.604	60.00	10002.56

ADD BASEFLOW TO SUBBASIN 2 USING A LINEAR DECAY FUNCTION OF THE FORM:

ADDED BF AT TIME T = SUM OF W(L) * BF(T-L) FOR L = 0 TO 2 ; WHERE

W(L): 0.3000 0.5000 0.2000

ADD FLOWS OF SUBBASINS 1, 2,

KINEMATIC WAVE ROUTING FOR SUBBASIN 3

TOTAL AREA = 45.00

OVERLAND FLOW ELEMENT 1

CHLNG = 4500.00
 SLOPE = 0.50000
 RCHAN = 0.30000
 PAREA = 100.000

MAIN CHANNEL

CHLNG = 8500.00
 SLOPE = 0.01330
 RCHAN = 0.04000
 SAREA = 45.000

ISHAPE = 5
 CHWDT = 85.000
 ZLNG = 1.000

FLOIC = 4.920

ROUTE UPSTREAM FLOW

COMPUTED KINEMATIC PARAMETERS

ELEMENT	ALPHA	M	DT (MIN)	DX (MT)
1	3.5120	1.667	60.00	1125.29
2	0.1068	1.654	60.00	4251.09

ADD BASEFLOW TO SUBBASIN 3 USING A LINEAR DECAY FUNCTION OF THE FORM:

ADDED BF AT TIME T = SUM OF W(L) * BF(T-L) FOR L = 0 TO 2 ; WHERE

W(L): 0.3000 0.5000 0.2000

VALUE OF OBJECTIVE FUNCTION (1) = 9.07377

DAY	HOUR	OBSERVED	COMPUTED	
1	1	4.920	0.589	C
1	2	4.920	1.453	C
1	3	5.570	2.542	C
1	4	5.570	4.419	C
1	5	5.570	5.923	E

Figure 1.8.A.5 (continuation)

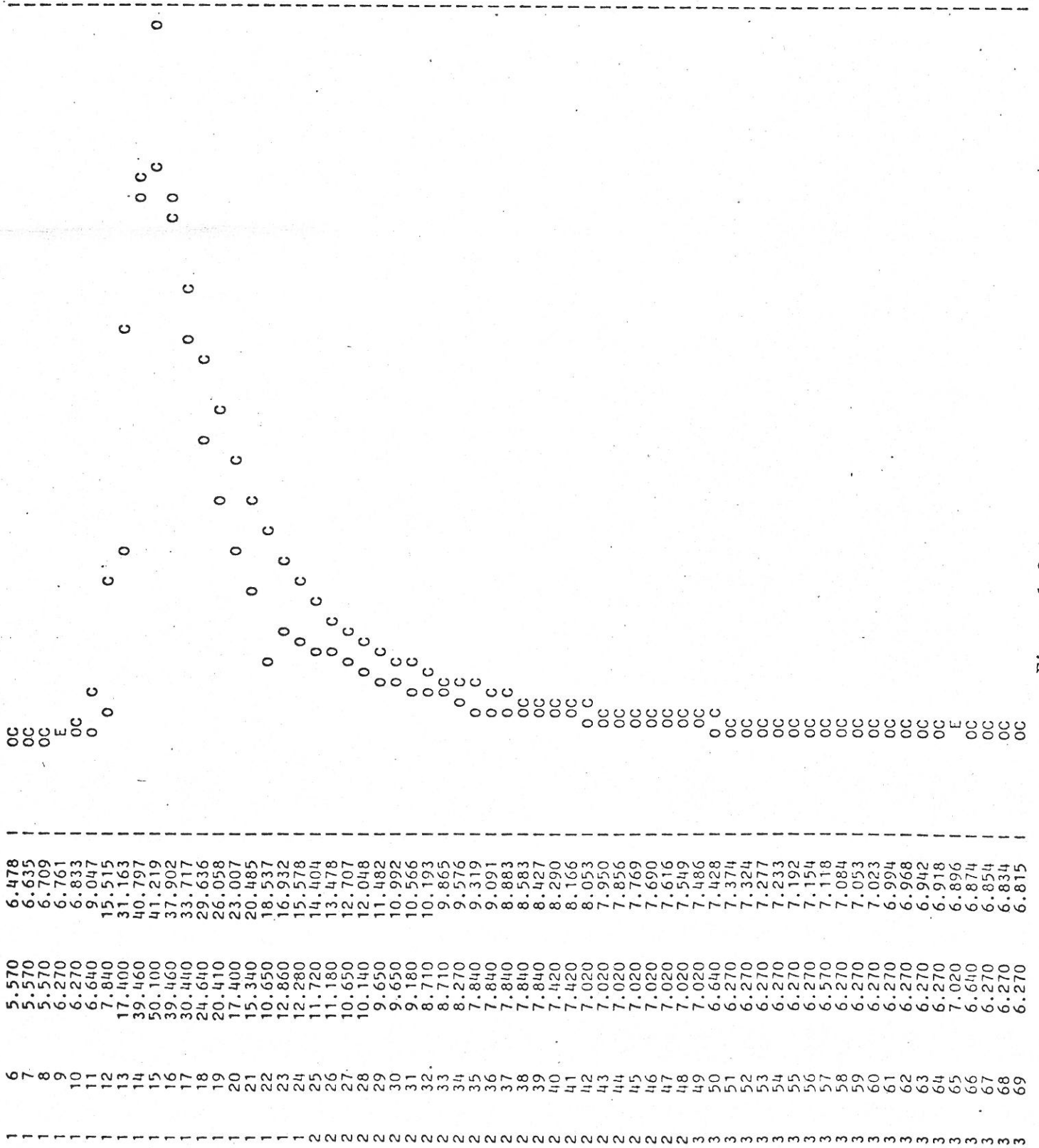


Figure 1.8.A.5 (continuation)

3	70	6.270	6.798	OC
3	71	6.270	6.781	OC
3	72	6.270	6.765	E

MINIMUM ORDINATE = 0.59
 MAXIMUM ORDINATE = 50.10

SYMBOLS USED:
 COMPUTED - C
 OBSERVED - O
 IF EQUAL - E

NORMAL TERMINATION

Figure 1.8.A.5 (continuation)

basis. This sample run is made where 5 SAC model parameters are optimized. In the printout, for the optimization results, the X(.) variable correspond to the model parameters optimized in the order they are inputted.

I-341

APPENDIX 1.8.B

PROGRAM LISTING OF SACKW MODEL

PROGRAM SACKW

1 (TAPE5,OUTPUT,TAPE7,TAPE8,TAPE9,TAPE10,TAPE6=OUTPUT)

C
C
C

MAINLINE PROGRAM FOR CALIBRATION AND FORECASTING

COMMON /CMIOP/ IOUA , IOUB ,
 1 NHOUR , NDATA ,
 2 NTIME, NHRP, NHRQ, NDTQ, INPQ, NBASIN, IWBLNC, IPFRQR, IPFRQQ
 COMMON /CMPRM/ NPARM , NPOP , IOBF ,
 1 MXITER , NSTEP , ERROR , IPROP ,
 2 INDX(50) , XMP(50) , XMIN(50) , XMAX(50) ,
 3 ESS(50),OBJFV
 COMMON /FDK00/ TAREA,TRMN,TRHR,METRC,IEL,NQ,MXNDX,MXNDT
 COMMON /FDK01/ CHLNG(5),SLOPE(5),RCMAN(5),PAREA(5),ISHAPE(5),
 1 CHWDT(5),ZLNG(5),ALPME(5),EMDEQ(5),DXROF(5),DXFLO(5),DTKWR(5),
 2 DXKWR(5),IRUPF(5),SAREA(5),FLOIC(5)
 COMMON /FDK02/ EXCSR(1224),Q(1224),
 1 QK(1224),QUB(1224),QBF(1224)
 COMMON /CMDAT/ RDT(1224),PDT(1224),QDT(1224)
 REAL LZTWM , LZFSM , LZFPM ,
 1 LZSK , LZPK , LZTWC , LZFSC ,
 2 LZFPC
 COMMON /CMSMP/ PXADJ , PEADJ , UZTWM ,
 1 UZFWM , UZK , PCTIM , ADIMP ,
 2 RIVA , RSERV , ZPERC , REXP ,
 3 LZTWM , LZFSM , LZFPM , LZSK ,
 4 LZPK , PFREE , SIDE , SAVED ,
 5 PCIAR
 COMMON /CMSMC/ UZTWC , UZFWC , LZTWC ,
 1 LZFSC , LZFPC , ADIMC , RSUM(7)
 COMMON /CMSMS/ SROT, SRECHT, SETT
 DIMENSION ALIST(4),TITLE(10) , FMT(10) , PNAME(50).
 DATA ALIST/5HRAIN , 5HETDAT,5HFLOW , 5HEND /

C

IOUA = 5
 IOUB = 6
 NBASIN=0
 NHOUR=0
 NHRP=0

C
C
C

READ TITLE OF RUN

90 READ (IOUA,200,END=91) TITLE
 NBASIN=NBASIN+1
 WRITE (IOUB,240) NBASIN, TITLE
 NHRQ=0
 INPQ=0

C
C
C
C
C
C
C
C

READ CONTROL PARAMETERS

NPOP = NUMBER OF PARAMETERS TO BE OPTIMIZED
 IUPDF = 1 IF UPDATED DATA FILE TO BE GENERATED, 0 OTHERWISE
 IWBLNC = WATER BALANCE COMPUTATION PERIOD
 IPFRQR = PRINTING FREQUENCY FOR RAINFALL
 IPFRQQ = PRINTING FREQUENCY FOR STREAMFLOW

READ (IOUA,210) NPOP,IUPDF,IWBLNC,IPFRQR,IPFRQQ

IF (NPOP.EQ.0) GO TO 100

READ OPTIMIZATION ROUTINE PARAMETERS IF NPOP .NE. 0

ERROR = CONVERGENCE CRITERION FOR RELATIVE OBJ FUNC IMPROVEMENT
IOBF = TYPE OF OBJECTIVE FUNCTION OPTION

NOTES ON TYPE OF OBJECTION FUNCTION OPTION FROM SEFE AND
BOUGHTON, JOUR. OF HYDROLOGY (NEW ZEALAND), VOL 21 NO 2, 1982.

IOBF = 1, $SUM[(QOBS-QCOM)**2]$. THIS FUNCTION GIVES MORE
WEIGHT TO LARGE DIFFERENCES THAN SMALL DIFFERENCES
LEADING TO BETTER ESTIMATE OF HIGH FLOWS.

IOBF = 2, $SUM[ABS(QOBS-QCOM)]$. THIS FUNCTION GIVES EQUAL
WEIGHT TO ALL RESIDUALS.

IOBF = 3, $SUM[((QOBS-QCOM)/QOBS)**2]$. THIS FUNCTION GIVES
EQUAL WEIGHT TO EQUAL PROPORTIONAL RESIDUALS AND ALSO
GREATER WEIGHT TO SMALLER ABSOLUTE RESIDUALS AT TIMES
OF LOW FLOW.

IOBF = 4, $SUM[(QOBS**(1/3)-QCOM**(1/3))**2]$. THIS FUNCTION
REDUCES THE EFFECT OF LARGE RESIDUALS THUS GIVING MORE
WEIGHT TO LOW FLOWS. THE CUBE-ROOT TRANSFORMATION OF THE
FUNCTION IS AIMED AT MAKING THE RESIDUALS MUTUALLY UNCOR-
RELATED, NORMALLY DISTRIBUTED AND THE LOG-LIKELIHOOD
FUNCTION OF THE RESIDUALS IS APPROXIMATELY QUADRATIC IN
THE PARAMETER VALUES IN THE NEIGHBORHOOD OF ITS MAXIMUM
SO THAT ITS CONTOURS ARE APPROXIMATELY ELLIPSOIDAL.

IOBF = 5, $SUM[(QOBS**(1/2)-QCOM**(1/2))**2]$. THIS FUNCTION
IS SIMILAR TO OBJECTIVE FUNCTION OPTION IOBF = 4 BUT IT
USES A SQUARE-ROOT TRANSFORMATION INSTEAD.

MXITER = MAXIMUM NUMBER OF ITERATIONS

NSTEP = 0 IF STEP SIZES REMAIN THE SAME, 1 OTHERWISE

IPROP = PRINTING FREQUENCY FOR STAGE EVALUATION

READ (IOUA,235) ERROR,IOBF,MXITER,NSTEP,IPROP

IOBF = MAX0(IOBF,1)

IOBF = MIN0(IOBF,5)

MXITER = MAX0(MXITER,1)

IPROP = MAX0(IPROP,1)

100 IF (IUPDF.NE.0) WRITE (IOUB,250)

IP = 0

IF (NPOP.EQ.0) GO TO 120

READ PARAMETERS TO BE OPTIMIZED, ONE CARD EACH CONTAINING:

PNAME(.) = NAME OF VARIABLE (COLUMNS 1-5)

DUM = 5-CHARACTER LABEL FOR USER IDENTIFICATION (COLUMNS 6-10)

XMP(.) = INITIAL PARAMETER FLOATING-POINT VALUE (COLUMNS 11-20)

XMIN(.) = PARAMETER LOWER BOUND (COLUMNS 21-30)

XMAX(.) = PARAMETER UPPER BOUND (COLUMNS 31-40)

ESS(.) = PARAMETER STEP SIZE (COLUMNS 41-50)

WRITE (IOUB,260)

DO 110 I = 1,NPOP

IP = IP + 1

READ (IOUA,220) PNAME(IP),DUM,XMP(IP),XMIN(IP),XMAX(IP),ESS(IP)

```

      IF (PNAME(IP).EQ.ALIST(4)) GO TO 130
110 WRITE (IOUB,280) PNAME(IP),DUM,XMP(IP),XMIN(IP),XMAX(IP),ESS(IP)
C
C   READ PARAMETERS MANUALLY CALIBRATED BY USER (INPUT VALUES),
C   ONE CARD CONTAINS SAME ITEMS AS ABOVE EXCLUDING XMIN(.),
C   XMAX(.) AND ESS(.)
C
120 IP = IP + 1
    READ (IOUA,220) PNAME(IP),DUM,XMP(IP),XMIN(IP),XMAX(IP)
    IF (PNAME(IP).EQ.ALIST(4)) GO TO 130
    IF (IP.EQ.(NPOP + 1)) WRITE (IOUB,270)
    WRITE (IOUB,280) PNAME(IP),DUM,XMP(IP)
    GO TO 120
130 IP = IP - 1
    NPARM = IP
    NPMC = NPARM - NPOP
    WRITE (IOUB,285) IOBF,MXITER,ERROR,NSSTEP,IPROP
    WRITE (IOUB,290) NPOP,NPMC
C
C   CALL SETPAR FOR SETTING-UP AND CHECKING PARAMETERS
C
C   CALL SETPAR (PNAME,NPMC)
C
C   READ KINEMATIC WAVE ROUTING PARAMETERS AND STORE IN TAPE10
C
    REWIND 10
135 READ (IOUA,205) DUM,TITLE
    WRITE(10,200) TITLE
    IF(DUM.EQ.ALIST(4)) GO TO 136
    GO TO 135
C
C   READ HYDROLOGIC DATA
C
136 READ(IOUA,205,END=170) RLIST
    DO 137 IL=1,4
    IF(RLIST.EQ.ALIST(IL)) GO TO 138
137 CONTINUE
    WRITE(IOUB,295) RLIST
    STOP
138 GO TO (141,142,160,170), IL
C
C   READ RAINFALL DATA
C
141 READ (IOUA,210) NHRP,IOUC,KFMT,IREW
    IF(IOUC.EQ.0) IOUC=5
    IF(IREW.NE.0) REWIND IOUC
    IF (KFMT.NE.0) READ (IOUA,200) (FMT(I),I = 1,10)
    IF(KFMT.EQ.0) READ(IOUC,230) (RDT(I),I=1,NDATA)
    IF(KFMT.NE.0) READ(IOUC,FMT) (RDT(I),I=1,NDATA)
    GO TO 136
C
C   READ EVAPOTRANSPIRATION DEMAND DATA
C
142 READ (IOUA,210) NHRP,IOUC,KFMT,IREW
    IF(IOUC.EQ.0) IOUC=5
    IF(IREW.NE.0) REWIND IOUC

```

```

NHRP = MAX0(NHRP,1)
NDTP=NDATA/NHRP
IF (KFMT.NE.0) READ (IOUA,200) (FMT(I),I = 1,10)
IF (KFMT.EQ.0) READ (IOUC,230) (PDT(I),I=1,NDTP)
IF (KFMT.NE.0) READ (IOUC,FMT) (PDT(I),I=1,NDTP)
GO TO 136

```

C
C
C

READ STREAMFLOW DATA, REQUIRED IN MODEL CALIBRATION RUNS

```

160 READ (IOUA,210) NHRQ,INPQ,IOUC,KFMT,IREW
IF (IOUC.EQ.0) IOUC=5
IF (IREW.NE.0) REWIND IOUC
NHRQ = MAX0(NHRQ,1)
NDTQ=NDATA/NHRQ
IF (INPQ.EQ.0) GO TO 136
IF (KFMT.NE.0) READ (IOUA,200) (FMT(I),I = 1,10)
IF (KFMT.EQ.0) READ (IOUC,230) (PDT(I),I=1,NDTQ)
IF (KFMT.NE.0) READ (IOUC,FMT) (PDT(I),I=1,NDTQ)
GO TO 136
170 IF (NPOP.EQ.0) GO TO 180
IF (INPQ.NE.0) GO TO 180
WRITE (IOUB,300)
STOP
180 IF (NHRP.NE.0) GO TO 183
WRITE (IOUB,296)
STOP
183 IF (NHRP.NE.0) GO TO 181
NHRP=NHRP
DO 140 I=1,NDATA
140 PDT(I)=0.0
181 IF (NHRQ.EQ.0) NHRQ = NHRP
NDTQ=NDATA/NHRQ
IF (NPOP.EQ.0) GO TO 190
IPRINT = 0
CALL OPTIM (IPRINT)
190 IPRINT = 1
CALL SACROUT (IUPDF,IPRINT)
IF (IUPDF.EQ.1) CALL UPDATE
GO TO 90
91 STOP

```

C
C

```

200 FORMAT (10A8)
205 FORMAT (A5,T1,10A8)
210 FORMAT (16I5)
220 FORMAT (2A5,7F10.0)
230 FORMAT (8F10.0)
235 FORMAT (F10.0,5I5)
240 FORMAT ('1'/5X,'BASIN',I2,' RUN: ',10A8/)
250 FORMAT (/2X,'NOTE: UPDATED DATA FILE IS GENERATED'/)
260 FORMAT (/2X,34HLIST OF PARAMETERS TO BE OPTIMIZED,//3X,5H NAME,10X
1,5HVALUE,6X,7HMINIMUM,4X,19HMAXIMUM STEP SIZES,/)
270 FORMAT (/2X,38HLIST OF PARAMETERS MANUALLY CALIBRATED,//3X,5H NAME
1,10X,5HVALUE,/)
280 FORMAT (3X,2A5,2X,4F12.5)
285 FORMAT (/2X,'OPTIMIZATION PARAMETERS :'/)

```

```

1 4X,'OBJECTIVE FUNCTION TYPE',T35,I5/
2 4X,'MAXIMUM NUMBER OF ITERATIONS',T35,I5/
3 4X,'CONVERGENCE CRITERION',T28,G12.5/
4 4X,'STEP SIZE UPDATE OPTION',T35,I5/
5 4X,'PRINTING FREQUENCY',T35,I5)
290 FORMAT (/2X,38HNUMBER OF PARAMETERS TO BE OPTIMIZED =,I4/2X,42HNUM
BER OF PARAMETERS MANUALLY CALIBRATED =,I3)
295 FORMAT(/3X,'***** EXECUTION TERMINATED - INPUT LIST ',A5,
1 ' IS UNRECOGNIZED STATEMENT *****')
300 FORMAT(/3X,'***** EXECUTION TERMINATED - OPTIMIZATION DESIRED BUT
1 NO STREAMFLOW DATA *****')
296 FORMAT(/3X,'***** EXECUTION TERMINATED - NO RAINFALL DATA *****')
END
SUBROUTINE SETPAR (PNAME,NPMC)

```

C
C
C

THIS ROUTINE CHECKS AND SETS-UP THE PARAMETERS

```

COMMON /CMIOP/      IOUA      , IOUB      ,
1      NHOOR      , NDATA      ,
2      NTIME,NHRP,NHRQ,NDTQ,INPQ,NEASIN,IWBLNC,IPFRQR,IPFRQQ
COMMON /CMPRM/      NPARM      , NPOP      , IOBF      ,
1      MXITER      , NSTEP      , ERROR      , IPROP      ,
2      INDX(50)      , XMP(50)      , XMIN(50)      , XMAX(50)      ,
3      ESS(50),OBJFV
COMMON /FDK00/ TAREA,TRMN,TRHR,METRC,IEL,NQ,MXNDX,MXNDT
COMMON /FDK01/ CHLNG(5),SLOPE(5),RCMAN(5),PAREA(5),ISHAPE(5),
1 CHWDT(5),ZLNG(5),ALPME(5),EMDEQ(5),DXROF(5),DXFLO(5),DTKWR(5),
2 DXKWR(5),IRUPF(5),SAREA(5),FLOIC(5)
COMMON /FDK02/ EXCSR(1224),Q(1224),
1 QK(1224),QUB(1224),QBF(1224)
COMMON /CMDAT/ RDT(1224),PDT(1224),QDT(1224)
REAL
1      LZSK      , LZTWM      , LZFSM      , LZFPM      ,
2      LZFK      , LZTWC      , LZFSC      ,
COMMON /CMSMP/      PXADJ      , PEADJ      , UZTWM      ,
1      UZFWM      , UZK      , PCTIM      , ADIMP      ,
2      RIVA      , RSERV      , ZPERC      , REXP      ,
3      LZTWM      , LZFSM      , LZFPM      , LZSK      ,
4      LZPK      , PFREE      , SIDE      , SAVED      ,
5      PCIAR
COMMON /CMSMC/      UZTWC      , UZFWC      , LZTWC      ,
1      LZFSC      , LZFPC      , ADIMC      , RSUM(7)
DIMENSION
      PNAME(50)      , FNAME(24)
DATA FNAME/5HPXADJ,5HPEADJ,5HPCTIM,5HADIMP,5HRIVA ,5HUZK ,5HLZSK
1,5HLZPK ,5HPFREE,5HRSERV,5HZPERC,5HREXP ,5HSIDE ,5HUZTWM,5HUZFWM,5
2HLZTWM,5HLZFSM,5HLZFPM,5HUZTWC,5HLZTWC,5HUZFWC,5HLZFSC,5HLZFPC,5HA
3DIMC/
UZTWM = -99.0
UZFWM = -99.0
LZTWM = -99.0
LZFSM = -99.0
LZFPM = -99.0
UZTWC = -99.0
DO 460 IP = 1,NPARM
DO 100 KP = 1,30
IF (PNAME(IP).EQ.FNAME(KP)) GO TO 110

```

```

100  CONTINUE
      WRITE (IOUB,500) PNAME(IP)
      STOP
110  GO TO (120,130,140,150,160,170,180,190,200,210,220,230,240,250,
1    260,270,280,290,300,310,320,330,340,350), KP

```

C
C
C

PARAMETER CHECKS

```

120  PXADJ = XMP(IP)
      XLO = 0.0001
      XUP = 10.0
      PXADJ = AMAX1(PXADJ,XLO)
      PXADJ = AMIN1(PXADJ,XUP)
      XMP(IP) = PXADJ
      INDX(IP) = 1
      IF (IP.GT.NPOP) GO TO 460
      GO TO 430
130  PEADJ = XMP(IP)
      XLO = 0.0001
      XUP = 10.0
      PEADJ = AMAX1(PEADJ,XLO)
      PEADJ = AMIN1(PEADJ,XUP)
      XMP(IP) = PEADJ
      INDX(IP) = 2
      IF (IP.GT.NPOP) GO TO 460
      GO TO 430
140  PCTIM = XMP(IP)
      PCTIM = AMAX1(PCTIM,0.0001)
      PCTIM = AMIN1(PCTIM,1.0)
      XMP(IP) = PCTIM
      INDX(IP) = 3
      IF (IP.GT.NPOP) GO TO 460
      GO TO 420
150  ADIMP = XMP(IP)
      ADIMP = AMAX1(ADIMP,0.0001)
      XUP = 1.0 - PCTIM
      ADIMP = AMIN1(ADIMP,XUP)
      XMP(IP) = ADIMP
      INDX(IP) = 4
      IF (IP.GT.NPOP) GO TO 460
      XLO = 0.0001
      GO TO 430
160  RIVA = XMP(IP)
      RIVA = AMAX1(RIVA,0.0001)
      RIVA = AMIN1(RIVA,1.0)
      XMP(IP) = RIVA
      INDX(IP) = 5
      IF (IP.GT.NPOP) GO TO 460
      GO TO 420
170  UZK = XMP(IP)
      UZK = AMAX1(UZK,0.0001)
      UZK = AMIN1(UZK,1.0)
      XMP(IP) = UZK
      INDX(IP) = 6
      IF (IP.GT.NPOP) GO TO 460
      GO TO 420

```



```

180  LZSK = XMP(IP)
      LZSK = AMAX1(LZSK,0.0001)
      LZSK = AMIN1(LZSK,1.0)
      XMP(IP) = LZSK
      INDX(IP) = 7
      IF (IP.GT.NPOP) GO TO 460
      GO TO 420
190  LZPK = XMP(IP)
      LZPK = AMAX1(LZPK,0.0001)
      LZPK = AMIN1(LZPK,1.0)
      XMP(IP) = LZPK
      INDX(IP) = 8
      IF (IP.GT.NPOP) GO TO 460
      GO TO 420
200  PFREE = XMP(IP)
      PFREE = AMAX1(PFREE,0.0001)
      PFREE = AMIN1(PFREE,1.0)
      XMP(IP) = PFREE
      INDX(IP) = 9
      IF (IP.GT.NPOP) GO TO 460
      GO TO 420
210  RSERV = XMP(IP)
      RSERV = AMAX1(RSERV,0.0001)
      RSERV = AMIN1(RSERV,1.0)
      XMP(IP) = RSERV
      INDX(IP) = 10
      IF (IP.GT.NPOP) GO TO 460
      GO TO 420
220  ZPERC = XMP(IP)
      INDX(IP) = 11
      IF (IP.GT.NPOP) GO TO 460
      GO TO 440
230  REXP = XMP(IP)
      INDX(IP) = 12
      IF (IP.GT.NPOP) GO TO 460
      GO TO 440
240  SIDE = XMP(IP)
      INDX(IP) = 13
      IF (IP.GT.NPOP) GO TO 460
      GO TO 440

```

C
C
C

INSURE THAT CAPACITIES ARE NOT EQUAL TO ZERO

```

250  UZTWM = XMP(IP)
      UZTWM = AMAX1(UZTWM,0.0001)
      XMP(IP) = UZTWM
      INDX(IP) = 14
      IF (IP.GT.NPOP) GO TO 460
      GO TO 440
260  UZFWM = XMP(IP)
      UZFWM = AMAX1(UZFWM,0.0001)
      XMP(IP) = UZFWM
      INDX(IP) = 15
      IF (IP.GT.NPOP) GO TO 460
      GO TO 440
270  LZTWM = XMP(IP)

```

```
LZTWM = AMAX1(LZTWM,0.0001)
XMP(IP) = LZTWM
INDX(IP) = 16
IF (IP.GT.NPOP) GO TO 460
GO TO 440
280 LZFSM = XMP(IP)
LZFSM = AMAX1(LZFSM,0.0001)
XMP(IP) = LZFSM
INDX(IP) = 17
IF (IP.GT.NPOP) GO TO 460
GO TO 440
290 LZFPM = XMP(IP)
LZFPM = AMAX1(LZFPM,0.0001)
XMP(IP) = LZFPM
INDX(IP) = 18
IF (IP.GT.NPOP) GO TO 460
GO TO 440
300 UZTWC = XMP(IP)
UZTWC = AMAX1(UZTWC,0.0)
XUP = UZTWM
IF (XUP.NE.-99.0) GO TO 305
XUP=XMAX(IP)
IF (XUP.NE.0.0.AND.XUP.GT.UZTWC) GO TO 305
WRITE(IOUB,560) PNAME(IP),UZTWC,XUP
STOP
305 UZTWC = AMIN1(UZTWC,XUP)
XMP(IP) = UZTWC
INDX(IP) = 19
IF (IP.GT.NPOP) GO TO 460
XLO = 0.0
GO TO 430
310 LZTWC = XMP(IP)
LZTWC = AMAX1(LZTWC,0.0)
XUP = LZTWM
IF (XUP.NE.-99.0) GO TO 315
XUP=XMAX(IP)
IF (XUP.NE.0.0.AND.XUP.GT.LZTWC) GO TO 315
WRITE(IOUB,560) PNAME(IP),LZTWC,XUP
STOP
315 LZTWC = AMIN1(LZTWC,XUP)
XMP(IP) = LZTWC
INDX(IP) = 20
IF (IP.GT.NPOP) GO TO 460
XLO = 0.0
GO TO 430
320 UZFWC = XMP(IP)
UZFWC = AMAX1(UZFWC,0.0)
XUP = UZFWM
IF (XUP.NE.-99.0) GO TO 325
XUP=XMAX(IP)
IF (XUP.NE.0.0.AND.XUP.GT.UZFWC) GO TO 325
WRITE(IOUB,560) PNAME(IP),UZFWC,XUP
STOP
325 UZFWC = AMIN1(UZFWC,XUP)
XMP(IP) = UZFWC
INDX(IP) = 21
```



```

      IF (IP.GT.NPOP) GO TO 460
      XLO = 0.0
      GO TO 430
330   LZFSC = XMP(IP)
      LZFSC = AMAX1(LZFSC,0.0)
      XUP = LZFSM
      IF(XUP.NE.-99.0) GO TO 335
      XUP=XMAX(IP)
      IF(XUP.NE.0.0.AND.XUP.GT.LZFSC) GO TO 335
      WRITE(IOUB,560) PNAME(IP),LZFSC,XUP
      STOP
335   LZFSC = AMINI(LZFSC,XUP)
      XMP(IP) = LZFSC
      INDX(IP) = 22
      IF (IP.GT.NPOP) GO TO 460
      XLO = 0.0
      GO TO 430
340   LZFPC = XMP(IP)
      LZFPC = AMAX1(LZFPC,0.0)
      XUP = LZFCM
      IF(XUP.NE.-99.0) GO TO 345
      XUP=XMAX(IP)
      IF(XUP.NE.0.0.AND.XUP.GT.LZFPC) GO TO 345
      WRITE(IOUB,560) PNAME(IP),LZFPC,XUP
      STOP
345   LZFPC = AMINI(LZFPC,XUP)
      XMP(IP) = LZFPC
      INDX(IP) = 23
      IF (IP.GT.NPOP) GO TO 460
      XLO = 0.0
      GO TO 430
350   ADIMC = XMP(IP)
      XLO = UZTWC
      IF(XLO.EQ.-99.0) XLO = 0.0001
      XUP = UZTWM + LZTWM
      IF(UZTWM.NE.-99.0.AND.LZTWM.NE.-99.0) GO TO 355
      XUP=XMAX(IP)
      IF(XUP.NE.0.0.AND.XUP.GT.ADIMC) GO TO 355
      WRITE(IOUB,560) PNAME(IP),ADIMC,XUP
      STOP
355   ADIMC = AMAX1(ADIMC,XLO)
      ADIMC = AMINI(ADIMC,XUP)
      XMP(IP) = ADIMC
      INDX(IP) = 24
      IF (IP.GT.NPOP) GO TO 460
      GO TO 430
420   XLO = 0.0001
      XUP = 1.0
430   XMIN(IP) = AMAX1(XMIN(IP),XLO)
      XMAX(IP) = AMINI(XMAX(IP),XUP)
440   IF (XMAX(IP).GT.XMIN(IP).AND.(XMAX(IP) - XMIN(IP)).NE.0.0) GO T
1     O 450
      WRITE (IOUB,510) PNAME(IP),XMIN(IP),XMAX(IP)
      STOP
450   IF (ESS(IP).NE.0.0) GO TO 460
      ESS(IP) = ABS(XMAX(IP) - XMIN(IP))/20.0

```

```

      WRITE (IOUB,520) PNAME(IP),ESS(IP)
460 CONTINUE
      IP = 0
      IF (NPOP.EQ.0) GO TO 480
      WRITE (IOUB,530)
      DO 470 I = 1,NPOP
        IP = IP + 1
470 WRITE (IOUB,550) IP,PNAME(IP),XMP(IP),XMIN(IP),XMAX(IP),ESS(IP)
480 IF (NPMC.EQ.0) RETURN
      WRITE (IOUB,540)
      DO 490 I = 1,NPMC
        IP = IP + 1
490 WRITE (IOUB,550) IP,PNAME(IP),XMP(IP)
      RETURN

```

C
C
C

```

500 FORMAT (/2X,10HPARAMETER ,A5,21H IS NOT IN DICTIONARY,//2X,20HEXEC
      IUTION TERMINATED,/)
510 FORMAT (/2X,10HPARAMETER ,A5,10H HAS MIN =,F10.3,10H AND MAX =,F10
      1.3,/2X,43HWHICH IS A VIOLATION - EXECUTION TERMINATED,/)
520 FORMAT (/2X,10HPARAMETER ,A5,34H WITH STEP SIZE ESS = 0.0 IS RESET
      1,/2X,8HTO ESS =,F10.5,29H TAKEN FROM ABS(MAX-MIN)/20.0,/)
530 FORMAT (/2X,41HPARAMETERS TO BE OPTIMIZED AFTER CHECKING,//3X,11H
      1 ITEM NAME,5X,5HVALUE,6X,7HMINIMUM,4X,19HMAXIMUM STEP SIZES,/)
540 FORMAT (/2X,45HPARAMETERS MANUALLY CALIBRATED AFTER CHECKING,//3X,
      110H ITEM NAME,5X,5HVALUE,/)
550 FORMAT (3X,I3,2X,A5,2X,4F12.5)
560 FORMAT(/2X,'PARAMETER ',A5,' =',G12.6,' AND MAX =',G12.6//
      1 2X,'EXECUTION TERMINATED. THE FOLLOWING ACTIONS MAY BE TAKEN: '/
      2 2X,'IF PARAMETER IS OPTIMIZED: '/
      3 4X,'-CHECK SPECIFIED XMAX'/
      4 4X,'-CHECK IF SOIL MOISTURE CONTENT IS CONSISTENT WITH CORREPPDIN
      5G CAPACITIES'/
      6 2X,'IF PARAMETER IS NOT OPTIMIZED: '/
      7 4X,'-INPUT CORRESPONDING CAPACITIES BEFORE CONTENTS, OTHERWISE'//
      8 4X,' SPECIFY XMAX AS IN OPTIMIZED PARAMETERS'//
      9 2X,'EXECUTION TERMINATED')
      END
      SUBROUTINE SACSMA (IPRINT)

```

C
C
C

THIS IS THE MAINLINE ROUTINE FOR SACRAMENTO SOIL-MOISTURE MODEL

```

COMMON /CMIOP/      IOUA      , IOUB      ,
1      NHOOR      , NDATA      ,
2      NTIME, NHRP, NHRQ, NDTQ, INPQ, NBASIN, IWBLNC, IPFRQR, IPFRQQ
COMMON /CMPRM/      NPARM      , NPOP      , IOBF      ,
1      MXITER      , NSTEP      , ERROR      , IPROP      ,
2      INDX(50)      , XMP(50)      , XMIN(50)      , XMAX(50)      ,
3      ESS(50), OBJFV
COMMON /FDK00/ TAREA, TRMN, TRHR, METRC, IEL, NQ, MXNDX, MXNDT
COMMON /FDK01/ CHLNG(5), SLOPE(5), RCMAN(5), PAREA(5), ISHAPE(5),
1 CHWDI(5), ZLNG(5), ALPME(5), EMDEQ(5), DXROF(5), DXFLO(5), DTKWR(5),
2 DXKWR(5), IRUPF(5), SAREA(5), FLOIC(5)
COMMON /FDK02/ EXCSR(1224), Q(1224),
1 QK(1224), QUB(1224), QBF(1224)

```

```

COMMON /CMDAT/ RDT(1224),PDT(1224),QDT(1224)
REAL
1 LZSK , LZPK , LZFSM , LZFPM ,
2 LZFPC , LZTWC , LZFSC ,
COMMON /CMSMP/ PXADJ , PEADJ , UZTWM ,
1 UZFWM , UZK , PCTIM , ADIMP ,
2 RIVA , RSERV , ZPERC , REXP ,
3 LZTWM , LZFSM , LZFPM , LZSK ,
4 LZPK , PFREE , SIDE , SAVED ,
5 PCIAR
COMMON /CMSMC/ UZTWC , UZFWC , LZTWC ,
1 LZFSC , LZFPC , ADIMC , RSUM(7)
COMMON /CMSMS/ SROT, SRECHT, SETT
DIMENSION EPDIST(24)

```

C

```
IF(NHRP.NE.NHOUR) GO TO 101
```

C

```
ASSUME UNIFORM ET-DISTRIBUTION
```

```
DEP = FLOAT(NHOUR)/FLOAT(NHRP)
```

```
NHRE=NHRP/NHOUR
```

```
DEP=1.0/FLOAT(NHRE)
```

```
DO 100 IHRE = 1,NHRE
```

```
100 EPDIST(IHRE) = DEP
```

C

C

```
STORE INITIAL CARRYOVER
```

```
101 UZTWC1 = UZTWC
```

```
UZFWC1 = UZFWC
```

```
LZTWC1 = LZTWC
```

```
LZFSC1 = LZFSC
```

```
LZFPC1 = LZFPC
```

```
ADIMC1 = ADIMC
```

C

C

```
COMPUTED SOIL-MOISTURE PARAMETERS
```

```
SAVED = RSERV * (LZFPM + LZFSM)
```

```
PCIAR = 1.0 - PCTIM - ADIMP
```

```
PBASE = LZFSM * LZSK + LZFPM * LZPK
```

C

C

```
NOTE: PBASE IS NOT USE IN ACTUAL COMPUTATIONS
```

C

C

```
INITIALIZE SUMS
```

```
SRECHT = 0.0
```

```
SROT = 0.0
```

```
SETT = 0.0
```

```
SPRT = 0.0
```

```
DO 110 I = 1,7
```

```
110 RSUM(I) = 0.0
```

C

C

```
PRINT INITIAL VALUES
```

```
IF (IPRINT.EQ.0) GO TO 120
```

```
WRITE (IOUB,170)
```

```
WRITE (IOUB,180) PXADJ, PEADJ, UZTWM, UZFWM, UZK, PCTIM, ADIMP, RIVA
```

```
WRITE (IOUB,190) PBASE, ZPERC, REXP, LZTWM, LZFSM, LZFPM, LZSK, LZPK, P
```

```
1 FREE, RSERV, SIDE
```

```
WRITE (IOUB,200) UZTWC, UZFWC, LZTWC, LZFSC, LZFPC, ADIMC
```

```
WRITE (IOUB,210)
```

C

C

```
BEGIN TIME LOOP
```

```

C
120 DELT=NHOUR/24.0
    NTIME=0
    DO 150 IHOURL = 1, NDATA
        PXV = PXADJ * RDT(IHOURL)
        NTIME=NTIME+NHOUR
        SPRT = SPRT + PXV
        IF (NHRP.NE.NHOUR) GO TO 130
        EP = PEADJ * PDT(IHOURL)
        GO TO 140
130     IHRP=(NTIME-1)/NHRP+1
        IHRE=IHOURL-(IHRP-1)*NHRE
        EP = PEADJ * EPDIST(IHRE) * PDT(IHRP)
140     CALL SMAONE (DELT, PXV, EP, IHOURL, IPRINT)
C
C     WATER BALANCE FOR DESIRED PERIOD
        IF (IPRINT.EQ.0) GO TO 150
        IF (MOD(IHOURL, IWBLNC).NE.0) GO TO 150
        WBAL = (UZTWC + UZFWC + LZTWC + LZFPC + LZFSC - UZTWC1 - UZFWC1
1     - LZTWC1 - LZFPC1 - LZFSC1) * PCJAR + (ADIMC - ADIMC1) * ADIMP
2     + SROT + SRECHT + SETT - SPRT
        WRITE (IOUB, 220) (RSUM(I), I = 1, 7)
        WRITE (IOUB, 230) WBAL
C
C     RESTORE INITIAL CARRYOVER
        UZTWC1 = UZTWC
        UZFWC1 = UZFWC
        LZTWC1 = LZTWC
        LZFSC1 = LZFSC
        LZFPC1 = LZFPC
        ADIMC1 = ADIMC
C
C     REINITIALIZE SUMS
        SRECHT = 0.0
        SROT = 0.0
        SETT = 0.0
        SPRT = 0.0
        DO 112 I = 1, 7
112     RSUM(I) = 0.0
150 CONTINUE
    RETURN
C
C
C
170 FORMAT (1H1/1X, 'SACRAMENTO SOIL-MOISTURE ACCOUNTING OPERATION'/)
180 FORMAT (1X, 40HPARAMETER VALUES - CAPACITIES ARE IN MM., //4X, 6HPX-A
    1DJ, 4X, 6HPE-ADJ, 5X, 5HUZTWM, 5X, 5HUZFWM, 7X, 3HUZK, 5X, 5HPCTIM, 5X, 5HADIM
    2P, 6X, 4HRIVA, /8F10.3//1X, 30HDAILY ET DIST. ASSUMED UNIFORM,/)
190 FORMAT (5X, 5HPBASE, 5X, 5HZPERC, 6X, 4HREXP, 5X, 5HLZTWM, 5X, 5HLZFSM, 5X, 5
    1HLZFPM, 6X, 4HLZSK, 6X, 5HLZPK, 5X, 5HPFREE, 5X, 5HRSERV, 6X, 4HSHIDE, /11F10
    2.3/)
200 FORMAT (/1X, 27H SOIL-MOISTURE CONTENTS (MM), //5X, 5HUZTWC, 5X, 5HUZFWC
    1, 5X, 5HLZTWC, 5X, 5HLZFSC, 5X, 5HLZFPC, 5X, 5HADIMC, /6F10.3/)
210 FORMAT (/1X, 71H DETAILED SOIL-MOISTURE ACCOUNTING OUTPUT
    1     UNITS ARE IN MM., //4X, 3HDAY, 1X, 2HHR, 2X, 5HUZTWC, 2X, 5HUZFWC, 2X, 5
    2HLZTWC, 2X, 5HLZFSC, 2X, 5HLZFPC, 2X, 5HADIMC, 3X, 4HPERC, 4X, 3HIMP, 4X, 3HDI

```

3R, 4X, 3HSUR, 4X, 3HINT, 4X, 3HSUP, 4X, 3HPRI, 2X, 6HTOT-RO, 2X, 6HET-DMD, 1X, 6
 4HACT-ET, 3X, 4HRAIN, /)
 220 FORMAT (/2X, 22HTOTAL CHANNEL INFLOW =, F12.4 /2X, 35HCOMPONENTS OF T
 TOTAL CHANNEL INFLOWS, /3X, 39HRUNOFF FROM PERMANENT IMPERVIOUS AREA
 2 =, F12.4/3X, 39HRUNOFF FROM TEMPORARY IMPERVIOUS AREA =, F12.4/3X, 35
 3HSURFACE RUNOFF WHEN UZFWS IS FULL =, F12.4/3X, 42HINTERFLOW FROM LA
 4TERAL DRAINAGE OF UZFWS =, F12.4/3X, 24HSUPPLEMENTARY BASEFLOW =, F12
 5.4/3X, 18HPRIMARY BASEFLOW =, F12.4)
 230 FORMAT (2X, 41HWATER BALANCE RESIDUAL (IDEALLY EQ 0.0) =, F10.5/
 END
 SUBROUTINE SMAONE (DELT, PXV, EP, IHOOR, IPRINT)

C
 C
 C
 C
 C
 C

THIS SUBROUTINE EXECUTES THE "SAC-SMA " OPERATION FOR ONE TIME
 PERIOD.

SUBROUTINE INITIALLY WRITTEN BY. . .

ERIC ANDERSON - HRL APRIL 1979 VERSION 1

COMMON /CMIOP/ IOUA , IOUB ,
 1 NHOOR , NDATA ,
 2 NTIME, NHRP, NHRQ, NDTQ, INPQ, NBASIN, IWBLNC, IPFRQR, IPFRQQ
 COMMON /CMPRM/ NPARM , NPOP , IOBF ,
 1 MXITER , NSTEP , ERROR , IPROP ,
 2 INDX(50) , XMP(50) , XMIN(50) , XMAX(50) ,
 3 ESS(50), OBJFV
 COMMON /FDK00/ TAREA, TRMN, TRHR, METRC, IEL, NQ, MXNDX, MXNDT
 COMMON /FDK01/ CHLNG(5), SLOPE(5), RCMAN(5), PAREA(5), ISHAPE(5),
 1 CHWDT(5), ZLNG(5), ALPME(5), EMDEQ(5), DXROF(5), DXFLO(5), DTKWR(5),
 2 DXKWR(5), IRUPF(5), SAREA(5), FLOIC(5)
 COMMON /FDK02/ EXCSR(1224), Q(1224),
 1 QK(1224), QUB(1224), QBF(1224)
 COMMON /CMDAT/ RDT(1224), PDT(1224), QDT(1224)
 REAL LZTWM , LZFSM , LZFPM ,
 1 LZSK , LZPK , LZTWC , LZFSC ,
 2 LZFPC
 COMMON /CMSMP/ PXADJ , PEADJ , UZTWM ,
 1 UZFWM , UZK , PCTIM , ADIMP ,
 2 RIVA , RSERV , ZPERC , REXP ,
 3 LZTWM , LZFSM , LZFPM , LZSK ,
 4 LZPK , PFREE , SIDE , SAVED ,
 5 PCJAR
 COMMON /CMSMC/ UZTWC , UZFWC , LZTWC ,
 1 LZFSC , LZFPC , ADIMC , RSUM(7)
 COMMON /CMSMS/ SROT, SRECHT, SETT

C
 C
 C
 C
 C
 C
 C

COMPUTE EVAPOTRANSPIRATION LOSS FOR THE TIME INTERVAL.

EDMND IS THE ET-DEMAND FOR THE TIME INTERVAL

EDMND = EP

COMPUTE ET FROM UPPER ZONE.

E1 = EDMND * (UZTWC/UZTWM)

RED = EDMND - E1

RED IS RESIDUAL EVAP DEMAND

UZTWC = UZTWC - E1

E2 = 0.0

IF (UZTWC.GE.0.) GO TO 110

```

C
C   E1 CAN NOT EXCEED UZTWC
    E1 = E1 + UZTWC
    UZTWC = 0.0
    RED = EDMND - E1
    IF (UZFWC.GE.RED) GO TO 100
C
C   E2 IS EVAP FROM UZFWC.
    E2 = UZFWC
    UZFWC = 0.0
    RED = RED - E2
    GO TO 120
100 E2 = RED
    UZFWC = UZFWC - E2
    RED = 0.0
110 IF ((UZTWC/UZTWM).GE.(UZFWC/UZFWM)) GO TO 120
C
C   UPPER ZONE FREE WATER RATIO EXCEEDS UPPER ZONE
C   TENSION WATER RATIO, THUS TRANSFER FREE WATER TO TENSION
    UZRAT = (UZTWC + UZFWC)/(UZTWM + UZFWM)
    UZTWC = UZTWM * UZRAT
    UZFWM = UZFWM * UZRAT
120 IF (UZTWC.LT.0.00001) UZTWC = 0.0
C
C   COMPUTE ET FROM THE LOWER ZONE.
C   COMPUTE ET FROM LZTWC (E3)
    E3 = RED * (LZTWC/(UZTWM + LZTWM))
    LZTWC = LZTWC - E3
    IF (LZTWC.GE.0.0) GO TO 130
C
C   E3 CAN NOT EXCEED LZTWC
    E3 = E3 + LZTWC
    LZTWC = 0.0
130 RATLZT = LZTWC/LZTWM
    RATLZ = (LZTWC + LZFPFC + LZFSC - SAVED)/(LZTWM + LZFPFM + LZFSM - S
LAVED)
    IF (RATLZT.GE.RATLZ) GO TO 140
C
C   RESUPPLY LOWER ZONE TENSION WATER FROM LOWER
C   ZONE FREE WATER IF MORE WATER AVAILABLE THERE.
    DEL = (RATLZ - RATLZT) * LZTWM
C
C   TRANSFER FROM LZFSC TO LZTWC.
    LZTWC = LZTWC + DEL
    LZFSC = LZFSC - DEL
    IF (LZFSC.GE.0.0) GO TO 140
C
C   IF TRANSFER EXCEEDS LZFSC THEN REMAINDER COMES FROM LZFPFC
    LZFPFC = LZFPFC + LZFSC
    LZFSC = 0.0
140 IF (LZTWC.LT.0.00001) LZTWC = 0.0
C
C   COMPUTE ET FROM ADIMP AREA.-E5
    E5 = E1 + (RED + E2) * ((ADIMC - E1 - UZTWC)/(UZTWM + LZTWM))
C
C   ADJUST ADIMC, ADDITIONAL IMPERVIOUS AREA STORAGE, FOR EVAPORATION.

```



```

ADIMC = ADIMC - E5
IF (ADIMC.GE.0.0) GO TO 150
C
C
E5 CAN NOT EXCEED ADIMC.
E5 = E5 + ADIMC
ADIMC = 0.0
150 E5 = E5 * ADIMP
C
C
E5 IS ET FROM THE AREA ADIMP.
C
C
COMPUTE PERCOLATION AND RUNOFF AMOUNTS.
TWX = PXV + UZTWC - UZTWM
C
C
TWX IS THE TIME INTERVAL AVAILABLE MOISTURE IN EXCESS
OF UZTW REQUIREMENTS.
IF (TWX.GE.0.0) GO TO 160
C
C
ALL MOISTURE HELD IN UZTW--NO EXCESS.
UZTWC = UZTWC + PXV
TWX = 0.0
GO TO 170
C
C
MOISTURE AVAILABLE IN EXCESS OF UZTW STORAGE.
160 UZTWC = UZTWM
170 ADIMC = ADIMC + PXV - TWX
C
C
COMPUTE IMPERVIOUS AREA RUNOFF.
ROIMP = PXV * PCTIM
C
C
ROIMP IS RUNOFF FROM THE MINIMUM IMPERVIOUS AREA.
C
C
INITIALIZE TIME INTERVAL SUMS.
SBF = 0.0
SSJR = 0.0
SIF = 0.0
SPERC = 0.0
SDRO = 0.0
SPBF = 0.0
C
C
DETERMINE COMPUTATIONAL TIME INCREMENTS FOR THE BASIC TIME
INTERVAL
NINC = 1.0 + 0.2 * (UZFWC + TWX)
C
C
NINC=NUMBER OF TIME INCREMENTS THAT THE TIME INTERVAL
IS DIVIDED INTO FOR FURTHER
C
C
SOIL-MOISTURE ACCOUNTING. NO ONE INCREMENT
C
C
WILL EXCEED 5.0 MILLIMETERS OF UZFWC+PAV
DINC = (1.0/NINC) * DELT
C
C
DINC=LENGTH OF EACH INCREMENT IN DAYS.
PINC = TWX/NINC
C
C
PINC=AMOUNT OF AVAILABLE MOISTURE FOR EACH INCREMENT.
C
C
COMPUTE FREE WATER DEPLETION FRACTIONS FOR
C
C
THE TIME INCREMENT BEING USED--BASIC DEPLETIONS
C
C
ARE FOR ONE DAY

```

```

DUZ = 1.0 - ((1.0 - UZK) * * DINC)
DLZP = 1.0 - ((1.0 - LZPK) * * DINC)
DLZS = 1.0 - ((1.0 - LZSK) * * DINC)
C
C START INCREMENTAL DO LOOP FOR THE TIME INTERVAL.
DO 300 I = 1, NINC
  ADSUR = 0.0
C
C COMPUTE DIRECT RUNOFF (FROM ADIMP AREA).
  RATIO = (ADIMC - UZTWC)/LZTWM
  IF (RATIO.LT.0.0) RATIO = 0.0
  ADDRO = PINC * (RATIO * * 2)
C
C ADDRO IS THE AMOUNT OF DIRECT RUNOFF FROM THE AREA ADIMP.
C
C COMPUTE BASEFLOW AND KEEP TRACK OF TIME INTERVAL SUM.
  BF = LZFC * DLZP
  LZFC = LZFC - BF
  IF (LZFC.GT.0.0001) GO TO 180
  BF = BF + LZFC
  LZFC = 0.0
180  SBF = SBF + BF
  SPBF = SPBF + BF
  BF = LZFC * DLZS
  LZFC = LZFC - BF
  IF (LZFC.GT.0.0001) GO TO 190
  BF = BF + LZFC
  LZFC = 0.0
190  SBF = SBF + BF
C
C COMPUTE PERCOLATION-IF NO WATER AVAILABLE THEN SKIP
  IF ((PINC + UZFC).GT.0.01) GO TO 200
  UZFC = UZFC + PINC
  GO TO 280
200  PERCM = LZFC * DLZP + LZFSM * DLZS
  PERC = PERCM * (UZFC/UZFM)
  DEFR = 1.0 - ((LZTWC + LZFC + LZFC)/(LZTWM + LZFC + LZFSM))
C
C DEFR IS THE LOWER ZONE MOISTURE DEFICIENCY RATIO
  PERC = PERC * (1.0 + ZPERC * (DEFR * * REXP))
C NOTE...PERCOLATION OCCURS FROM UZFC BEFORE PAV IS ADDED.
  IF (PERC.LT.UZFC) GO TO 210
C
C PERCOLATION RATE EXCEEDS UZFC.
  PERC = UZFC
C
C PERCOLATION RATE IS LESS THAN UZFC.
210  UZFC = UZFC - PERC
C
C CHECK TO SEE IF PERCOLATION EXCEEDS LOWER ZONE DEFICIENCY.
  CHECK = LZTWC + LZFC + LZFC + PERC - LZTWM - LZFC - LZFSM
  IF (CHECK.LE.0.0) GO TO 220
  PERC = PERC - CHECK
  UZFC = UZFC + CHECK
220  SPERC = SPERC + PERC
C

```


C SPERC IS THE TIME INTERVAL SUMMATION OF PERC
 C
 C COMPUTE INTERFLOW AND KEEP TRACK OF TIME INTERVAL SUM.
 C NOTE...PINC HAS NOT YET BEEN ADDED
 C DEL = UZFWC * DUZ
 C SIF = SIF + DEL
 C UZFWC = UZFWC - DEL
 C
 C DISTRIBUTE PERCOLATED WATER INTO THE LOWER ZONES
 C TENSION WATER MUST BE FILLED FIRST EXCEPT FOR THE PFREE AREA.
 C PERCT IS PERCOLATION TO TENSION WATER AND PERCF IS PERCOLATION
 C GOING TO FREE WATER.
 C PERCT = PERC * (1.0 - PFREE)
 C IF ((PERCT + LZTWC).GT.LZTWM) GO TO 230
 C LZTWC = LZTWC + PERCT
 C PERCF = 0.0
 C GO TO 240
 230 PERCF = PERCT + LZTWC - LZTWM
 C LZTWC = LZTWM
 C
 C DISTRIBUTE PERCOLATION IN EXCESS OF TENSION
 C REQUIREMENTS AMONG THE FREE WATER STORAGE.
 C 240 PERCF = PERCF + PERC * PFREE
 C IF (PERCF.EQ.0.0) GO TO 260
 C HPL = LZFPF/(LZFPF + LZFSM)
 C
 C HPL IS THE RELATIVE SIZE OF THE PRIMARY STORAGE
 C AS COMPARED WITH TOTAL LOWER ZONE FREE WATER STORAGE.
 C RATLP = LZFPF/LZFPF
 C RATLS = LZFSF/LZFSM
 C
 C RATLP AND RATLS ARE CONTENT TO CAPACITY RATIOS, OR
 C IN OTHER WORDS, THE RELATIVE FULLNESS OF EACH STORAGE
 C FRACP = (HPL * 2.0 * (1.0 - RATLP))/((1.0 - RATLP) + (1.0 - RAT
 1 LS))
 C
 C FRACP IS THE FRACTION GOING TO PRIMARY.
 C IF (FRACP.GT.1.0) FRACP = 1.0
 C PERCP = PERCF * FRACP
 C PERCS = PERCF - PERCP
 C
 C PERCP AND PERCS ARE THE AMOUNT OF THE EXCESS
 C PERCOLATION GOING TO PRIMARY AND SUPPLEMENTAL
 C STORGES, RESPECTIVELY.
 C LZFSF = LZFSF + PERCS
 C IF (LZFSF.LE.LZFSM) GO TO 250
 C PERCS = PERCS - LZFSF + LZFSM
 C LZFSF = LZFSM
 250 LZFPF = LZFPF + (PERCF - PERCS)
 C
 C CHECK TO MAKE SURE LZFPF DOES NOT EXCEED LZFPF.
 C IF (LZFPF.LE.LZFPF) GO TO 260
 C EXCESS = LZFPF - LZFPF
 C LZTWC = LZTWC + EXCESS
 C LZFPF = LZFPF
 C


```

C   DISTRIBUTE PINC BETWEEN UZFWC AND SURFACE RUNOFF.
260  IF (PINC.EQ.0.0) GO TO 280
C
C   CHECK IF PINC EXCEEDS UZFWM
      IF ((PINC + UZFWC).GT.UZFWM) GO TO 270
C
C   NO SURFACE RUNOFF
      UZFWC = UZFWC + PINC
      GO TO 280
C
C   COMPUTE SURFACE RUNOFF (SUR) AND KEEP TRACK OF TIME INTERVAL SUM.
270  SUR = PINC + UZFWC - UZFWM
      UZFWC = UZFWM
      SSUR = SSUR + SUR * PCIAR
      ADSUR = SUR * (1.0 - ADDRO/PINC)
C
C   ADSUR IS THE AMOUNT OF SURFACE RUNOFF WHICH COMES
C   FROM THAT PORTION OF ADIMP WHICH IS NOT
C   CURRENTLY GENERATING DIRECT RUNOFF.  ADDRO/PINC
C   IS THE FRACTION OF ADIMP CURRENTLY GENERATING
C   DIRECT RUNOFF.
      SSUR = SSUR + ADSUR * ADIMP
C
C   ADIMP AREA WATER BALANCE — SDRO IS THE 6 HR SUM OF
C   DIRECT RUNOFF.
280  ADIMC = ADIMC + PINC - ADDRO - ADSUR
      IF (ADIMC.LE.(UZTWM + LZTWM)) GO TO 290
      ADDRO = ADDRO + ADIMC - (UZTWM + LZTWM)
      ADIMC = UZTWM + LZTWM
290  SDRO = SDRO + ADDRO * ADIMP
      IF (ADIMC.LT.0.00001) ADIMC = 0.0
300 CONTINUE
C
C   END OF INCREMENTAL DO LOOP.
C
C   COMPUTE SUMS AND ADJUST RUNOFF AMOUNTS BY THE AREA OVER
C   WHICH THEY ARE GENERATED.
      EUSED = E1 + E2 + E3
C
C   EUSED IS THE ET FROM PCIAR WHICH IS 1.0-ADIMP-PCTIM
      SIF = SIF * PCIAR
C
C   SEPARATE CHANNEL COMPONENT OF BASEFLOW
C   FROM THE NON-CHANNEL COMPONENT
      TBF = SBF * PCIAR
C
C   TBF IS TOTAL BASEFLOW
      BFCC = TBF * (1.0/(1.0 + SIDE))
C
C   BFCC IS BASEFLOW, CHANNEL COMPONENT
      BFP = SPBF * PCIAR/(1.0 + SIDE)
      BFS = BFCC - BFP
      IF (BFS.LT.0.0) BFS = 0.0
      BFNCC = TBF - BFCC
C
C   BFNCC IS BASEFLOW, NON-CHANNEL COMPONENT

```

```

3  ESS(50),OBJFV
COMMON /FDK00/ TAREA,TRMN,TRHR,METRC,IEL,NQ,MXNDX,MXNDT
COMMON /FDK01/ CHLNG(5),SLOPE(5),RCMAN(5),PAREA(5),ISHAPE(5),
1 CHWDT(5),ZLNG(5),ALPME(5),EMDEQ(5),DXROF(5),DXFLO(5),DTKWR(5),
2 DXKWR(5),IRUPF(5),SAREA(5),FLOIC(5)
COMMON /FDK02/ EXCSR(1224),Q(1224),
1 QK(1224),QUB(1224),QBF(1224)
COMMON /CMDAT/ RDT(1224),PDT(1224),QDT(1224)
REAL
1 LZSK , LZTWM , LZFSM , LZFPM ,
2 LZPK , LZTWC , LZFSC ,
COMMON /CMSMP/ PXADJ , PEADJ , UZTWM ,
1 UZFWM , UZK , PCTIM , ADIMP ,
2 RIVA , RSERV , ZPERC , REXP ,
3 LZTWM , LZFSM , LZFPM , LZSK ,
4 LZPK , PFREE , SIDE , SAVED ,
5 PCIAR
COMMON /CMSMC/ UZTWC , UZFWC , LZTWC ,
1 LZFSC , LZFPC , ADIMC , RSUM(7)
DIMENSION
1 PH(50) , EINT(50) , SA(50) , B(50,50) ,
2 V0(50,50) , V1(50,50)

```

C

WRITE (IOUB,500)

C

DEFINITION OF SOME TERMS

C

M = -1 SIGNIFIES MINIMIZE OBJECTIVE FUNCTION

C

NP = NUMBER OF PARAMETERS TO BE OPTIMIZED

C

NC = NUMBER OF CONIRAINIS

C

MXITER = MAXIMUM NUMBER OF ITERATIONS

C

ERROR = CONVERGENCE CRITERION SUCH THAT ABS((F1-F0)/F1) LT ERROR

C

IPROP = PRINTING FREQUENCY BETWEEN STAGE EVALUATION

M = - 1

NP = NPOP

NC = NP

C

IF (MXITER.EQ.0) MXITER = 10

ERROR = AMAX1(ERROR,0.000001)

IPROP = MAX0(IPROP,1)

C

C

INITIALIZE COUNTERS

KIPR = IPROP - 1

KITER = 0

ISW = 0

INIT = 0

KOUNT = 0

ICNVG = 0

F1 = 0.0

DO 100 K = 1,NC

100 AL(K) = (XMAX(K) - XMIN(K)) * 0.0001

DO 120 I = 1,NP

DO 110 J = 1,NP

110 V0(I,J) = 0.0

120 V0(I,I) = 1.0

DO 130 K = 1,NP

130 EINT(K) = ESS(K)

C

```

140 DO 150 J = 1, NP
      IF (NSTEP.EQ.0) ESS(J) = EINT(J)
      SA(J) = 2.0
150 D(J) = 0.0
      FBEST = F1
160 I = 1
      IF (INIT.EQ.0) GO TO 200
170 DO 180 K = 1, NP
180 XMP(K) = XMP(K) + ESS(I) * V0(I, K)
      DO 190 K = 1, NC
190 H(K) = F0
C
200 KOUNT = KOUNT + 1
      CALL SACROUT (0, IPRINT)
      F1 = M * OBJFV
      IF (ISW.EQ.0) F0 = F1
      ISW = 1
      IF (ABS((FBEST - F1)/F1).GT.ERROR) GO TO 210
      ICNVG = 1
      GO TO 460
C
210 J = 1
220 XV = XMP(J)
      XL = XMIN(J)
      XU = XMAX(J)
      IF (XV.LE.XL) GO TO 430
      IF (XV.GE.XU) GO TO 430
      IF (F1.LT.F0) GO TO 430
      IF (XV.LT.(XL + AL(J))) GO TO 230
      IF (XV.GT.(XU - AL(J))) GO TO 230
      H(J) = F0
      GO TO 290
C
230 BW = AL(J)
      IF (XV.LE.XL.OR.XU.LE.XV) GO TO 240
      IF (XL.LT.XV.AND.XV.LT.(XL + BW)) GO TO 250
      IF ((XU - BW).LT.XV.AND.XV.LT.XU) GO TO 260
      PH(J) = 1.0
      GO TO 290
C
240 PH(J) = 0.0
      GO TO 280
250 PW = (XL + BW - XV)/BW
      GO TO 270
260 PW = (XV - XU + BW)/BW
270 PH(J) = 1.0 - 3.0 * PW + 4.0 * PW * PW - 2.0 * PW * PW * PW
280 F1 = H(J) + (F1 - H(J)) * PH(J)
C
290 IF (J.EQ.NC) GO TO 300
      J = J + 1
      GO TO 220
C
300 INIT = 1
      IF (F1.LT.F0) GO TO 430
      D(I) = D(I) + ESS(I)
C

```

```

C   NOTE : SCALING FACTOR FOR STEP SIZE INCREASE ALPHA IS SET AT,
      ALPHA = 2.0
      ESS(I) = ALPHA * ESS(I)
      F0 = F1
      IF (SA(I).GE.1.5) SA(I) = 1.0
      GO TO 310
C
430 IF (INIT.EQ.0) GO TO 460
      DO 440 IC = 1,NP
440 XMP(IC) = XMP(IC) - ESS(I) * V0(I,IC)
C
C   NOTE : SCALING FACTOR FOR STEP SIZE REDUCTION BETA IS SET AT,
      BETA = 0.5
      ESS(I) = - BETA * ESS(I)
      IF (SA(I).LT.1.5) SA(I) = 0.0
C
310 DO 320 K = 1,NP
      IF (SA(K).GE.0.5) GO TO 450
320 CONTINUE
C
C   AXES ROTATION
C
      DO 330 IR = 1,NP
      DO 330 IC = 1,NP
330 V1(IR,IC) = 0.0
      DO 350 IR = 1,NP
      DO 350 IC = 1,NP
      DO 340 K = IR,NP
340 V1(IR,IC) = V1(IR,IC) + D(K) * V0(K,IC)
350 B(IR,IC) = V1(IR,IC)
      SBM1 = 0.0
      DO 360 IC = 1,NP
      SBM1 = SBM1 + B(1,IC) * B(1,IC)
360 CONTINUE
      SBM1 = SQRT(SBM1)
      DO 370 IC = 1,NP
370 V0(1,IC) = B(1,IC)/SBM1
      DO 400 IR = 2,NP
      IRI = IR - 1
      DO 400 IC = 1,NP
      SUMB = 0.0
      DO 390 KR = 1,IRI
      SUMA = 0.0
      DO 380 KC = 1,NP
380 SUMA = SUMA + V1(IR,KC) * V0(KR,KC)
390 SUMB = SUMB + SUMA * V0(KR,IC)
400 B(IR,IC) = V1(IR,IC) - SUMB
      DO 420 IR = 2,NP
      SBM2 = 0.0
      DO 410 K = 1,NP
410 SBM2 = SBM2 + B(IR,K) * B(IR,K)
      SBM2 = SQRT(SBM2)
      DO 420 IC = 1,NP
420 V0(IR,IC) = B(IR,IC)/SBM2
      KITER = KITER + 1
      KIPR = KIPR + 1

```

```

IF (KIPR.EQ.IPROP) GO TO 460
GO TO 140
C
450 IF (I.EQ.NP) GO TO 160
I = I + 1
GO TO 170
C
460 WRITE (IOUB,510)
WRITE (IOUB,520) KITER,F0,SBM1,SBM2
WRITE (IOUB,530) KOUNT
WRITE (IOUB,540)
C
PRINT CURRENT VALUES OF XMP
C
WRITE (IOUB,550) (K,XMP(K),K = 1,NP)
C
KIPR = 0
IF (INIT.EQ.0) GO TO 470
IF (ICNVG.EQ.1) GO TO 480
IF (KITER.GE.MXITER) GO TO 480
GO TO 140
C
470 WRITE (IOUB,560)
480 CONTINUE
WRITE (IOUB,570)
DO 490 J = 1,NP
490 WRITE (IOUB,580) (J,I,V0(J,I),I = 1,NP)
WRITE (IOUB,590)
WRITE (IOUB,600) (J,ESS(J),J = 1,NP)
RETURN
C
C
C
500 FORMAT (1HL/10X,46HOPTIMIZATION BY ROSENBROCK HILLCLIMB PROCEDURE,
1/)
510 FORMAT (/2X,5HSTAGE,8X,8HFUNCTION,12X,8HPROGRESS,9X,16HLATERAL PR
LOGRESS)
520 FORMAT (1H ,I5,3E20.8)
530 FORMAT (/2X,33HNUMBER OF FUNCTION EVALUATIONS = ,I8)
540 FORMAT (/2X,28HVALUES OF X(.) AT THIS STAGE)
550 FORMAT (3(4X,2HX(,I2,4H) = ,E15.6))
560 FORMAT (/2X,52HTHE STARTING POINTS MUST NOT VIOLATE THE CONSTRAINT
1S,/2X,26HIT APPEARS TO HAVE DONE SO,/)
570 FORMAT (/2X,29HFINAL DIRECTION VECTOR MATRIX)
580 FORMAT (3(4X,2HV(,I2,1H,,I2,4H) = ,F10.8))
590 FORMAT (/2X,16HFINAL STEP SIZES)
600 FORMAT (3(4X,2HS(,I2,4H) = ,E15.6))
END
FUNCTION FPOBJ(QOBS,QCOM)
C
C
C
ROUTINE TO EVALUATE EACH SUMMATION TERM OF OBJECTIVE FUNCTION
COMMON /CMPRM/
1 MXITER , NPARM , NPOP , IOBF ,
2 INDX(50) , NSTEP , ERROR , IPROP ,
3 ESS(50),OBJFV , XMP(50) , XMIN(50) , XMAX(50) ,

```

```

      GO TO (100,110,120,130,140), IOBF
100  FPOBJ = (QOBS - QCOM) * * 2
      RETURN
110  FPOBJ = ABS(QOBS - QCOM)
      RETURN
120  IF (QOBS.EQ.0.0) QOBS = 0.0000001
      FPOBJ = ((QOBS - QCOM)/QOBS) * * 2
      RETURN
130  FPOBJ = (QOBS * * 0.3333333 - QCOM * * 0.3333333) * * 2
      RETURN
140  FPOBJ = (QOBS * * 0.5 - QCOM * * 0.5) * * 2
      RETURN
      END
      SUBROUTINE SACROUT (IUPDF,IPRINT)

```

C
C
C
C

THIS ROUTINE CONTROLS THE TIMING SEQUENCE FOR SACRAMENTO SOIL-
MOISTURE ACCOUNTING AND STREAMFLOW ROUTING

```

COMMON /CMIOP/      IOUA      , IOUB      ,
1      NHOUR      , NDATA      ,
2      NTIME, NHRP, NHRQ, NDTQ, INPQ, NBASIN, IWBLNC, IPFRQR, IPFRQQ
COMMON /CMPRM/      NPARM      , NPOP      , IOBF      ,
1      MXITER      , NSTEP      , ERROR      , IPROP      ,
2      INDX(50)      , XMP(50)      , XMIN(50)      , XMAX(50)      ,
3      ESS(50), OBJFV
COMMON /FDK00/ TAREA, TRMN, TRHR, METRC, IEL, NQ, MXNDX, MXNDT
COMMON /FDK01/ CHLNG(5), SLOPE(5), RCMAN(5), PAREA(5), ISHAPE(5),
1 CHWDT(5), ZLNG(5), ALPME(5), EMDEQ(5), DXROF(5), DXFLO(5), DTKWR(5),
2 DXKWR(5), IRUPF(5), SAREA(5), FLOIC(5)
COMMON /FDK02/ EXCSR(1224), Q(1224),
1 QK(1224), QUB(1224), QBF(1224)
COMMON /CMDAT/ RDT(1224), PDT(1224), QDT(1224)
REAL
1      LZSK      , LZPK      , LZFSM      , LZFPM      ,
2      LZFPC
COMMON /CMSMP/      PXADJ      , PEADJ      , UZTWM      ,
1      UZFWM      , UZK      , PCTIM      , ADIMP      ,
2      RIVA      , RSERV      , ZPERC      , REXP      ,
3      LZTWM      , LZFSM      , LZFPM      , LZSK      ,
4      LZPK      , PFREE      , SIDE      , SAVED      ,
5      PCIAR
COMMON /CMSMC/      UZTWC      , UZFWC      , LZTWC      ,
1      LZFSC      , LZFPC      , ADIMC      , RSUM(7)
DIMENSION      QPLOT(81)
DATA CHAR0, CHAR1, CHAR2, CHAR3/1H ,1HC,1HO,1HE/
DO 400 IP = 1, NPARM
      KP = INDX(IP)
      GO TO (100,110,120,130,140,150,160,170,180,190,200,210,220,230,
1 240,250,260,270,280,290,300,310,320,330), KP
100  PXADJ = XMP(IP)
      GO TO 400
110  PEADJ = XMP(IP)
      GO TO 400
120  PCTIM = XMP(IP)
      GO TO 400
130  ADIMP = XMP(IP)

```



```
ADIMP = AMINL (ADIMP, (1.0 - PCTIM))
IF (IP.GT.NPOP) GO TO 400
XMAX(IP) = AMINL (XMAX(IP), (1.0 - PCTIM))
GO TO 400
140 RIVA = XMP(IP)
IF (IP.GT.NPOP) GO TO 400
RIVA = AMAXL (RIVA, XMIN(IP))
RIVA = AMINL (RIVA, XMAX(IP))
GO TO 400
150 UZK = XMP(IP)
IF (IP.GT.NPOP) GO TO 400
UZK = AMAXL (UZK, XMIN(IP))
UZK = AMINL (UZK, XMAX(IP))
GO TO 400
160 LZSK = XMP(IP)
IF (IP.GT.NPOP) GO TO 400
LZSK = AMAXL (LZSK, XMIN(IP))
LZSK = AMINL (LZSK, XMAX(IP))
GO TO 400
170 LZPK = XMP(IP)
IF (IP.GT.NPOP) GO TO 400
LZPK = AMAXL (LZPK, XMIN(IP))
LZPK = AMINL (LZPK, XMAX(IP))
GO TO 400
180 PFREE = XMP(IP)
IF (IP.GT.NPOP) GO TO 400
PFREE = AMAXL (PFREE, XMIN(IP))
PFREE = AMINL (PFREE, XMAX(IP))
GO TO 400
190 RSERV = XMP(IP)
IF (IP.GT.NPOP) GO TO 400
RSERV = AMAXL (RSERV, XMIN(IP))
RSERV = AMINL (RSERV, XMAX(IP))
GO TO 400
200 ZPERC = XMP(IP)
GO TO 400
210 REXP = XMP(IP)
GO TO 400
220 SIDE = XMP(IP)
GO TO 400
230 UZTWM = XMP(IP)
GO TO 400
240 UZFWM = XMP(IP)
GO TO 400
250 LZTWM = XMP(IP)
GO TO 400
260 LZFSM = XMP(IP)
GO TO 400
270 LZFPM = XMP(IP)
GO TO 400
280 UZTWC = XMP(IP)
IF (IP.GT.NPOP) GO TO 400
XMAX(IP) = AMINL (XMAX(IP), UZTWM)
GO TO 400
290 LZTWC = XMP(IP)
IF (IP.GT.NPOP) GO TO 400
```

```

XMAX(IP) = AMINI(XMAX(IP),LZTWM)
GO TO 400
300 UZFWC = XMP(IP)
IF (IP.GT.NPOP) GO TO 400
XMAX(IP) = AMINI(XMAX(IP),UZFWM)
GO TO 400
310 LZFSC = XMP(IP)
IF (IP.GT.NPOP) GO TO 400
XMAX(IP) = AMINI(XMAX(IP),LZFSC)
GO TO 400
320 LZFPC = XMP(IP)
IF (IP.GT.NPOP) GO TO 400
XMAX(IP) = AMINI(XMAX(IP),LZFPC)
GO TO 400
330 ADIMC = XMP(IP)
IF (IP.GT.NPOP) GO TO 400
XMIN(IP) = AMAXI(XMIN(IP),UZTWC)
XMAX(IP) = AMINI(XMAX(IP),(UZTWC + LZTWM))
400 CONTINUE
CALL SACSMA (IPRINT)
IF (IUPDF.EQ.0) GO TO 480
DO 470 IP = 1,NPARM
  KP = INDX(IP)
  GO TO (470,470,470,470,470,470,470,470,470,470,470,470,470,470,470,
1 470,470,470,470,410,420,430,440,450,460), KP
410 XMP(IP) = UZTWC
GO TO 470
420 XMP(IP) = LZTWC
GO TO 470
430 XMP(IP) = UZFWC
GO TO 470
440 XMP(IP) = LZFSC
GO TO 470
450 XMP(IP) = LZFPC
GO TO 470
460 XMP(IP) = ADIMC
470 CONTINUE
480 NQ=NDATA
TRHR=NHOUR
CALL KINWAVE (IPRINT)
DO 500 I = 1,81
500 QPLOT(I) = CHAR0
OBJFV = 0.0
QMIN = 9999999.0
QMAX = - QMIN
ANQ=0.0
IF (NHOUR.NE.NHRQ) GO TO 520
DO 510 IHOURL = 1,NDATA
  QOUTI = Q(IHOURL)
  QMIN = AMINI(QMIN,QOUTI)
  QMAX = AMAXI(QMAX,QOUTI)
  IF (INPQ.EQ.0) GO TO 510
  QDTI = QDT(IHOURL)
  IF(QDTI.LT.0.0) GO TO 510
  ANQ=ANQ+1.0
  QMIN = AMINI(QMIN,QDTI)

```

```

      QMAX = AMAX1(QMAX,QDTI)
      OBJFV = OBJFV + FPOBJ(QDTI,QOUTI)
510 CONTINUE
      IF(INPQ.NE.0) OBJFV=OBJFV/ANQ
      GO TO 560
520 NRQ = NHRQ/NHOUR
      DN = NRQ
      IHOURL = 0
      DO 540 IHRQ = 1,NDTQ
          SUM = 0.0
          DO 530 IRQ = 1,NRQ
              IHOURL = IHOURL + 1
530      SUM = SUM + Q(IHOURL)
          QOUTI = SUM/DN
          Q(IHRQ) = QOUTI
          QMIN = AMIN1(QMIN,QOUTI)
          QMAX = AMAX1(QMAX,QOUTI)
          IF (INPQ.EQ.0) GO TO 540
          QDTI = QDT(IHRQ)
          IF(QDTI.LT.0.0) GO TO 540
          ANQ=ANQ+1.0
          QMIN = AMIN1(QMIN,QDTI)
          QMAX = AMAX1(QMAX,QDTI)
          OBJFV = OBJFV + FPOBJ(QDTI,QOUTI)
540 CONTINUE
      IF(INPQ.NE.0) OBJFV=OBJFV/ANQ
560 IF (IPRINT.EQ.0) RETURN
      IF (INPQ.NE.0) WRITE (IOUB,650) IOBF,OBJFV
      IF (INPQ.EQ.0) WRITE (IOUB,660)
      QCON = 80.0/(QMAX - QMIN)
      NIMQ=0
      DO 610 IHRQ = 1,NDTQ
          NIMQ=NIMQ+NHRQ
          IDAY=(NIMQ-1)/24+1
          QOUTI = Q(IHRQ)
          NQO = 1 + IFIX((QOUTI - QMIN) * QCON)
          QPLOT(NQO) = CHAR1
          NQD = NQO
          IF (INPQ.EQ.0) GO TO 600
          QDTI = QDT(IHRQ)
          IF (QDTI.LT.0.0) GO TO 570
          NQD = 1 + IFIX((QDTI - QMIN) * QCON)
          IF (NQO.NE.NQD) QPLOT(NQD) = CHAR2
          IF (NQO.EQ.NQD) QPLOT(NQD) = CHAR3
570 IF(MOD(NIMQ,IPFRQO).EQ.0)
      1 WRITE (IOUB,670) IDAY,NIMQ,QDTI,QOUTI,(QPLOT(I),I = 1,81)
          QPLOT(NQO) = CHAR0
          QPLOT(NQD) = CHAR0
          GO TO 610
600 IF(MOD(NIMQ,IPFRQO).EQ.0)
      1 WRITE (IOUB,680) IDAY,NIMQ,QOUTI,(QPLOT(I),I = 1,81)
          QPLOT(NQO) = CHAR0
610 CONTINUE
      WRITE (IOUB,690) QMIN,QMAX
      WRITE (IOUB,700) CHAR1,CHAR2,CHAR3
      RETURN

```

C
C

```

650 FORMAT (/5X,29HVALUE OF OBJECTIVE FUNCTION (,11,3H) =,F15.5//4X, 3
      12HDAY HOUR OBSERVED COMPUTED I,81( 1H-), 1HI)
660 FORMAT (//4X, 32HDAY HOUR OBSERVED COMPUTED I,81( 1H-), 1HI)
670 FORMAT (2X,I4,I6,1X,2F10.3,2X, 1HI,81A1, 1HI)
680 FORMAT (2X,I4,I6,7X, 1H*,3X,F10.3,2X, 1HI,81A1, 1HI)
690 FORMAT (35X, 1HI,81( 1H-), 1HI,//5X, 18HMINIMUM ORDINATE =,F10.
      12/5X, 18HMAXIMUM ORDINATE =,F10.2)
700 FORMAT (/5X, 13HSYMBOLS USED:,/6X, 11HCOMPUTED - ,A1/6X, 11HOBSERV
      LED - ,A1/6X, 11HIF EQUAL - ,A1///1X, 18HNORMAL TERMINATION,/)
      END
      SUBROUTINE KINWAVE (IPRINT)

```

C
C
C
C

THIS SUBROUTINE CONTROL THE TIMING AND SEQUENCING OF
KINEMATIC WAVE ROUTING METHODOLOGY

```

COMMON /CMIOP/          IOUA          , IOUB          ,
1      NHOUR            ,NDATA          ,
2      NTIME,NHRP,NHRQ,NDTQ,INPQ,NBASIN,IWBLNC,IPFRQR,IPFRQQ
COMMON /FDK00/ TAREA,TRMN,TRHR,MEIRC,IEL,NQ,MXNDX,MXNDT
COMMON /FDK01/ CHLNG(5),SLOPE(5),RCMAN(5),PAREA(5),ISHAPE(5),
1 CHWDT(5),ZLNG(5),ALPME(5),EMDEQ(5),DXROF(5),DXFLO(5),DTKWR(5),
2 DXKWR(5),IRUPF(5),SAREA(5),FLOIC(5)
COMMON /FDK02/ EXCSR(1224),Q(1224),
1 QK(1224),QUB(1224),QBF(1224)
DIMENSION DOPER(4),QS(3672),WBF(50)
DATA DOPER/5HROUTE,5HADD ,5HBASEF,5HEND /
DATA MEIRC/1/,MXNDX/51/,MXNDT/1224/

```

C

```
IF(IPRINT.NE.0) WRITE(IOUB,99)
```

C

```
CFQ=1.0
```

C

C

C

```
CONVERSION FROM SQ MI-IN TO CFS
```

```
CATOQ=645.333/TRHR
```

```
IF (MEIRC.LE.0) GO TO 01
```

```
CFQ=35.31
```

C

C

C

```
CONVERSION FROM SQ KM-MM TO CUBIC METERS PER SECOND
```

```
CATOQ=1.0/(3.6*TRHR)
```

C

```
01 TRMN=TRHR*60.0
```

```
REWIND 10
```

```
NSQ=0
```

```
NSB=0
```

```
45 READ(10,100,END=90) ROPER,IOPER
```

```
DO 05 IOP=1,4
```

```
IF(DOPER(IOP).EQ.ROPER) GO TO 10
```

```
05 CONTINUE
```

```
WRITE(IOUB,105) ROPER
```

```
STOP
```

```
10 GO TO (25,50,70,90), IOP
```

C

C KINEMATIC WAVE ROUTING
C

```

25 DO 30 I=1,5
   CHLNG(I)=0.0
   SLOPE(I)=0.0
   RCMAN(I)=0.0
   PAREA(I)=0.0
   ISHAPE(I)=0
   CHWDT(I)=0.0
   ZLNG(I)=0.0
   ALPME(I)=0.0
   EMDEQ(I)=0.0
   DXROF(I)=0.0
   DXFLO(I)=0.0
   DTKWR(I)=0.0
   DXKWR(I)=0.0
   FLOIC(I)=0.0
   SAREA(I)=0.0
   IRUPF(I)=0
30 CONTINUE
   NSB=NSB+1
   SBA=0.0
   IEL=2
   READ(10,110) TAREA
   READ(10,110) CHLNG(1),SLOPE(1),RCMAN(1),PAREA(1)
   IF(IPRINT.EQ.0) GO TO 33
   WRITE(IOUB,115) NSB, TAREA
   IELP=1
   WRITE(IOUB,116) IELP,CHLNG(1),SLOPE(1),RCMAN(1),PAREA(1)
33 IF(PAREA(1).GE.99.5) GO TO 35
   READ(10,110) CHLNG(2),SLOPE(2),RCMAN(2),PAREA(2)
   IF(IPRINT.EQ.0) GO TO 35
   IELP=2
   WRITE(IOUB,116) CHLNG(2),SLOPE(2),RCMAN(2),PAREA(2)
35 IEL=IEL+1
   READ(10,110) CHLNG(IEL),SLOPE(IEL),RCMAN(IEL),SAREA(IEL)
   READ(10,110) ASHAPE,CHWDT(IEL),ZLNG(IEL),FLOIC(IEL),ARUPF
   IF(SAREA(IEL).EQ.0.0) SAREA(IEL)=TAREA
   ISHAPE(IEL)=ASHAPE
   IRUPF(IEL)=ARUPF
   IF(IPRINT.EQ.0) GO TO 36
   IELP=IEL-2
   IF(SAREA(IEL).NE.TAREA) WRITE(IOUB,117) IELP
   IF(SAREA(IEL).EQ.TAREA) WRITE(IOUB,118)
   WRITE(IOUB,119) CHLNG(IEL),SLOPE(IEL),RCMAN(IEL),SAREA(IEL),
1 ISHAPE(IEL),CHWDT(IEL),ZLNG(IEL),FLOIC(IEL)
   IF(IRUPF(IEL).EQ.1) WRITE(IOUB,121)
36 SBA=SBA+SAREA(IEL)
   IF(SBA.LT.TAREA) GO TO 35
   NSQ=NSQ+1
   IF(IRUPF(IEL).NE.1) GO TO 37
   IF(NSQ.GE.2) GO TO 39
   DO 41 N=1,NQ
41 QK(N)=QUB(N)
   GO TO 37
39 ISQ=(NSQ-2)*NQ

```

```

DO 42 N=1,NQ
42 QK(N)=QS(ISQ+N)
37 CALL KINOFF (CFQ,IPRINT)
   ISQ=(NSQ-1)*NQ
   DO 40 N=1,NQ
40 QS(ISQ+N)=Q(N)
   GO TO 45
C
C   ADD PREVIOUS (IOPER) FLOWS
C
50 NS1=NSQ-IOPER+1
   NSB1=NSB-IOPER+1
   IF(IPRINT.NE.0) WRITE(IOUB,120) (ISB,ISB=NSB1,NSB)
   ISQ1=(NS1-1)*NQ
   DO 55 N=1,NQ
   SQA=0.0
   DO 60 IS=NS1,NSQ
   ISQ=(IS-1)*NQ
60 SQA=SQA+QS(ISQ+N)
55 QS(ISQ1+N)=SQA
   NSQ=NS1
   GO TO 45
C
C   ADD BASEFLOW
C
70 IWBF=IOPER+1
   READ(10,110) (WBF(I),I=1,IWBF)
   IF(IPRINT.NE.0) WRITE(IOUB,125) NSB, IOPER, (WBF(I),I=1,IWBF)
   ISQ=(NSQ-1)*NQ
   DO 75 N=1,NQ
   SBF=0.0
   DO 80 I=1,IWBF
   L=I-1
   IF(N.LE.L) GO TO 80
   SBF=SBF+WBF(I)*QBF(N-L)
80 CONTINUE
   SBF=SBF*CATOQ*TAREA+QS(ISQ+N)
75 QS(ISQ+N)=SBF
   GO TO 45
C   STORE FINAL COMPUTED FLOWS IN Q(.)
90 ISQ=(NSQ-1)*NQ
   DO 85 N=1,NQ
   Q(N)=QS(ISQ+N)
   QK(N)=Q(N)
   IF(IPRINT.NE.0) QUB(N)=Q(N)
85 CONTINUE
   RETURN
99 FORMAT(1H1)
100 FORMAT(A5,I5)
105 FORMAT(/3X,'***** EXECUTION TERMINATED - INPUT LIST ',A5,' IS AN
   LUNRECOGNIZED OPERATION'/3X,'IN KINEMATIC WAVE ROUTING *****')
115 FORMAT(/1X,'KINEMATIC WAVE ROUTING FOR SUBBASIN',I3/
   1 /2X,'TOTAL AREA =',F10.2)
116 FORMAT(/1X,'OVERLAND FLOW ELEMENT',I2//2X,'CHLNG =',F10.2/
   1 2X,'SLOPE =',F10.5/2X,'RCMAN =',F10.5/2X,'PAREA =',F10.3/)
117 FORMAT(/1X,'COLLECTOR CHANNEL',I3)

```

```

118 FORMAT(/1X,'MAIN CHANNEL')
119 FORMAT(/2X,'CHLNG =' ,F10.2/2X,'SLOPE =' ,F10.5/2X,'RCMAN =' ,
1 F10.5/2X,'SAREA =' ,F10.3//2X,'ISHAPE=' ,I10/2X,'CHWDT. =' ,
2 F10.3/2X,'ZLNG =' ,F10.3//2X,'FLOIC =' ,F10.3)
121 FORMAT(/1X,'ROUTE UPSTREAM FLOW'/)
110 FORMAT(5G10.4)
120 FORMAT(/1X,'ADD FLOWS OF SUBBASINS ' ,I3,10(' ,',I3))
125 FORMAT(/1X,'ADD BASEFLOW TO SUBBASIN' ,I3,' USING A LINEAR DECAY FU
UNCTION OF THE FORM: '//2X,'ADDED BF AT TIME T = SUM OF W(L) * BF(T-
2L) FOR L = 0 TO' ,I3,' ; WHERE'//2X,'W(L): ' ,8G11.4,10(/8X,8G11.4))
END
SUBROUTINE KINOFF (CFQ,IPRINT).

```

C
C
C
C
C
C
C
C
C
C
C
C
C
C
C
C
C
C
C
C
C

```

SUBROUTINE 'KINOFF' DETERMINES THE SUB-AREA RUNOFF HYDROGRAPH
THROUGH THE FINITE DIFFERENCE/KINEMATIC WAVE METHODOLOGY.
SUBROUTINE 'KINOFF' CALLS THE FOLLOWING SUBROUTINES

```

```

SUBROUTINE 'ROFGRD'
SUBROUTINE 'FDKRUT'
SUBROUTINE 'FLOGRD'

```

```

SUBROUTINE 'FRMMFC'
SUBROUTINE 'TOMTRC'

```

```

'KINOFF' USES THE EXCESS RAINFALL FROM THE 'EXCSR()' ARRAY AND
THE ASSUMPTION OF AN INITIALLY DRY SURFACE TO PRODUCE A RUNOFF
HYDROGRAPH - THIS HYDROGRAPH IS PLACED IN THE 'Q' ARRAY.

```

```

COMMON /CMIOP/          IOUA          , IOUB          ,
1  NHOUR                , NDATA
2  NTIME, NHRP, NHRQ, NDTQ, INPQ, NBASIN, IWBLNC, IPFRQR, IPFRQQ
COMMON /FDK00/ TAREA, TRMN, TRHR, METRC, IEL, NQ, MXNDX, MXNDT
COMMON /FDK01/ CHLNG(5), SLOPE(5), RCMAN(5), PAREA(5), ISHAPE(5),
1 CHWDT(5), ZLNG(5), ALPME(5), EMDEQ(5), DXROF(5), DXFLO(5), DTKWR(5),
2 DXKWR(5), IRUPF(5), SAREA(5), FLOIC(5)
COMMON /FDK02/ EXCSR(1224), Q(1224),
1 QK(1224), QUB(1224), QBF(1224)

```

C
C

```

DIMENSION QLAT(1225), QUPST(1225), ADS(51),
1 RNOFF1(1225), RNOFF2(1225), RUNOFF(1225), FLDFLO(1225)

```

C

```

REAL MC, MS

```

C

C

C

C

```

IF(METRC.NE.0) CALL FRMMFC

```

C

C

C

C

C

C

```

DETERMINE THE KINEMATIC PARAMETERS ALPHA & M FOR CATCHMENTS #1
BYPASS CALCULATIONS FOR CATCHMENT #2 IF CATCHMENT #1 REPRESENTS
MORE THAN 99.5% OF THE CURRENT SUB-BASIN

```

```

MC=5./3.
EMDEQ(1)=MC
EMDEQ(2)=MC

```



```

ALPME(2)=0.
ALPH1=(1.49/RCMAN(1))*SLOPE(1)**0.5
ALPME(1)=ALPH1
IF (PAREA(1).GT.99.5) ALPH2=ALPH1
IF (PAREA(1).GT.99.5) GO TO 1030
ALPH2=(1.49/RCMAN(2))*SLOPE(2)**0.5
ALPME(2)=ALPH2

```

```

1030 CONTINUE

```

```

C
C
C     SUBROUTINE 'ROFGRD' COMPUTES THE NUMBER AND MAGNITUDE OF THE
C     DX AND DT INTERVALS IN THE FINITE DIFFERENCE GRID FOR
C     OVERLAND FLOW.
C

```

```

CALL ROFGRD (NDX1,NDX2,NDT,KINERR,IPRINT)
IF (KINERR.EQ.0) GO TO 1050

```

```

C
C     NO EXCESS,
C     SET RUNOFF TO ZERO,
C     ROUTE UPSTREAM HYDROGRAPH IF REQUESTED,
C     OTHERWISE SKIP TO END OF ROUTING
C

```

```

KINERR=0
DO 1040 I=1,NDT
RUNOFF(I)=0.
FLDFLO(I)=0.

```

```

1040 CONTINUE
IF (IRUPF(IEL).NE.1) GO TO 1340
JEL=IEL-1
GO TO 1180

```

```

C
C
C     THE FOLLOWING DEFINES THE APPLIED RAINFALL (QLAT) AND SETS THE
C     UPSTREAM INFLOW ARRAY (QUPST) EQUAL TO ZERO. INITIAL CONDITIONS
C     ARE ASSUMED TO BE THAT OF A DRY SURFACE, AS INDICATED BY SETTING
C     ALL ELEMENTS OF ARRAY ADS(.) EQUAL TO ZERO.
C

```

```

1050 JR=(NDT-1)/(NQ-1)
RJR=JR
TEMP=1./(RJR*12.*DTKWR(1))
QUPST(1)=0.
QLAT(1)=EXCSR(1)*TEMP
M=1
DO 1060 J=1,NQ
TMP=EXCSR(J)*TEMP
DO 1060 K=1,JR
M=M+1
QLAT(M)=TMP
QUPST(M)=0.0
1060 CONTINUE
DO 1070 J=1,NDX1
1070 ADS(J)=0.0
SUMA=0.

```

```

C
C
C     'FDKRUT' IS THE FINITE DIFFERENCE SOLUTION SCHEME. IN THIS

```



```

C     INSTANCE IT IS USED TO COMPUTE THE OVERLAND FLOW HYDROGRAPH.
C
CALL FDKRUT (NDX1,NDT,1,SUMA,ADS,QLAT,QUPST,RNOFF1)
C
C
IF (PAREA(1).GT.99.5) GO TO 1120
C
C     DISTRIBUTE EXCESS FROM SECOND STRIP,
C     AND COMPUTE RUNOFF FROM STRIP
C
TEMP=1./(RJR*12.*DTKWR(2))
QUPST(1)=0.
QLAT(1)=EXCSR(1)*TEMP
M=1
DO 1100 J=2,NQ
TMP=EXCSR(J)*TEMP
DO 1090 K=1,JR
M=M+1
QLAT(M)=TMP
QUPST(M)=0.
1090 CONTINUE
1100 CONTINUE
DO 1110 J=1,NDX2
ADS(J)=0.
1110 CONTINUE
CALL FDKRUT (NDX2,NDT,2,SUMA,ADS,QLAT,QUPST,RNOFF2)
C
C     THE EFFECTIVE STREAM LENGTH IS COMPUTED BASED ON GIVEN OVERLAND
C     FLOW DISTANCES AND DRAINAGE AREA OF THE SUB-AREA. THE RUNOFF
C     HYDROGRAPH IS ADJUSTED BY THE RATIO OF EFFECTIVE TO ACTUAL
C     STREAM LENGTH.
C
C     USE AREA SERVED BY COLLECTOR SYSTEM
1120 IF (SAREA(3).LE.0.0) SAREA(3)=TAREA
IF (IEL.EQ.3) SAREA(3)=TAREA
ASZE=5280.*5280.*SAREA(3)
AREAL=PAREA(1)*ASZE/100.
EL1=AREAL/CHLNG(1)
EL2=0.0
IF (PAREA(1).GT.99.5) GO TO 1140
EL2=(ASZE-AREAL)/CHLNG(2)
EL3=(EL1*PAREA(1))/100.+(EL2*PAREA(2))/100.
DO 1130 J=1,NDT
1130 RUNOFF(J)=(RNOFF1(J)*EL1+RNOFF2(J)*EL2)/EL3
GO TO 1160
1140 EL3=EL1
DO 1150 J=1,NDT
1150 RUNOFF(J)=RNOFF1(J)
1160 DO 1170 J=1,NDT
1170 RUNOFF(J)=RUNOFF(J)*EL3/CHLNG(3)
C
C     LOOP FOR MULT COLLECTOR CHANNELS
JEL=2
1180 JEL=JEL+1
C
C     ALPHA AND M FOR THE STREAM BASED ON CROSS-SECTION SHAPE.

```

```

C
J=ISHAPE(JEL)
GO TO (1190,1200,1210,1220,1230),J
C
C
C
C
C      CIRCULAR CROSS-SECTION
1190 MS=5./4.
    ARG=1./6.
    ALPHS=CHWDT(JEL)**ARG*SLOPE(JEL)**0.5*0.804/RCMAN(JEL)
    GO TO 1240
C
C
C      TRIANGULAR CROSS-SECTION
1200 MS=4./3.
    ARG=1./3.
    ALPHS=(ZLNG(JEL)/(ZLNG(JEL)**2+1.))**ARG*
1    SQRT(SLOPE(JEL))*0.94/RCMAN(JEL)
    GO TO 1240
C
C
C      SQUARE CROSS-SECTION
1210 MS=4.0/3.
    ALPHS=SLOPE(JEL)**0.5*0.72/RCMAN(JEL)
    GO TO 1240
C
C
C      RECTANGULAR CROSS-SECTION
1220 MS=5./3.
    ARG=-2./3.
    ALPHS=CHWDT(JEL)**ARG*SLOPE(JEL)**0.5*1.49/RCMAN(JEL)
    GO TO 1240
C
C
C      TRAPEZOIDAL CROSS-SECTION
1230 Y1=0.5
    Y2=5.0
    W=CHWDT(JEL)
    Z=ZLNG(JEL)
    ZZ=2.*SQRT(1.+Z*Z)
    A1=Y1*(W+Z*Y1)
    A2=Y2*(W+Z*Y2)
    Q1=A1**1.6667*(1./(W+Y1*ZZ))**0.6666667
    Q2=A2**1.6667*(1./(W+Y2*ZZ))**0.6666667
    MS=(ALOG10(Q2)-ALOG10(Q1))/(ALOG10(A2)-ALOG10(A1))
    ALP=Q2/(A2**MS)
    ALPHS=1.49/RCMAN(JEL)*SQRT(SLOPE(JEL))*ALP
1240 ALPME(JEL)=ALPHS
    EMDEQ(JEL)=MS
C
C
C      STREAM LATERAL INFLOW IS SET EQUAL TO THE ADJUSTED RUNOFF
C      HYDROGRAPH, UPSTREAM INFLOW IS SET EQUAL TO ZERO, AND THE
C      INITIAL CONDITIONS ARE DERIVED FROM THE INITIAL DISCHARGE
C      SPECIFIED
C
DO 1250 J=1,NDT

```

```

      QLAT(J)=RUNOFF(J)
1250 QUPST(J)=0.0
      IF (IRUPF(JEL).NE.1) GO TO 1300
C
C   DEFINE UPSTREAM FLOWS QK(.)
      K=0
      K=K+1
      QUPST(1)=QK(K)*CFQ
      L=1
      DO 1280 I=2,NQ
      TMP1=QK(K)
      K=K+1
      TMP2=(QK(K)-TMP1)/RJR
      DO 1270 J=1,JR
      L=L+1
      QUPST(L)=(TMP1+TMP2*FLOAT(J))*CFQ
1270 CONTINUE
1280 CONTINUE
C
C   SET RUNOFF TO SUM OF UPSTREAM AND LATERAL INFLOW,
C   THEN USE FLOGRD TO COMPUTE DELTA X
C
      DO 1290 J=1,NDT
      RUNOFF(J)=QUPST(J)+QLAT(J)
1290 CONTINUE
C
C   FLOGRD COMPUTES THE MAGNITUDE AND NUMBER OF DX INTERVALS IN THE
C   FINITE DIFFERENCE GRID FOR THE KINEMATIC STREAM ROUTING OF THE
C   RUNOFF HYDROGRAPH.
C
1300 CALL FLOGRD (NDT,RUNOFF,NDX,JEL,KINERR,IPRINT)
      IF (KINERR.NE.0) RETURN
C   SET INITIAL CONDITIONS TO FIRST FLOW
      IF (FLOIC(JEL).LE.0.) FLOIC(JEL)=QUPST(1)
      RMS=1./MS
      IF (FLOIC(JEL).LE.0.0) AINIT=0.0
      IF (FLOIC(JEL).LE.0.0) GO TO 1310
      AINIT=(FLOIC(JEL)/ALPHS)**RMS
1310 DO 1320 J=1,NDX
1320 ADS(J)=AINIT
C
      SUMA=FLOAT(NDX)*AINIT
C   FDKRUT IS USED HERE TO COMPUTE THE SUB-AREA OUTFLOW HYDROGRAPH
C
      CALL FDKRUT (NDX,NDT,JEL,SUMA,ADS,QLAT,QUPST,FLDFLO)
C
      IF (JEL.GE.IEL) GO TO 1340
      AX=SAREA(JEL)
      IF (AX.LE.0.0) AX=TAREA
C   RESET RUNOFF AS FLDFLO RATIOED BY AREA SERVED
      TA=SAREA(JEL+1)
      IF (TA.LE.0.0) TA=TAREA
      IF (JEL+1.EQ.IEL) TA=TAREA
      DO 1330 I=1,NDT
1330 RUNOFF(I)=FLDFLO(I)*TA/AX/CHLNG(JEL+1)
      GO TO 1180

```

1340 CONTINUE

C
C
C

STORE FLOWS CORRESPONDING TO STANDARD TIME INTERVAL IN Q(.

L=0

DO 1355 J=1,NDT,JR

TMP=FLDFLO(J)

IF (TMP.LE.0.0) TMP=0.0

L=L+1

Q(L)=TMP

1355 CONTINUE

C
C
C

SUBROUTINE 'TOMTRC' CONVERTS ENGLISH UNITS TO METRIC UNITS.

IF (METRC.LE.0) GO TO 1360

CALL TOMTRC

1360 CONTINUE

C

IF (IPRINT.EQ.0) RETURN

C
C
C
C

PRINT INPUT DATA AND COMPUTED KINEMATIC PARAMETERS FOR THE
CURRENT SUB-AREA BEING SIMULATED.

IF (METRC.EQ.0) WRITE(IOUB,1370)

IF (METRC.NE.0) WRITE(IOUB,1371)

IELP=0

DO 1390 I=1,IEL

IF (I.EQ.2 .AND. PAREA(1).GT.99.5) GO TO 1390

DTKWR(I)=DTKWR(I)/60.

IF (METRC.NE.0) DXKWR(I)=DXKWR(I)/3.28

IELP=IELP+1

WRITE(IOUB,1380) IELP,ALPME(I),EMDEQ(I),DTKWR(I),DXKWR(I)

1390 CONTINUE

C

RETURN

1370 FORMAT(/ 1X, 29HCOMPUTED KINEMATIC PARAMETERS//2X, 50HELEMENT
1ALPHA M DT (MIN) DX (FT))

1371 FORMAT(/ 1X, 29HCOMPUTED KINEMATIC PARAMETERS//2X, 50HELEMENT
1ALPHA M DT (MIN) DX (MT))

1380 FORMAT(2X,I4,3X,F10.4,F9.3,F12.2,F12.2)

END

SUBROUTINE FDKRUT (NDX,NDT,IL,SUM,A,QLAT,QUPST,DSCHRG)

C
C
C
C
C
C
C
C
C
C
C

SUBROUTINE 'FDKRUT' IS THE FINITE DIFFERENCE SOLUTION SCHEME.
GENERATES OVERLAND FLOW RUNOFF HYDROGRAPHS OR STREAM DISCHARGE
HYDROGRAPHS.

THIS SUBROUTINE IS CALLED BY SUBROUTINES 'KINOFF'.

FDKRUT REQUIRES A LATERAL INFLOW HYDROGRAPH, AN UPSTREAM INFLOW
HYDROGRAPH AND A SET OF INITIAL CONDITIONS. THESE QUANTITIES ARE
CALCULATED IN THE CALLING SUBROUTINES AND PASSED TO 'FDKRUT'.

COMMON /FDK01/ CHLNG(5),SLOPE(5),RCMAN(5),PAREA(5),ISHAPE(5),
1 CHWDT(5),ZLNG(5),ALPME(5),EMDEQ(5),DXROF(5),DXFLO(5),DTKWR(5),

2. DXKWR(5), IRUPF(5), SAREA(5), FLOIC(5)

DIMENSION A(51), QLAT(1225), QUPST(1225), DSCHRG(1225)
REAL M

ASSIGN VALUES TO KINEMATIC PARAMETERS

M=EMDEQ(IL)
ALPH=ALPME(IL)
RM=1./M
EXM=M-1.
NDT1=NDT+1

START OF COMPUTATION LOOP TO GENERATE HYDROGRAPH.

DO 1060 I=2,NDT1
K=I-1
ASAV1=A(1)
QNEW=QUPST(K)
IF (QNEW.LT.0.0) QNEW=0.0
IF (QNEW.LE.0.0) A(1)=0.0
IF (QNEW.LE.0.0) GO TO 1000
A(1)=(QNEW/ALPH)**RM

1000 ASAVE=A(1)
ALAT=QLAT(K)*DTKWR(IL)
ANDX=A(NDX)
IF (ANDX.LT.1.0E-20) ANDX=0.0

ESTIMATE SPEED OF 'DISTURBANCE' TO SELECT APPROPRIATE COMPUTATION
SCHEME.

WAVEQ=(QNEW-ALPH*ANDX**M)/CHLNG(IL)+ALAT+SUM/FLOAT(NDX)
CELER=ALPH*M*WAVEQ**EXM
SUM=A(1)
IF (CELER.GT.(DXKWR(IL)/DTKWR(IL))) GO TO 1030

STANDARD FORM OF FINITE DIFFERENCE

DO 1020 J=2,NDX
AZ=(ASAVE+A(J))/2.
IF (AZ.LT.1.0E-20) AZ=0.
IF (AZ.LT.1.0E-20) THETA=0.0
IF (AZ.LT.1.0E-20) GO TO 1010
THETA=(ALPH*M*DTKWR(IL)/DXKWR(IL))*AZ**EXM
1010 ANEW=ALAT+A(J)-THETA*(A(J)-ASAVE)
IF (ANEW.LT.0.0) ANEW=0.0
ASAVE=A(J)
A(J)=ANEW
1020 SUM=SUM+ANEW
GO TO 1050

CONSERVATION FORM OF FINITE DIFFERENCE

1030 QLATCV=ALAT*DXKWR(IL)/DTKWR(IL)
DO 1040 J=2,NDX
QNEW=QNEW+QLATCV-((DXKWR(IL)/DTKWR(IL))*(A(J-1)-ASAV1))

```

IF (QNEW.LT.0.0) QNEW=0.0
ASAVL=A(J)
A(J)=(QNEW/ALPH)**RM

```

```

1040 SUM=SUM+A(J)

```

```

C
C
C
C

```

```

'A(NDX)' IS CONCENTRATION AT DOWNSTREAM LIMIT OF ELEMENT BEING
SIMULATED. DISCHARGE IS GIVEN BY KINEMATIC WAVE FORMULA.

```

```

1050 IF (A(NDX).LT.1.0E-20) A(NDX)=0.
DSCHRG(K)=ALPH*A(NDX)**M

```

```

1060 CONTINUE

```

```

C
C

```

```

RETURN
END

```

```

SUBROUTINE FLOGRD (NDT,RUNOFF,NDX,IL,KINERR,IPRINT)

```

```

C
C
C
C
C
C
C
C
C
C
C
C
C
C
C
C
C

```

```

SUBROUTINE 'FLOGRD' COMPUTES THE NUMBER AND MAGNITUDE OF THE
DX INTERVALS USED IN THE STREAM ROUTING PORTION OF THE SUB-AREA
RUNOFF COMPUTATIONS BY THE KINEMATIC ROUTINES.

```

```

THIS SUBROUTINE IS CALLED BY SUBROUTINE 'KINOFF'.

```

```

'FLOGRD' USES THE SAME NUMBER AND MAGNITUDE FOR THE DT
INTERVAL FOR STREAM ROUTING AS WAS COMPUTED IN 'ROFGRD' FOR
CATCHMENT OVERLAND FLOW ROUTING. THE NUMBER OF DX INTERVALS
IS THEN COMPUTED AS A FUNCTION OF DT AND ADJUSTED SO THAT
THERE ARE AT LEAST 2 BUT NO MORE THAN MXNDX DX INTERVALS.
MXNDX MAY BE CHANGED IN DATA STATEMENT IN SUBROUTINE
KINWAVE AND VARIABLES ADS(MXNDX+1) AND A(MXNDX+1) IN
SUBROUTINES KINOFF AND FDKRUT, RESPECTIVELY SHOULD HAVE
CORRESPONDING DIMENSIONS CHANGED.

```

```

COMMON /CMIOP/ IOUA , IOUB ,
1 NHOOR , NDATA ,
2 NTIME, NHRP, NHRQ, NDTQ, INPQ, NBASIN, IWBLNC, IPFRQR, IPFRQQ
COMMON /FDK00/ TAREA, TRMN, TRHR, METRC, IEL, NQ, MXNDX, MXNDT
COMMON /FDK01/ CHLNG(5), SLOPE(5), RCMAN(5), PAREA(5), ISHAPE(5),
1 CHWDT(5), ZLNG(5), ALPME(5), EMDEQ(5), DXROF(5), DXFLO(5), DTKWR(5),
2 DXKWR(5), IRUPF(5), SAREA(5), FLOIC(5)

```

```

C
C
C
C
C

```

```

DIMENSION RUNOFF(1225)

```

```

REAL MS

```

```

ASSIGN VALUES TO KINEMATIC PARAMETERS.

```

```

ALPHS=ALPME(IL)
MS=EMDEQ(IL)
DMS=(MS-1.)/MS
EMS=MS-1.
RMS=1./MS

```

```

C
C
C

```

```

DETERMINE AVERAGE RATE OF LATERAL INFLOW.

```

```

PEAK=0.

```

```

RATE=0.
DURATN=0.
DO 900 I=1,NDT
IF (RUNOFF(I).GT.PEAK) PEAK=RUNOFF(I)
900 CONTINUE
FMIN=.01*PEAK
DO 1000 I=1,NDT
C
IF (RUNOFF(I).LT.FMIN) GO TO 1000
DURATN=DURATN+DTKWR(1)
RATE=RATE+RUNOFF(I)*CHLNG(IL)
1000 CONTINUE
KINERR=0
IF (DURATN.GT.0.) GO TO 1020
C
C
C
C     PROBABLE ERROR IN PREVIOUS SUBROUTINE - PRINT DIAGNOSTIC
C     MESSAGE AND RETURN.
C
KINERR=1
IF (IPRINT.NE.0) WRITE(IOUB,1010)
RETURN
C
1020 AVGIN=RATE/DURATN
TC=((CHLNG(IL)/(3.2*ALPHS*MS*AVGIN**EMS))**RMS*3.2
DIMT=DTKWR(1)/TC
C
C
C     COMPUTE DX FOR STREAM ROUTING.
C
DX1=CHLNG(IL)*DIMT*3.2**DMS*MS**DMS
DTKWR(IL)=DTKWR(1)
DXFLO(IL)=DX1
C
C
C     ADJUST DX TO BE INTEGER MULTIPLE OF STREAM LENGTH BETWEEN 2 & 5
C
DXMAX=CHLNG(IL)/2.
DXMIN=CHLNG(IL)/FLOAT(MXNDX-1)
IF (DX1.GE.DXMAX) GO TO 1030
IF (DX1.LE.DXMIN) GO TO 1040
NDX=CHLNG(IL)/DX1
NDX=NDX+1
RNDX=NDX-1
DXKWR(IL)=CHLNG(IL)/RNDX
RETURN
1030 NDX=3
DXKWR(IL)=CHLNG(IL)/2.
RETURN
1040 NDX=MXNDX
DXKWR(IL)=CHLNG(IL)/FLOAT(MXNDX-1)
RETURN
C
1010 FORMAT('//2X,'NOTE: NO RUNOFF FROM CATCHMENT(S) '/')
END
SUBROUTINE FRMMTC
C
C
C     SUBROUTINE 'FRMMTC' CONVERTS THE REQUIRED INPUT DATA TO KINOFF

```


C FROM METRIC TO ENGLISH UNITS WHEN METRC IS EQUAL TO 1.
 C THIS SUBROUTINE IS CALLED FROM SUBROUTINE 'KINOFF'
 C

COMMON /FDK00/ TAREA,TRMN,TRHR,METRC,IEL,NQ,MXNDX,MXNDT
 COMMON /FDK01/ CHLNG(5),SLOPE(5),RCMAN(5),PAREA(5),ISHAPE(5),
 1 CHWDT(5),ZLNG(5),ALPME(5),EMDEQ(5),DXROF(5),DXFLO(5),DTKWR(5),
 2 DXKWR(5),IRUPF(5),SAREA(5),FLOIC(5)
 COMMON /FDK02/ EXCSR(1224),Q(1224),
 1 QK(1224),QUB(1224),QBF(1224)

C
 C CONVERT EXCESS FROM MM TO INCHES

DO 1000 J=1,NQ
 EXCSR(J)=0.03937*EXCSR(J)

1000 CONTINUE

C
 C CONVERT CHANNEL LENGTH, WIDTH, AND SERVICE AREA

DO 1010 I=1,5
 CHLNG(I)=CHLNG(I)/0.3048
 CHWDT(I)=CHWDT(I)/0.3048
 SAREA(I)=SAREA(I)/2.590

1010 CONTINUE

C
 C TAREA=TAREA/2.589988

C
 C RETURN
 C END

SUBROUTINE ROFGRD (NDX1,NDX2,NDT,KINERR,IPRINT)

C
 C SUBROUTINE 'ROFGRD' COMPUTES THE NUMBER OF DX & DT INTERVALS IN
 C THE FINITE DIFFERENCE GRID FOR OVERLAND FLOW. IT ALSO CALCULATES
 C THE MAGNITUDE OF EACH INTERVAL IN FEET OR SECONDS.
 C

C THIS SUBROUTINE IS CALLED FROM SUBROUTINE 'KINOFF'.
 C

C 'ROFGRD' REQUIRES THE NUMBER OF COMPUTATION INTERVALS 'NQ'.
 C THE NUMBER OF DT INTERVALS GIVEN BY 'ROFGRD' WILL
 C BE NO LESS THAN 'NQ' BUT NO MORE THAN 1224. IF MORE THAN 'NQ'
 C INTERVALS ARE USED THE NUMBER OF INTERVALS WILL BE AN INTEGER
 C MULTIPLE OF 'NQ'.
 C

C THE NUMBER OF DX INTERVALS WILL BE A FUNCTION OF THE VALUE OF
 C DT BUT WILL BE NO LESS THAN 2 NOR MORE THAN 50.
 C

COMMON /CMIOP/ IOUA , IOUB ,
 1 NHOOR , NDATA ,
 2 NTIME, NHRP, NHRQ, NDTQ, INPQ, NBASIN, IWBLNC, IPFRQR, IPFRQQ
 COMMON /FDK00/ TAREA,TRMN,TRHR,METRC,IEL,NQ,MXNDX,MXNDT
 COMMON /FDK01/ CHLNG(5),SLOPE(5),RCMAN(5),PAREA(5),ISHAPE(5),
 1 CHWDT(5),ZLNG(5),ALPME(5),EMDEQ(5),DXROF(5),DXFLO(5),DTKWR(5),
 2 DXKWR(5),IRUPF(5),SAREA(5),FLOIC(5)
 COMMON /FDK02/ EXCSR(1224),Q(1224),
 1 QK(1224),QUB(1224),QBF(1224)

C
 C REAL MC

C
 C ASSIGN VALUES TO KINEMATIC PARAMETERS AND TO 'NDX2'

C

```

MC=EMDEQ(1)
RMC=1./MC
EMC=MC-1.
DMC=(MC-1.)/MC
DT2=1.0E20
ALPH1=ALPME(1)
ALPH2=ALPME(2)
NDX2=0

```

C
C
C
C

DETERMINE RAINFALL VOLUME AND AVERAGE INTENSITY

```

PEAK=0.
AMT1=0.
AMT2=0.
AVGRN1=0.
AVGRN2=0.
DURTN1=0.
DURTN2=0.
DO 900 I=1,NQ
IF (EXCSR(I).GT.PEAK) PEAK=EXCSR(I)
900 CONTINUE
EMIN=.1*PEAK
DO 1000 I=1,NQ
IF (EMIN.GT.EXCSR(I)) GO TO 1000
AMT1=AMT1+EXCSR(I)
DURTN1=DURTN1+TRMN
1000 CONTINUE
IF (PAREA(1).GT.99.5) GO TO 950
AMT2=AMT1
AVGRN2=AVGRN1
DURTN2=DURTN1
950 KINERR=0
IF ((AMT1+AMT2).GT.0.0.AND.(DURTN1+DURTN2).GT.0.0) GO TO 1020

```

C
C
C
C

NO EXCESS RAIN OR PROBABLE INPUT ERROR - PRINT DIAGNOSTIC
MESSAGE AND RETURN

```

KINERR=1
IF (KINERR.EQ.1.AND.IPRINT.NE.0) WRITE(IOUB,1010)
DT=TRMN*60.
DTKWR(1)=DT
DTKWR(2)=DT
NDT=NQ
RETURN

```

C
C
C
C

COMPUTE AVERAGE EXCESS IN FEET PER SECOND

```

1020 IF (DURTN1.GT.0.) AVGRN1=AMT1/(720.*DURTN1)
IF (DURTN2.GT.0.) AVGRN2=AMT2/(720.*DURTN2)

```

C
C
C
C

SET VALUES FOR ESTIMATING DX & DT.

SCALEQ=1.0
 SCALEX=1.0
 XQ=AVGRN1*SCALEQ

C
 C
 C
 C
 C
 C
 C

DETERMINE THE MAXIMUM NUMBER OF DT INTERVALS AND THE MINIMUM
 VALUE OF DT. FOR MORE INTERVALS OR SMALLER DT INCREASE
 THE DIMENSIONS OF ARRAYS QLAT(.), QUPST(.), RNOFF1(.),
 RNOFF2(.), RUNOFF(.), FLDFLO(.), AND DSCHRG(.) WHICH ARE
 CURRENTLY SET TO 1225 (=MXNDT, SEE DATA STATEMENT IN 'KINWAVE')

INDEX=(MXNDT-1)/(NQ-1)
 DTMIN=TRMN*60./FLOAT(INDEX)

C
 C
 C
 C

DETERMINE DX & DT VALUES BASED ON KINEMATIC WAVE AND APPLIED
 RAINFALL.

IF (AVGRN1.GT.0.) GO TO 1050
 DT1=TRMN*60.
 GO TO 1060

1050 XL=CHLNG(1)*SCALEX
 DT1=(XL/(3.2*ALPH1*MC*(XQ**EMC)))*RMC
 DT1=DT1*3.2

1060 DXROF(1)=DT1

C

IF (PAREA(1).GT.99.5) GO TO 1090

C

IF (AVGRN2.GT.0.) GO TO 1070
 DT2=TRMN*60.
 GO TO 1080

1070 XQ=AVGRN2*SCALEQ
 XL=CHLNG(2)*SCALEX
 DT2=(XL/(3.2*ALPH2*MC*(XQ**EMC)))*RMC
 DT2=DT2*3.2

1080 DXROF(2)=DT2

C

C

C

C

C

C

C

C

C

C

C

C

C

C

C

C

C

C

C

C

C

C

C

C

C

C

C

C

ADJUST COMPUTED DT TO BE AN INTEGER MULTIPLE OF 'NQ' AND
 COMPATIBLE WITH CURRENT DIMENSIONS.

INDEX=TRMN*60./TRYT+0.5
 IF (INDEX.LT.1) INDEX=1
 DT=TRMN*60./FLOAT(INDEX)
 NDT=INDEX*(NQ-1)+1

DTKWR(1)=DT
 DTKWR(2)=DT

C
C
C
C

DETERMINE DX FOR THE FIRST CATCHMENT ELEMENT.

```

DIMT=DT/DT1
DX1=CHLNG(1)*DIMT*3.2**DMC*MC**DMC
DXFLO(1)=DX1

```

C
C
C
C

ADJUST DX TO BE AN INTEGER MULTIPLE OF CATCHMENT LENGTH BETWEEN 2 AND MXNDX (SEE COMMENTS IN SUBROUTINE FLOGRD).

```

DXMIN=CHLNG(1)/FLOAT(MXNDX-1)
DXMAX=CHLNG(1)/2.
IF (DX1.GE.DXMAX) GO TO 1120
IF (DX1.LE.DXMIN) GO TO 1130
NDX1=CHLNG(1)/DX1
NDX1=NDX1+1
RNDX1=NDX1-1
DXKWR(1)=CHLNG(1)/RNDX1
GO TO 1140

```

1120 NDX1=3

```

DXKWR(1)=CHLNG(1)/2.
GO TO 1140

```

1130 NDX1=MXNDX

```

DXKWR(1)=CHLNG(1)/FLOAT(MXNDX-1)

```

C
C
C
C

DETERMINE DX FOR SECOND CATCHMENT ELEMENT.

1140 CONTINUE

```

IF (PAREA(1).GT.99.5) RETURN
DIMT=DT/DT2
DX2=CHLNG(2)*DIMT*3.2**DMC*MC**DMC
DXFLO(2)=DX2

```

C
C
C
C
C

ADJUST DX TO BE AN INTEGER MULTIPLE OF CATCHMENT LENGTH BETWEEN 2 AND MXNDX (SEE COMMENTS IN SUBROUTINE FLOGRD).

```

DXMIN=CHLNG(2)/FLOAT(MXNDX-1)
DXMAX=CHLNG(2)/2.
IF (DX2.GE.DXMAX) GO TO 1150
IF (DX2.LE.DXMIN) GO TO 1160
NDX2=CHLNG(2)/DX2
NDX2=NDX2+1
RNDX=NDX2-1
DXKWR(2)=CHLNG(2)/RNDX
RETURN

```

1150 NDX2=3

```

DXKWR(2)=CHLNG(2)/2.
RETURN

```

1160 NDX2=MXNDX

```

DXKWR(2)=CHLNG(2)/FLOAT(MXNDX-1)
RETURN

```

C

1010 FORMAT(//2X, 'NOTE: NO RAINFALL EXCESS FOR SUBAREA'/8X,

```
1 'RUNOFF SET EQUAL TO ZERO.'/)
```

C

```
END
SUBROUTINE TOMTRC
```

C

C

C

C

C

```
SUBROUTINE 'TOMTRC' CONVERTS THE ORIGINAL INPUT DATA TO 'KINOFF'
THE COMPUTED HYDROGRAPH TO METRIC UNITS
THIS SUBROUTINE IS CALLED FROM SUBROUTINE 'KINOFF'.
```

```
COMMON /FDK00/ TAREA,TRMN,TRHR,METRC,IEL,NQ,MXNDX,MXNDT
COMMON /FDK01/ CHLNG(5),SLOPE(5),RCMAN(5),PAREA(5),ISHAPE(5),
1 CHWDT(5),ZLNG(5),ALPME(5),EMDEQ(5),DXROF(5),DXFLO(5),DTKWR(5),
2 DXKWR(5),IRUPF(5),SAREA(5),FLOIC(5)
COMMON /FDK02/ EXCSR(1224),Q(1224),
1 QK(1224),QUB(1224),QBF(1224)
```

C

C

```
CFS=(0.0254*12.)**3
DO 1000 J=1,NQ
EXCSR(J)=EXCSR(J)/0.03937
Q(J)=Q(J)*CFS
```

```
1000 CONTINUE
```

```
DO 1010 I=1,5
CHLNG(I)=CHLNG(I)*.3048
CHWDT(I)=CHWDT(I)*.3048
SAREA(I)=SAREA(I)*2.590
```

```
1010 CONTINUE
```

```
TAREA=TAREA*2.590
```

C

```
RETURN
END
```

~eor

APPENDIX 1.8.C

PROGRAM INPUT REQUIREMENTS AND DESCRIPTION

Data Set 1. Model run information

This data set is contained in one record and read in variable TITLE(.) using Fortran FORMAT(10A8). No data set identifier is specified.

Data Set 2. Control parameters

This data set is also contained in one record to read the following variables using FORMAT(5I5):

NPOP - number of SAC model parameters to be optimized

IUPDF - generates updated input file

IWBLNC - number of time steps water balance are made

IPFRQR - printing frequency of SAC results

IPFRQQ - printing frequency of streamflow

This data set has no record identifier.

Data Set 3. Optimization parameters

This data set is needed if optimization is desired by specifying NPOP not equal to zero. The values of the variables below are read using FORMAT (F10.0,4I5) with no record identifier.

ERROR - error criterion of relative difference of old and new objective function value

IOBF - type of objective function

MXITER - maximum number of iterations

NSTEP - step size update option, equals 1 if updated, otherwise equal to 0

IPROP - printing frequency of results at each stage evaluation

Note: For the option and types of objective functions available, see comment statement in program listing.

Data Set 4. Soil-moisture accounting model parameters

In this data set, the 24 SAC model parameters are read which could be any order. The value of each parameter is contained in one record with a record identifier. If a parameter is to be optimized, the record string should include some specified minimum and maximum values as well as the step size since these latter three values are required in the optimization routine. For each record, the following variables are read using FORMAT (2A5,4F10.0)

PNAME(.) - variable name of model parameter which is also the record identifier

DUM - dummy variable which is not used in the program but simply read for "echo" printing purposes

XMP(.) - value of parameter

XMIN(.) - lower bound of parameter

XMAX(.) - upper bound of parameter

ESS(.) - parameter step size

A record containing the word "END" should be placed at the bottom of this data set.

The variable name of the 24 model parameters are listed below with brief descriptions of each. These variable names are also used as identifiers and read in PNAME(.).

1. PXADJ - precipitation adjustment factor
2. PEADJ - evapotranspiration demand adjustment factor
3. PCTIM - fraction of permanent impervious area

4. ADIMP - fraction of impervious area when all tension storage water are met
5. RIVA - fraction of basin covered by streams, lakes and riparian vegetation
6. UZK - upper zone for water storage depletion coefficient
7. LZSK - lower zone supplementary storage depletion coefficient
8. LZPK - lower zone primary storage depletion coefficient
9. PFREE - fraction of percolated water transmitted directly to the lower zone free water
10. RSERV - fraction of lower zone free water unavailable for transpiration purposes
11. ZPERC - proportionality constant in increasing percolation from saturated to dry condition
12. REXP - exponent defining curvature in percolation curve with changes in the lower zone soil moisture deficiency
13. SIDE - portion of baseflow not observed in the channel
14. UZTWM - upper zone tension water storage content
15. UZFWM - upper zone free water storage content
16. LZTWM - lower zone tension water storage content
17. LZTSM - lower zone supplementary water storage content
18. LZFPM - lower zone primary water storage content
19. UZTWC - upper zone tension water storage capacity
20. UZFWC - upper zone free water storage capacity
21. LZTWC - lower zone tension water storage capacity
22. LZFSC - lower zone supplementary water storage capacity
23. LZFPC - lower zone primary water storage capacity

24. ADIMC - additional impervious area storage capacity usually
taken as $UZTWM + LZTWM$

There are two rules to be followed in inputting the above data set.

1. If some model parameters are to be optimized, they must be placed on top of the others.

2. The model parameters containing the initial soil-moisture contents should be placed after those containing the soil-moisture capacities. This is done when the soil-moisture content parameters are not optimized since otherwise the maximum values are specified.

Data Set 5. Kinematic wave routing parameters

The manner in which the kinematic wave model parameters are inputted is based on the physical configuration of the basin. For this purpose, three types of operations have been designated which accommodates practically any basin configuration. These three operations are "ROUTE", "ADD" and "BASEF" which are also used as record identifiers. In inputting a parameter set corresponding to an operation, the first record contains either one of the three operations with an integer variable which are read into ROPER and IOPER using FORMAT (A5, I5). This record may be followed by some input parameters depending on the type of operation as described below.

ROUTE Operation. This operation computes the flow hydrograph from overland plane to channel outlet. As many as two, overland flow planes and collector channels plus a main channel are used to represent a subbasin (second-level). The input parameter sets and sequence are as follow:

Parameter Set 1: Basin Area

TAREA - total subbasin area which is read as (F10.0)

Parameter Set 2: Overland flow element (one record and read as 4F10.10)

CHLNG(.) - overland flow length

SLOPE(.) - slope

RCMAN(.) - roughness coefficient

PAREA(.) - percent of subbasin area

If PAREA(1) is less than 99.5% a second overland-flow element is expected and the same parameters above are required for this second one.

Parameter Set 3: Collector and Main Channel (two records using 5F10.0)

Record 1: CHLNG(.) - channel length

SLOPE(.) - slope

RCMAN(.) - roughness coefficient

SAREA(.) - contributing area

Record 2: ASHAPE(.) - shape of channel used

CHWDT(.) - channel width

ZLNG(.) - side slope

FLOIC(.) - initial flows at outlet

ARUPF(.) - indicator if upstream flow from another subbasin is routed (applicable only in the main channel)

At least one collector channel and the main channel must be specified in a subbasin. A second collector channel is specified if SAREA(.) of the first collector channel is less than the total area

TAREA(.). However, the program assumes that there is only one collector channel if SAREA(1) is inputted as zero.

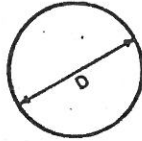
Variable ASHAPE(.) specifies the shape of collector or main channel. These are shown in Figure 1.8.C.1. The variable ASHAPE(.) is set equal to: 1.0 for circular, 2.0 for triangular, 3.0 for square, 4.0 for rectangular, and 5.0 for trapezoidal. Depending on channel shape specified, either ZLNG(.) or CHWDT(.) may be left blank. In the case of a circular channel, the diameter D is specified in CHWDT(.). Only in the case of triangular and trapezoidal channel shapes where ZLNG(.) is required. However, CHWDT(.) is not required for triangular channels.

Depending on the storage space (array dimension) fixed in the program, several ROUTE operations can be issued for a subwatershed (first-level partitioning). For example, if there are three subbasins (at second-level partitioning) that are separated from each other three ROUTE operations have to be made and possibly "ADD" the three routed flows after.

ADD operation. This operation adds the previous IOPER flow hydrograph processed in the "ROUTE" operation. The "ADD" operation can only be used after two or more flow hydrographs have already been computed.

BASE operation. This operation adds the baseflow component to the surface flow hydrograph at a subbasin outlet. Usually this operation is done after a "ROUTE" operation. The baseflow hydrograph component is computed using a linear weighted function of current and previous IOPER baseflow runoff components (obtained in SAC). One or more records should follow the "BASEF" record which contains the weights WBF(I), for $I = 1, \dots, IOPER + 1$ which is read using FORMAT (5G10.0).

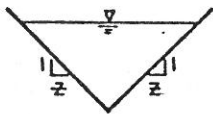
CIRCULAR



$$\alpha = \frac{.804}{n} S^{1/2} D^{1/6}$$

$$m = 5/4$$

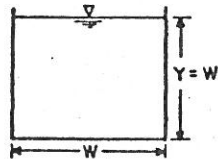
TRIANGULAR



$$\alpha = \frac{0.94}{n} S^{1/2} \left(\frac{z}{1+z^2} \right)^{1/3}$$

$$m = 4/3$$

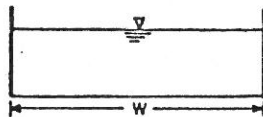
SQUARE



$$\alpha = \frac{.72}{n} S^{1/2}$$

$$m = 4/3$$

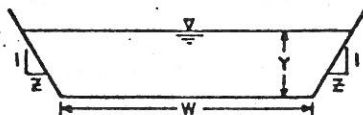
RECTANGULAR



$$\alpha = \frac{1.49}{n} S^{1/2} W^{-2/3}$$

$$m = 5/3$$

TRAPEZOIDAL



$$Q = \frac{1.49}{n} S^{1/2} A^{5/3} \left(\frac{1}{W+2Y\sqrt{1+z^2}} \right)^{2/3}$$

Figure 1.8.C.1. Kinematic wave parameters for various channel shapes.

The usage, sequencing and combinations of ROUTE, ADD and BASEF to represent a certain basin or subbasin configuration will be further explained in the sample model application.

As in data set 4, a record with "END" should be placed at the bottom of this data set 5.

Data Set 6. Hydrologic input data

Three types of hydrologic data that may be required in the model are rainfall, evapotranspiration (ET) demand, and streamflow. Rainfall data is always required in running the model. The ET data when not inputted is assumed to be zero. The streamflow data may not be inputted unless model parameter optimization is desired. Any hydrologic input data is read by issuing an identification record with identifiers "RAIN", "ETDAT" or "FLOW". Each record is followed by some input control parameters as follow.

RAIN input. This is followed by a record with control parameters (read as 5I5):

NHOUR - time interval in hours

NDATA - number of data points

IOUC - input device unit number

KFMT - read format option, if KFMT = 0, read format in (8F10.0), otherwise it is read in the next record

IREW - data file rewind option, if IREW is not = 0, the data file is rewound

If KFMT is not equal to 0 another record is expected containing the read format FMT(.) which is read as (10A8).

ETDAT input. To read the ET demand data, the following control parameters are read in the next record (read as 4I5):

NHRP - time interval in hours
 IOUC - input device unit number
 KFMT - read format option
 IREW - data file rewind option

As in "RAIN" input, if KFMT is not equal to 0, a record containing the read format follows.

FLOW input. Inputting the streamflow requires the following parameters:

NHRQ - time interval in hours
 INPQ - indicates if streamflow is available (if not available INPQ = 0, otherwise, INPQ is nonzero)
 IOUC - read device unit number
 KFMT - read format option
 IREW - data file rewind option

If a FLOW record is not issued, the streamflows are computed with the same time interval as rainfall. If one is interested in a different time interval (specified in NHRQ), and no streamflow data is available, the FLOW record can be issued with INPQ nonzero.

Some items above need further clarification as follows:

1. Given the number of data points NDATA for rainfall, the number of data points required for ET demand or flow, if inputted are: NDATA/NHRP and NDATA/NHRQ, respectively.

2. The read device unit number IOUC provides a control if a hydrologic data is to be read from another file. If IOUC is set equal to 5 or 0, the data is read as part of all other data sets (control parameters, model parameters, etc.). In this case the hydrologic data should be placed after the "Data Set 6" records.

3. The data file rewind option specified in IREW is desirable when two or more subwatersheds (first-level partitioning) with different hydrologic time series are stored in a file where values for the different subwatersheds at each time period (sampling time) are contained in one record. For example, an areally averaged rainfall for the first subwatershed can be read first with say, FORMAT (F10.0). Then the rainfall for the second subwatershed can be read, say as FORMAT (10X,F10.0) after rewinding the same file.

At the bottom of this data set 6, an "END" record should be placed.

1.9 STOCHASTIC GENERATION OF STREAMFLOWS AND TURBINE OPERATING HOURS

1.9.1 INTRODUCTION

This report presents the stochastic modeling and data generation of monthly and weekly streamflows in the Nizao Basin, Dominican Republic and monthly turbine operating hours for Valdesia reservoir. For streamflow modeling three gaging stations were selected for analysis, namely: Palo de Caja, Paso del Ermitaño and Rancho Arriba. Prior to data analysis, the gaps in the historical data were filled-in and short records were extended to improve the reliability of statistical parameters to be used in modeling. The stochastic models used for both monthly and weekly streamflows were developed following standard (currently used) modeling procedures. The modeling and generation of monthly turbine operating hours were based on the monthly flows of Paso del Ermitaño.

The report presented herein is divided into four major sections, namely: 1) description of hydrologic data used, 2) filling-in and extension of historical data, 3) stochastic modeling of streamflows, 4) streamflow data generation, and, 5) modeling and generation of turbine operating hours. Some general remarks are given at the end of this report as well as literature cited and appendices.

1.9.2 DESCRIPTION OF HYDROLOGIC DATA USED

Daily data from four streamflow stations and two rainfall stations provided by Dominican Republic were used in this study. The six gaging stations are listed in Table 1.9.1 and their corresponding years of records. The locations of these stations are shown in Fig. 1.9.1

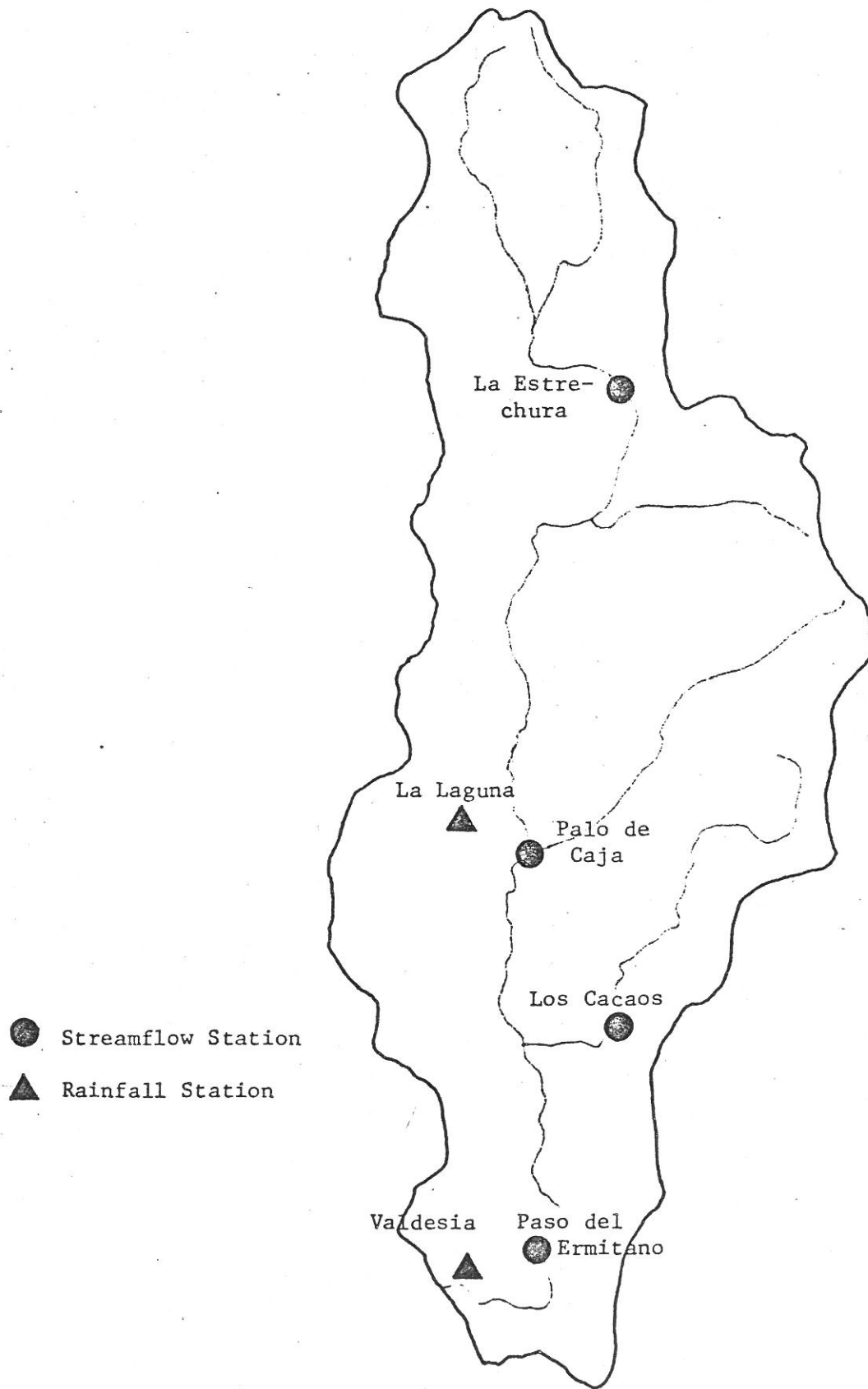


Figure 1.9.1. Map showing Nizao basin and gaging stations used in this study.

indicated by station names. For purposes of this study, the monthly data were derived from averaging the daily data of each month. The weekly data were derived from averaging daily data every seven days starting in January 1 (e.g., January 1-7 is first week, January 8-14 is second week, etc.) except for the last or 52nd week of the year (i.e., December 26 to 31) which is an eight-day average. In the case of leap years, the 9th weeks data is average from eight days comprising February 26, 27, 28 and 29, and March 1, 2, 3 and 4. Time series plots on a monthly basis for each station are given in Figs. 1.9.2 through 1.9.7. The weekly time series plots are not given due to space limitations. It has been seen however, that the basic seasonal and other time series patterns observed in the monthly time series are likewise exhibited in the weekly series.

Table 1.9.1. List of gaging stations used in this study.

Station Name	Location		Period of Record	Type of Data
	Latitude	Longitude		
Palo de Caja	18°33'17"	70°22'52"	Sept. 1956 - July, 1979	Streamflow
Paso del Ermitaño	18°26'02"	70°15'43"	Dec. 1967 - Oct. 1975	Streamflow
Rancho Arriba	18°42'58"	70°27'59"	Mar. 1959 - Oct. 1966	Streamflow
El Cacao*	18°31'44"	70°17'59"	Jan. 1962 - Dec. 1981	Streamflow
La Laguna	18°32'30"	70°24'45"	Jan. 1963 - Dec. 1979	Rainfall
Valdesia	18°24'30"	70°16'50"	Feb. 1963 - Aug. 1984	Rainfall

*Records fragmentary between 1966 to 1979.

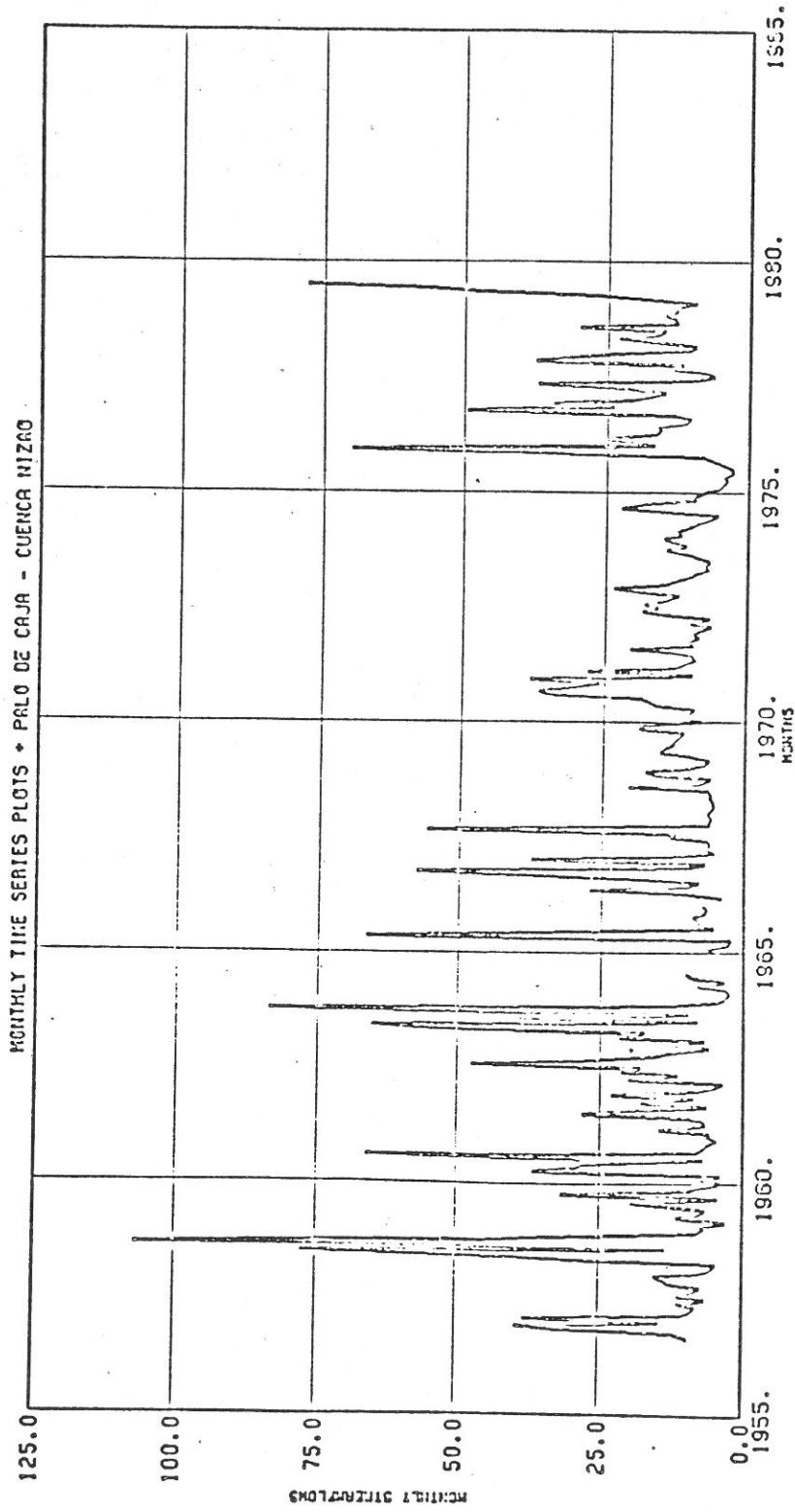


Figure 1.9.2. Time series plots of monthly streamflows of Palo de Caja.

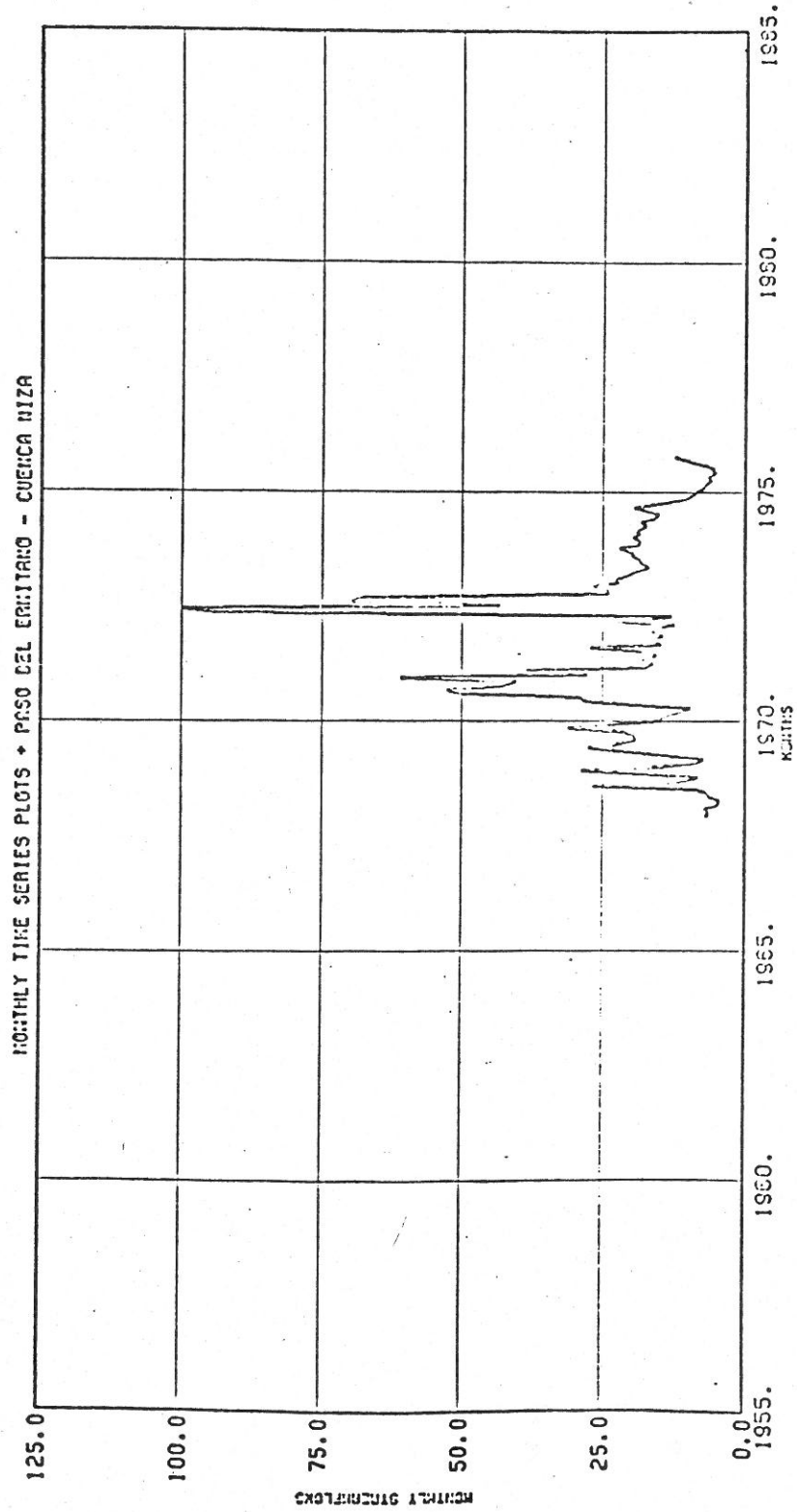


Figure 1.9.3. Time series plots of monthly streamflows of Paso del Ermitaño.

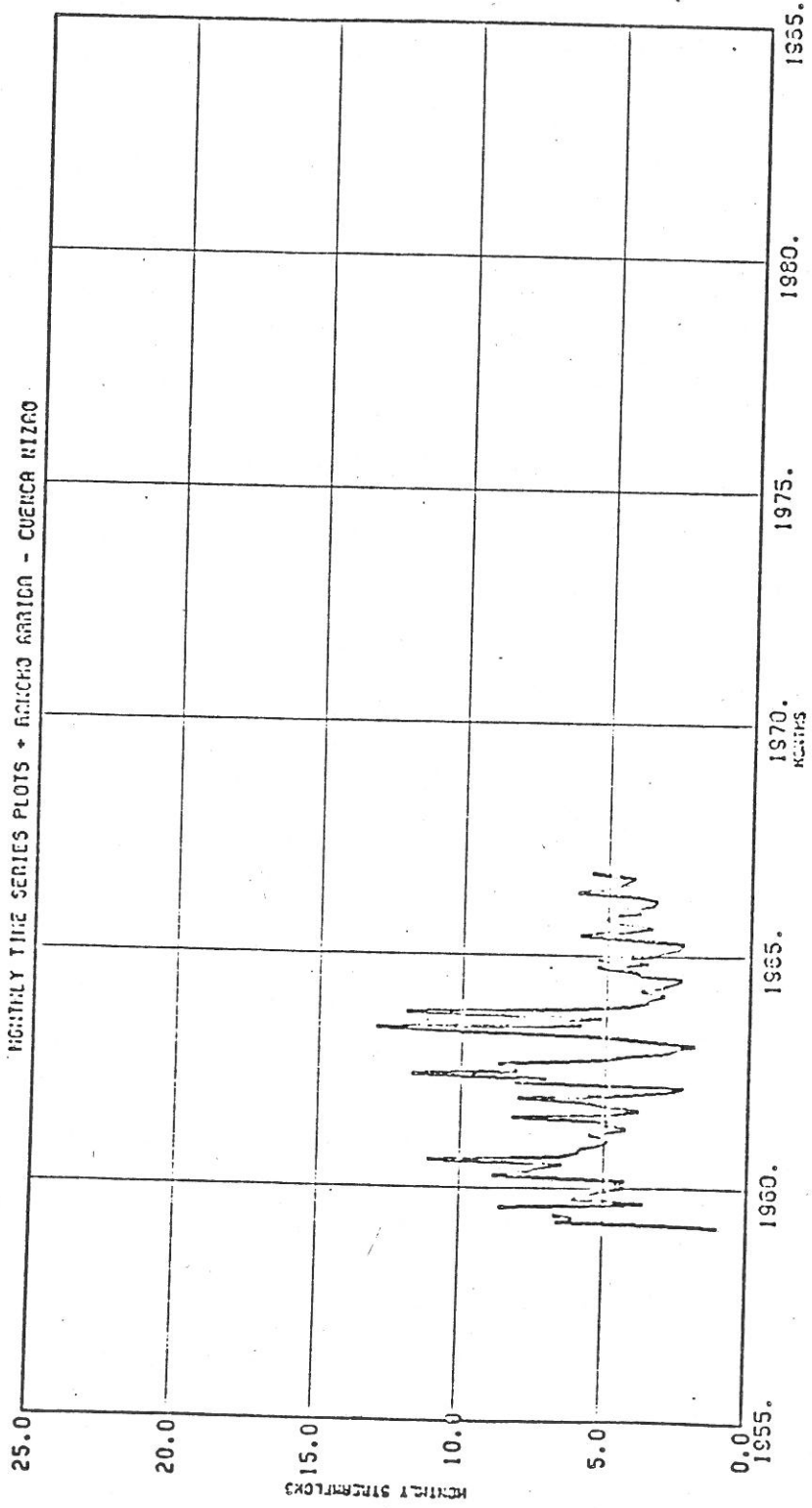


Figure 1.9.4. Time series plots of monthly streamflows of Rancho Arriba.

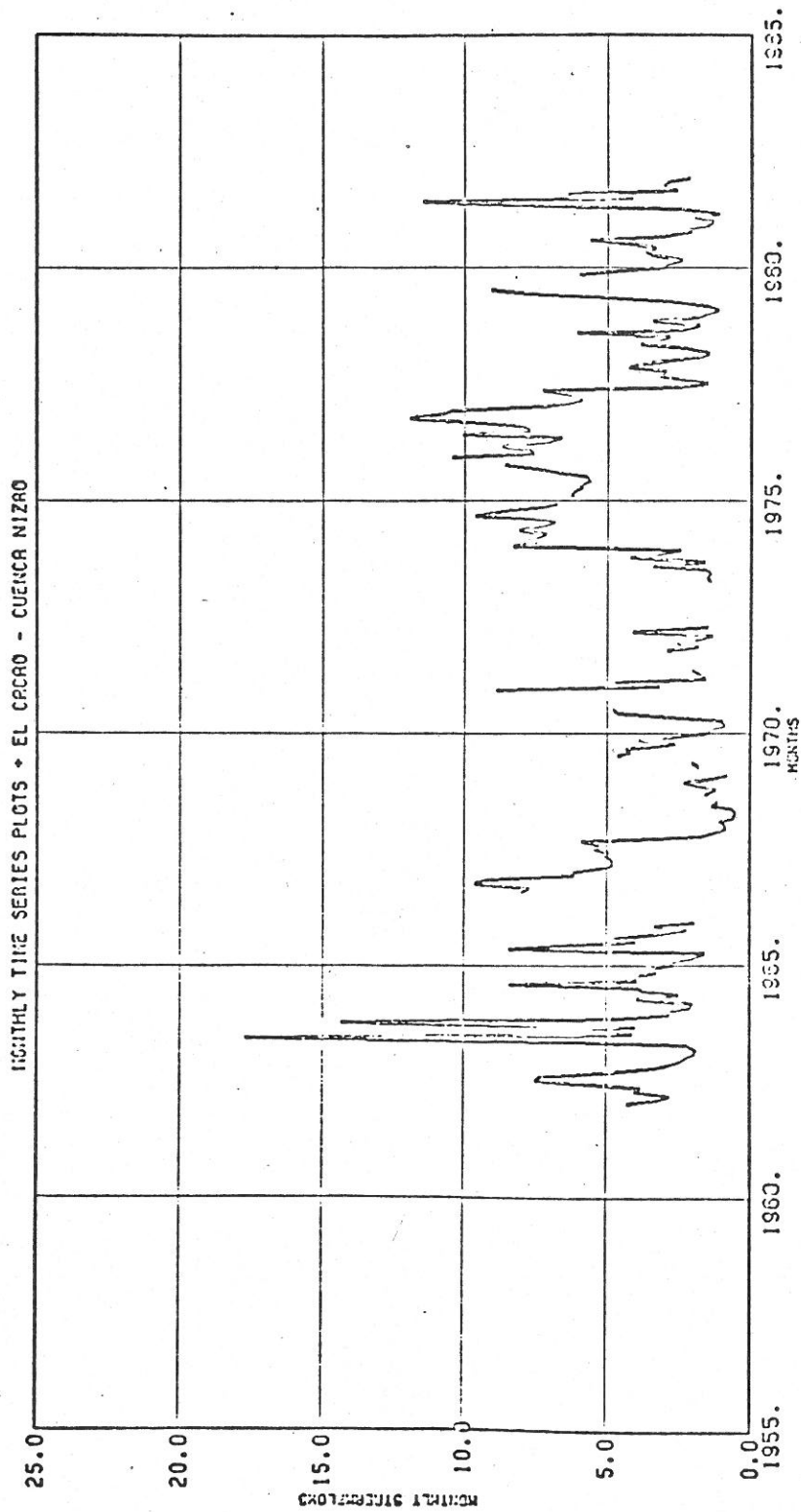


Figure 1.9.5. Time series plots of monthly streamflows of El Cacao.

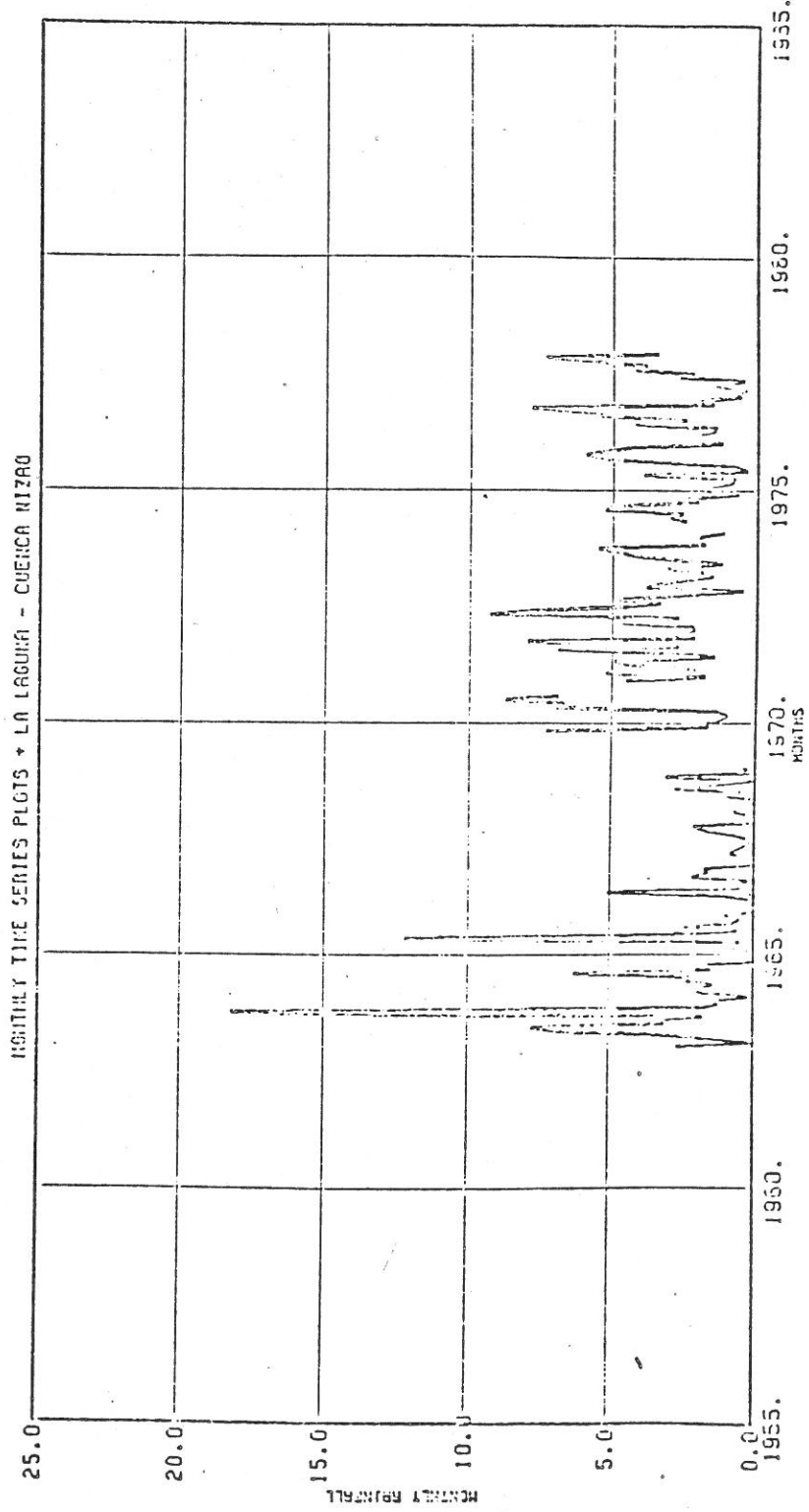


Figure 1.9.6. Time series plots of monthly rainfall of La Laguna.

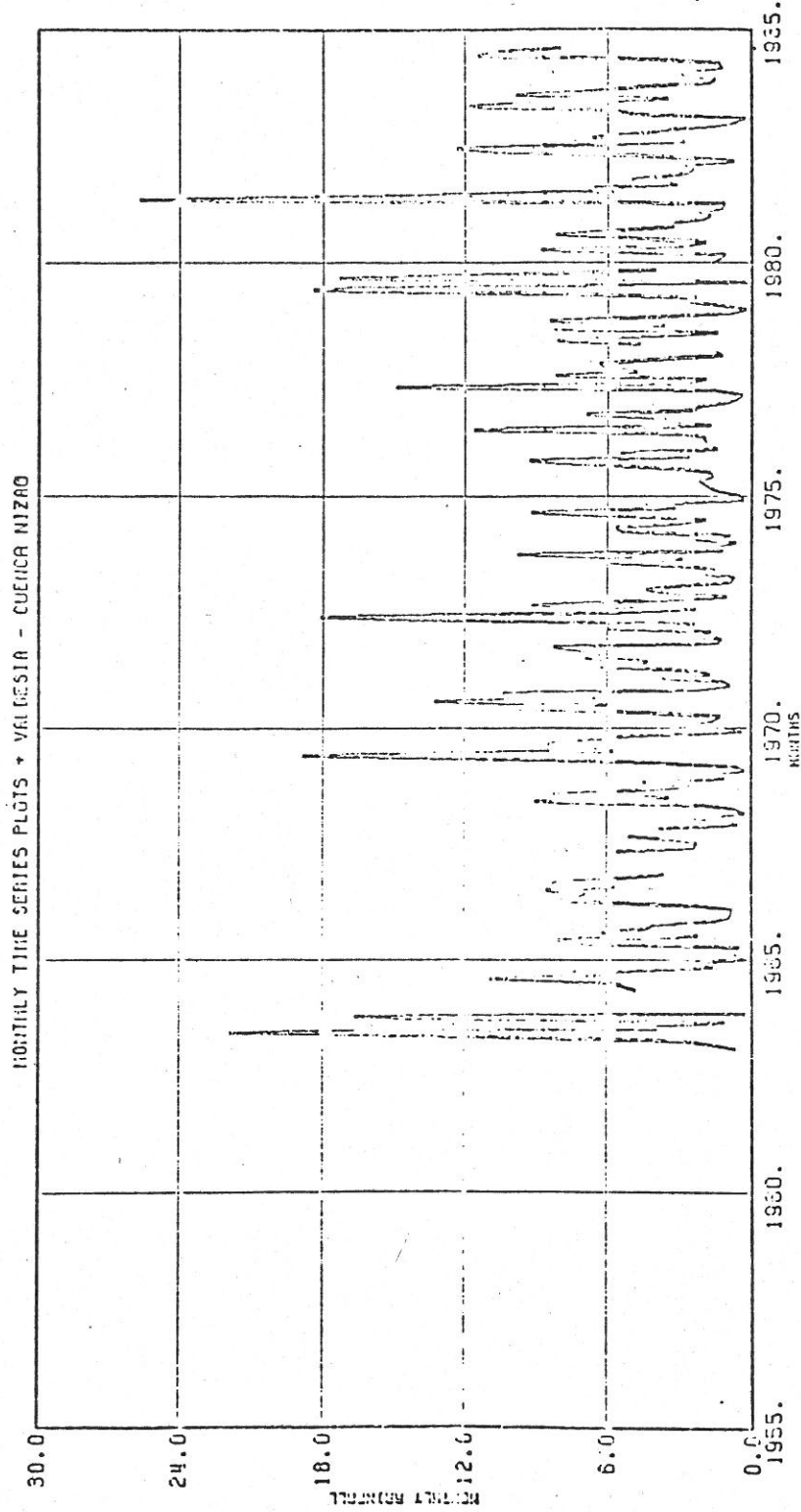


Figure 1.9.7. Time series plots of monthly rainfall of Valdesia.

1.9.3 FILLING-IN AND EXTENSION OF HISTORICAL DATA

To improve the reliability of statistical parameters such as the means, variances, skewness, autocorrelations and cross correlations which will be used for stochastic modeling, the historical data of each station are filled-in and/or extended first based on other stations. The monthly and weekly streamflows of Palo de Caja contain only some data gaps in years 1964 and 1966 thus requiring only minor filling-in of missing data. However, both Paso del Ermitaño and Rancho Arriba require major data extension due to their shortness of records. The succeeding subsections present filling-in of missing data of Palo de Caja and extension of records of Paso del Ermitaño and Rancho Arriba respectively.

1.9.3.1 Filling-in of Missing Data of Palo de Caja

On a monthly basis, the data of months August, October and December of year 1964 and February 1966 are missing. On a weekly basis, weeks 33, 34, 35, 42, 43, 44, 51 and 52 of year 1964 and weeks 6 and 7 of 1966 are missing. It is decided that a simple or multiple linear regression model be used in filling-in the missing data of Palo de Caja based on neighboring stations within the Nizao basin. As a requirement to linear regression models, the regression variables must be nominally normally distributed. Thus, prior to fitting these models, the streamflows of Palo de Caja and the time series variables of other stations are suitably transformed to become normal. Five normalizing transformations were tried, namely: 1) square-root, 2) cube root, 3) logarithmic, 4) Wilson-Hilferty, and 5) combined logarithmic and Wilson-Hilferty transformations. Details of the normalization schemes and results for.

each station are given in Appendix 1.9.A. Using the transformed data of each station, the overall monthly and weekly cross correlations between stations were computed to form the basis of choosing the regressor variables. The computed cross correlations for monthly and weekly data are given in Tables 1.9.2 and 1.9.3, respectively. The choice of the regressor variables is also dependent on the availability of data from these stations with respect to those periods of missing data of Palo de Caja and the coefficient of correlation of the fitted regression model.

On the above basis, the monthly missing values of Palo de Caja are filled-in using the bivariate regression model given by

$$Z_{\nu,\tau}^{(1)} = -0.038 + 0.358 Z_{\nu,\tau}^{(5)} + 0.257 Z_{\nu,\tau}^{(6)} + 0.814 \epsilon_{\nu,\tau} \quad (1.9.1)$$

where $Z_{\nu,\tau}^{(1)}$ is the log-Wilson-Hilferty domain streamflow of Palo de Caja of year ν and month τ ; $Z_{\nu,\tau}^{(5)}$ and $Z_{\nu,\tau}^{(6)}$ are the Wilson-Hilferty domain rainfalls of La Laguna and Valdesia, respectively; and $\epsilon_{\nu,\tau}$ is an added noise term which is identically and independently distributed standard normal random deviates. The addition of the noise term $\epsilon_{\nu,\tau}$ is necessary because otherwise the variance of the filled-in data may be reduced (Salas, et al., 1980). Equation (1.9.1) has a multiple correlation coefficient (R) equal to 0.540. Note that in Table 1.9.1, the overall monthly cross-correlation between Palo de Caja and Paso del Ermitaño is the highest. But during those months where Palo de Caja has missing values, the values of Paso del Ermitaño are also missing. The filled-in monthly values of Palo de Caja in the original domain of streamflows are obtained using the inverse log-Wilson-Hilferty transformation (see Appendix 1.9.A or section 1.9.6.1).

Table 1.9.3. Overall weekly cross-correlations of transformed data between stations.

	Palo de Caja	Paso del Ermitaño	Rancho Arriba	El Cacao	La Laguna	Valdesia
Palo de Caja (LWH)	1.000 (1187)*					
Paso del Ermitaño (LWH)	0.766 (417)	1.000 (417)				
Rancho Arriba (LWH)	0.507 (394)	-- (0)	1.000 (404)			
El Cacao (LWH)	0.371 (796)	0.296 (334)	0.477 (225)	1.000 (916)		
La Laguna (WH)	0.287 (714)	0.398 (359)	0.176 (205)	0.228 (626)	1.000 (724)	
Valdesia (WH)	0.190 (795)	0.240 (471)	0.124 (176)	0.180 (800)	0.479 (666)	1.000 (1062)

*Numbers enclosed in parentheses are total number of concurrent observations.

-- denotes no concurrent record between these two stations.

Note: LWH or WH indicates whether data for that station was transformed either by combined log-Wilson-Hilferty transformation or Wilson-Hilferty transformation, respectively.

For filling-in the weekly values of Palo de Caja, the following simple linear regression model is used:

$$Z_{\nu,\tau}^{(1)} = a(\tau,\nu) + b(\tau,\nu) Z_{\nu,\tau}^{(3)} + c(\tau,\nu) \epsilon_{\nu,\tau} \quad (1.9.2)$$

where $Z_{\nu,\tau}^{(1)}$ and $Z_{\nu,\tau}^{(3)}$ are the weekly log-Wilson-Hilferty domain streamflows of Palo de Caja and Rancho Arriba, respectively and the regression models parameters as a function of year ν and week τ are computed either on a week-to-week basis or overall-weekly basis depending on which of the two give the highest coefficient of correlation (R). Note again that Palo de Caja and Paso del Ermitaño has the highest correlation as shown in Table 1.8.3, but values of Paso del Ermitaño are missing during these weeks when Palo de Caja had missing values. The model coefficients of Eq. (1.9.2) and corresponding coefficient of correlations (R) for each missing year (ν) and weeks (τ) are given in Table 1.9.4. As in monthly filling-in of missing data, the original domain of weekly streamflows of Palo de Caja are obtained using the inverse log-Wilson-Hilferty transformation. Effectively now, the monthly and weekly data available for Palo de Caja is from 1957 to 1978 (1956 and 1979 are excluded since they are incomplete) which is a total of 22 years.

1.9.3.2 Extension of Records of Paso del Ermitaño and Rancho Arriba

For extending the records of Paso del Ermitaño and Rancho Arriba, three model forms are tentatively prescribed in the linear and normal domain of time series variables. These models are: i) nonseasonal multiple regression model, ii) seasonal multiple regression model, and

Table 1.9.4. Regression model parameters of Equation (1.9.2) for filling-in weekly missing values of Palo de Caja.

Year, ν	Week, τ	$a(\nu, \tau)$	$b(\nu, \tau)$	$c(\nu, \tau)$	$R(\nu, \tau)$
1964	33	-0.456	1.143	1.696	0.554
	34	0.363	1.079	0.310	0.971
	35	0.259	0.815	0.576	0.856
	42	0.360	0.809	1.072	0.579
	43	-0.197	0.936	0.776	0.813
	44	-0.324	0.613	0.964	0.507
	51	-0.324	0.613	0.964	0.507
	52	-0.324	0.613	0.964	0.507
1966	6	-0.324	0.613	0.964	0.507
	7	-0.843	0.608	0.882	0.644

iii) nonseasonal bivariate first-order autoregressive model. For the case of nonseasonal models, it is assumed that any inherent seasonality in the data is adequately removed by seasonal standardization which is included in the normalizing transformation used (see Appendix 1.9.A). In the multiple regression models, the choice of regressors were made on the basis of best regression correlation coefficients using one or two stations.

With these three model forms, sample extensions of monthly data were then performed. On this basis, it is found that the nonseasonal bivariate first-order autoregressive model best preserves the means, variances, autocorrelations and cross-correlations of the data. Thus, the short records of Paso del Ermitaño and Rancho Arriba for both monthly and weekly levels were extended using a bivariate first-order

autoregressive model with Palo de Caja. In equation form, this model is written as:

$$\begin{bmatrix} z_{\nu,\tau}^{(1)} \\ z_{\nu,\tau}^{(2)} \end{bmatrix} = \begin{bmatrix} a_{11} & a_{12} \\ a_{21} & a_{22} \end{bmatrix} \begin{bmatrix} z_{\nu,\tau-1}^{(1)} \\ z_{\nu,\tau-1}^{(2)} \end{bmatrix} + \begin{bmatrix} b_{11} & 0 \\ b_{21} & b_{22} \end{bmatrix} \begin{bmatrix} \epsilon_{\nu,\tau}^{(1)} \\ \epsilon_{\nu,\tau}^{(2)} \end{bmatrix} \quad (1.9.3a)$$

For purposes of data extension, the equation above is written as:

$$\begin{aligned} z_{\nu,\tau}^{(2)} = & \frac{b_{21}}{b_{11}} z_{\nu,\tau}^{(1)} + \left(a_{21} - \frac{a_{11}b_{21}}{b_{11}} \right) z_{\nu,\tau-1}^{(1)} \\ & + \left(a_{22} - \frac{b_{21}a_{12}}{b_{11}} \right) z_{\nu,\tau-1}^{(2)} + b_{22} \epsilon_{\nu,\tau}^{(2)} \end{aligned} \quad (1.9.3b)$$

where $z_{\nu,\tau}^{(2)}$ is the data of either Paso del Ermitaño or Rancho Arriba, $z_{\nu,\tau}^{(1)}$ is the data of Palo de Caja, $\epsilon_{\nu,\tau}^{(2)}$ is a standard normal random number, and the a's and b's are model parameters.

Since the data of Paso del Ermitaño and Rancho Arriba are to be extended forward in time and backward in time to coincide with those of Palo de Caja (see Table 1.9.1 for periods of available records), the model in Eqs. (1.9.3) are applied likewise forward and backward in time as the case requires. This scheme is possible for linear-normal time series models owing to the notions of time-reversibility and distributional symmetry in the linear and normal domain of time series variables.

The model parameters of Eqs. (1.9.3) are estimated using the method of moments (Salas, et al.; 1980). Tables 1.9.5 and 1.9.6 show the model parameters for monthly and weekly levels, respectively. Referring to

Table 1.9.5. Model parameters for monthly data extensions.

Forward in Time		A		B	
Paso del Ermitaño and	-0.0459	0.6325	0.8023	0.0	
Palo de Caja	0.0214	0.8250	0.3574	0.4040	
Rancho Arriba and	0.5817	-0.2896	0.8531	0.0	
Palo de Caja	-0.0757	0.5493	0.5819	0.6253	
Backward in Time		A		B	
Paso del Ermitaño and	-0.2035	0.8324	0.7300	0.0	
Palo de Caja	-0.1783	0.9825	0.2863	0.4440	
Ranch Arriba and	0.4710	-0.0408	0.8909	0.0	
Palo de Caja	-0.3245	0.6601	0.5396	0.5988	

Table 1.9.6. Model parameters for weekly data extensions.

Forward in Time		A		B	
Paso del Ermitaño and	0.5755	0.1963	0.6765	0.0	
Palo de Caja	0.0537	0.7764	0.3423	0.4620	
Rancho Arriba and	0.7617	-0.0725	0.6853	0.0	
Palo de Caja	-0.0382	0.7756	0.4206	0.4997	
Backward in Time		A		B	
Paso del Ermitaño and	0.5553	0.2226	0.6731	0.0	
Palo de Caja	0.0273	0.7965	0.3404	0.4644	
Ranch Arriba and	0.7343	-0.0176	0.6881	0.0	
Palo de Caja	-0.0932	0.8030	0.4164	0.4977	

Eq. (1.9.3b), the extension of records is done by generating the normal standard random noise $\epsilon_{\nu, \tau}^{(2)}$ to obtain $z_{\nu, \tau}^{(2)}$, then followed by back transformation to the original domain of flows. Due to the addition of the random term, a single series is only one possible (equally likely) sequence that may have occurred. In view of this, fifty data extensions are made. The intent here is simply to utilize the fifty extended

series for improving the estimates of the parameters to be used in modeling and data generation.

For each extended series, the seasonal statistics are then computed constituting a total of 50 samples of statistics for each season. From these, the averages and standard errors of each statistic are determined. The averages in this case are considered as the improved estimates representing the seasonal statistic while their corresponding standard errors indicate the degree of variation of each statistic from the averages. The computed averages and standard errors for monthly and weekly statistics are given in Appendix 1.9.B.

A reliability test for the improvement of the means and variances was performed after extending the records of Paso del Ermitaño and Rancho Arriba. The test is that the variances of historical means or variances should be greater than those of the extended ones under the null hypothesis that these statistics are improved. The variances of historical means and variances for seasonally autocorrelated processes are given respectively as:

$$\text{var}(\bar{x}_\tau) = \frac{S_\tau^2}{N} \left[1 + \frac{2}{N} \sum_{\nu=1}^{N-1} (N - \nu) \rho_\tau(w\nu) \right] \quad (1.9.4a)$$

and

$$\text{var}(S_\tau^2) = 2 S_\tau^4 \left[\frac{N}{N-1} - 1 \right]^{-1} \left[1 + \sum_{\nu=1}^2 \rho_\tau^2(w\nu) \right] \quad (1.9.4b)$$

where S_τ^2 is the historical variance at season τ , N is the number of years of record, w is the number of seasons and $\rho_\tau(w\nu)$ is the

seasonal autocorrelation at lag $w\nu$. Assuming that the historical series follows a seasonal first-order autoregressive process, the

$$\rho_{\tau}(w\nu) = \rho_{\tau}(1) \rho_{\tau-1}(1) \dots \rho_{\tau-w\nu+1}(1)$$

in which $\rho_{\tau}(1)$ is the lag-1 autocorrelation at season τ . The variances of the extended series statistics are the square of the standard errors computed from each statistic based on fifty series extensions.

Results from these tests show that for all monthly means and variances, all extended statistics are improved. For the weekly statistics, a maximum of 2 weeks out of 52 weeks failed the test. For all practical purposes, the extension of records has definitely improved the reliability of statistical parameters.

1.9.4 STOCHASTIC MODELING OF STREAMFLOWS

The stochastic models for both monthly and weekly streamflows adopted herein to be eventually used in data generation belong to the family of multivariate linear models. Similar to the models used in extension, the models are developed by first normalizing and standardizing each series. Thereafter the correlation structure of the residual series is derived to form the basis of the stochastic model. In normalization, the combination of logarithmic and Wilson-Hilferty transformations is used for both monthly and weekly streamflows. For Wilson-Hilferty transformation, the seasonal skewness coefficients (in the logarithmic domain of flows) are Fourier fitted functions using the first two harmonics for monthly skews and first four harmonics for

weekly skews. The skewness coefficients were also corrected for bias under the assumption that the process in the log domain are approximately gamma distributed. Details of the Fourier series fitting of the seasonal skewness coefficients are given in Appendix 1.9.C.

Three alternative stochastic model formulations were tried in this study. The first model, referred to here as "MODEL A," is a contemporaneous seasonal mixed autoregressive-moving average model (ARMA) which involves fitting first appropriate seasonal univariate ARMA models to each series. The model residuals of each series are then obtained and fitted to a zero-order multivariate model. In equation form, the univariate seasonal ARMA model can be generally written as:

$$z_{\nu,\tau}^{(s)} = \sum_{i=1}^p \phi_{\tau,i}^{(s)} z_{\nu,\tau-i}^{(s)} + \sum_{j=1}^q \theta_{\tau,j}^{(s)} e_{\nu,\tau-j}^{(s)} + e_{\nu,\tau}^{(s)} \quad (1.9.5)$$

where $z_{\nu,\tau}^{(s)}$ is the normalized and standardized series of station s , year ν and season τ ; $e_{\nu,\tau}^{(s)}$ is the residual series independent in time but the residuals of each station are contemporaneously correlated with each; and $\phi_{\tau,i}^{(s)}$ and $\theta_{\tau,j}^{(s)}$ are model parameters of orders p and q , respectively. For both monthly and weekly streamflows, the model parameters are estimated based on Fourier fitted functions of the seasonal sample autocorrelations. Details of the Fourier series fitting of autocorrelations for monthly and weekly levels of the three stations are given in Appendix 1.9.C. The autocorrelations fitted with the first two harmonics were used for monthly models and the fitted first four harmonics were used for weekly models. The zero-order multivariate model written for three stations takes the form below:

$$\begin{bmatrix} e_{\nu,\tau}^{(1)} \\ e_{\nu,\tau}^{(2)} \\ e_{\nu,\tau}^{(3)} \end{bmatrix} = \begin{bmatrix} b_{11} & 0 & 0 \\ b_{21} & b_{22} & 0 \\ b_{31} & b_{32} & b_{33} \end{bmatrix} \begin{bmatrix} \epsilon_{\nu,\tau}^{(1)} \\ \epsilon_{\nu,\tau}^{(2)} \\ \epsilon_{\nu,\tau}^{(3)} \end{bmatrix} \quad (1.9.6)$$

where $\epsilon_{\nu,\tau}^{(s)}$'s are identically and independently distributed normal deviates and the b's are nonseasonal model parameters. The second model tried in this study uses both Eqs. (1.9.5) and (1.9.6) except that in Eq. (1.9.6), the model parameter b's are allowed to vary seasonally. This latter model is referred to here as "MODEL B." The third model, referred to as "MODEL C," is a nonseasonal vector first-order autoregressive model written for three stations as:

$$\begin{bmatrix} e_{\nu,\tau}^{(1)} \\ e_{\nu,\tau}^{(2)} \\ e_{\nu,\tau}^{(3)} \end{bmatrix} = \begin{bmatrix} a_{11} & a_{12} & a_{13} \\ a_{21} & a_{22} & a_{23} \\ a_{31} & a_{32} & a_{33} \end{bmatrix} \begin{bmatrix} z_{\nu,\tau-1}^{(1)} \\ z_{\nu,\tau-1}^{(2)} \\ z_{\nu,\tau-1}^{(3)} \end{bmatrix} + \begin{bmatrix} b_{11} & 0 & 0 \\ b_{21} & b_{22} & 0 \\ b_{31} & b_{32} & b_{33} \end{bmatrix} \begin{bmatrix} \epsilon_{\nu,\tau}^{(1)} \\ \epsilon_{\nu,\tau}^{(2)} \\ \epsilon_{\nu,\tau}^{(3)} \end{bmatrix} \quad (1.9.7)$$

where the a's are nonseasonal model parameters and the other notations are defined as in Eqs. (1.9.5) and (1.9.6).

For the three tentative models above, the model parameters were estimated using the method of moments presented in Salas, et al., 1980. In choosing the appropriate model order of univariate seasonal ARMA in Eq. 1.9.5, the computer program UMOSE1 developed by Salas and Smith (1981) was used. A seasonal ARMA(1,0) or AR(1) model is found adequate and parsimonious to describe both monthly and weekly flows.

1.9.5 STREAMFLOW DATA GENERATION

1.9.5.1 Data Generation Scheme

For the three alternative stochastic models given previously, the data generation commences by generating the sequence $\epsilon_{\nu, \tau}^{(s)}$ using the Box-Muller formula written as (Salas, et al., 1980):

$$\epsilon_1 = [2 \ln (1/u_1)]^{1/2} \cos (2\pi u_2) \quad (1.9.8a)$$

and

$$\epsilon_2 = [2 \ln (1/u_1)]^{1/2} \sin (2\pi u_2) \quad (1.9.8b)$$

where u_1 and u_2 are two independent uniformly distributed (0,1) random numbers. Note that two random numbers can be generated at one time. Then these generated values are applied to either Eq. (1.9.6) or (1.9.7). For models A and B, Eq. (1.9.6) is used with nonseasonal or seasonal b 's respectively to obtain $e_{\nu, \tau}^{(s)}$, followed by Eq. (1.9.5) to arrive at $z_{\nu, \tau}^{(2)}$. For model C, Eq. (1.9.7) is used to arrive at $z_{\nu, \tau}^{(s)}$.

Having obtained the sequence $z_{\nu, \tau}^{(s)}$ for any model, the backward Wilson-Hilferty transformation is applied using the inverse of Eq. (1.9.A.6) such that

$$x'_{\nu, \tau}^{(s)} = \frac{2}{G_{\tau}^{(s)}(x)} \left\{ \frac{G_{\tau}^{(s)}(x)}{6} \left[z_{\nu, \tau}^{(s)} - \frac{G_{\tau}^{(s)}(x)}{6} \right] + 1 \right\}^3 - \frac{2}{G_{\tau}^{(s)}(x)} \quad (1.9.9)$$

in which $G_{\tau}^{(s)}(x) \neq 0$, where $G_{\tau}^{(s)}(x)$ is the seasonal Fourier fitted skewness coefficient in the log domain. The variable $x'_{\nu, \tau}^{(s)}$ is further transformed by

$$x_{\nu, \tau}^{(s)} = \begin{cases} \max \left[x_{\nu, \tau}'^{(s)}, -2/G_{\tau}^{(s)}(x) \right] & \text{if } G_{\tau}^{(s)}(x) > 0 \\ \min \left[x_{\nu, \tau}'^{(s)}, -2/G_{\tau}^{(s)}(x) \right] & \text{if } G_{\tau}^{(s)}(x) < 0 \end{cases}$$

and

$$x_{\nu, \tau}^{(s)} = z_{\nu, \tau}^{(s)} \quad \text{if } G_{\tau}^{(s)}(x) = 0 \quad (1.9.10)$$

where $x_{\nu, \tau}^{(s)}$ is the generated data in the log domain. Finally, the data $y_{\nu, \tau}^{(s)}$ in the actual or original domain is obtained from

$$w_{\nu, \tau}^{(s)} = \bar{w}_{\tau}^{(s)} + S_{\tau}^{(s)}(w) x_{\nu, \tau}^{(s)} \quad (1.9.11)$$

and

$$y_{\nu, \tau}^{(s)} = \exp \left[w_{\nu, \tau}^{(s)} \right] \quad (1.9.12)$$

where $\bar{w}_{\tau}^{(s)}$ and $S_{\tau}^{(s)}(w)$ are the seasonal mean and standard deviation in the log-domain of flows.

The program GENSEA developed by Salas and Smith (1981) was utilized for data generation which was slightly modified for purposes here.

1.9.5.2 Analysis of Generated Data

A total of 50 samples of size 22 years each of monthly and weekly streamflows were generated using the three alternative model formulations above. The best monthly and weekly models were then selected based on the comparison of historical and generated seasonal means, standard deviations, skewness coefficients, autocorrelations, and cross-correlations. For each generated sample, the above mentioned statistical properties are computed, then the arithmetic averages and standard errors for all samples are determined. Given in Appendix 1.9.D are plots of the monthly and weekly historical and generated statistics

of the three models in the original domain of flows. Only the computed averages are shown for clarity in presentation. Based on these plots, the best model selected for both monthly and weekly streamflows is MODEL B (i.e., univariate seasonal first-order autoregressive process with seasonal multivariate zero-order station-to-station dependence).

Given in Appendix 1.9.E are the monthly and weekly, historical and generated statistics of MODEL B in the three domains of streamflows, namely: original domain, log domain, and log-Wilson-Hilferty domain. Plotted along the generated average statistics are the positive and negative one-standard errors relative to these averages. In general, the results show that the historical statistics are satisfactorily reproduced in the different domain of flows. Notice that in almost all cases, the confidence bands of the generated statistics encloses the historical statistics. It may be noted also that in the different domain of flows, the mean, standard deviations are best reproduced in the log-domain while the auto- and cross-correlations are best reproduced in the log-Wilson-Hilferty domain. This is only logical since such statistical properties are parameterized in the model at these corresponding domain of flows.

1.9.6 MODELING AND GENERATION OF TURBINE OPERATING HOURS

A preliminary analysis done in this study showed that the monthly inflows to Valdesia reservoir is significantly correlated to the monthly turbine operating hours of the said reservoir based on 9 years of data covering the period of 1976 to 1984. On this basis, it is decided to model and generate the turbine operating hours monthly time series based on reservoir inflows. Due to the proximity of Paso del Ermitaño and

Valdesia reservoir, it is assumed that the said reservoir inflows are the same streamflows as those of Paso del Ermitaño. Subsequently, the turbine operating hours can be generated by solely using the generated streamflows of Paso del Ermitaño. However, since the available turbine operating hours (i.e., 1976-1984 period) has no corresponding recorded streamflow at Paso del Ermitaño, the monthly inflow data at Valdesia reservoir are used to derive the stochastic model for data generation.

The model selected for purposes here is a bivariate, contemporaneous, first-order autoregressive model with monthly parameters. Specifically, the model takes the following form:

$$\begin{bmatrix} z_{\nu,\tau}^{(1)} \\ z_{\nu,\tau}^{(2)} \end{bmatrix} = \begin{bmatrix} a_{11,\tau} & 0 \\ 0 & a_{22,\tau} \end{bmatrix} \begin{bmatrix} z_{\nu,\tau-1}^{(1)} \\ z_{\nu,\tau-1}^{(2)} \end{bmatrix} + \begin{bmatrix} b_{11,\tau} & 0 \\ b_{21,\tau} & b_{22,\tau} \end{bmatrix} \begin{bmatrix} \epsilon_{\nu,\tau}^{(1)} \\ \epsilon_{\nu,\tau}^{(2)} \end{bmatrix} \quad (1.9.13)$$

where $z_{\nu,\tau}^{(1)}$ is the Paso del Ermitaño streamflows, $z_{\nu,\tau}^{(2)}$ is the turbine operating hours, and, a's and b's are model parameters. The model above is similar to the model used for data extension (see Eq. 1.9.3a, Section 1.9.3.2) except for the seasonality of its model parameters and the parametric matrix of a's which renders the model as contemporaneous (Salas et al., 1985). For generating the turbine operating hours $z_{\nu,\tau}^{(2)}$, Eq. (1.9.13) can be rewritten as

$$z_{\nu,\tau}^{(2)} = \frac{b_{21,\tau}}{b_{11,\tau}} z_{\nu,\tau}^{(1)} - \frac{a_{11,\tau} b_{21,\tau}}{b_{11,\tau}} z_{\nu,\tau-1}^{(1)} + a_{22,\tau} z_{\nu,\tau-1}^{(2)} + b_{22,\tau} \epsilon_{\nu,\tau}^{(2)} \quad (1.9.14)$$

The time series variables $z_{\nu,\tau}^{(1)}$ and $z_{\nu,\tau}^{(2)}$ are monthly normalized and standardized using the Wilson-Hilferty transformation (see Appendix

1.9.A). The use of Wilson-Hilferty (WH) transformation alone for Paso del Ermitaño instead of in combination with logarithmic transformation as done previously is simply for convenience and consistency with respect to the type of transformation used for turbine hours. Besides, either the log-WH or plain WH transformations are valid for Paso del Ermitaño monthly streamflows as shown in Table 1.9.A.1.

The monthly model parameters estimated for Eq. (1.9.14) are given in Table 1.9.7. Data generation of turbine operating hours follows in the same manner as the extension of records in section 1.9.3.2. Once, $z_{\nu, \tau}^{(2)}$ in the Wilson-Hilferty domain of turbine hours are obtained, the backtransformation to its original domain is performed as in section 1.9.5.1 using Eqs. (1.9.9), (1.9.10) and (1.9.11), i.e., without the anti-log operation.

Table 1.9.7. Monthly parameters of turbine operating hours model in Eq. (1.9.14).

Month, τ	$a_{11, \tau}$	$a_{22, \tau}$	$b_{11, \tau}$	$b_{21, \tau}$	$b_{22, \tau}$
1	0.6694	0.9753	0.7429	0.1002	0.1969
2	0.3083	0.1667	0.9513	0.2410	0.9561
3	0.5255	0.2728	0.8508	0.4014	0.8744
4	0.5900	0.6270	0.8074	-0.2026	0.7522
5	0.4592	0.0780	0.8883	0.5654	0.8211
6	0.4547	0.5428	0.8906	0.6946	0.4721
7	0.5435	0.5640	0.8394	0.3352	0.7547
8	0.3396	-0.1422	0.9406	0.4205	0.8961
9	0.1236	0.4279	0.9923	0.6525	0.6254
10	0.3649	0.5630	0.9310	0.3378	0.7543
11	0.1527	0.4804	0.9883	0.2200	0.8490
12	0.4888	0.2435	0.8724	0.7023	0.6690

As in streamflow data generation, 50 samples of 22 years of monthly turbine operation hours were generated. A comparison of important statistical parameters between historical and generated turbine hours.

was also performed. Results of the comparison are shown in Appendix 1.9.F which show satisfactory reproduction of monthly means, standard deviations skewness coefficients, autocorrelations and cross-correlations.

1.9.7 FINAL REMARKS

Considerable data analysis and manipulations have been done prior to data modeling and generation. First of all, the missing data of Palo de Caja has to be filled-in using an appropriate regression model. Then data extension has to be performed for Palo del Ermitaño and Rancho Arriba. In data extension, 50 series extensions are made where in principle, any one of the extended series can be used for deriving the statistical parameters for modeling and data generation. However, since our objective is to improve the reliability of these statistical estimates, the 50 series extensions were used to represent the statistical parameters in terms of averages.

As a requirement for linear normal models in this study, the combination of logarithmic and Wilson-Hilferty transformation is proven to be effective in normalizing the streamflow. It may be noted that the Wilson-Hilferty transformation is developed on the basis of the Pearson Type III (gamma) distribution. Thus, it can be said that streamflow data follows a log-Pearson Type III distribution.

In time series modeling of the extended data, three alternative stochastic models were found. The main reason for tentatively prescribing three model forms is for us to select the best model in terms of model parsimony (i.e., economy of parameters) and overall modeling efficiency without compromising the adequacy and ability of the

final model adopted to represent the time series process studied at hand. The paper by Salas, et al., 1985 may be consulted for further elaborations on alternative multivariate models similar to the ones used here, and applications of multivariate models in general.

REFERENCES

- Matalas, N. C. 1967. Mathematical assessment of synthetic hydrology. Water Resources Research, Vol. 3, No. 4, 4th Quarter, pp. 931-945.
- McGinnis, D. F., Jr. and W. H. Sammons. 1970. discussion of "Daily streamflow simulation" by K. Payne, W. D. Neumann and K. D. Kerri. Jour. of the Hydraulics Div., Proc. ASCE, Vol. 96, No. HY5, May, pp. 1201-1206.
- Salas, J. D., J. W. Delleur, V. Yevjevich and W. L. Lane. 1980. Applied Modeling of Hydrologic Time Series. Water Resources Publ., Littleton, CO.
- Salas, J. D. and R. A. Smith. 1981. Computer Programs for Modeling and Generation of Hydrologic Time Series. Dept. of Civil Engineering, Colorado State University, Fort Collins, CO, July.
- Salas, J. D., G. Q. Tabios and P. Bartolini. 1985. Approaches to multivariate modeling of water resources time series. Water Resources Bulletin, AWRA, Vol. 21, No. 4, August, pp. 683-708.
- Stedinger, J. R. and M. R. Taylor. 1982. Synthetic streamflow generation, 1. Model verification and validation. Water Resources Research, Vol. 18, No. 4, August, pp. 909-918.
- Yevjevich, V. and J. T. B. Obeysekera. 1984. Estimation of skewness of hydrologic variables. Water Resources Research, Vol. 20, No. 7, July, pp. 935-943.

APPENDIX 1.9.A

STANDARDIZATION AND NORMALIZATION

The purpose of standardization is to remove the seasonalities in the means and variances of the data. Removal of such seasonalities may be made by using the raw estimates of the means and variances or their corresponding smoothed estimates by say, Fourier series fitting. For this study, the raw estimates of the semi-monthly means and variances are used. As mentioned in section 1.9.3, a transformation is made to render the data normal. Currently, there are various transformations of frequent use in hydrology such as square-root, cube-root, logarithmic, Wilson-Hilferty or log-Wilson-Hilferty transformations. Depending on the type of normalization (transformation) used, standardization is applied before or after transformation. The five normalizing transformations mentioned above were used in this study.

1) Square-root transformation:

Denoting the original time series data by $Y_{\nu, \tau}$ of year ν and season τ , the square-root transformation is done by

$$X_{\nu, \tau} = \sqrt{Y_{\nu, \tau}} \quad (1.9.A.1)$$

for all $\nu = 1, \dots, n$ years and $\tau = 1, \dots, w$ seasons. Then standardization follows given by:

$$Z_{\nu, \tau} = \frac{X_{\nu, \tau} - \bar{X}_{\tau}}{S_{\tau}(x)} \quad (1.9.A.2)$$

where \bar{X}_τ and $S_\tau(x)$ are the semi-monthly mean and standard deviation of the square-root transformed data.

2) Cube-root transformation:

In this case, the original series $Y_{\nu,\tau}$ is transformed using the equation

$$X_{\nu,\tau} = Y_{\nu,\tau}^{1/3} \quad (1.9.A.3)$$

then, standardization is performed using Eq. (1.8.A.2) where \bar{X}_τ and $S_\tau(x)$ are evaluated using the cube-root transformed data.

3) Logarithmic transformation:

The raw series $Y_{\nu,\tau}$ is transformed by

$$X_{\nu,\tau} = \log(Y_{\nu,\tau}) \quad (1.9.A.4)$$

where \log stands for the base-e logarithms. This is followed by standardization as in Eq. (1.9.A.2).

4) Wilson-Hilferty transformation:

The original series $Y_{\nu,\tau}$ is first standardized as in Eq. (1.9.A.2) by

$$X_{\nu,\tau} = \frac{Y_{\nu,\tau} - \bar{Y}_\tau}{S_\tau(Y)} \quad (1.9.A.5)$$

where \bar{Y}_τ and $S_\tau(Y)$ are the semi-monthly mean and standard deviation of the series $\{Y_{\nu,\tau}\}$. The Wilson-Hilferty transformation is given by (Matala, 1967)

$$Z_{\nu,\tau} = \frac{6}{G_\tau(x)} \left\{ \left[\frac{G_\tau(x) X'_{\nu,\tau}}{2} + 1 \right]^{1/3} - 1 \right\} + \frac{G_\tau(x)}{6} \quad (1.9.A.6)$$

which is valid for $G_\tau(x) \neq 0$, where $G_\tau(x)$ is the semi-monthly skewness coefficient of $X_{\nu,\tau}$ and $X'_{\nu,\tau}$ is given by McGinnis and Sammons (1970) as

$$X'_{\nu,\tau} = \begin{cases} \max[X_{\nu,\tau}, -2/G_\tau(x)] & \text{if } G_\tau(x) \geq 0 \\ \min[X_{\nu,\tau}, -2/G_\tau(x)] & \text{if } G_\tau(x) < 0 \end{cases} \quad (1.9.A.7)$$

If $G_\tau(x) = 0$ no transformation is necessary then $Z_{\nu,\tau} = X_{\nu,\tau}$.

5) Log-Wilson-Hilferty transformation:

This transformation is a combination of logarithmic transformation (item 3) and Wilson-Hilferty transformation (item 4). For the sake of clarity, let us rewrite some of the equations above. First the original data $Y_{\nu,\tau}$ is transformed as in Eq. (1.9.A.4) as

$$W_{\nu,\tau} = \log(Y_{\nu,\tau}) \quad (1.9.A.8)$$

Then, standardization is performed using

$$X_{\nu,\tau} = \frac{W_{\nu,\tau} - \bar{W}_\tau}{S_\tau(w)} \quad (1.9.A.9)$$

where \bar{W}_τ and $S_\tau(w)$ are the seasonal means and standard deviations of the logarithmic transformed sequence $(W_{\nu,\tau})$. Then, the Wilson-Hilferty transformation is performed using Eqs. (1.9.A.6) and (1.9.A.7) to arrive at $Z_{\nu,\tau}$.

The criteria for selecting the type of normalizing transformation to be used in this study is the skewness test for normality. This test assumes that if the observations are independent and sampled from the normal distribution, then, the sample skewness coefficient must fall within the $(1-\alpha)$ confidence limits

$$[-u_{1-\alpha/2} \sqrt{6/n}, u_{1-\alpha/2} \sqrt{6/n}] \quad (1.9.A.10)$$

where $u_{1-\alpha/2}$ is the $1-\alpha/2$ quantile of the standard normal distribution and n is the sample size. Since we have a total of 12 monthly and 52 weekly skewness coefficients computed after each transformation, the relative number of skews within the confidence limits are counted. The greatest number of passes forms the basis of choice of the suitable transformation used here.

The results of these analysis are given in Tables 1.9.A.1 and 1.9.A.2 for monthly and weekly levels, respectively. From these results it can be concluded that the streamflow data of Palo de Caja, Paso del Ermitaño, Rancho Arriba and El Cacao be normalized using the log-Wilson-Hilferty transformation, while the rainfall data of La Laguna and Valdesia be normalized using Wilson-Hilferty transformation.

Table 1.9.A.1. Relative scores of passing and failing the skewness test on a monthly basis using the five normalizing transformations including no transformation.

	No Trans- formation	Square- Root	Cube- Root	Logar- ithmic	Wilson- Hilferty	Log- Wilson- Hilferty
Palo de Caja						
Pass	0	3	6	10	8	12*
Fail	12	9	6	2	4	0
Paso del Ermitaño						
Pass	5	9	10	12	11	12*
Fail	7	3	2	0	1	0
Rancho Arriba						
Pass	10	10	10	10	11	11*
Fail	2	2	2	2	1	1
El Cacao						
Pass	6	12	11	9	12	12*
Fail	6	0	1	3	0	0
La Laguna						
Pass	7	11	8	2	12*	5
Fail	5	1	4	10	0	7
Valdesia						
Pass	5	10	11	9	12*	10
Fail	7	2	1	3	0	2

*Indicate transformation used.

Table 1.9.A.2. Relative scores of passing and failing the skewness test on a weekly basis using the five normalizing transformations including no transformation.

	No Trans- formation	Square- Root	Cube- Root	Logar- ithmic	Wilson- Hilferty	Log- Wilson- Hilferty
<hr/>						
Palo de Caja						
Pass	1	30	28	30	28	48*
Fail	51	22	24	22	24	4
<hr/>						
Paso del Ermitaño						
Pass	33	47	47	48	47	52*
Fail	19	5	5	4	5	0
<hr/>						
Rancho Arriba						
Pass	36	43	48	43	49	50*
Fail	16	9	4	9	3	2
<hr/>						
El Cacao						
Pass	29	46	45	48	46	52*
Fail	23	6	7	4	6	0
<hr/>						
La Laguna						
Pass	10	41	50	41	51*	34
Fail	42	11	2	11	1	18
<hr/>						
Valdesia						
Pass	7	36	44	37	46*	32
Fail	45	16	8	15	6	20

*Indicate transformation used.

APPENDIX 1.9.B

HISTORICAL AND EXTENDED SERIES STATISTICS OF MONTHLY AND
WEEKLY DATA OF PASO DEL ERMITAÑO AND RANCHO ARRIBA

Note: In the figures given below, the extended series statistics are averages computed from the fifty series extensions and correspondingly the positive and negative one-standard errors relative to these averages.

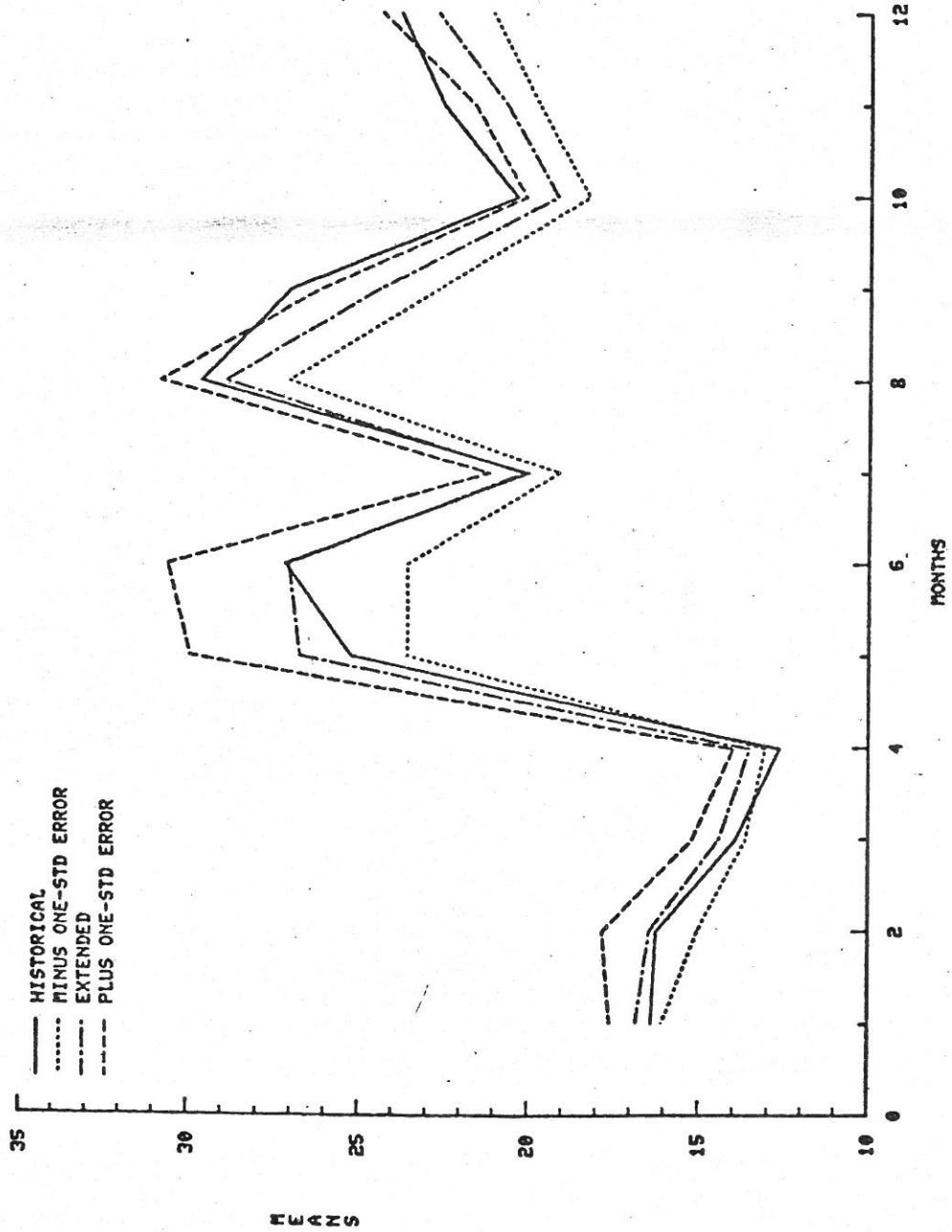


Figure 1.9.B.1. Historical and extended series monthly means of Paso del Ermitaño in original domain of flows.

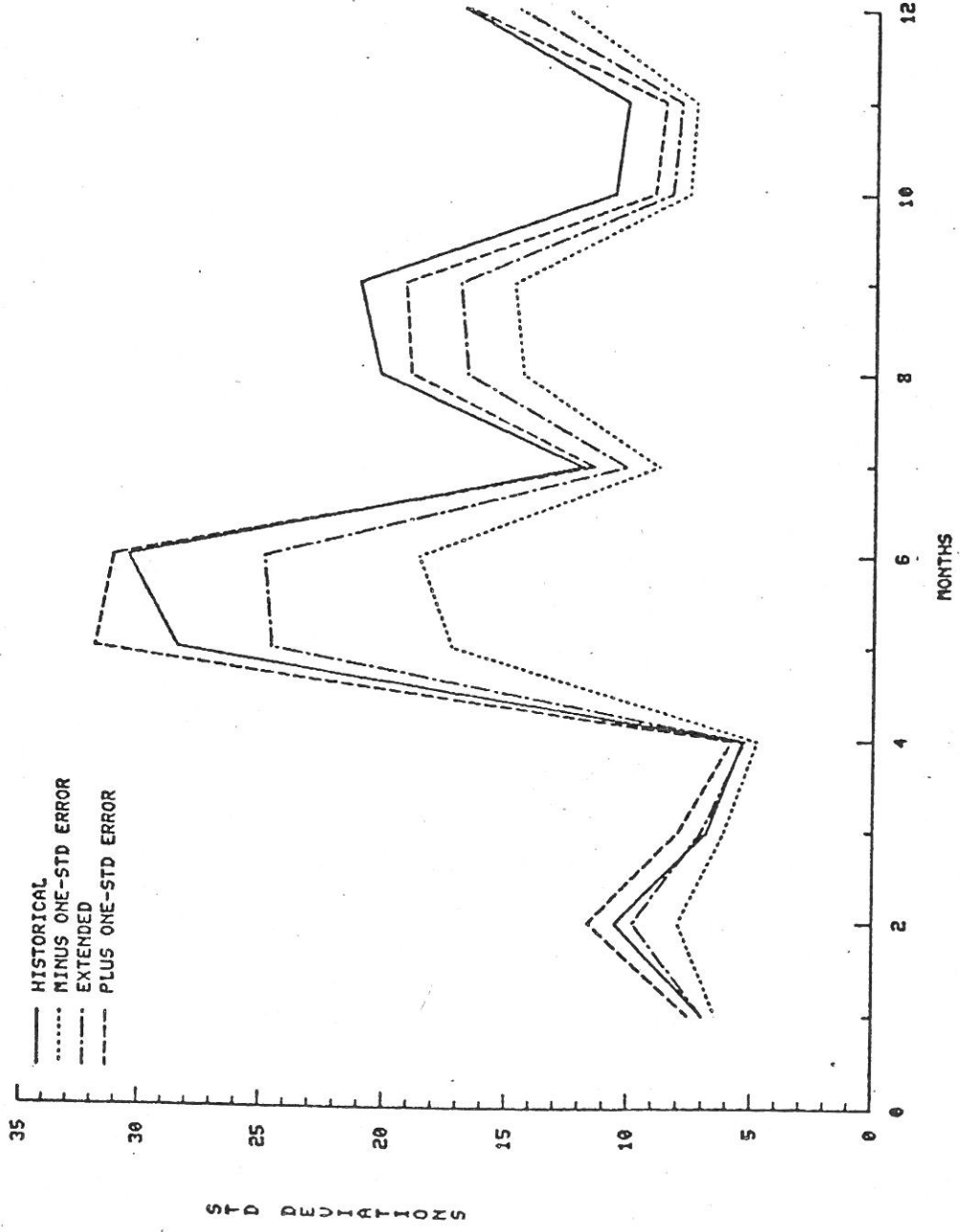


Figure 1.9.B.2. Historical and extended series monthly standard deviations of Paso del Ermitaño in original domain of flows.

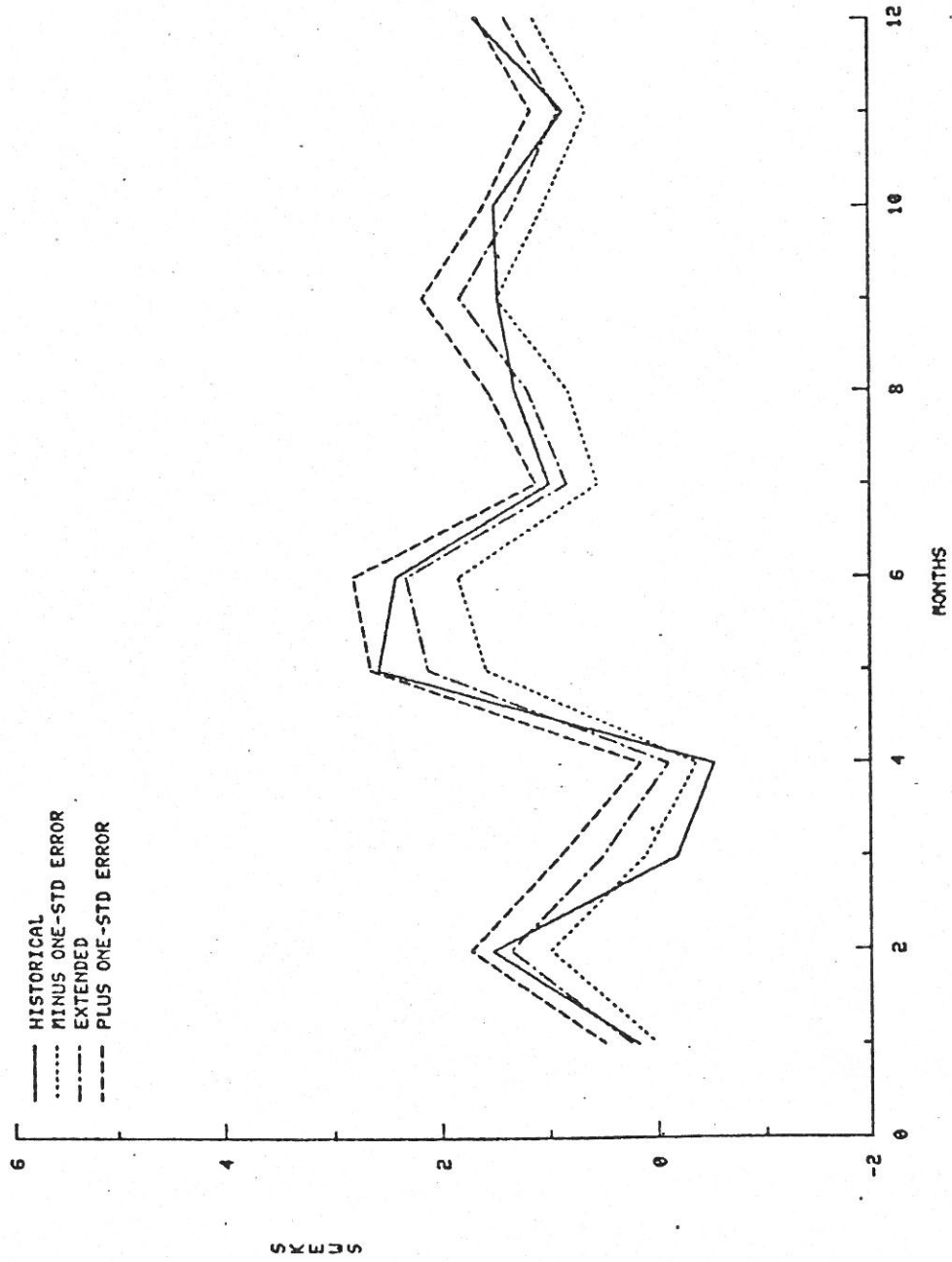


Figure 1.9.B.3. Historical and extended series monthly skewness coefficients of Paso del Ermitaño in original domain of flows.

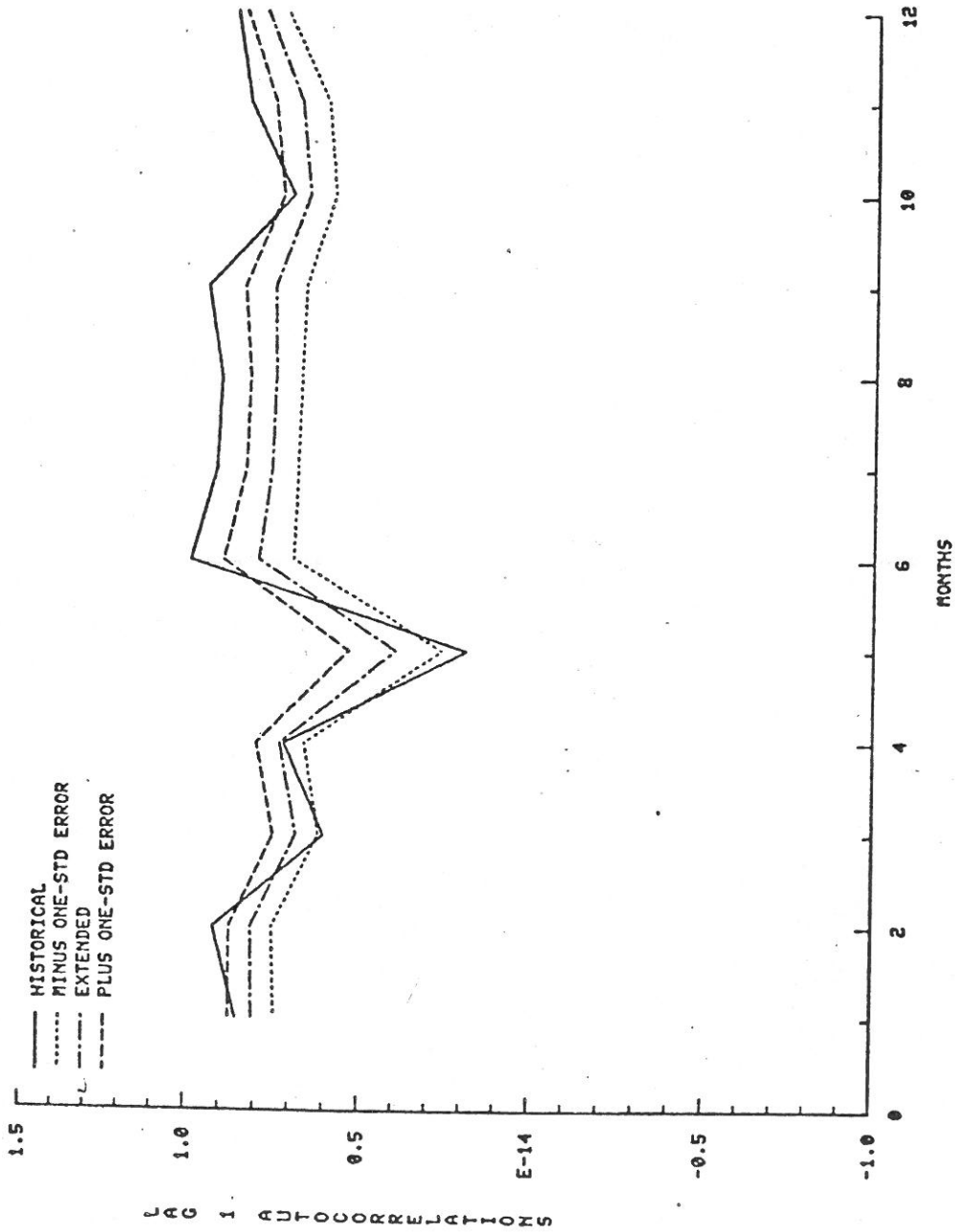


Figure 1.9.B.4. Historical and extended series monthly lag-1 autocorrelation coefficients of Paso del Ermitaño in original domain of flows.

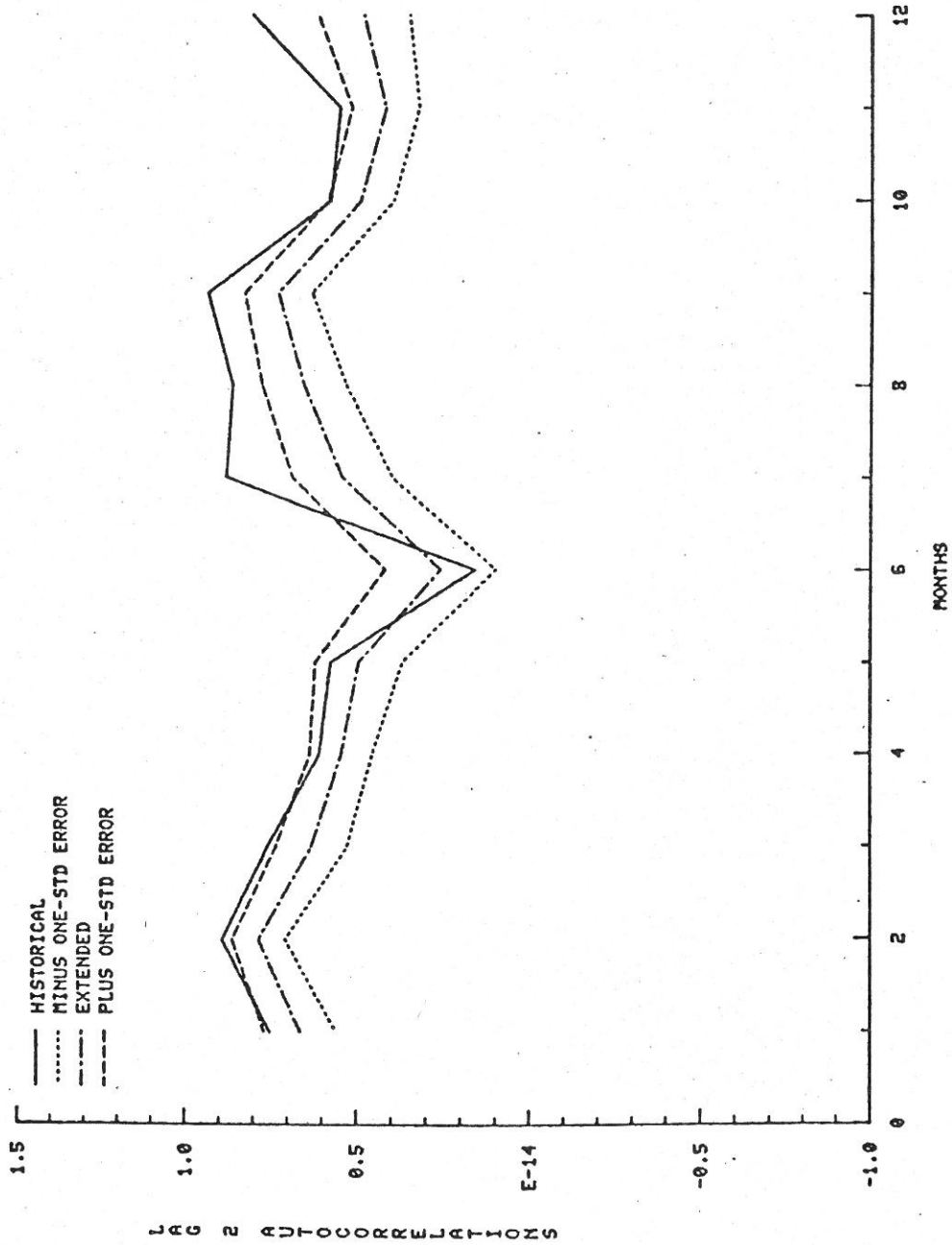


Figure 1.9.B.5. Historical and extended series monthly lag-2 autocorrelation coefficients of Paso del Ermitaño in original domain of flows.

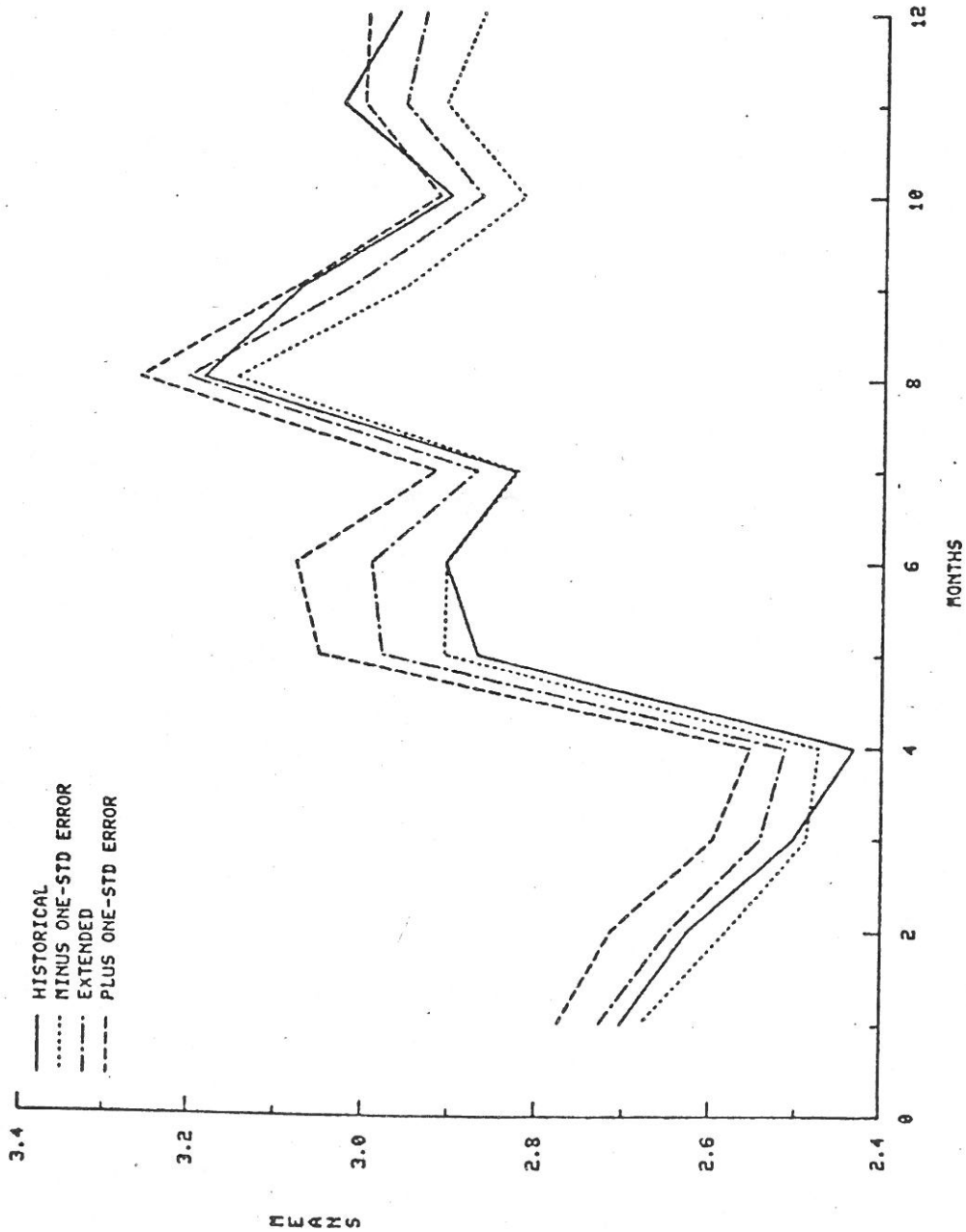


Figure 1.9.B.6. Historical and extended series monthly means of Paso del Ermitaño in logarithmic domain of flows.

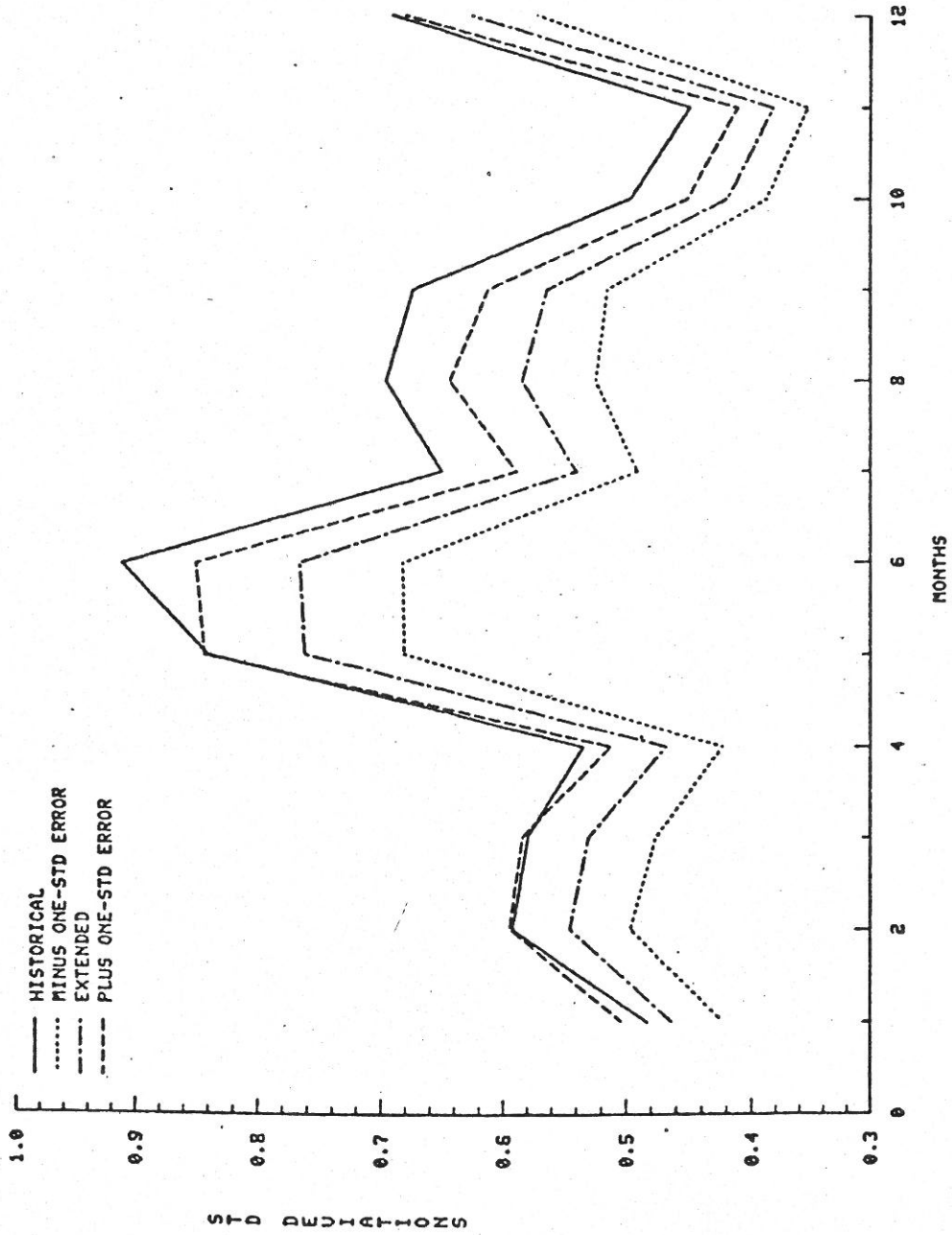


Figure 1.9.B.7. Historical and extended series monthly standard deviations of Paso del Ermitaño in logarithmic domain of flows.

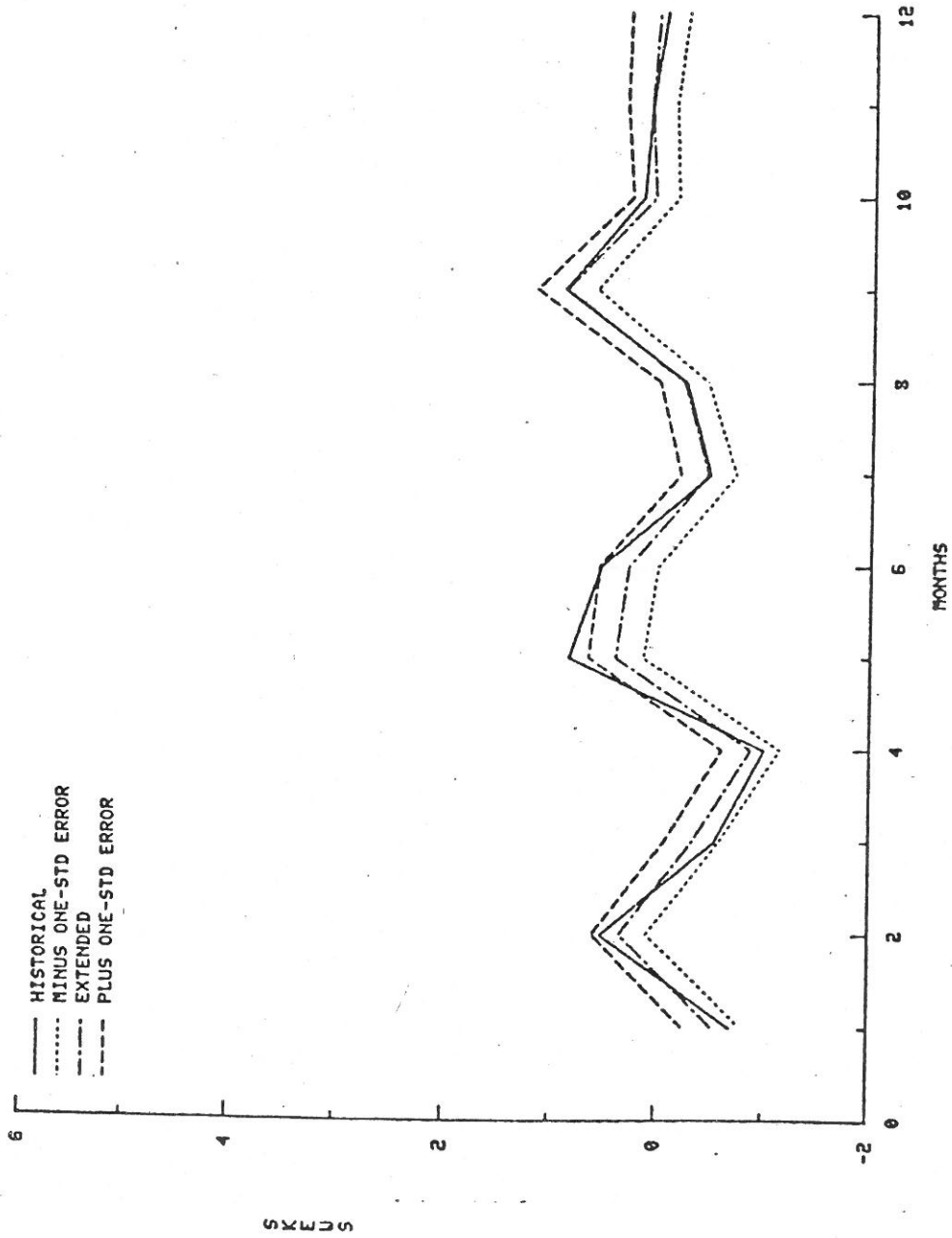


Figure 1.9.B.8. Historical and extended series monthly skewness coefficients of Paso del Ermitaño in logarithmic domain of flows.

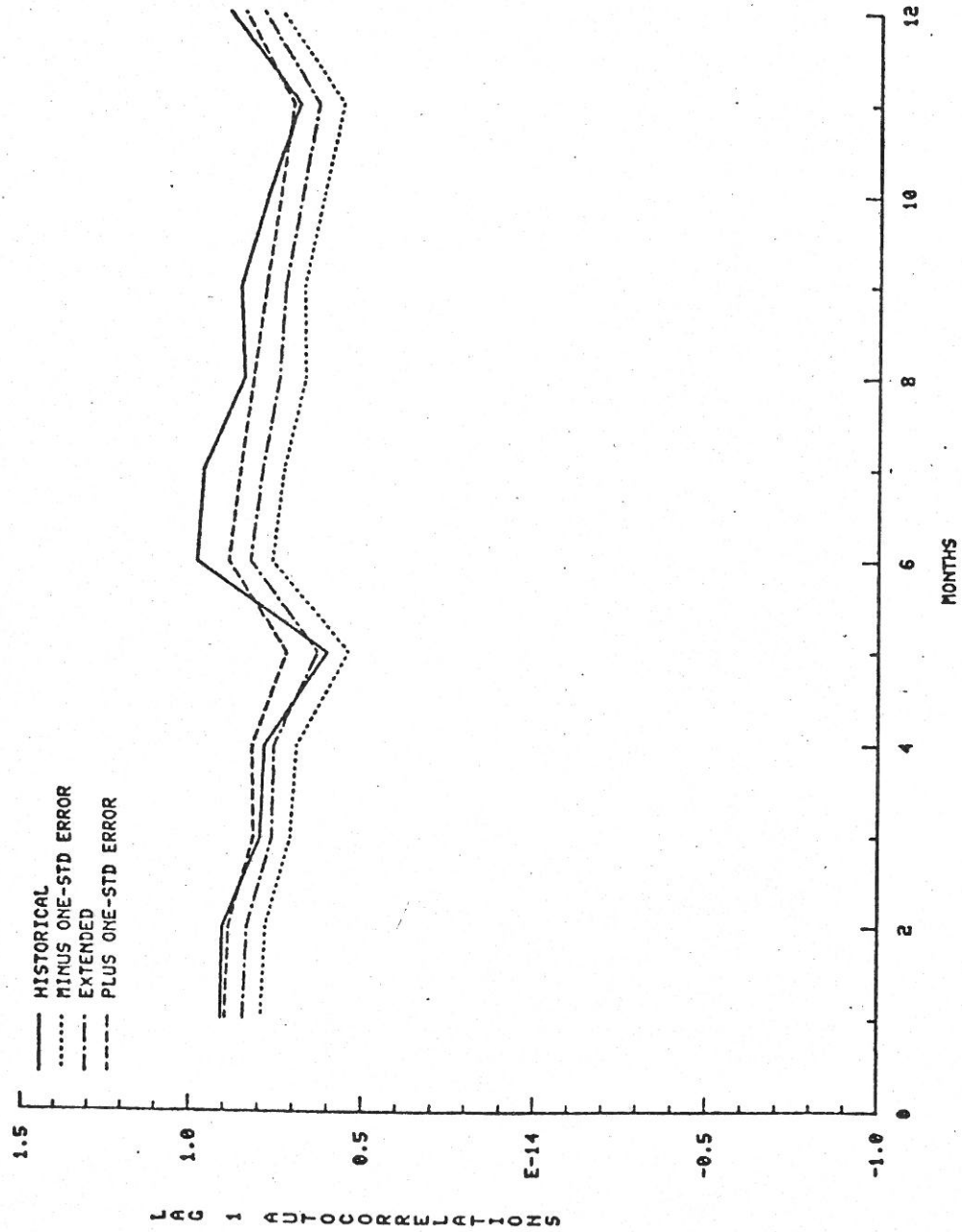


Figure 1.9.B.9. Historical and extended series monthly lag-1 autocorrelation coefficients of Paso del Ermitaño in logarithmic domain of flows.

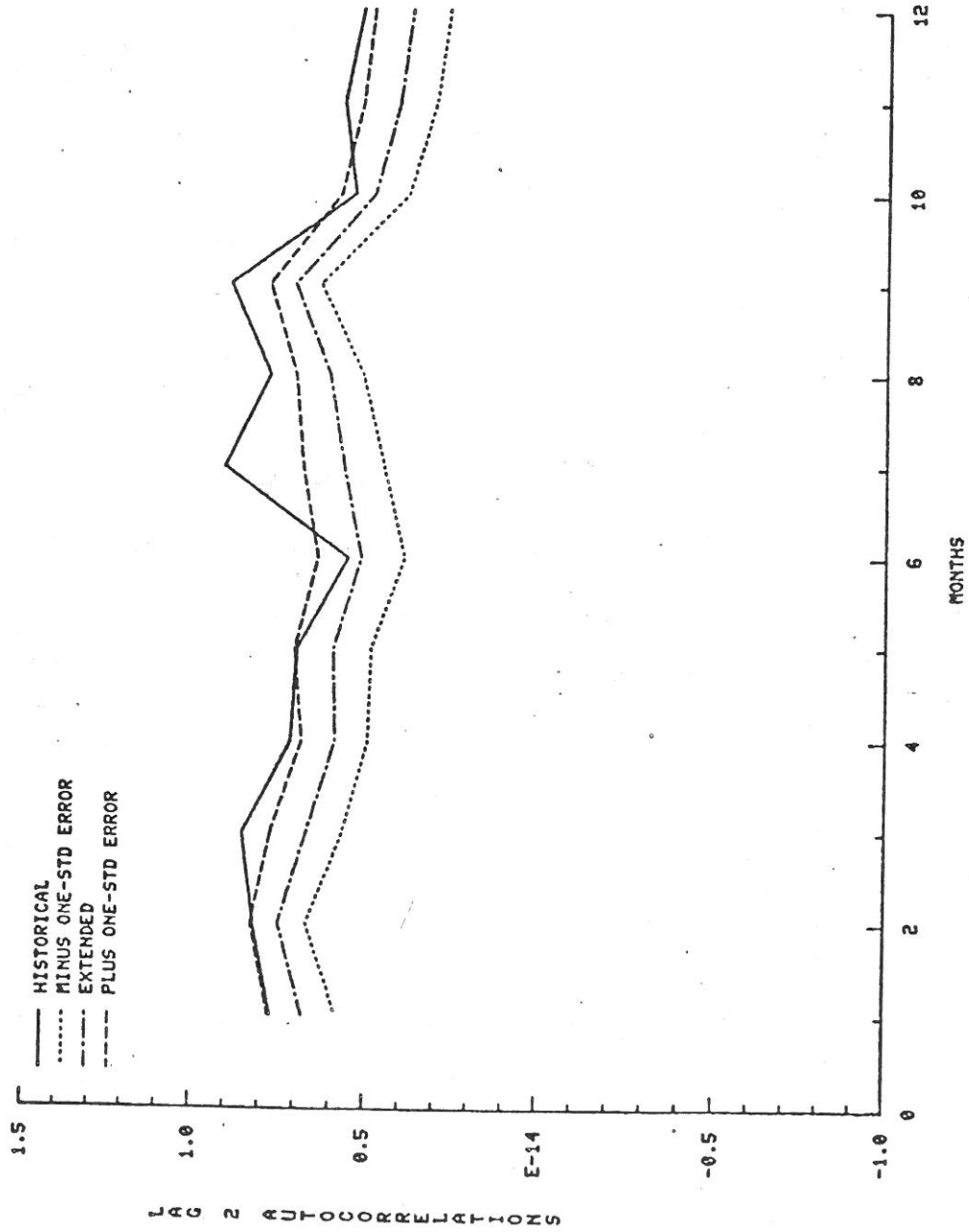


Figure 1.9.B.10. Historical and extended series monthly lag-2 autocorrelation coefficients of Paso del Ermitaño in logarithmic domain of flows.

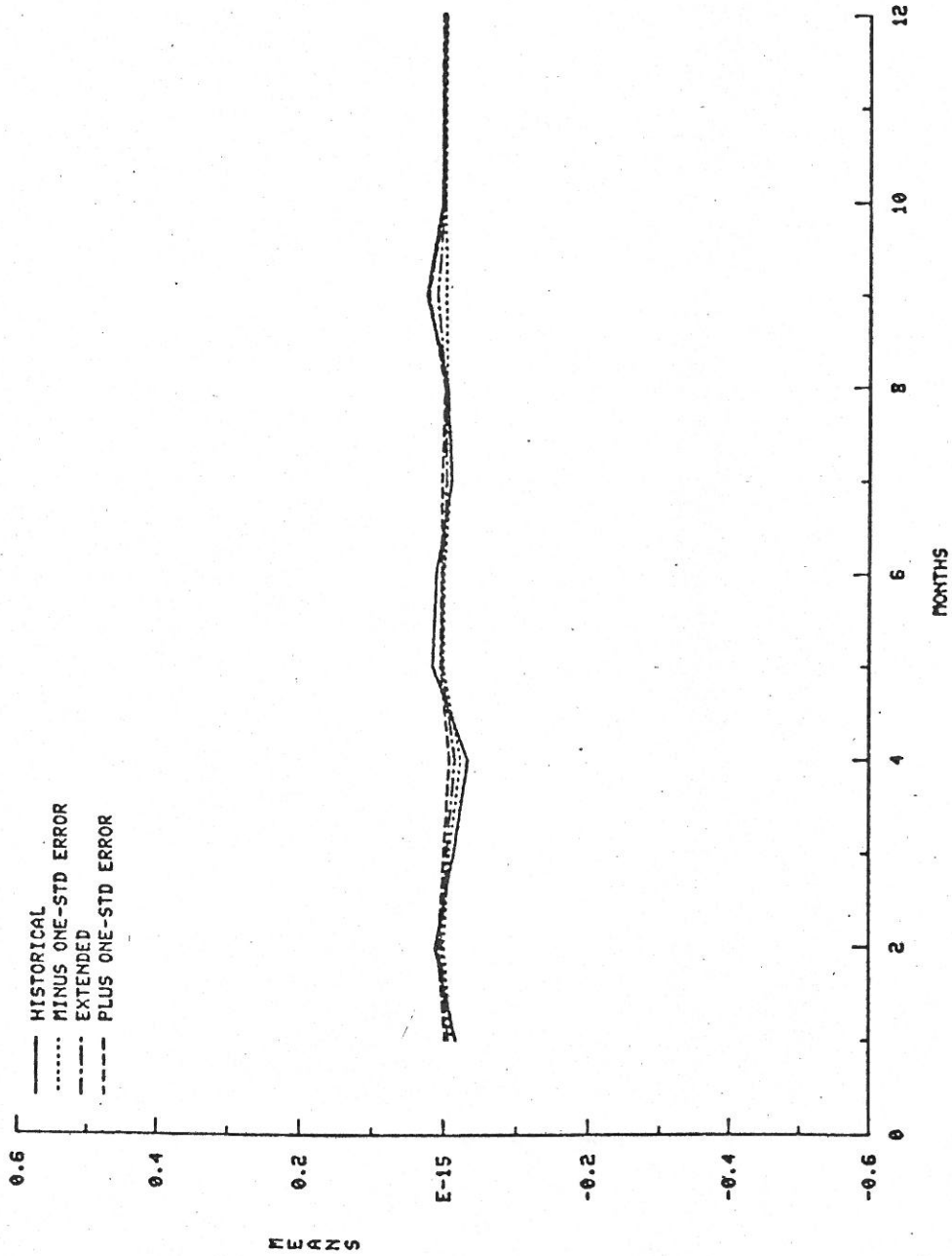


Figure 1.9.B.11. Historical and extended series monthly means of Paso del Ermitaño in log-Wilson-Hilferty domain of flows.

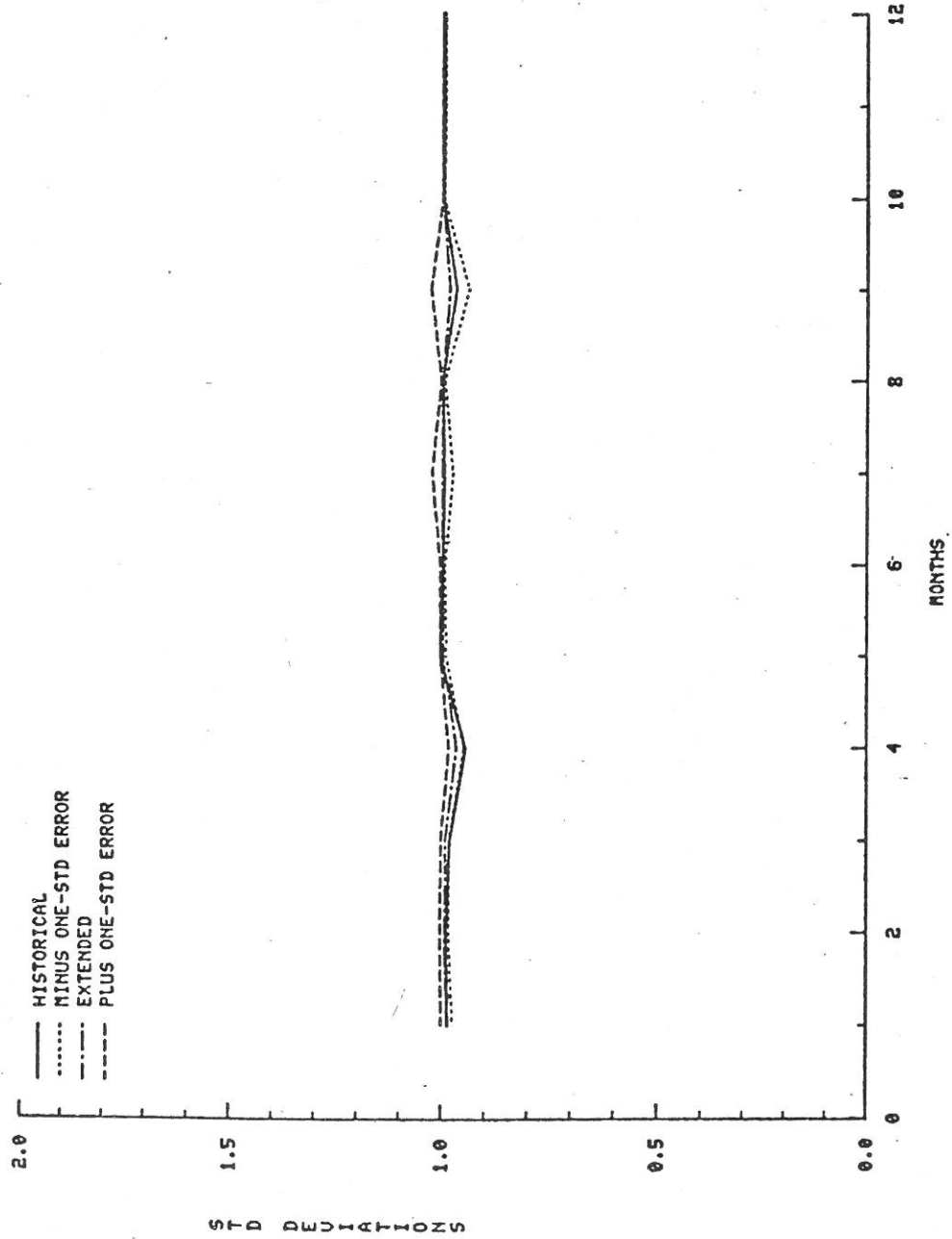


Figure 1.9.B.12. Historical and extended series monthly standard deviations of Paso del Ermitaño in log-Wilson-Hilferty domain of flows.

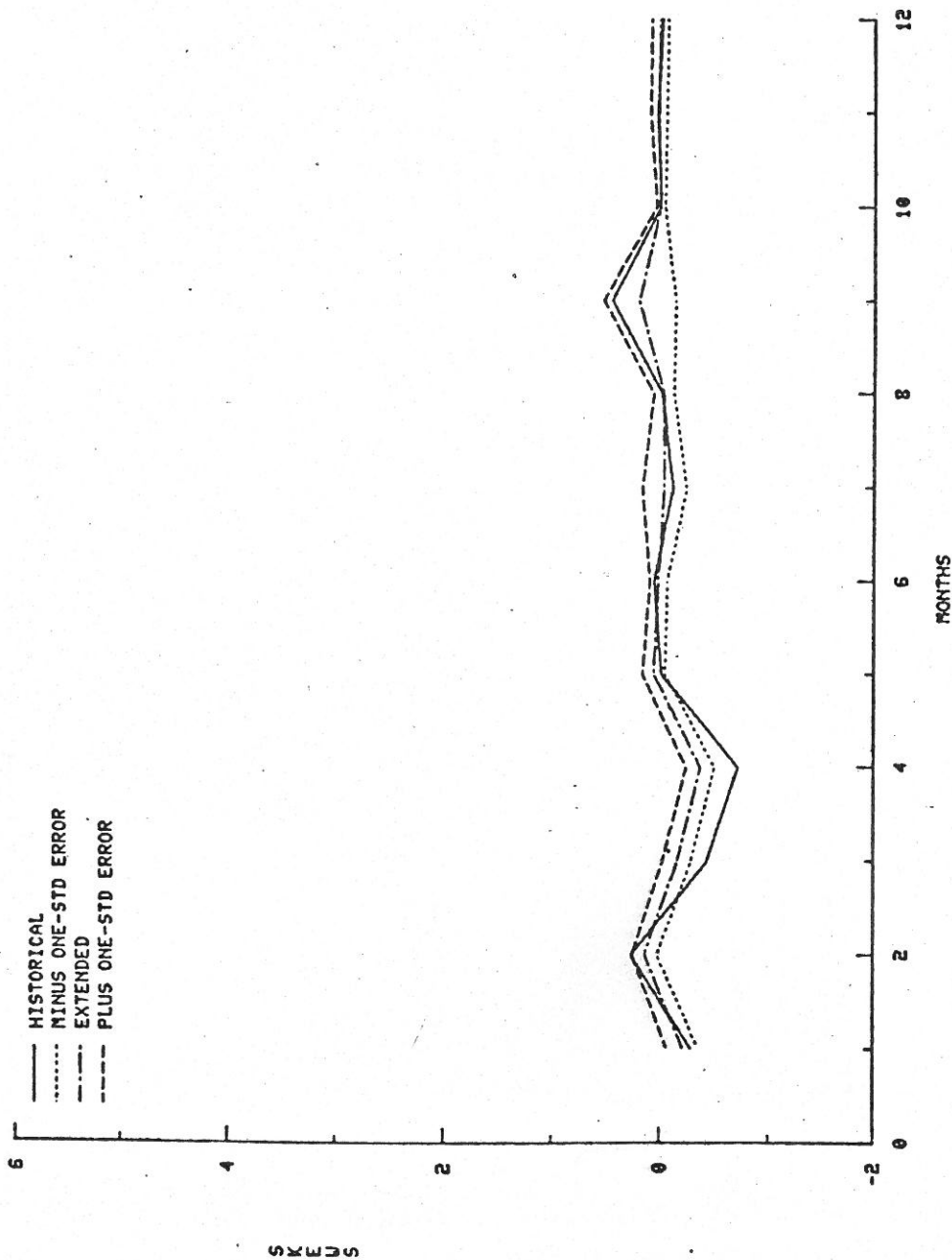


Figure 1.9.B.13. Historical and extended series monthly skewness coefficients of Paso del Ermitaño in log-Wilson-Hilferty domain of flows.

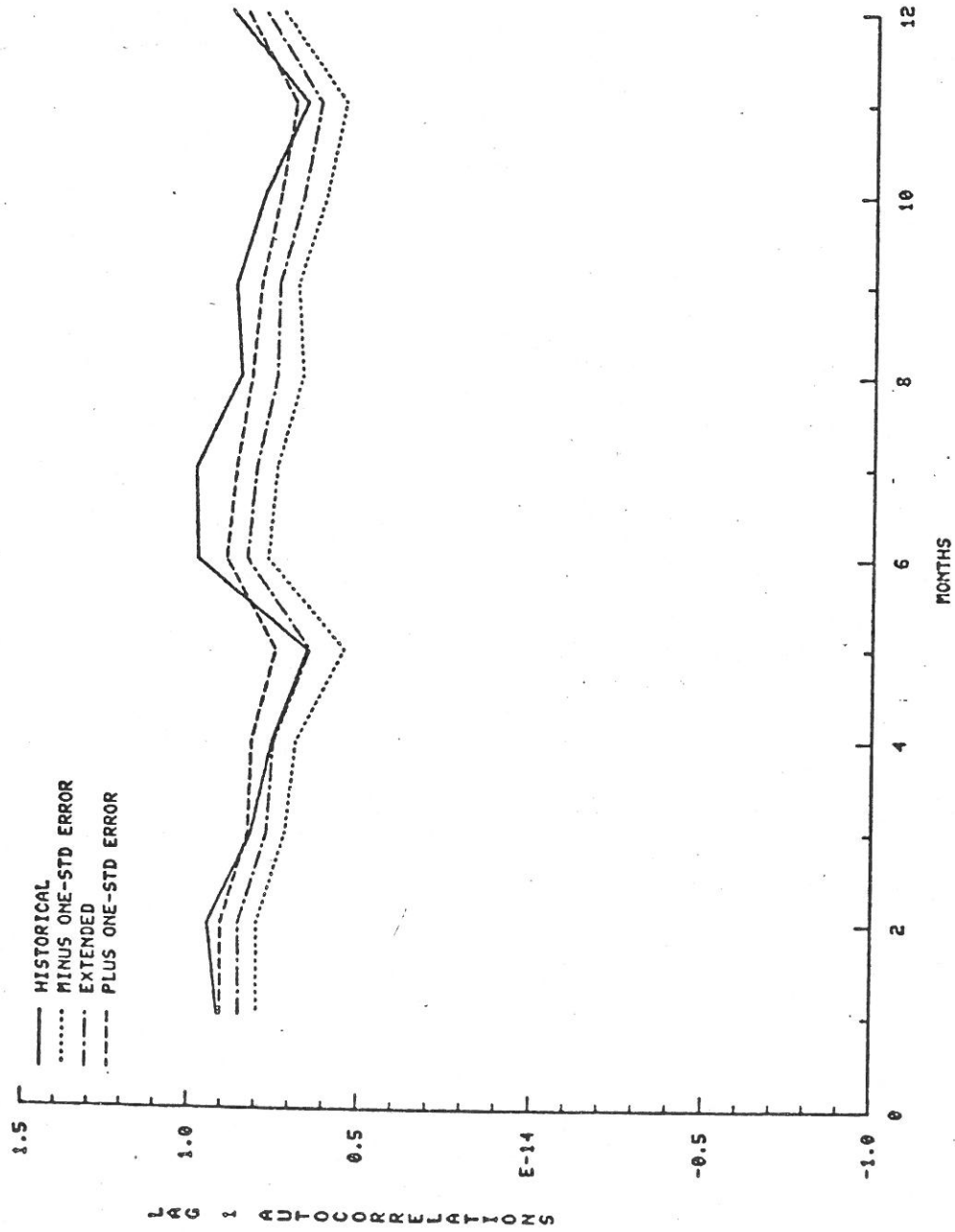


Figure 1.9.B.14. Historical and extended series monthly lag-1 autocorrelation coefficients of Paso del Ermitaño in log-Wilson-Hilferty domain of flows.

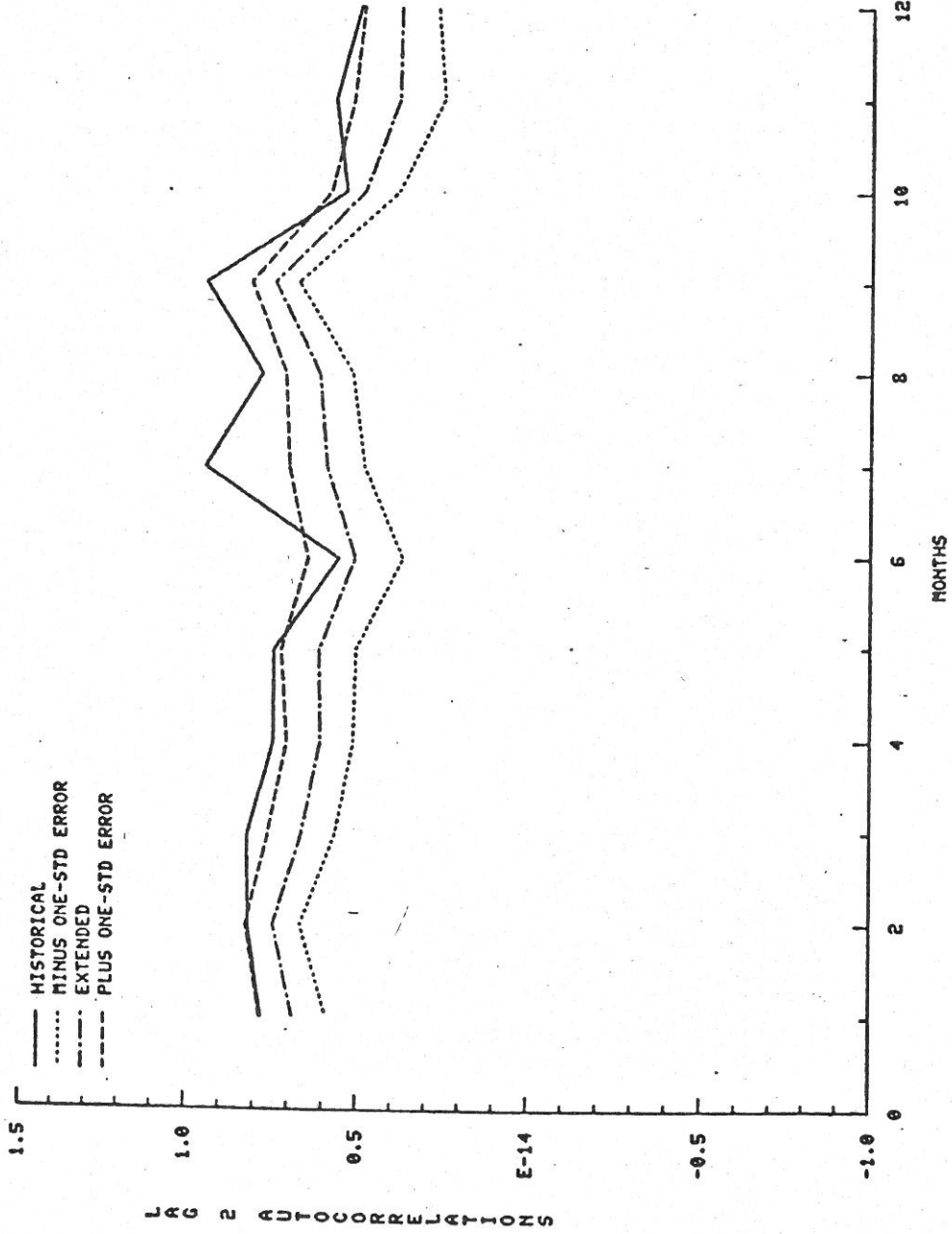


Figure 1.9.B.15. Historical and extended series monthly lag-2 autocorrelation coefficients of Paso del Ermitaño in log-Wilson-Hilferty domain of flows.

UNSATISFIED EXT P - 011666
EXCHANGE PA

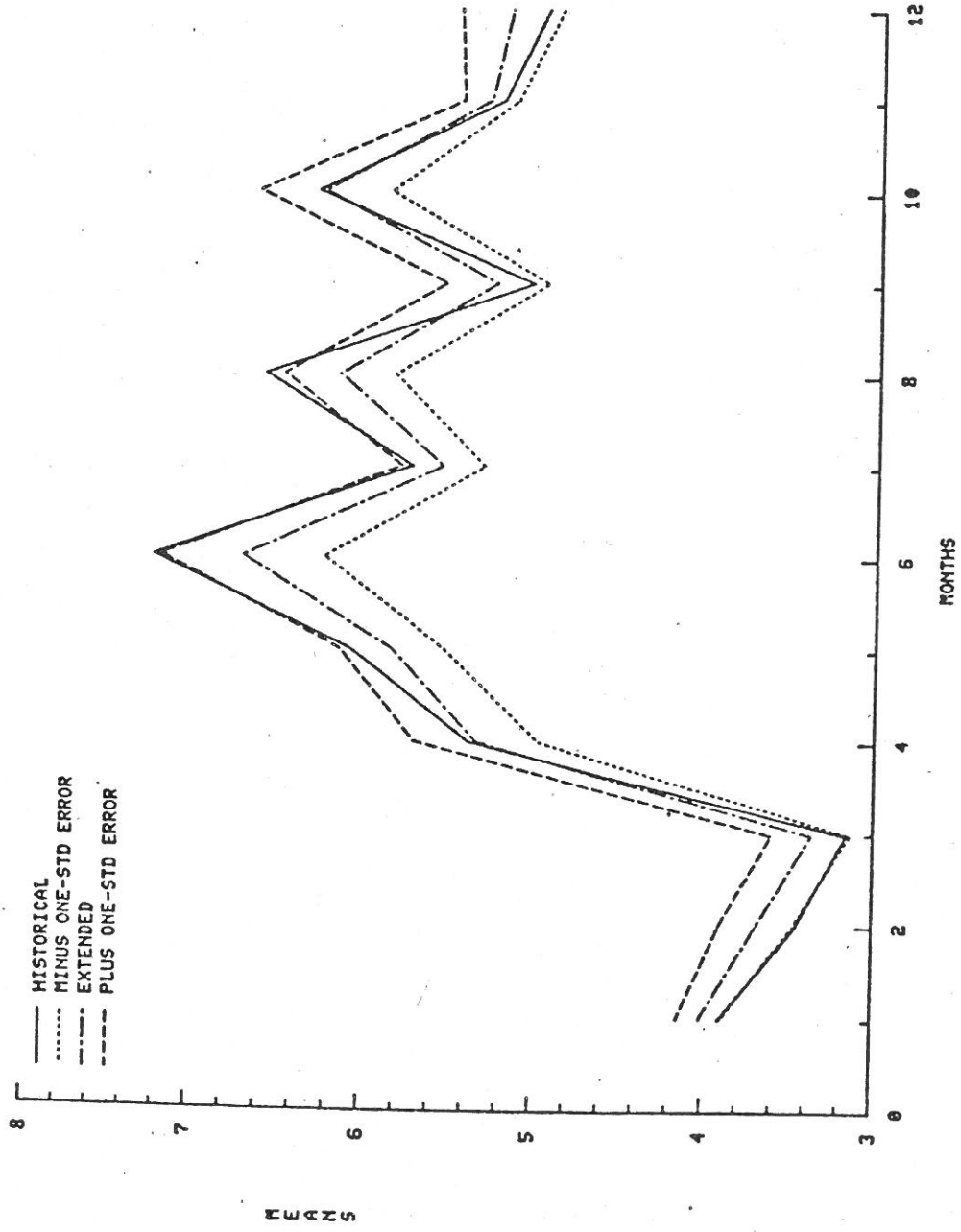


Figure 1.9.B.16. Historical and extended series monthly means of Rancho Arriba in original domain of flows.

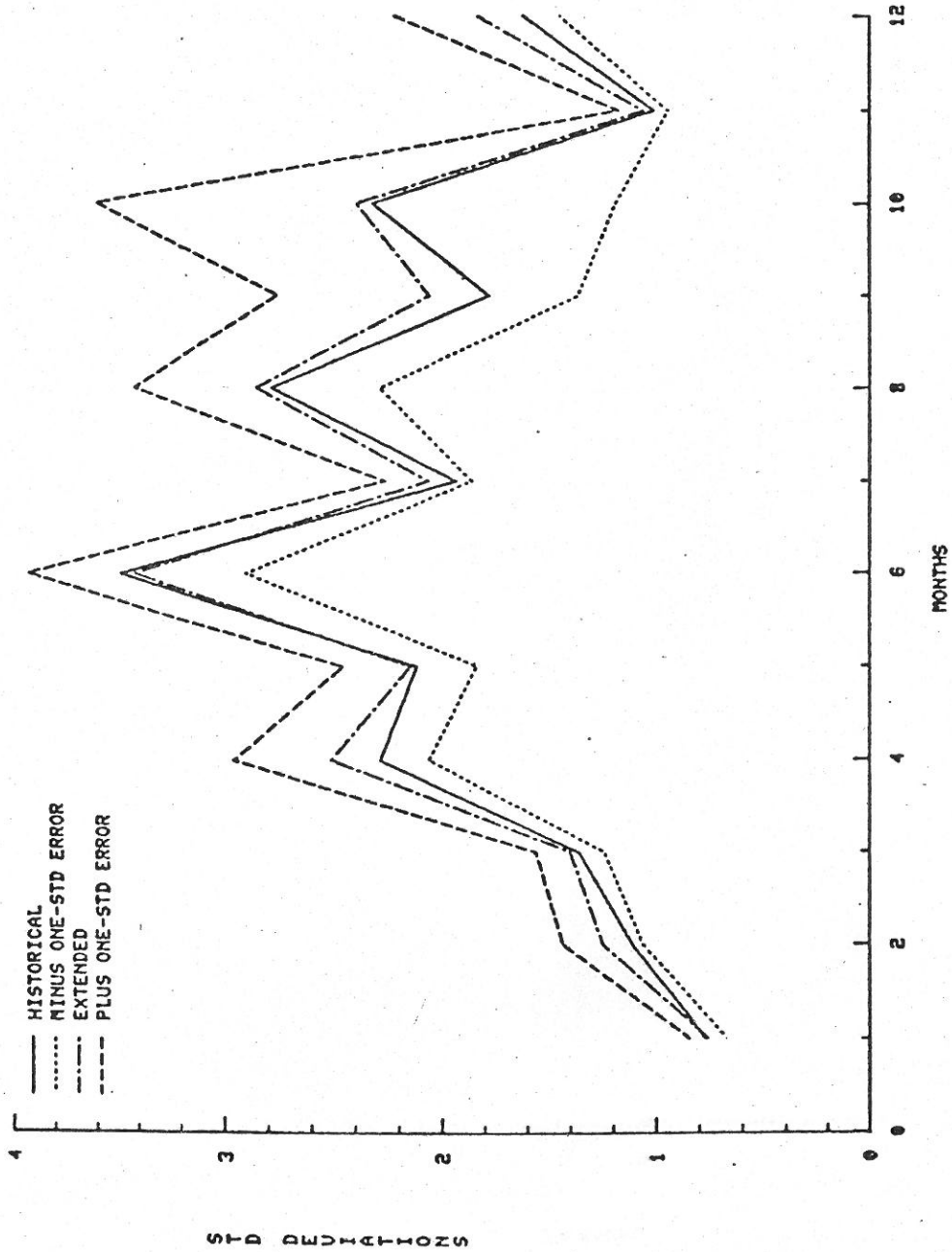


Figure 1.9.B.17. Historical and extended series monthly standard deviations of Rancho Arribain original domain of flows.

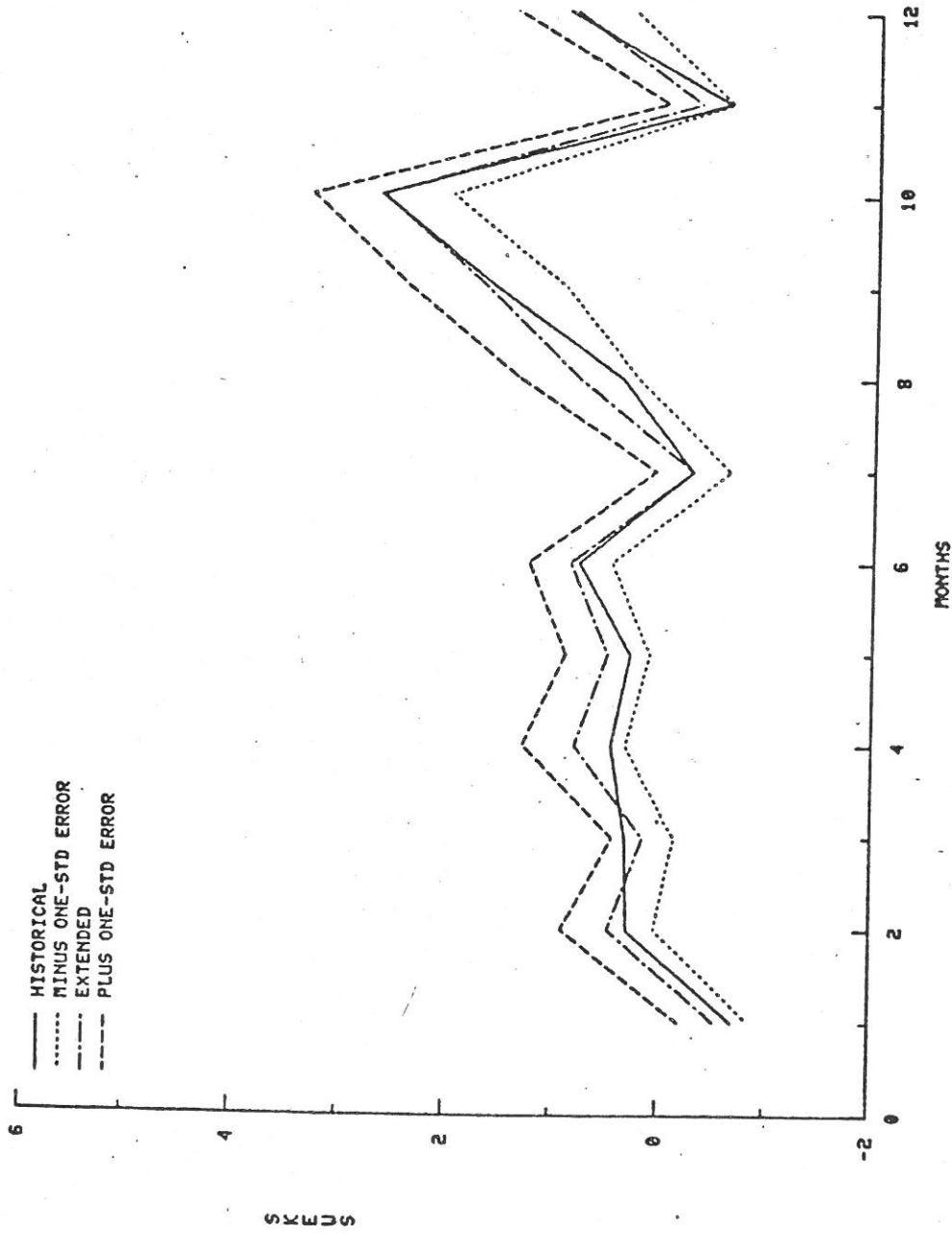


Figure 1.9.B.18. Historical and extended series monthly skewness coefficients of Rancho Arriba in original domain of flows.

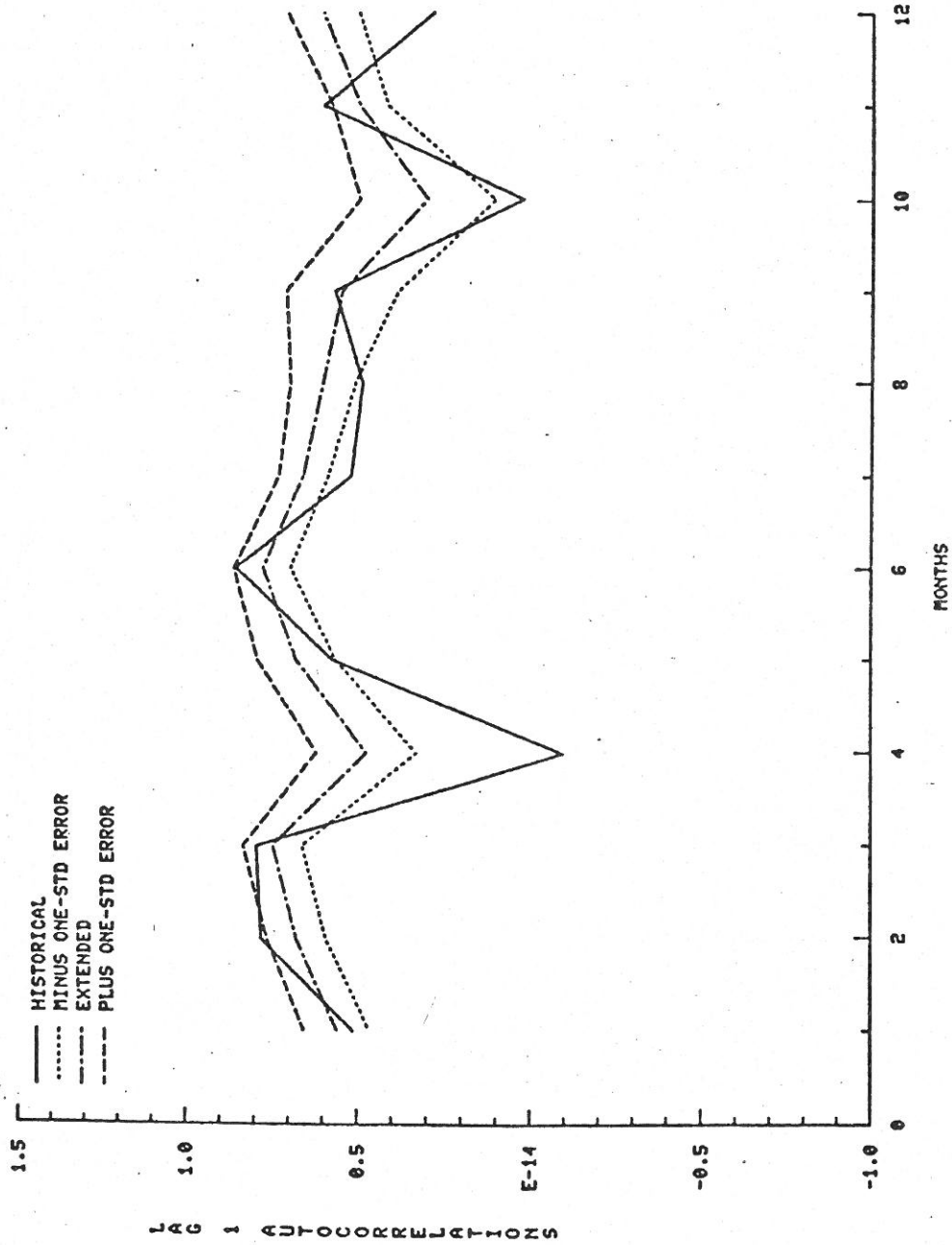


Figure 1.9.B.19. Historical and extended series monthly lag-1 autocorrelation coefficients of Rancho Arriba in original domain of flows.

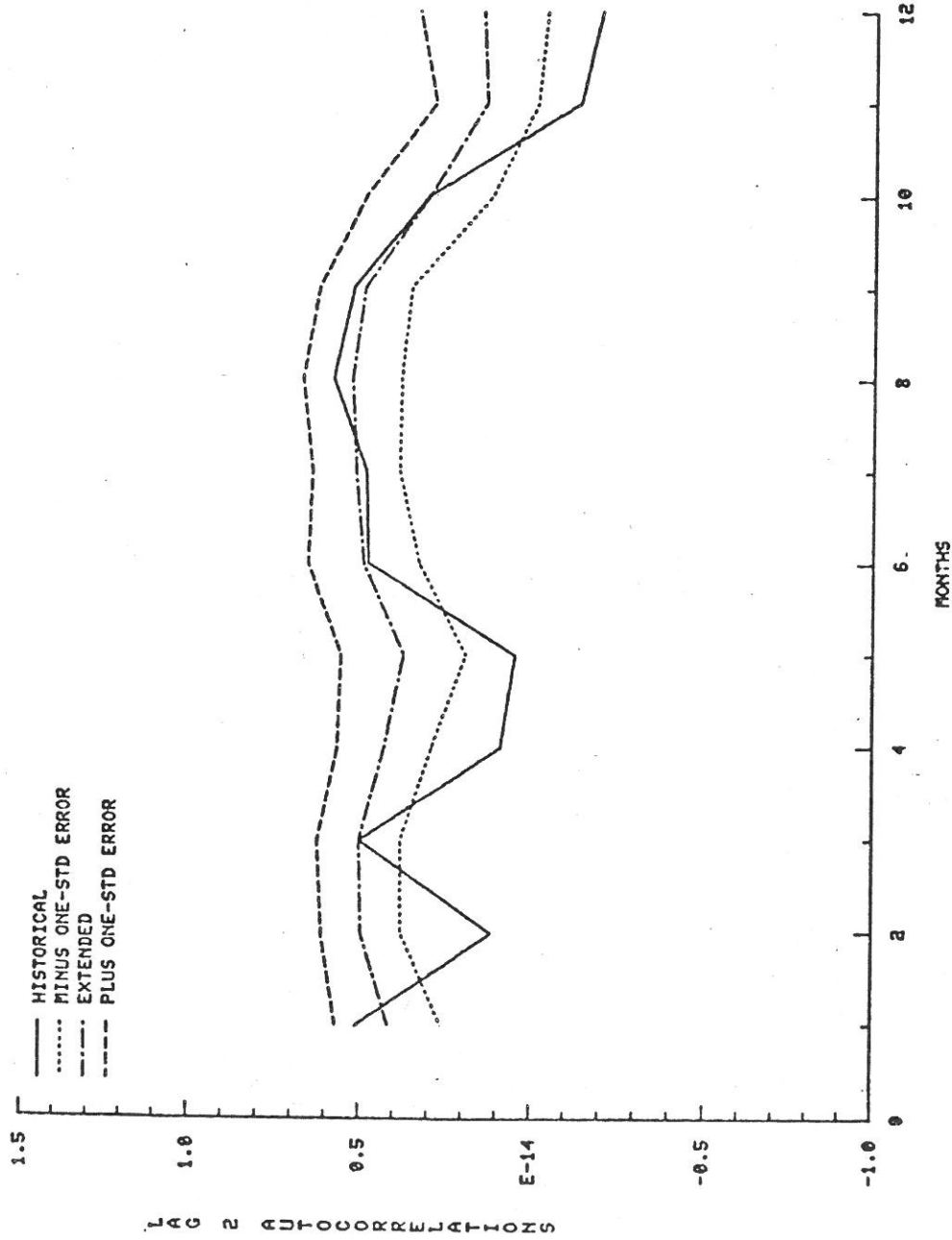


Figure 1.9.B.20. Historical and extended series monthly lag-2 autocorrelation coefficients of Rancho Arriba in original domain of flows.

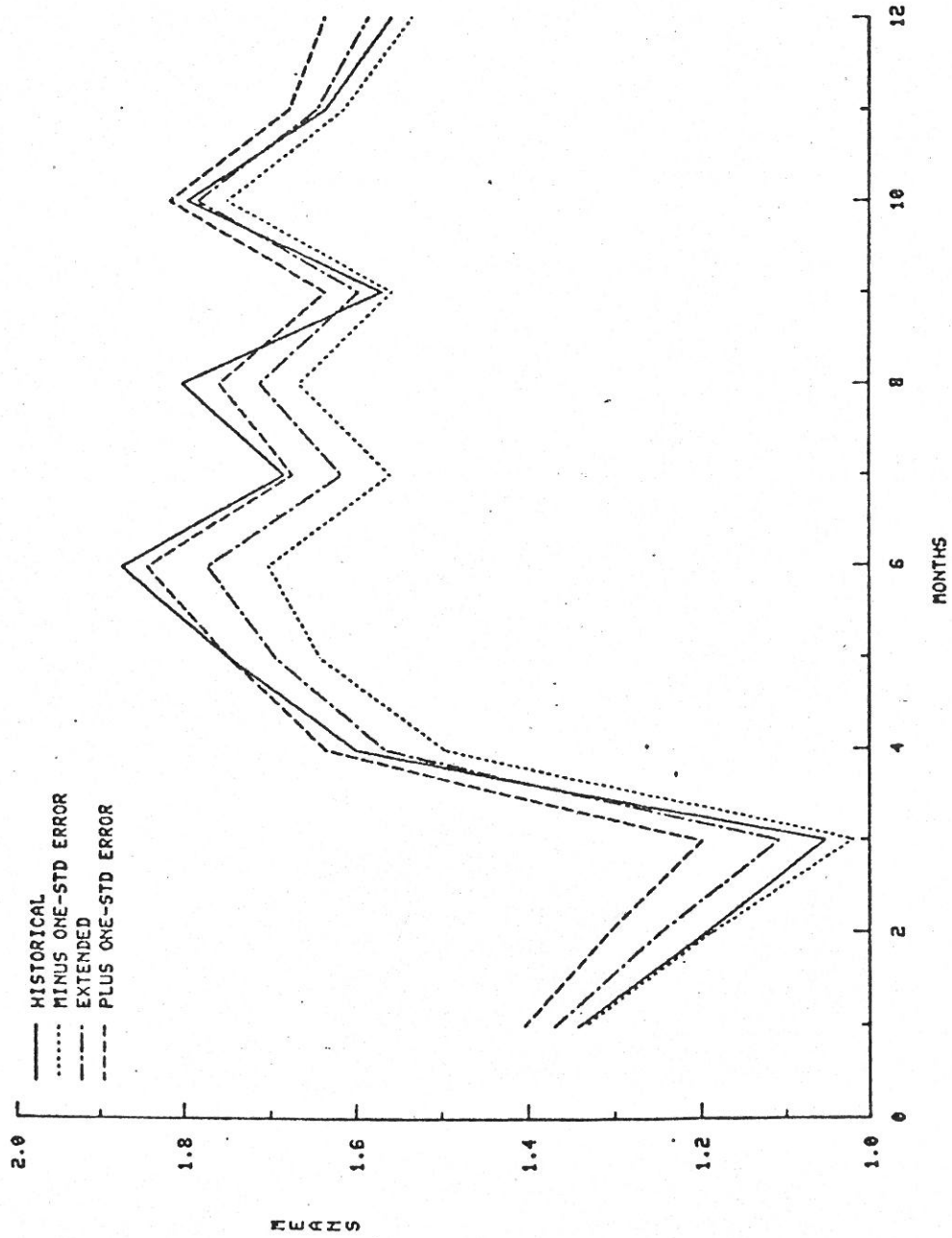


Figure 1.9.B.21. Historical and extended series monthly means of Rancho Arriba in logarithmic domain of flows.

URSAT

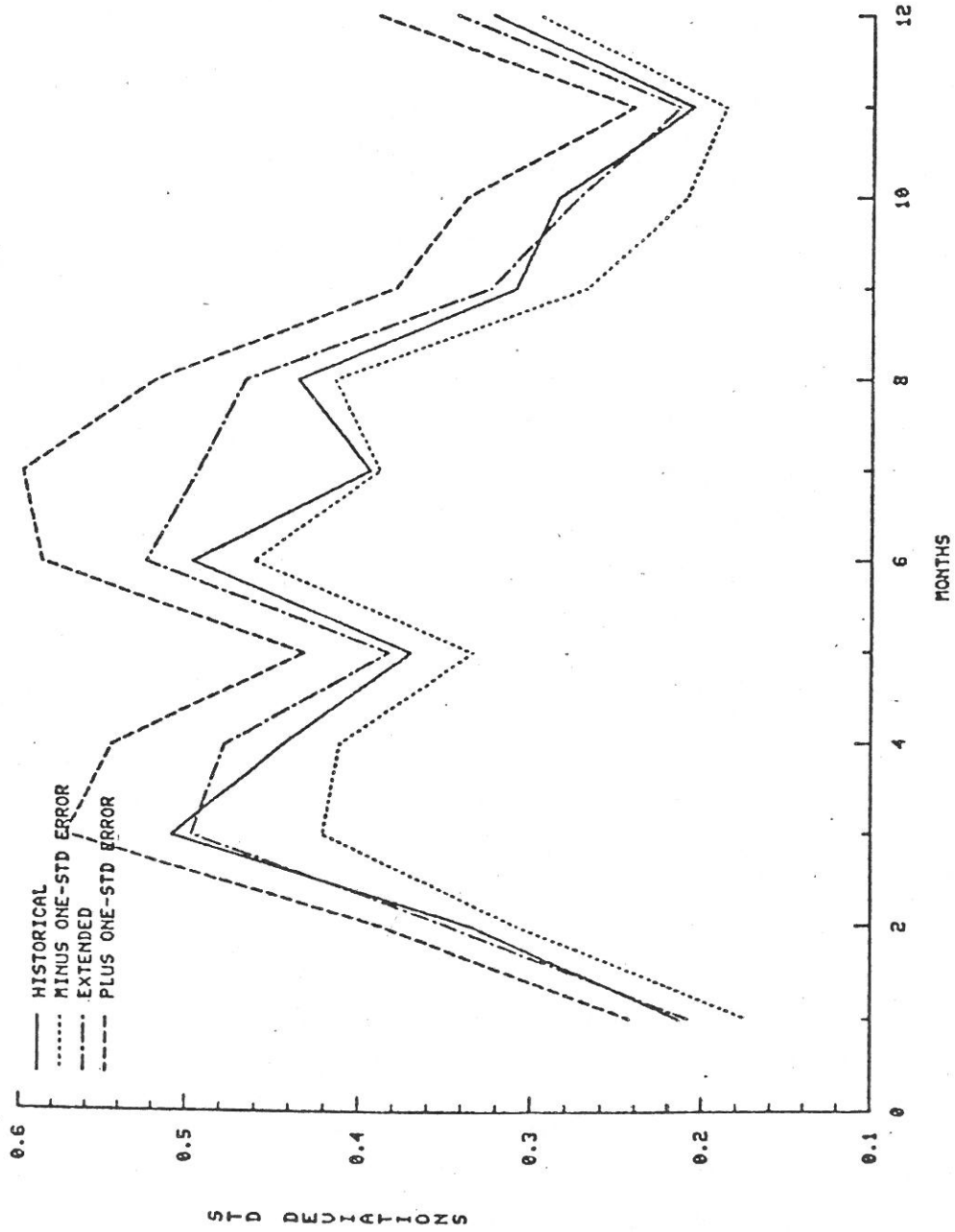


Figure 1.9.B.22. Historical and extended series monthly standard deviations of Rancho Arriba in logarithmic domain of flows.

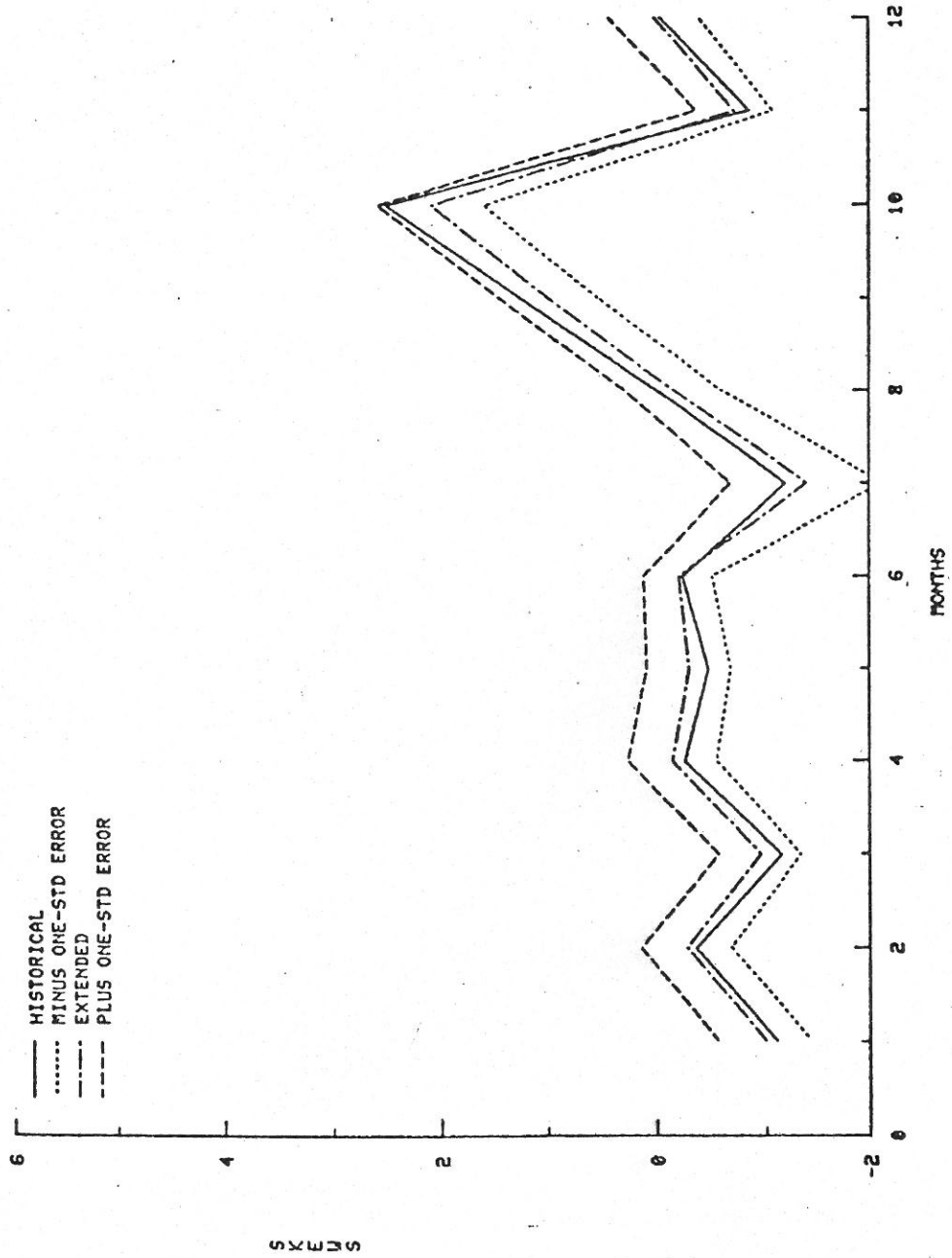


Figure 1.9.B.23. Historical and extended series monthly skewness coefficients of Rancho Arriba in logarithmic domain of flows.

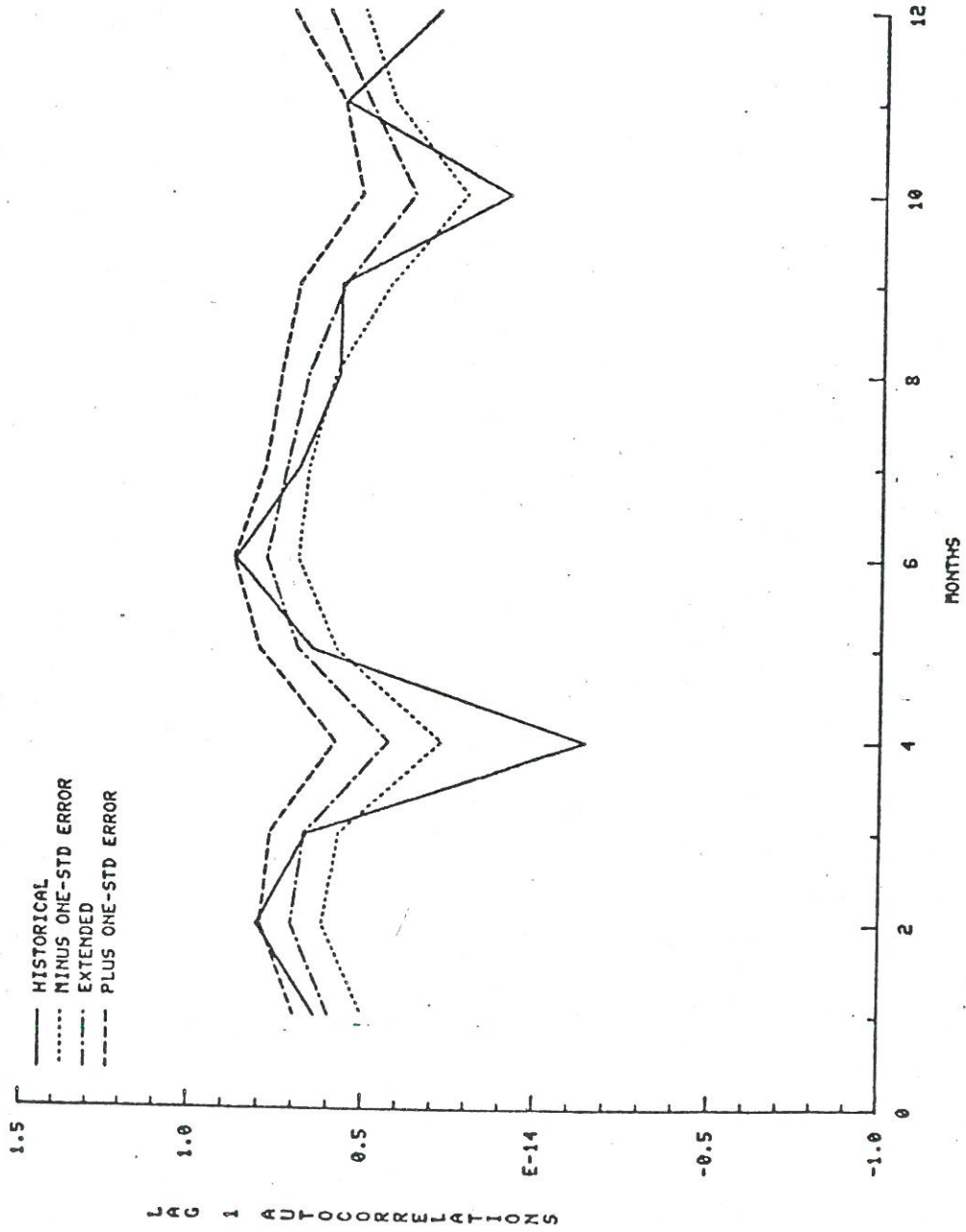


Figure 1.9.B.24. Historical and extended series monthly lag-1 autocorrelation coefficients of Rancho Arriba in logarithmic domain of flows.

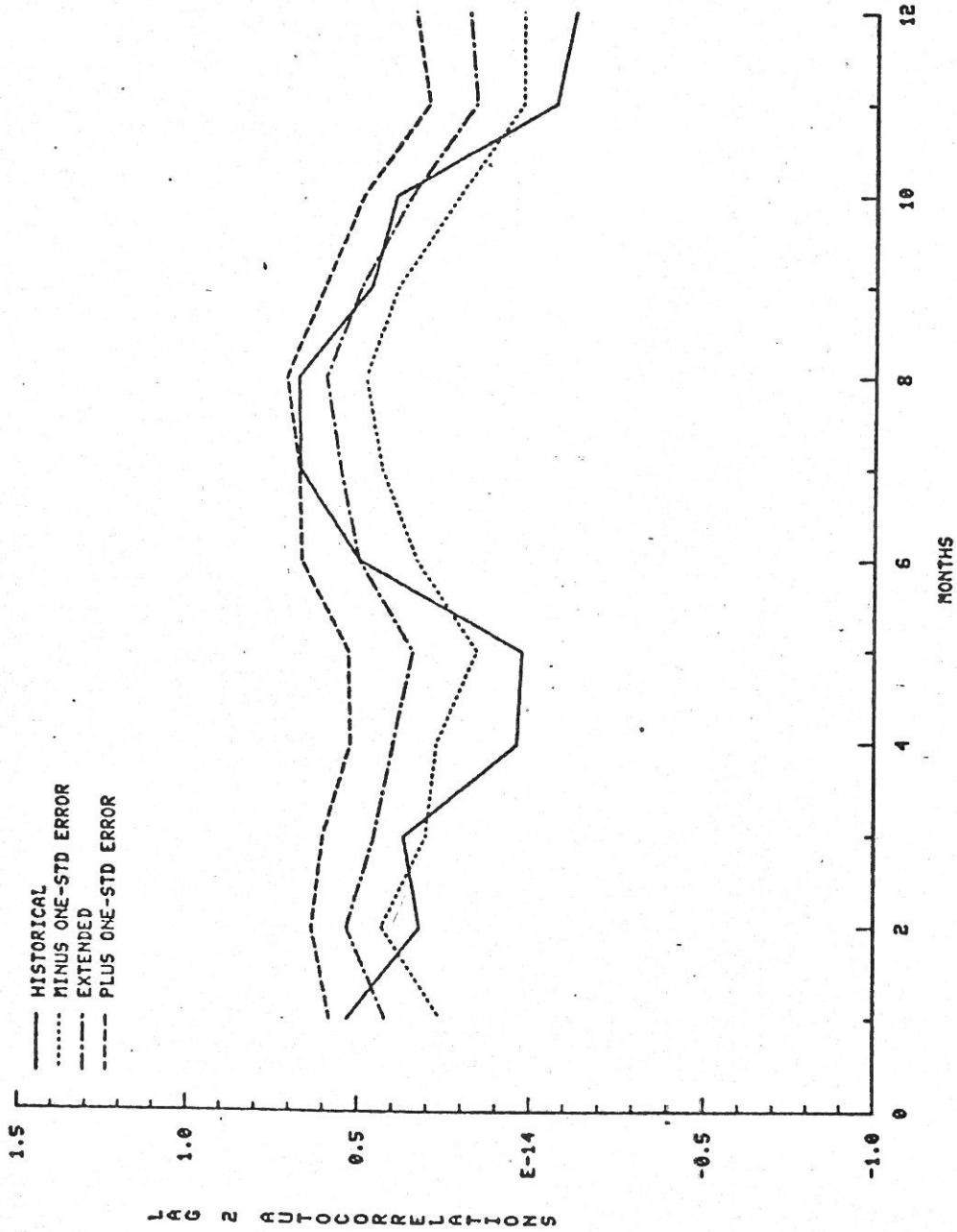


Figure 1.9.B.25. Historical and extended series monthly lag-2 autocorrelation coefficients of Rancho Arriba in logarithmic domain of flows.

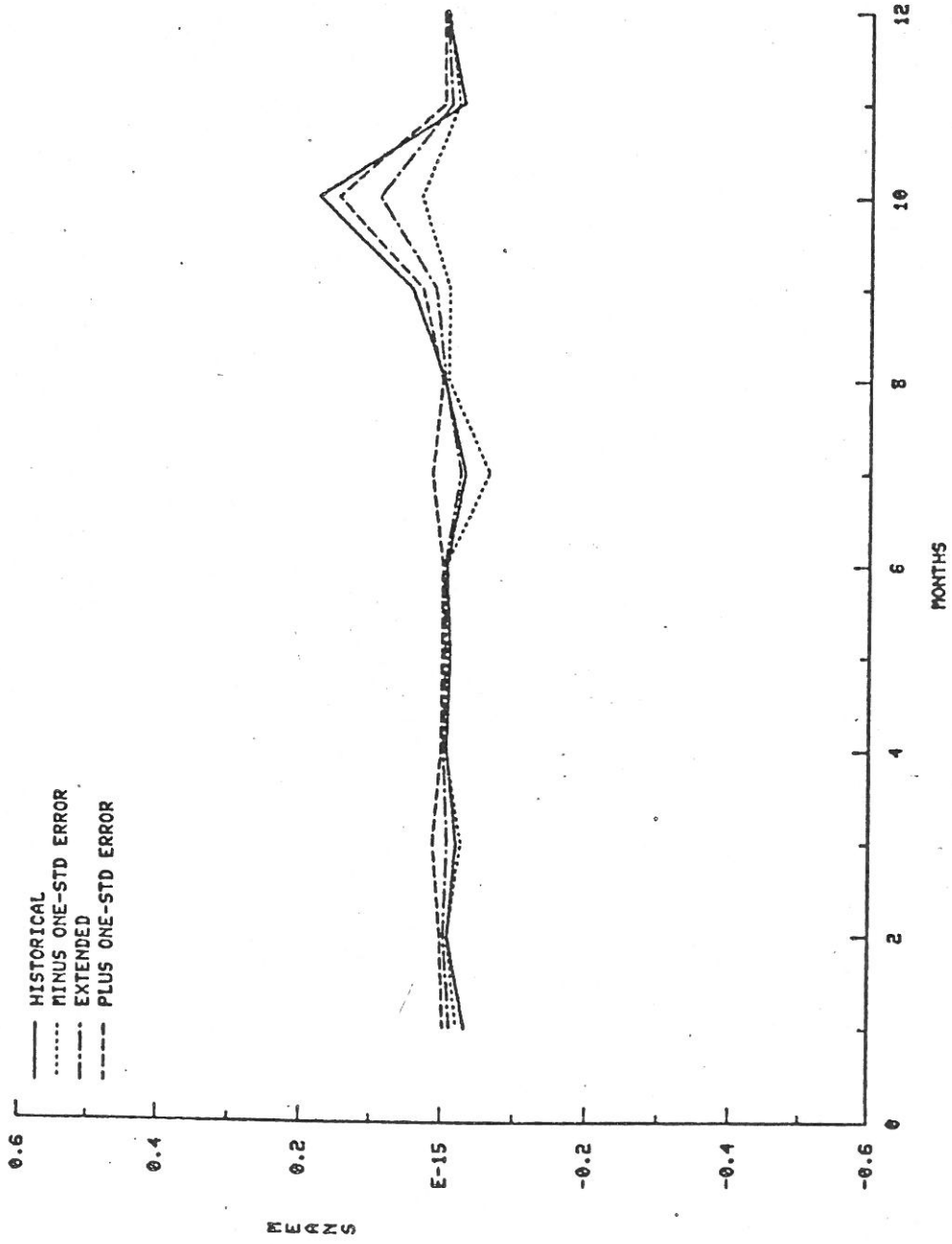


Figure 1.9.B.26. Historical and extended series monthly means of Rancho Arriba in log-Wilson-Hilferty domain of flows.

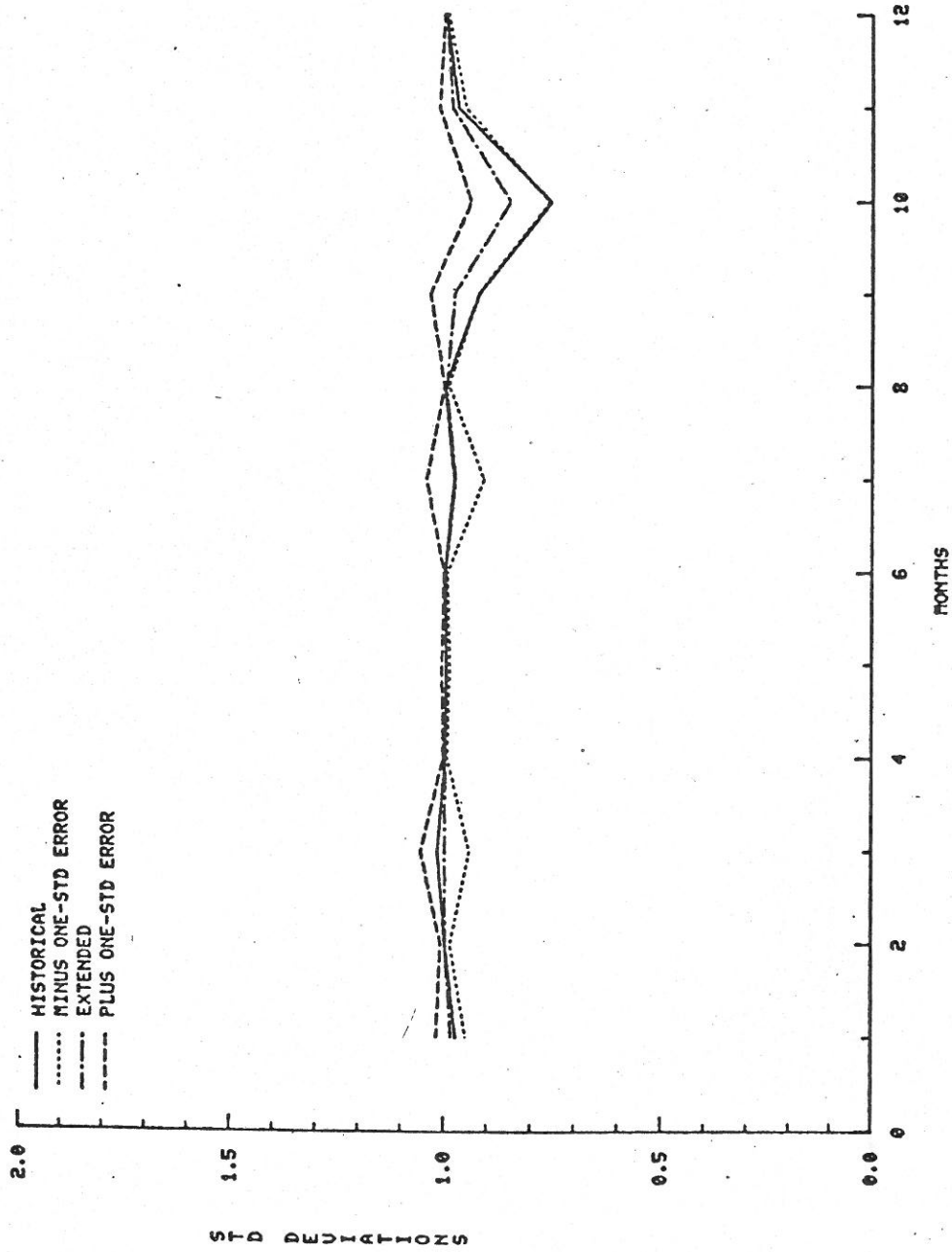


Figure 1.9.B.27. Historical and extended series monthly standard deviations of Rancho Arriba in log-Wilson-Hilferty domain of flows.

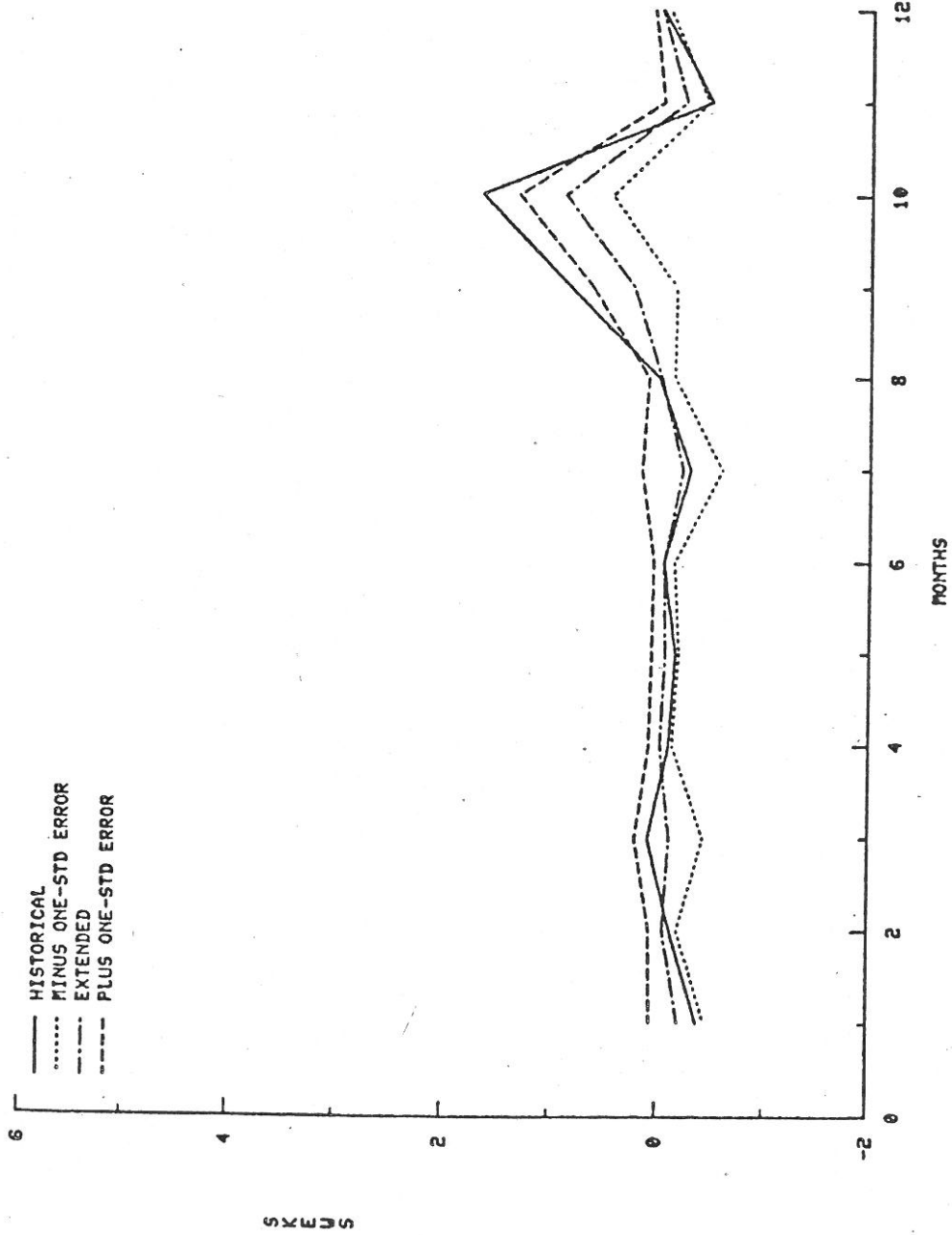


Figure 1.9.B.28. Historical and extended series monthly skewness coefficients of Rancho Arriba in log-Wilson-Hilferty domain of flows.

UNSATISFIED EXT P •

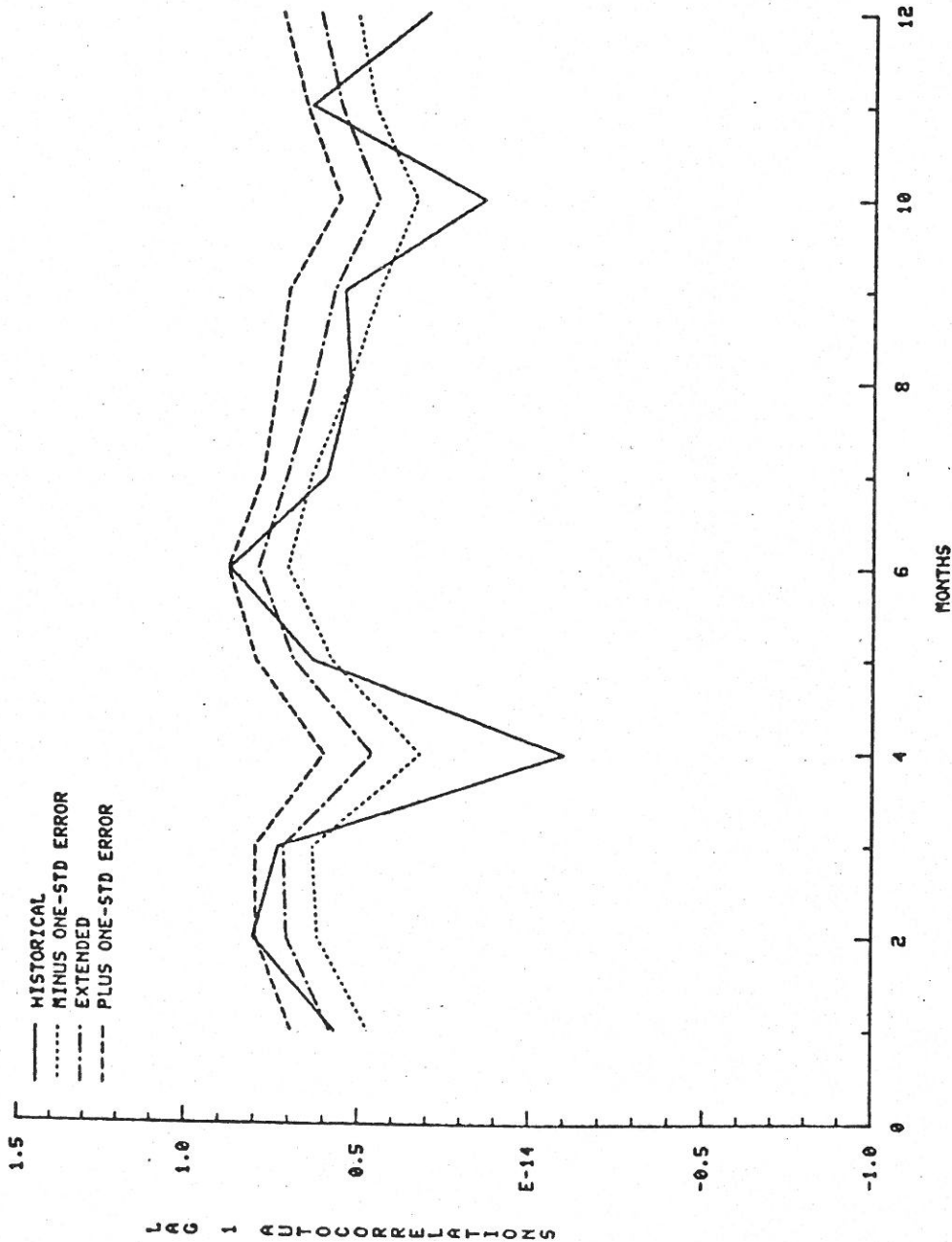


Figure 1.9.B.29. Historical and extended series monthly lag-1 autocorrelation coefficients of Rancho Arriba in log-Wilson-Hilferty domain of flows.

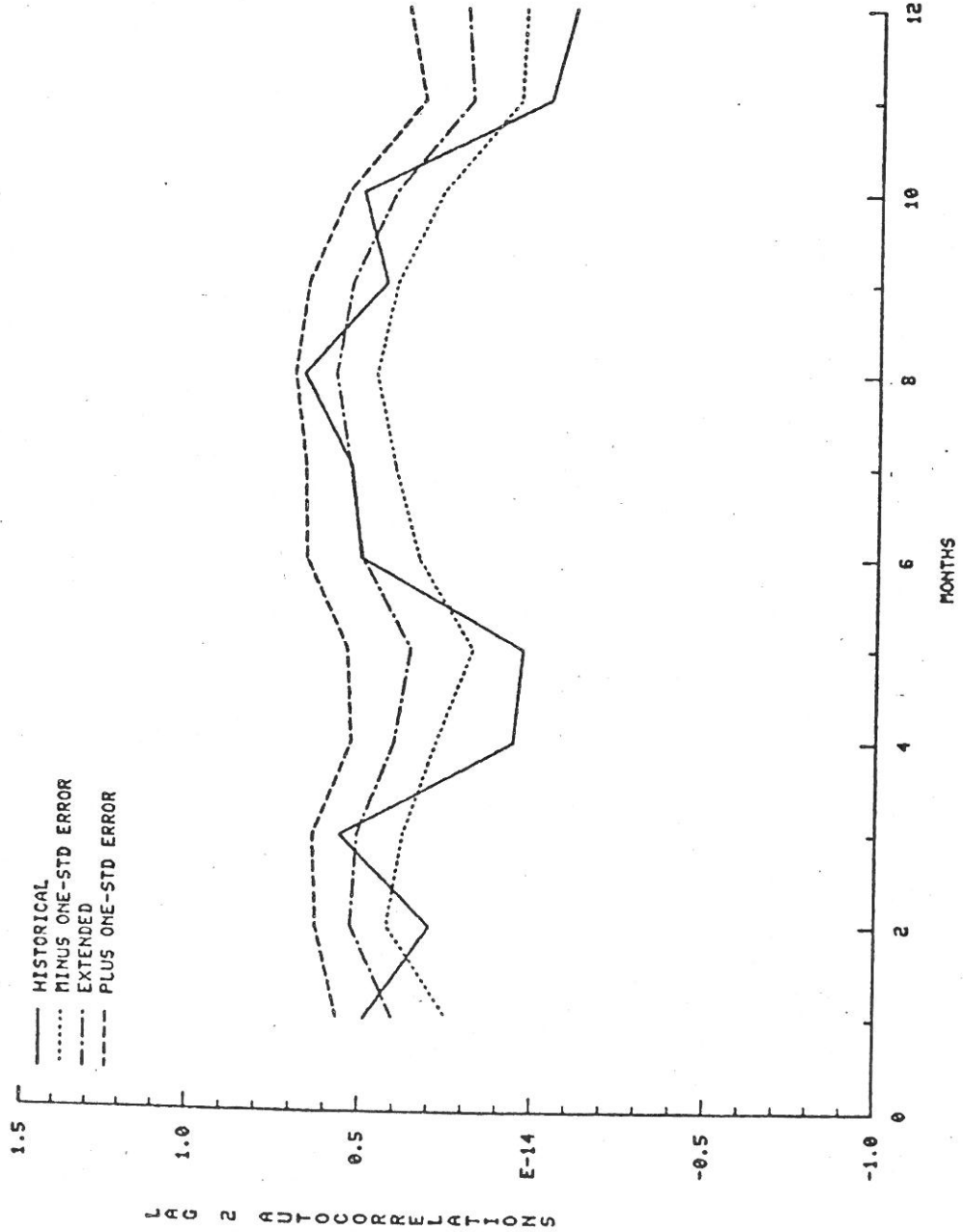


Figure 1.9.B.30. Historical and extended series monthly lag-2 autocorrelation coefficients of Rancho Arriba in log-Wilson-Hilferty domain of flows.

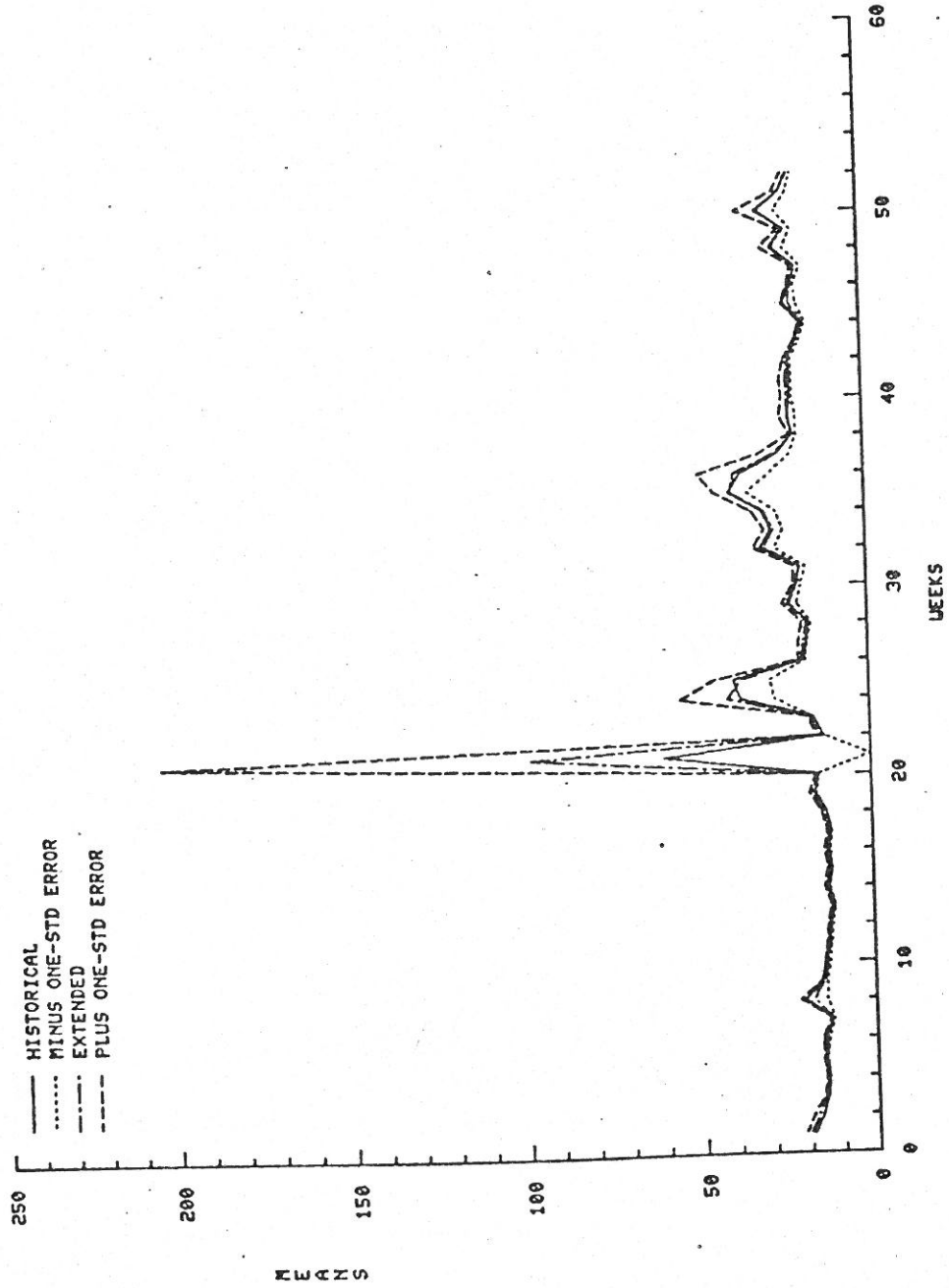


Figure 1.9.B.31. Historical and extended series weekly means of Paso del Ermitaño in original domain of flows.

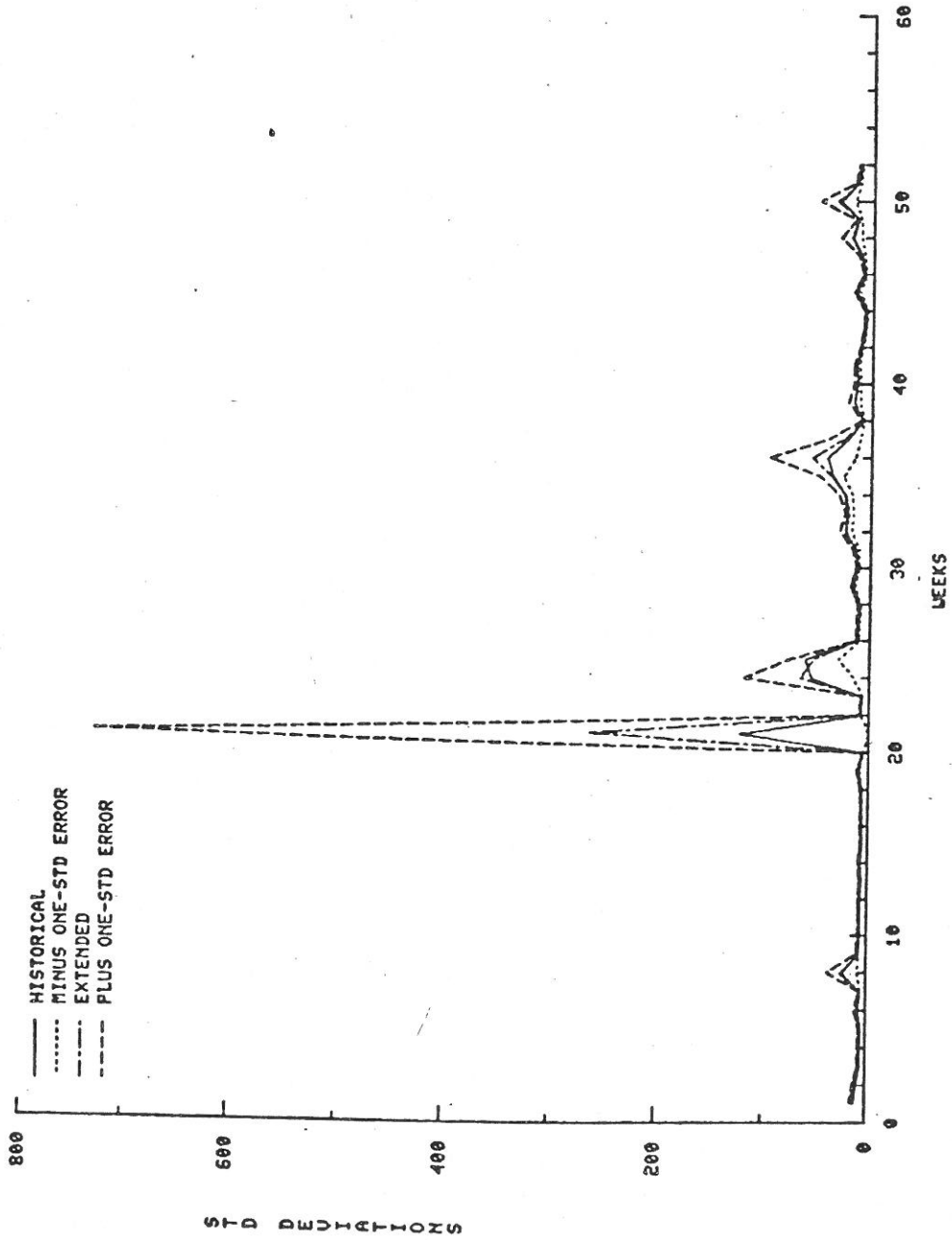


Figure 1.9.B.32. Historical and extended series weekly standard deviations of Paso del Ermitaño in original domain of flows.

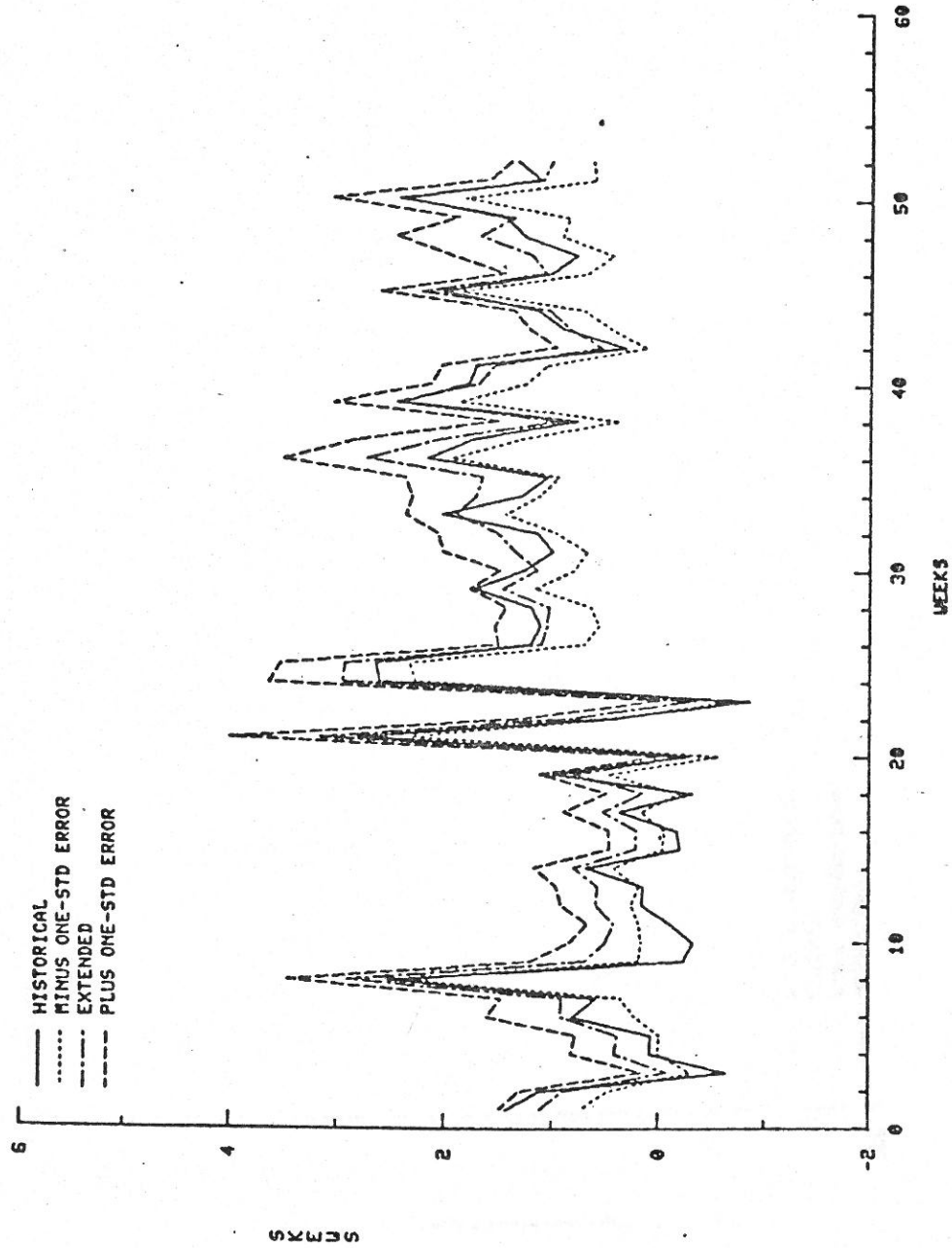


Figure 1.9.B.33. Historical and extended series weekly skewness coefficients of Paso del Ermitaño in original domain of flows.

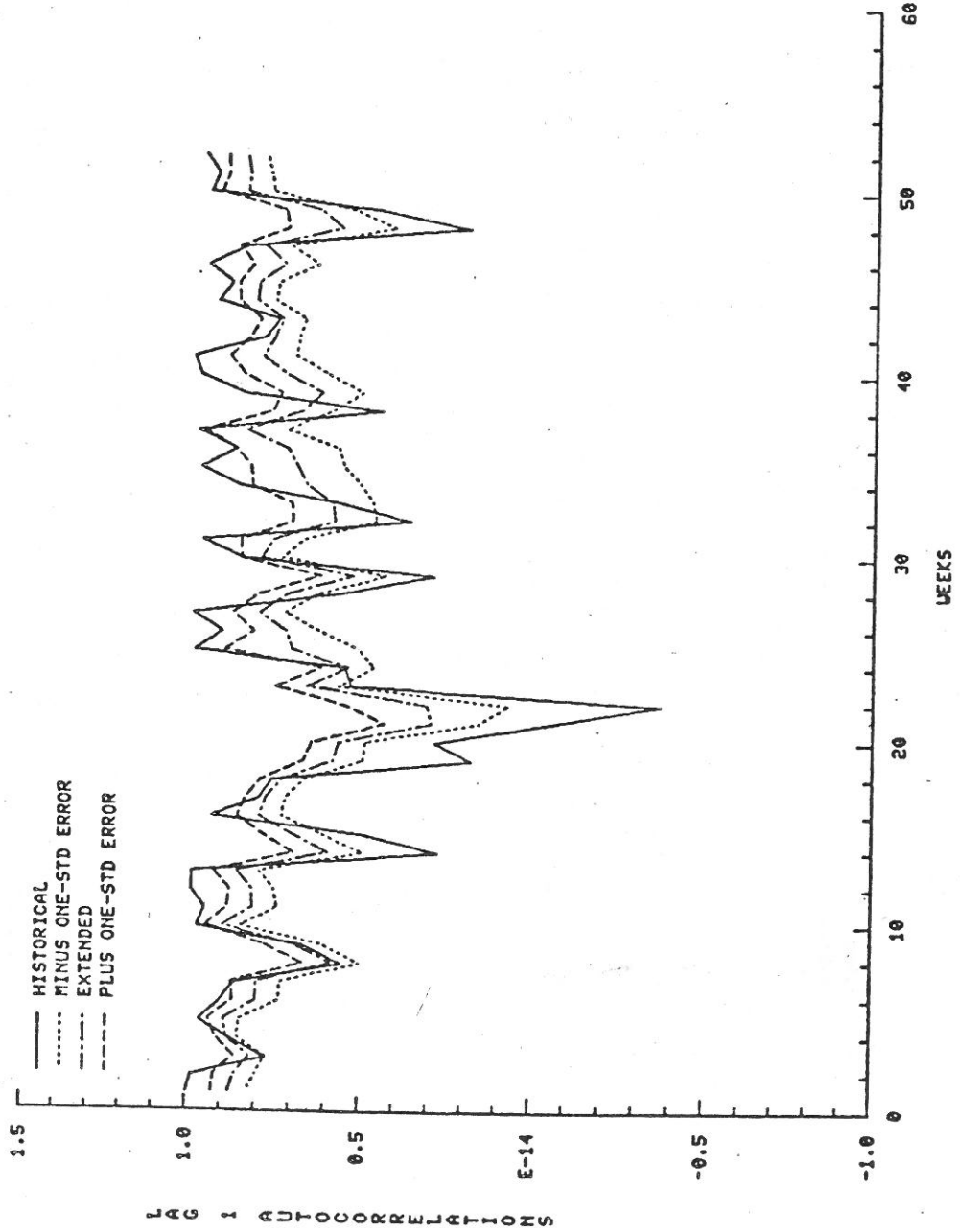


Figure 1.9.B.34. Historical and extended series weekly lag-1 autocorrelation coefficients of Paso del Ermitaño in original domain of flows.

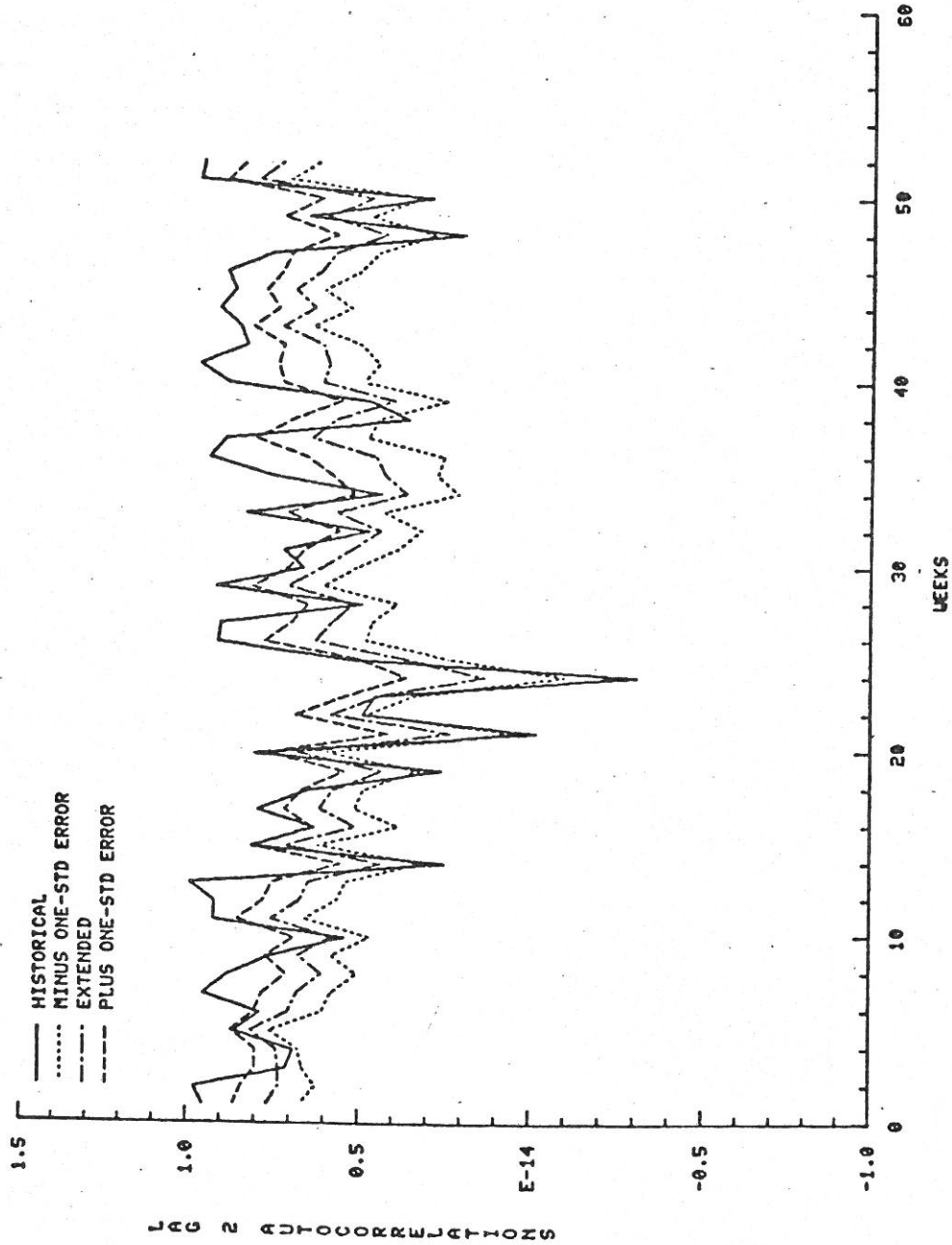


Figure 1.9.B.35. Historical and extended series weekly lag-2 autocorrelation coefficients of Paso del Ermitaño in original domain of flows.

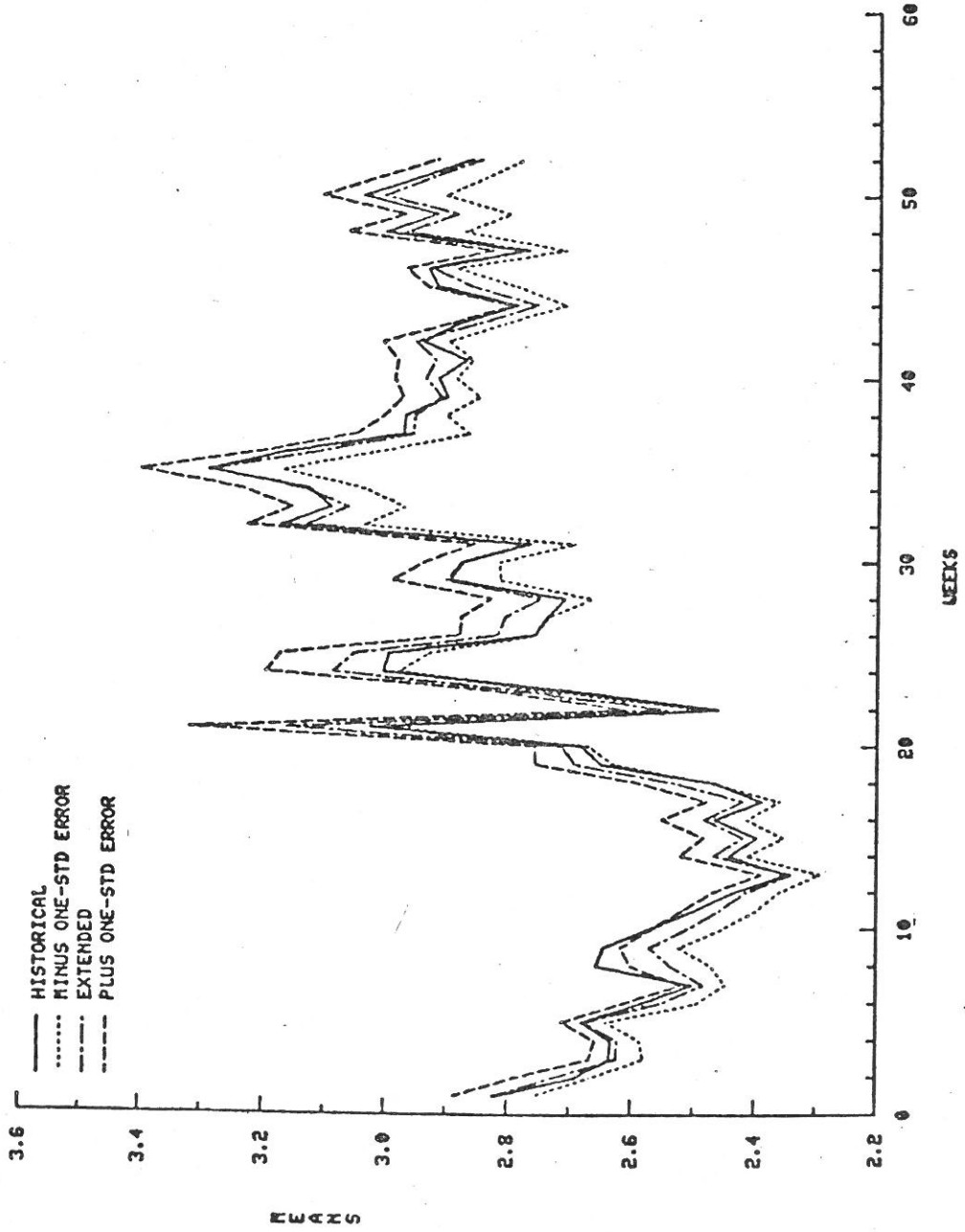


Figure 1.9.B.36. Historical and extended series weekly means of Paso del Ermitaño in logarithmic domain of flows.

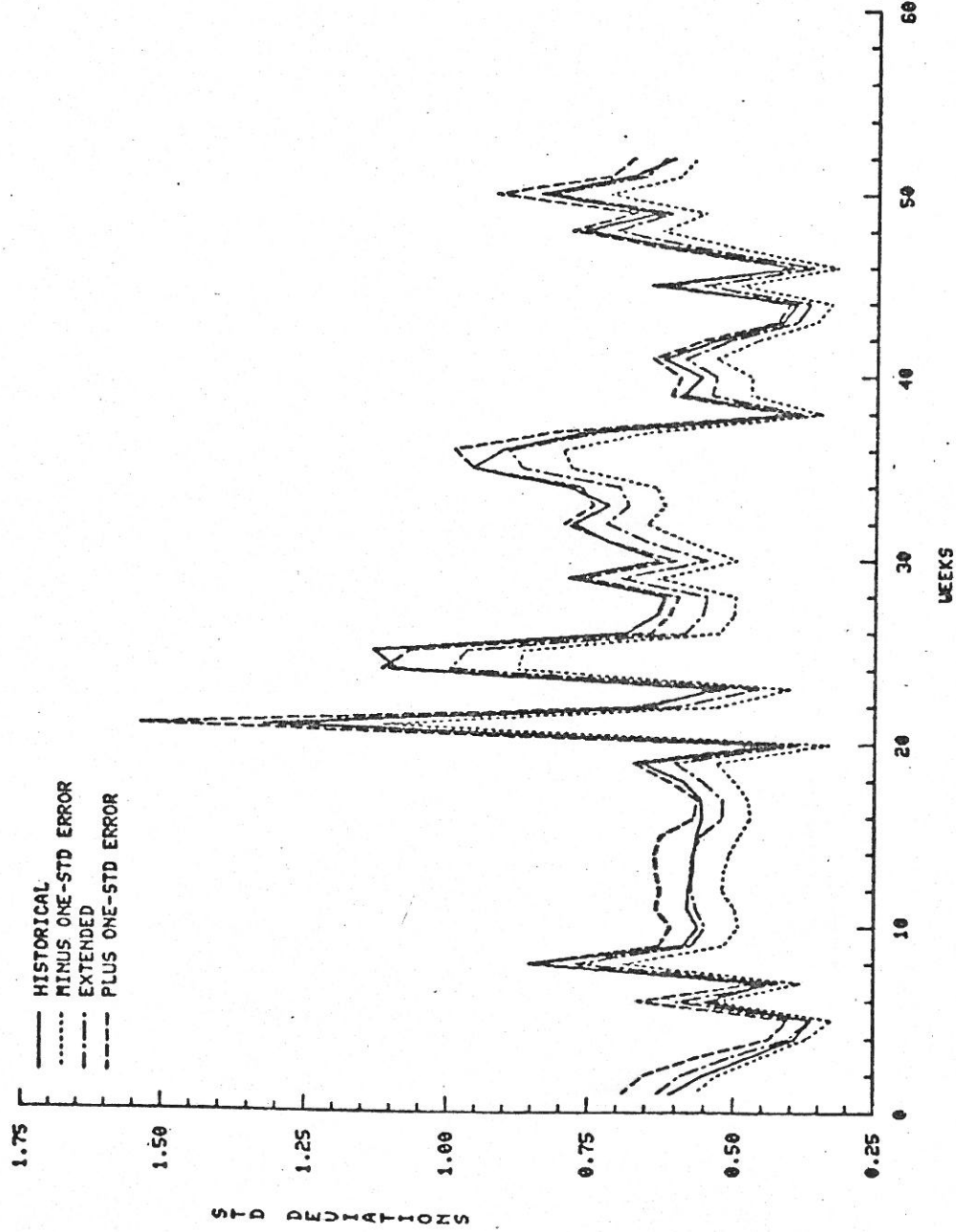


Figure 1.9.B.37. Historical and extended series weekly standard deviations of Paso del Ermitaño in logarithmic domain of flows.

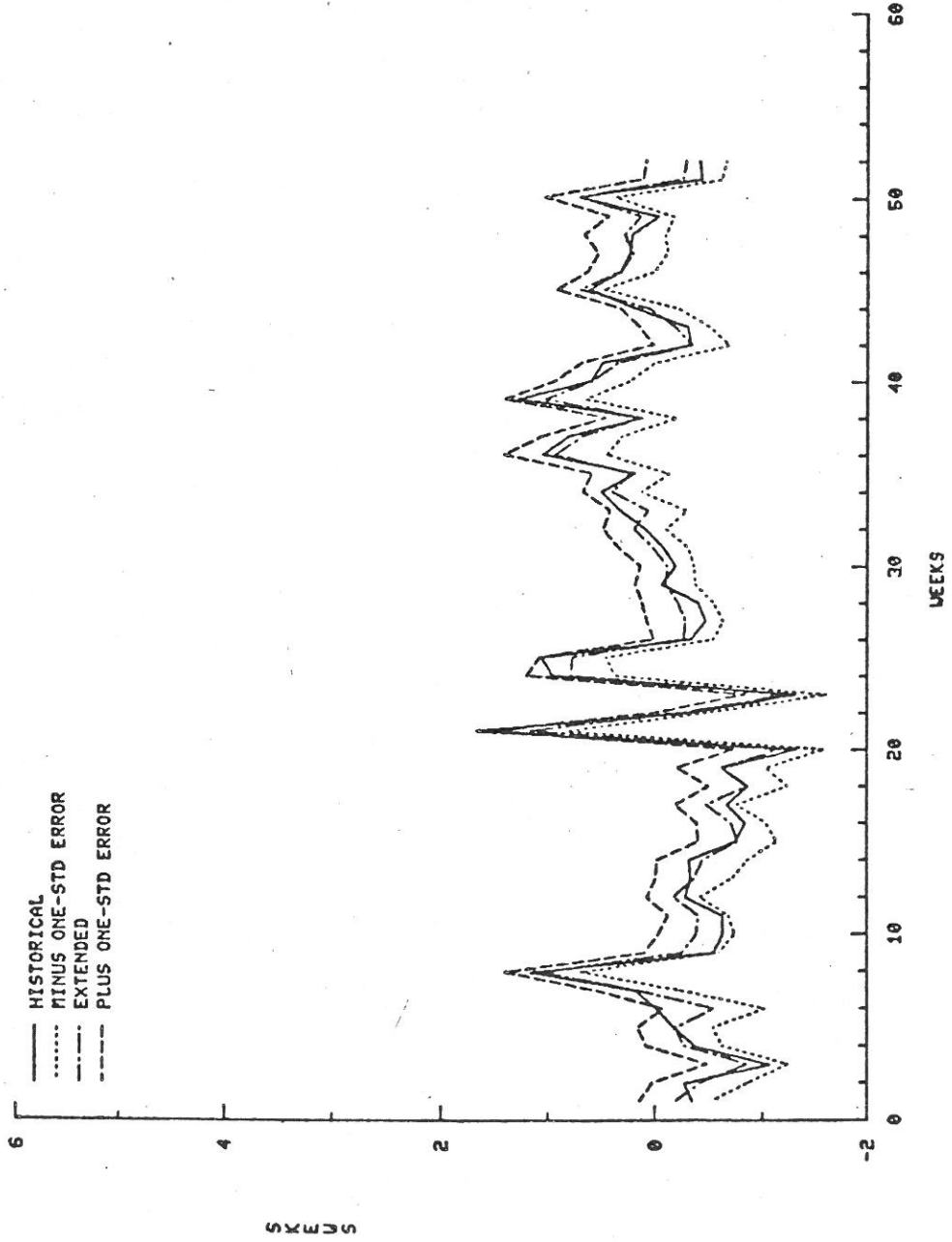


Figure 1.9.B.38. Historical and extended series weekly skewness coefficients of Paso del Ermitaño in logarithmic domain of flows.

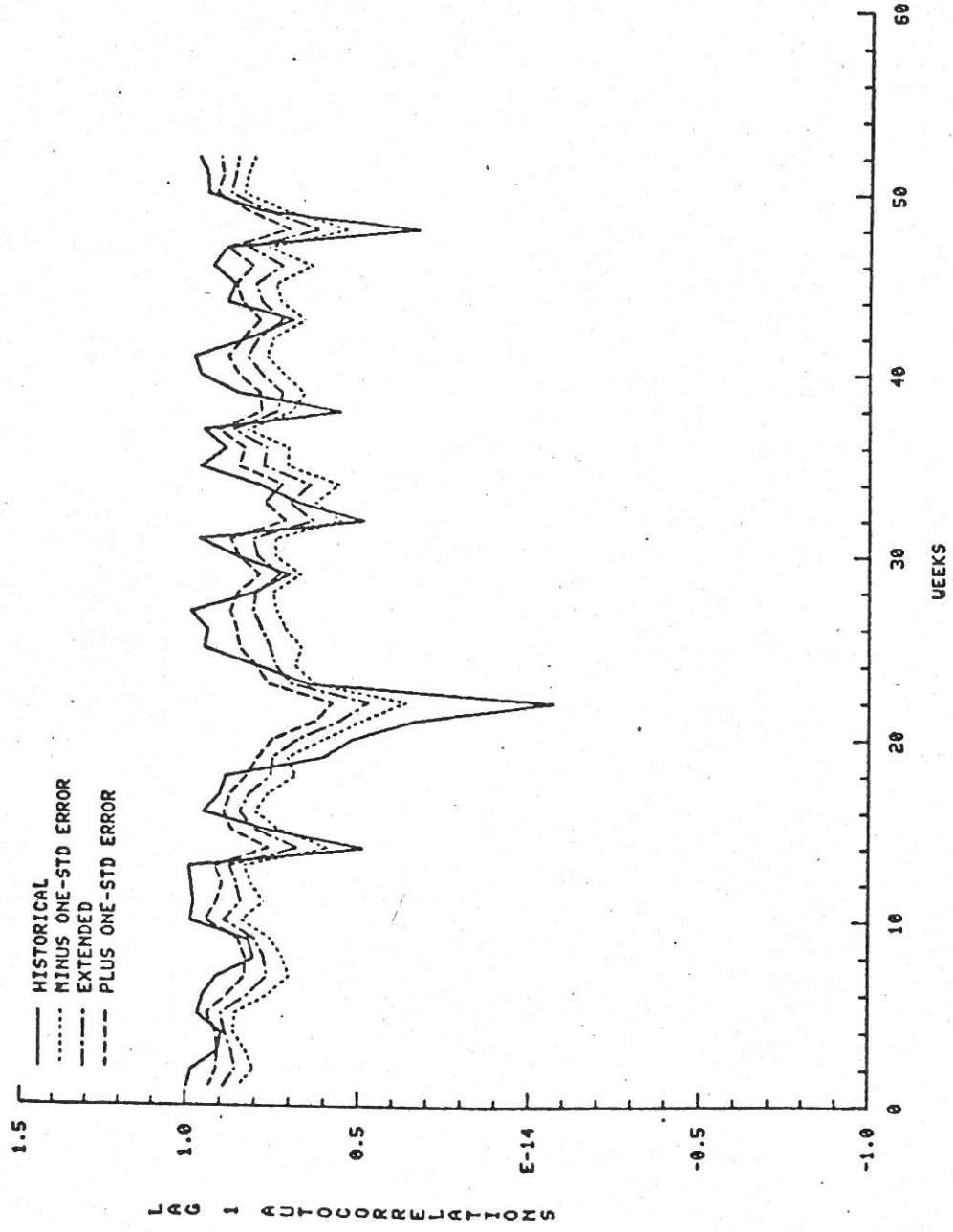


Figure 1.9.B.39. Historical and extended series weekly lag-1 autocorrelation coefficients of Paso del Ermitaño in logarithmic domain of flows.

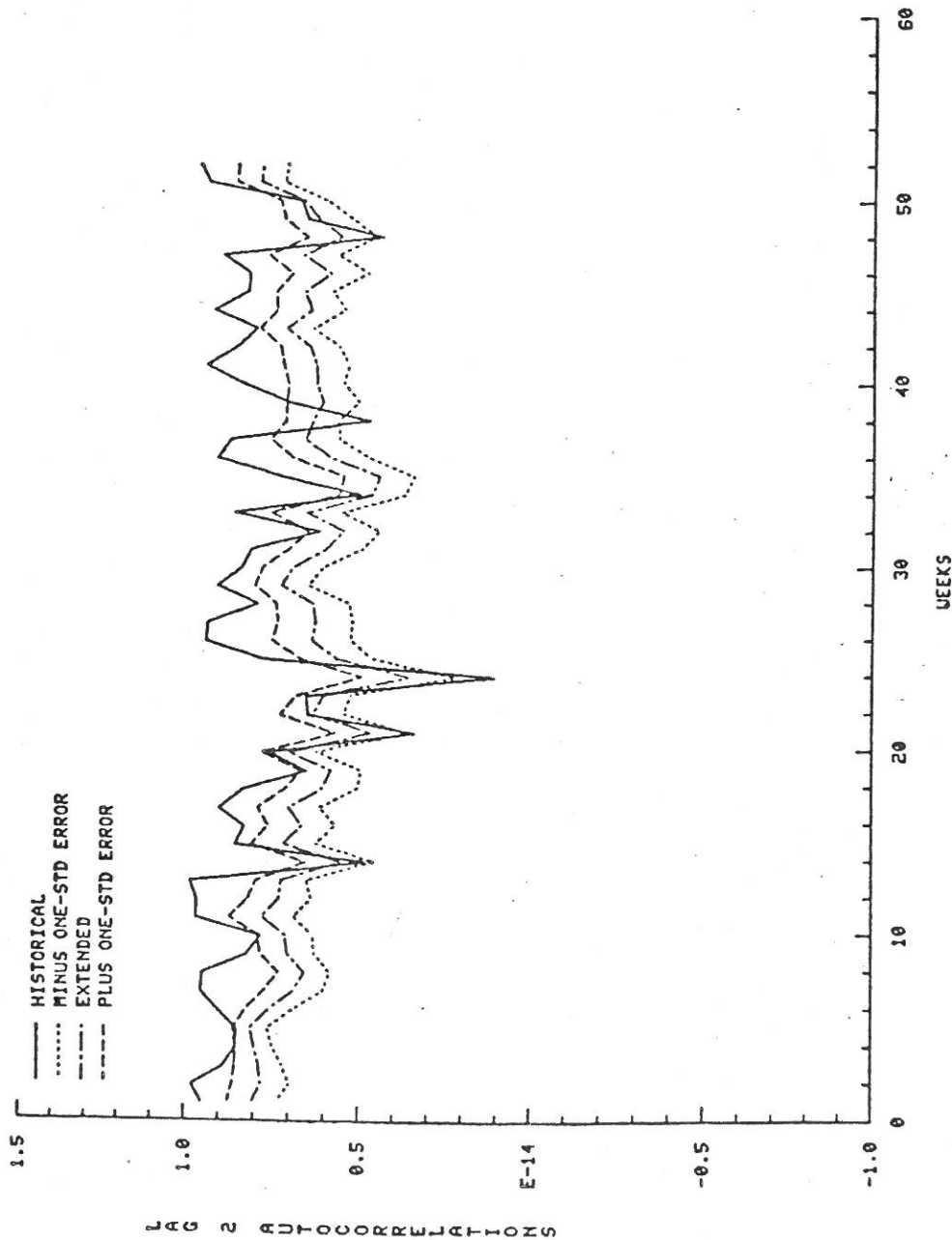


Figure 1.9.B.40. Historical and extended series weekly lag-2 autocorrelation coefficients of Paso del Ermitaño in logarithmic domain of flows.

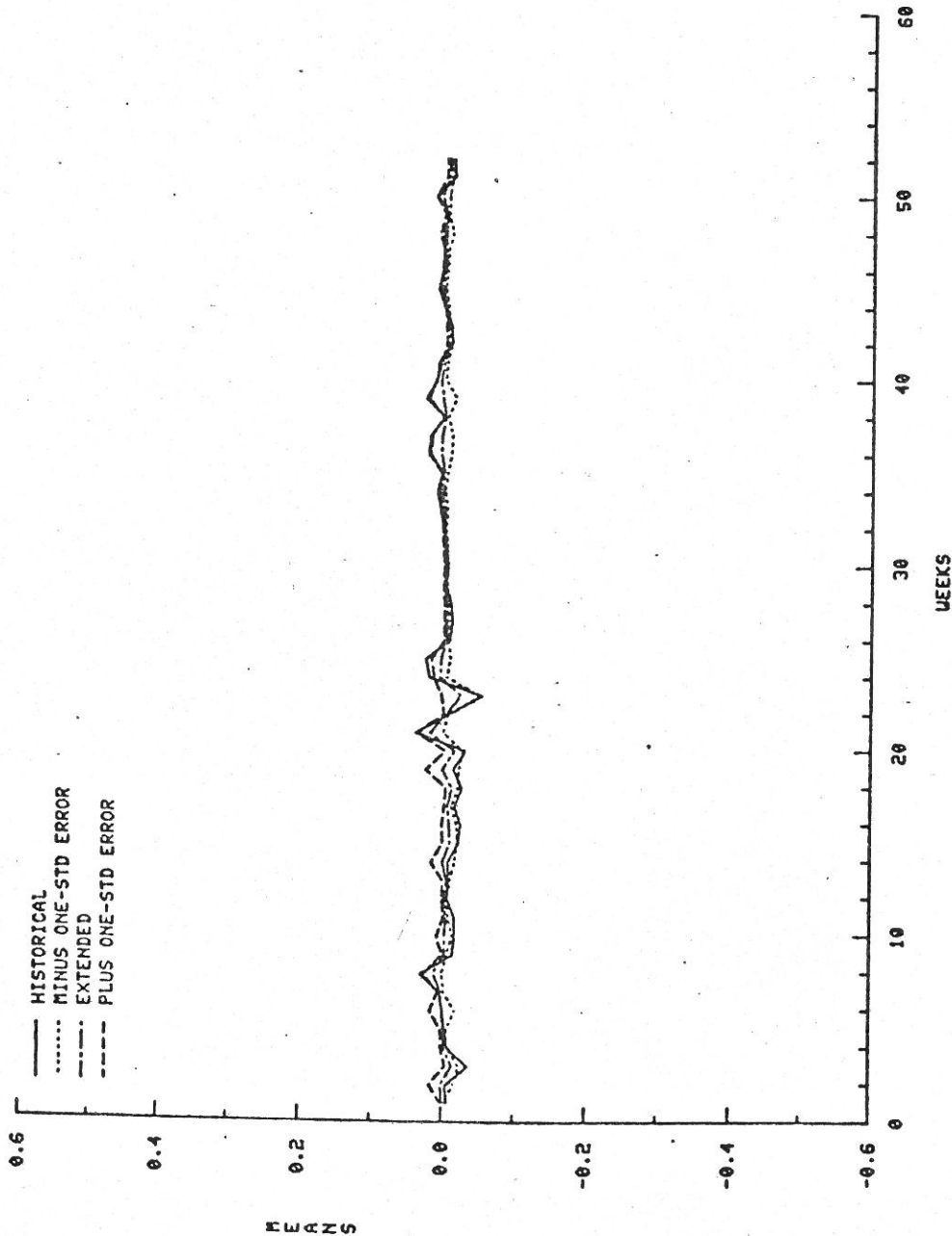


Figure 1.9.B.41. Historical and extended series weekly means of Paso del Ermitaño in log-Wilson-Hilferty domain of flows.

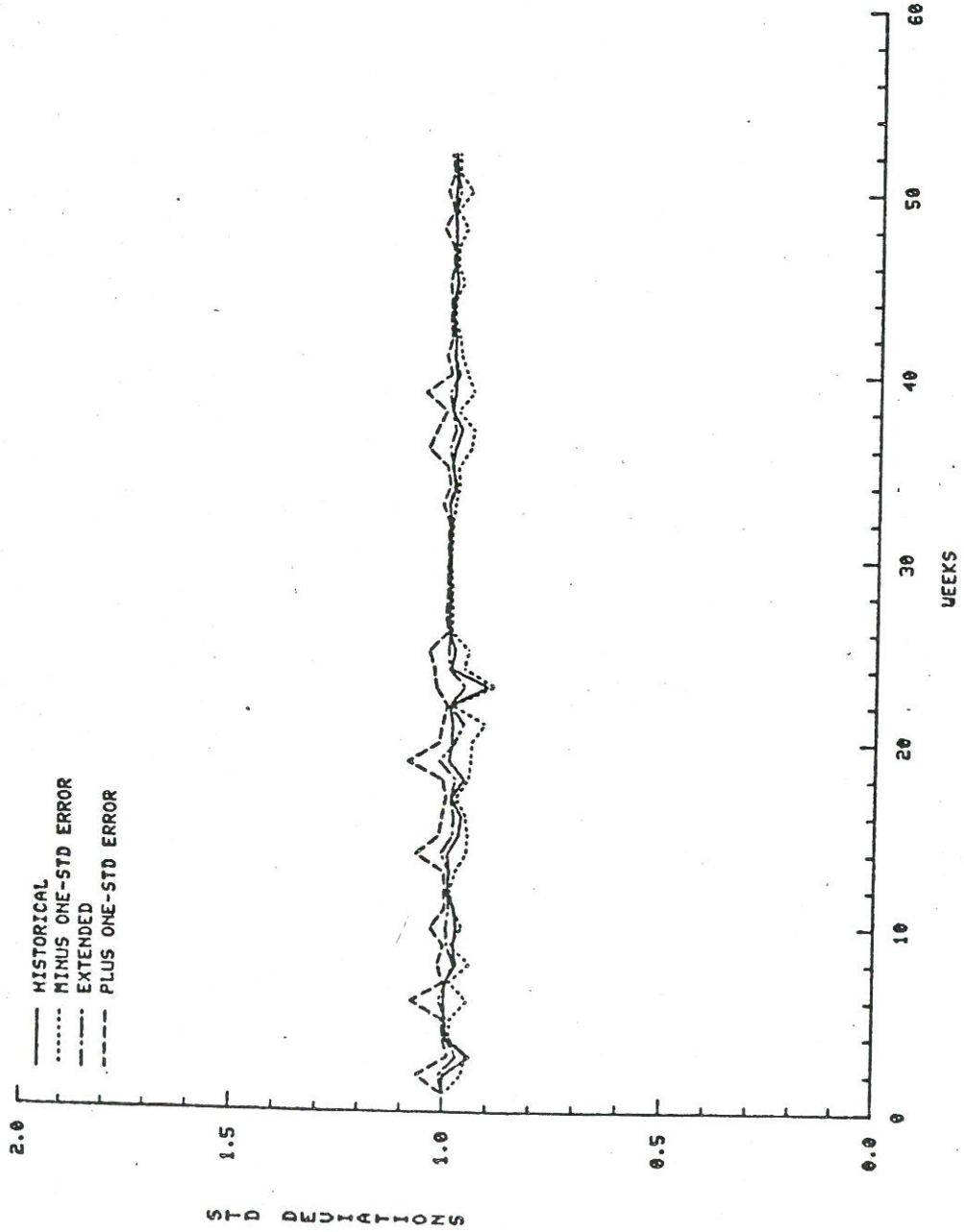


Figure 1.9.B.42. Historical and extended series weekly standard deviations of Paso del Ermitaño in log-Wilson-Hilferty domain of flows.

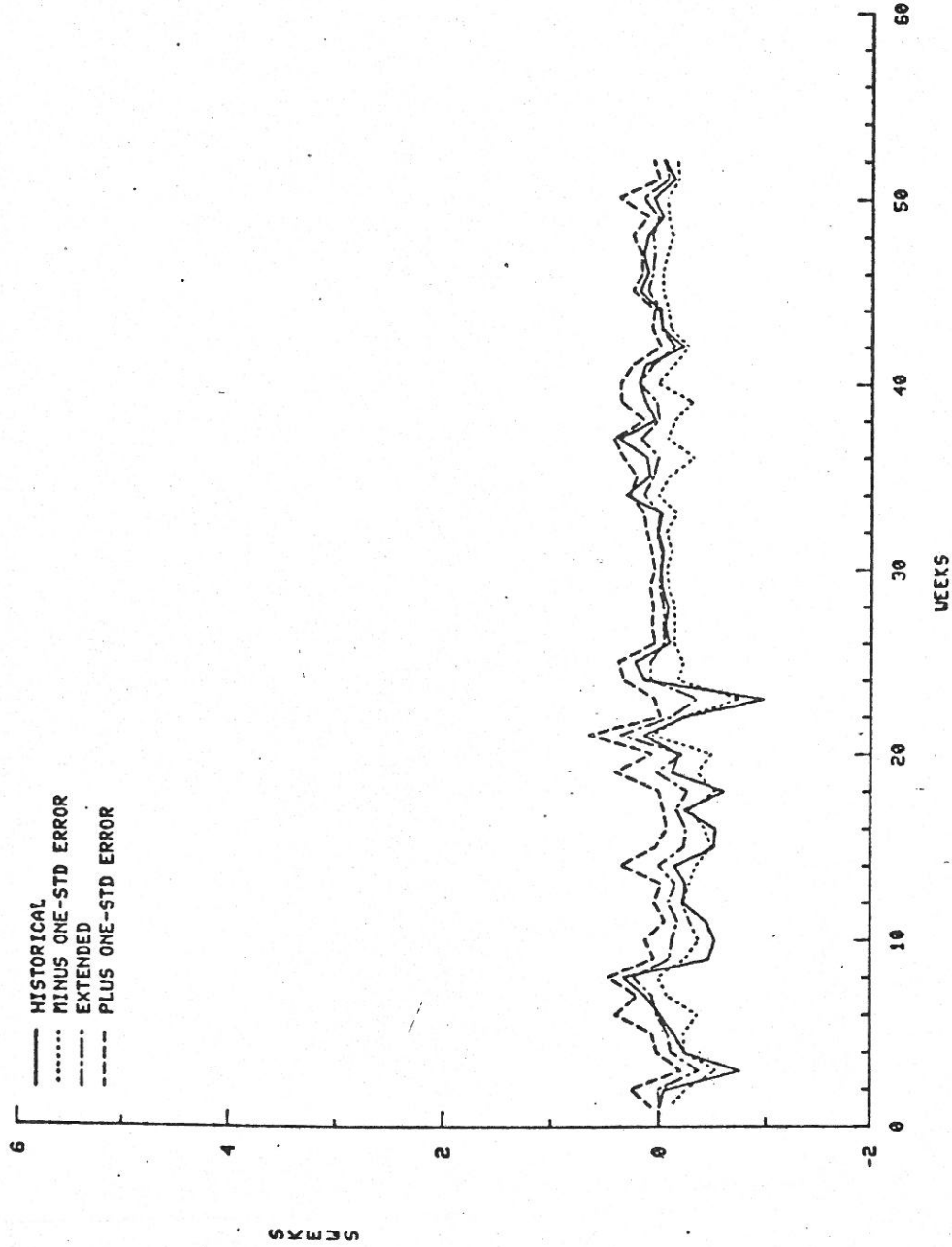


Figure 1.9.B.43. Historical and extended series weekly skewness coefficients of Paso del Ermitaño in log-Wilson-Hilferty domain of flows.

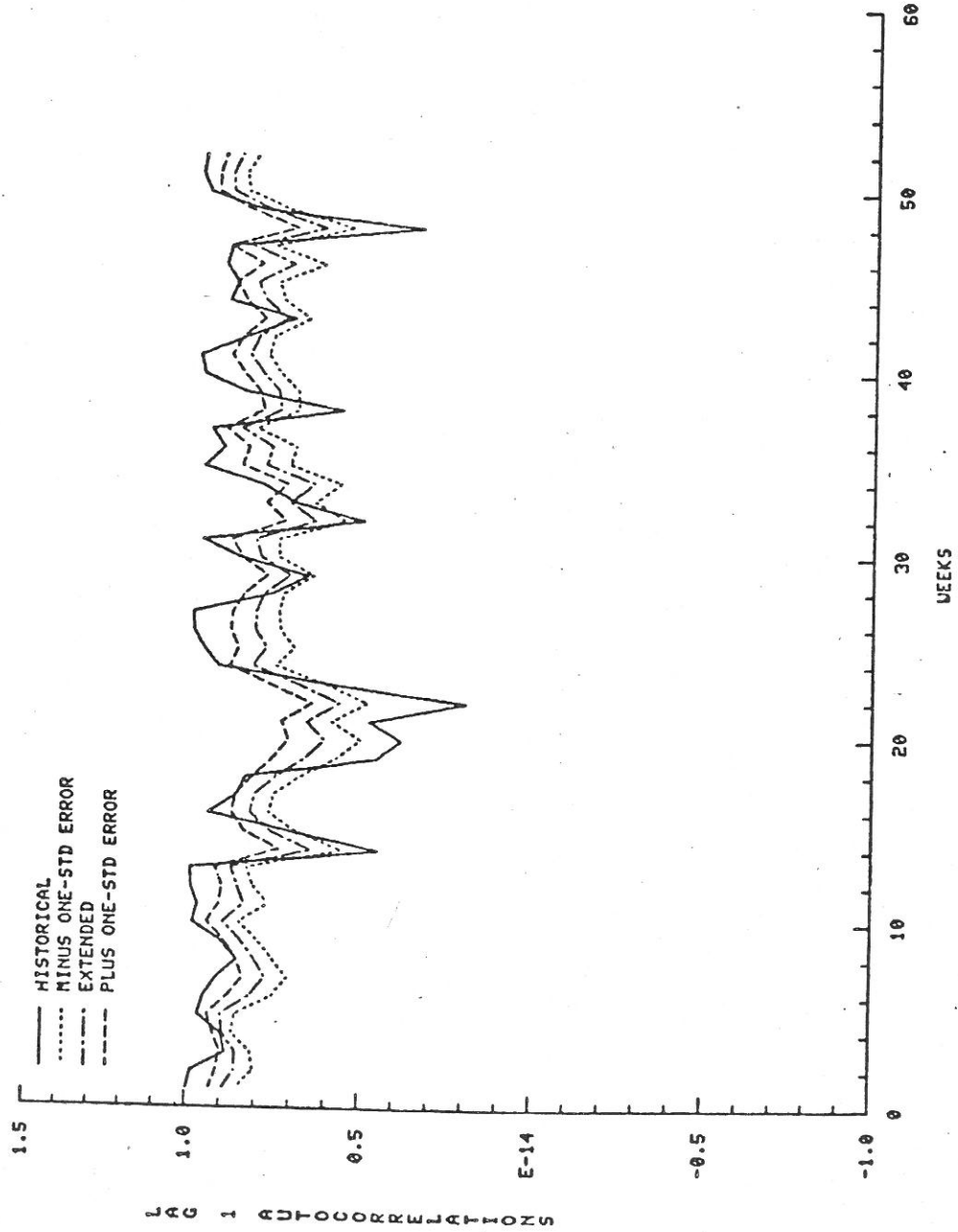


Figure 1.9.B.44. Historical and extended series weekly lag-1 autocorrelation coefficients of Paso del Ermitaño in log-Wilson-Hilferty domain of flows.

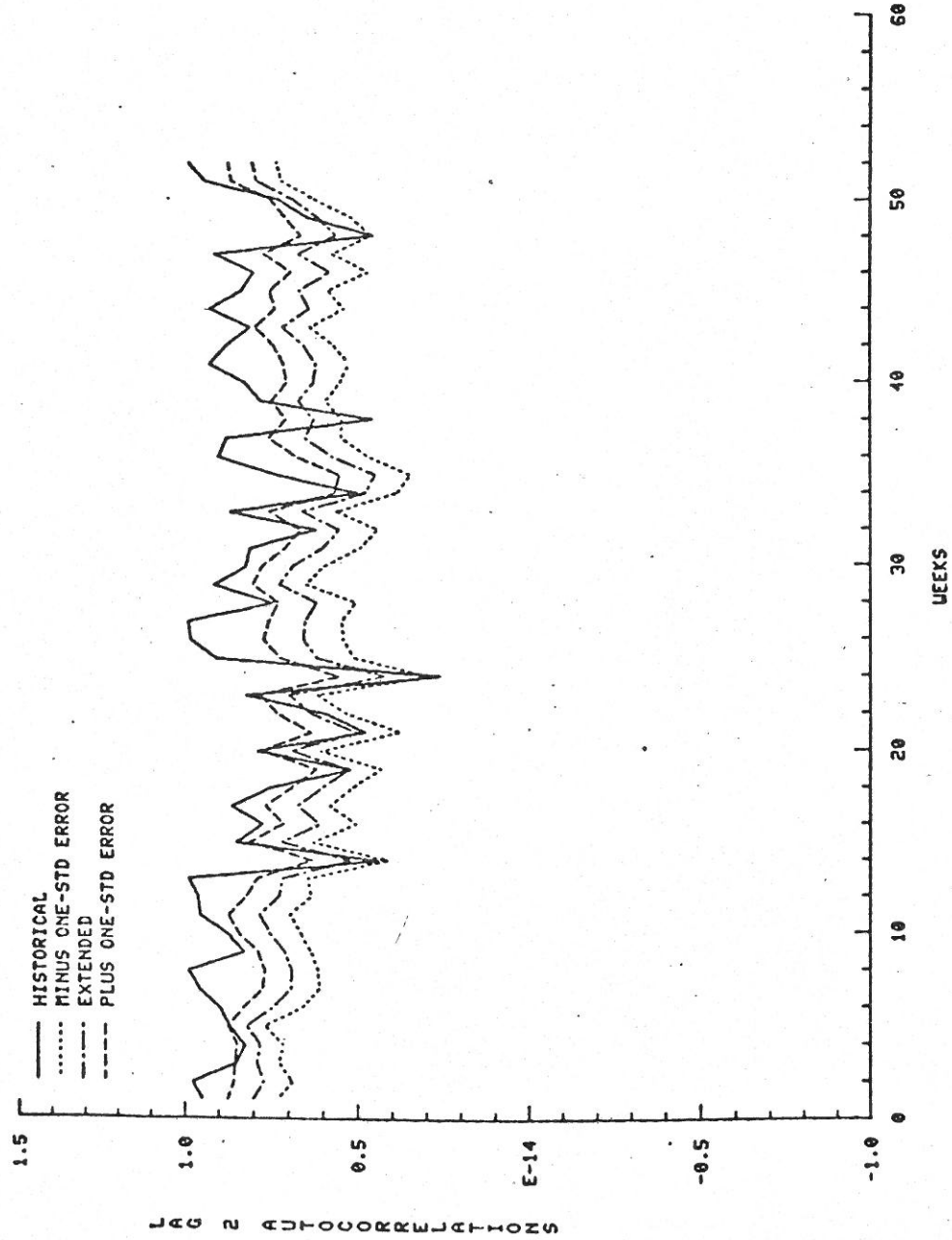


Figure 1.9.B.45. Historical and extended series weekly lag-2 autocorrelation coefficients of Paso del Ermitaño in log-Wilson-Hilferty domain of flows.

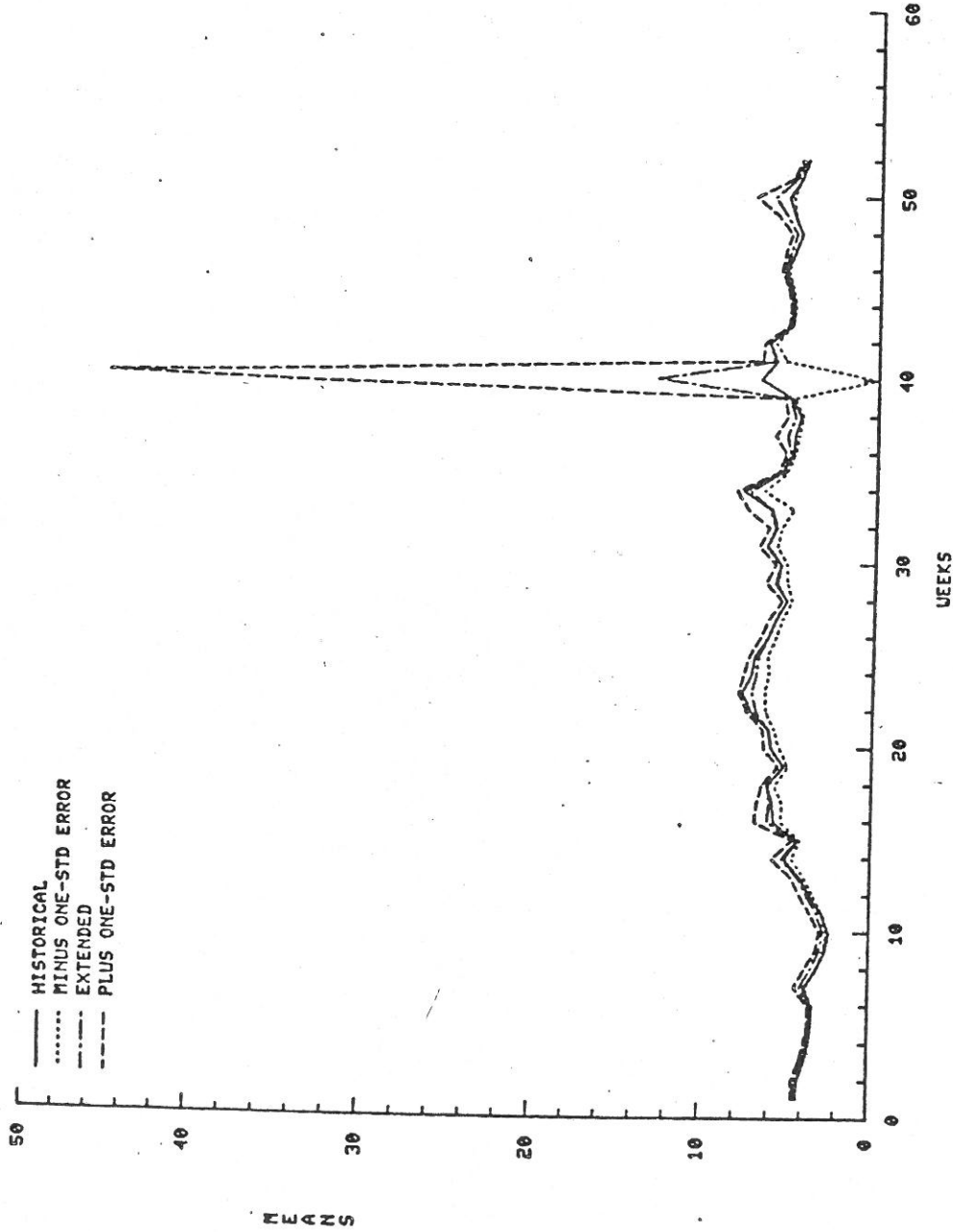


Figure 1.9.B.46. Historical and extended series weekly means of Rancho Arriba in original domain of flows.

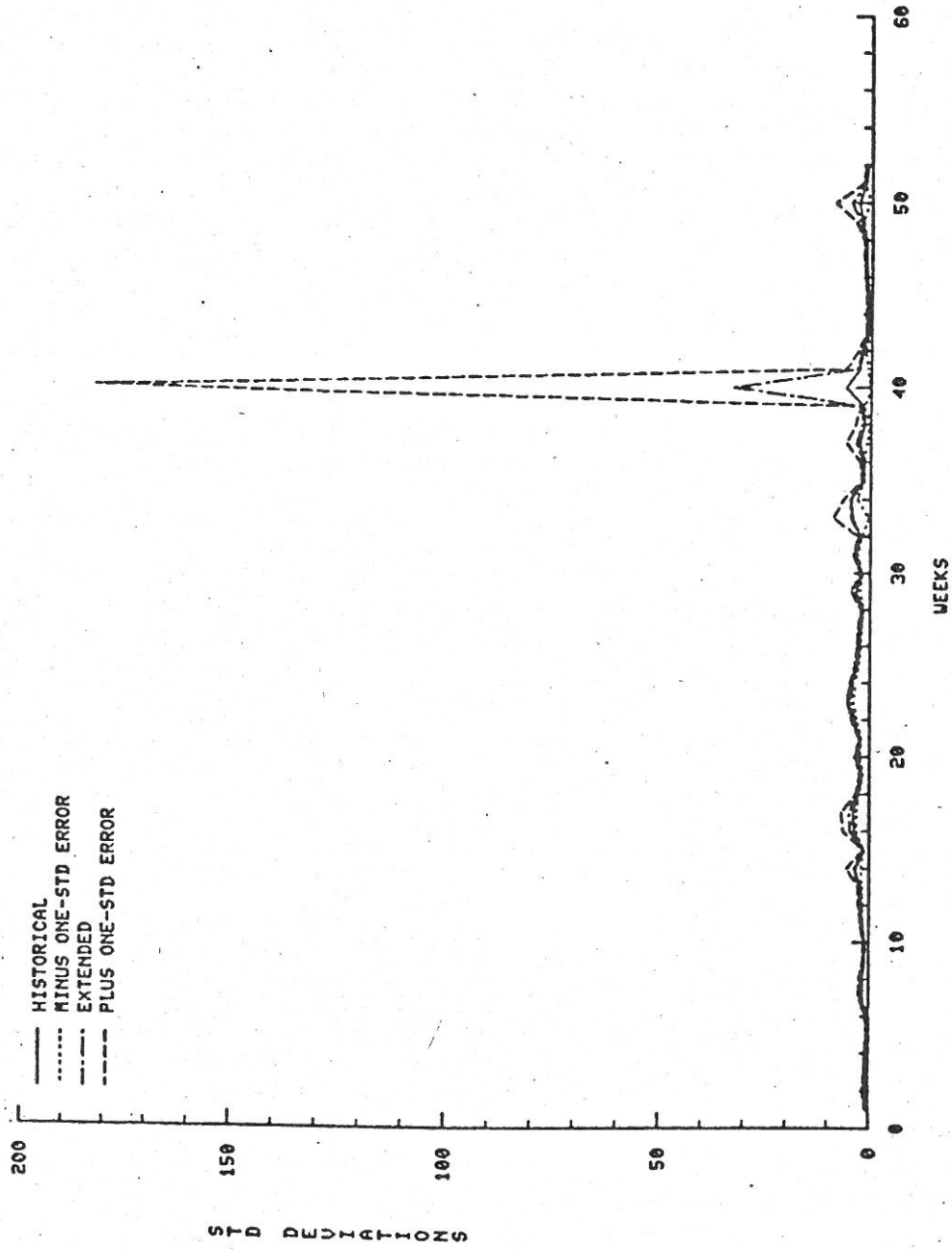


Figure 1.9.B.47. Historical and extended series weekly standard deviations of Rancho Arribain original domain of flows.

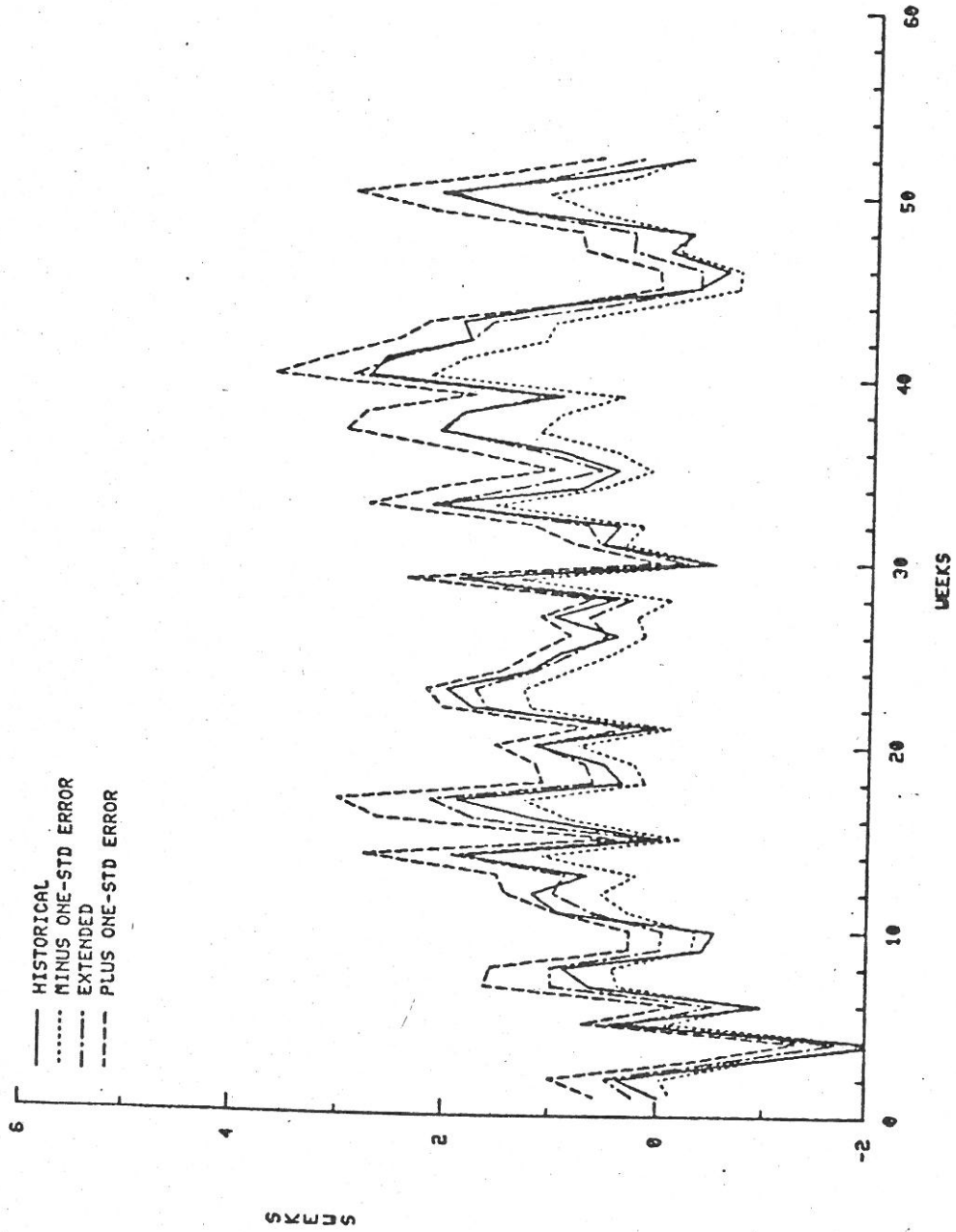


Figure 1.9.B.48. Historical and extended series weekly skewness coefficients of Rancho Arriba in original domain of flows.

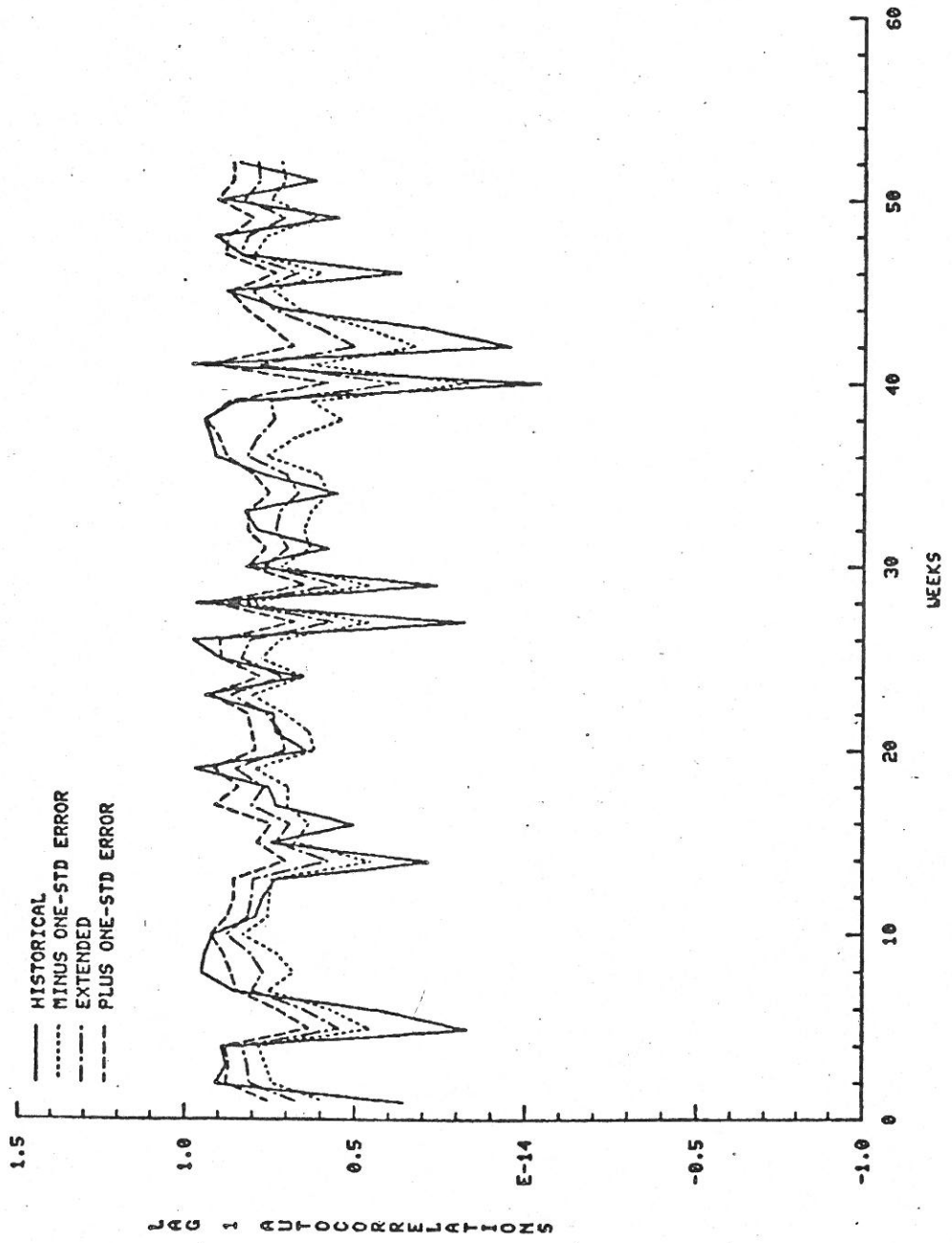


Figure 1.9.B.49. Historical and extended series weekly lag-1 autocorrelation coefficients of Rancho Arriba in original domain of flows.

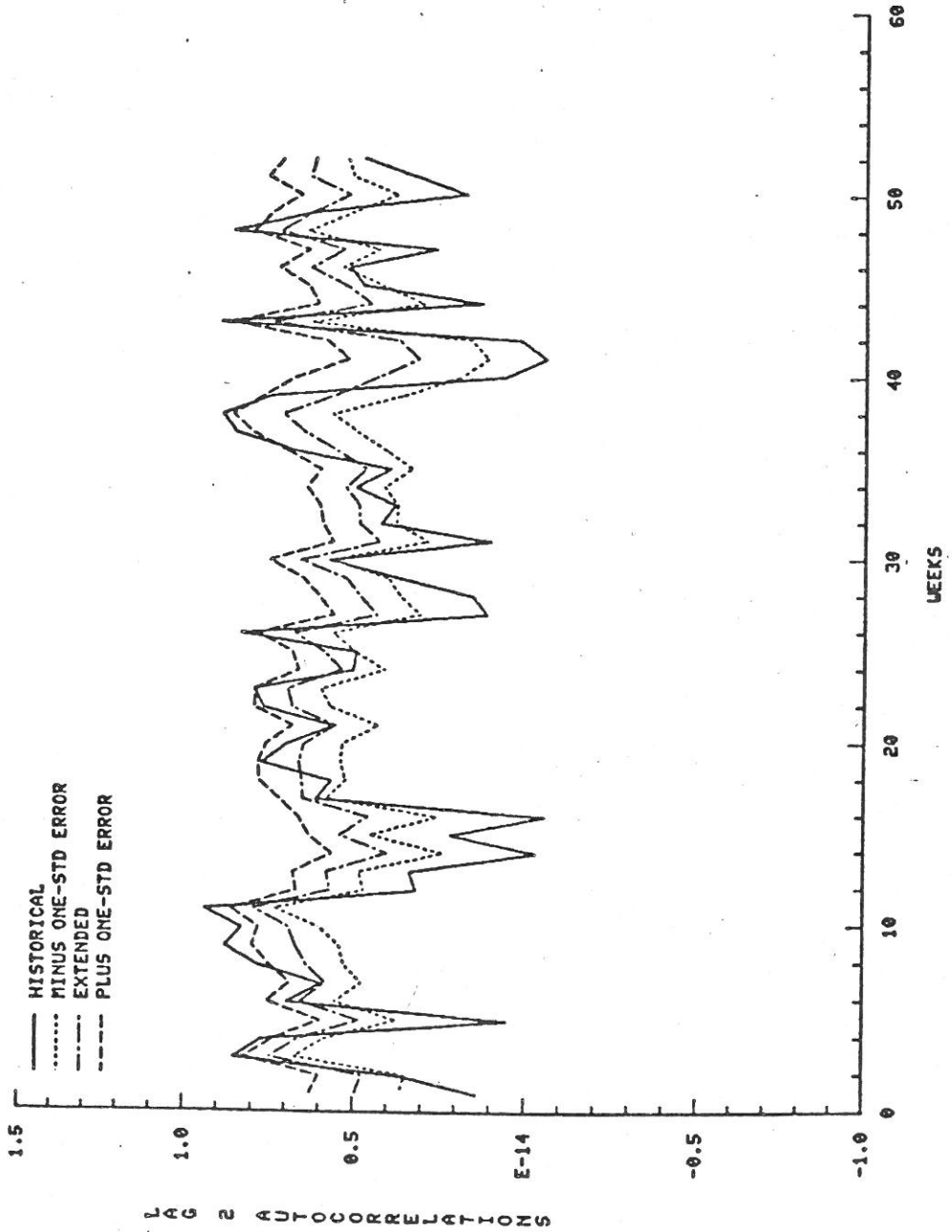


Figure 1.9.B.50. Historical and extended series weekly lag-2 autocorrelation coefficients of Rancho Arriba in original domain of flows.

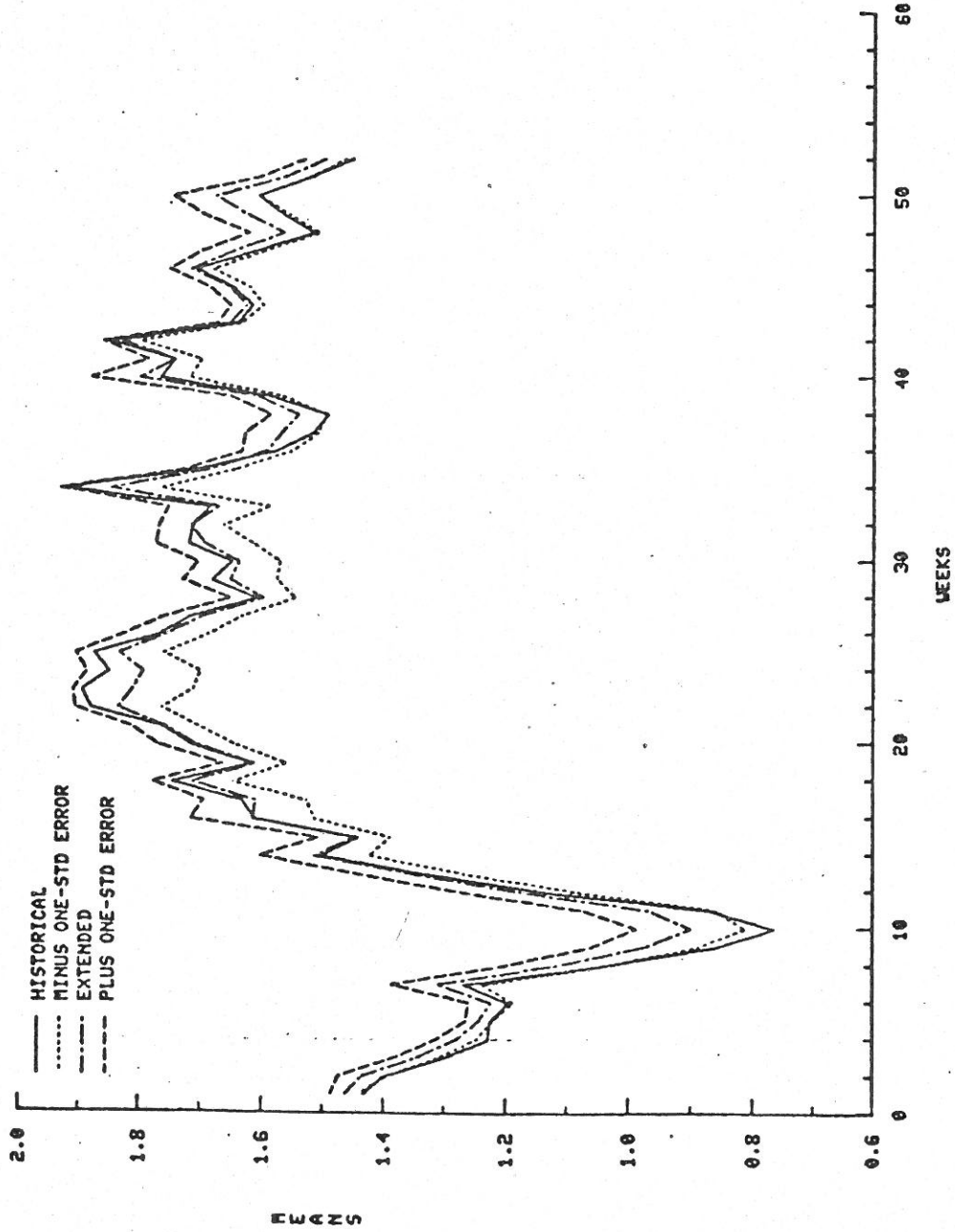


Figure 1.9.B.51. Historical and extended series weekly means of Rancho Arriba in logarithmic domain of flows.

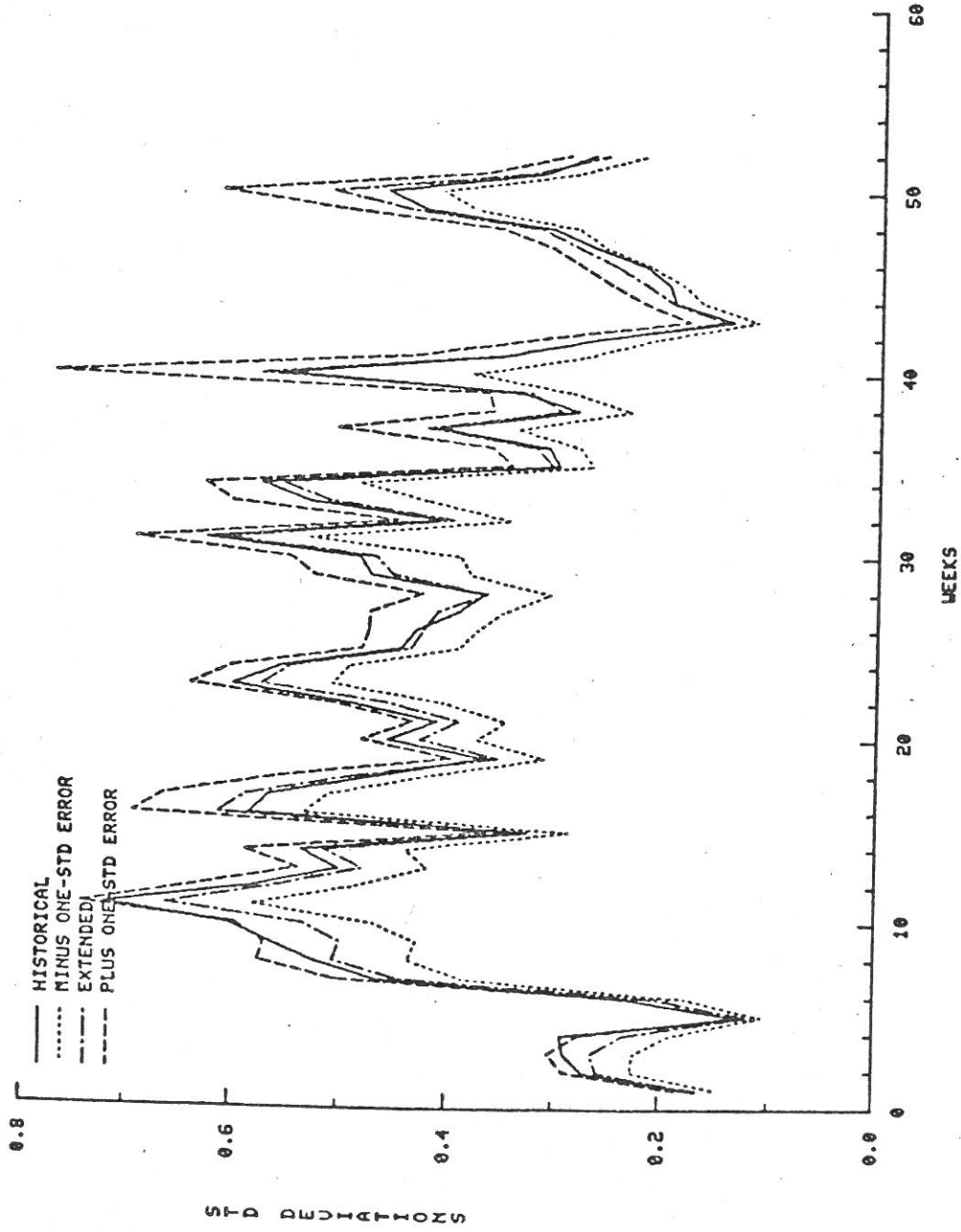


Figure 1.9.B.52. Historical and extended series weekly standard deviations of Rancho Arriba in logarithmic domain of flows.

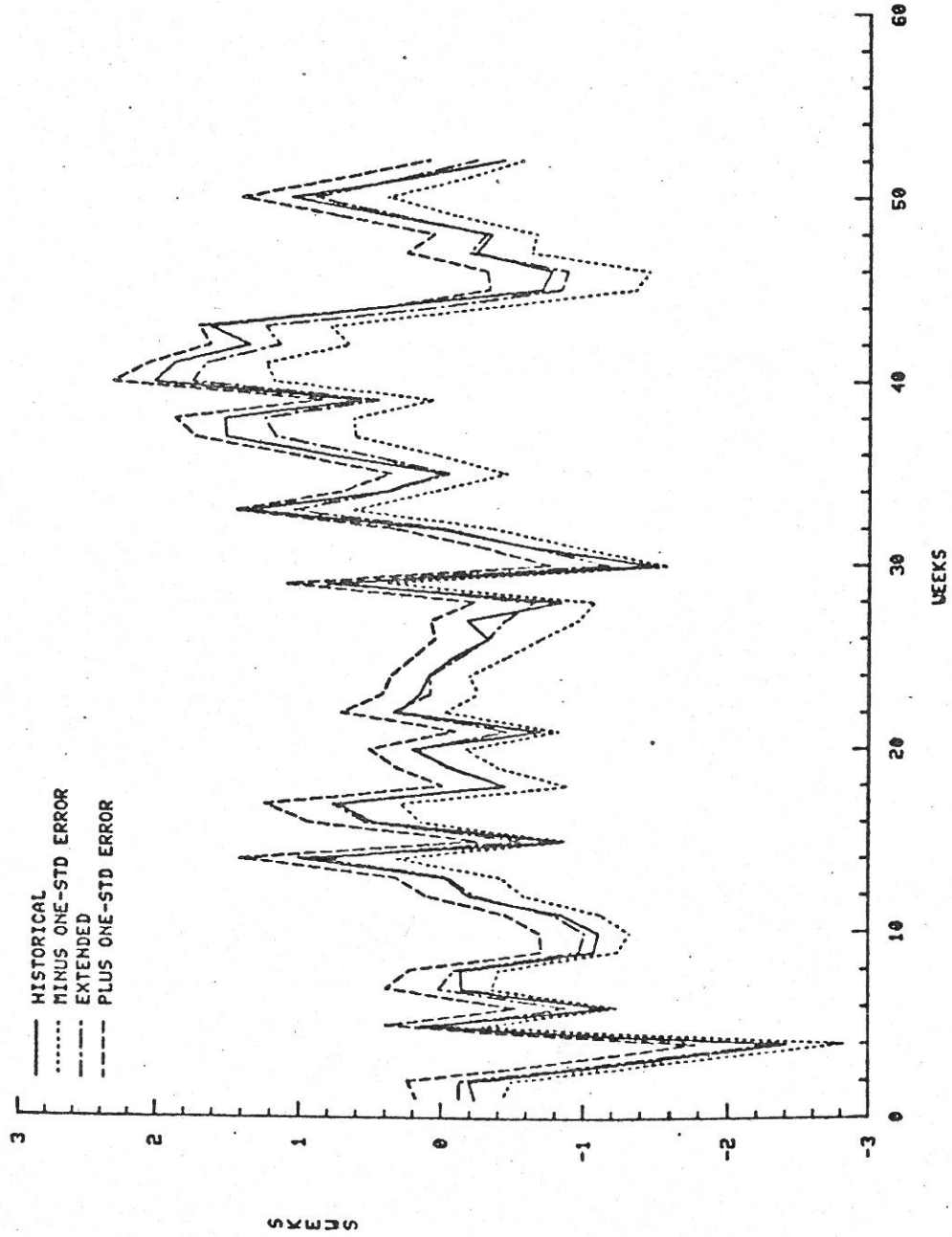


Figure 1.9.B.53. Historical and extended series weekly skewness coefficients of Rancho Arriba in logarithmic domain of flows.

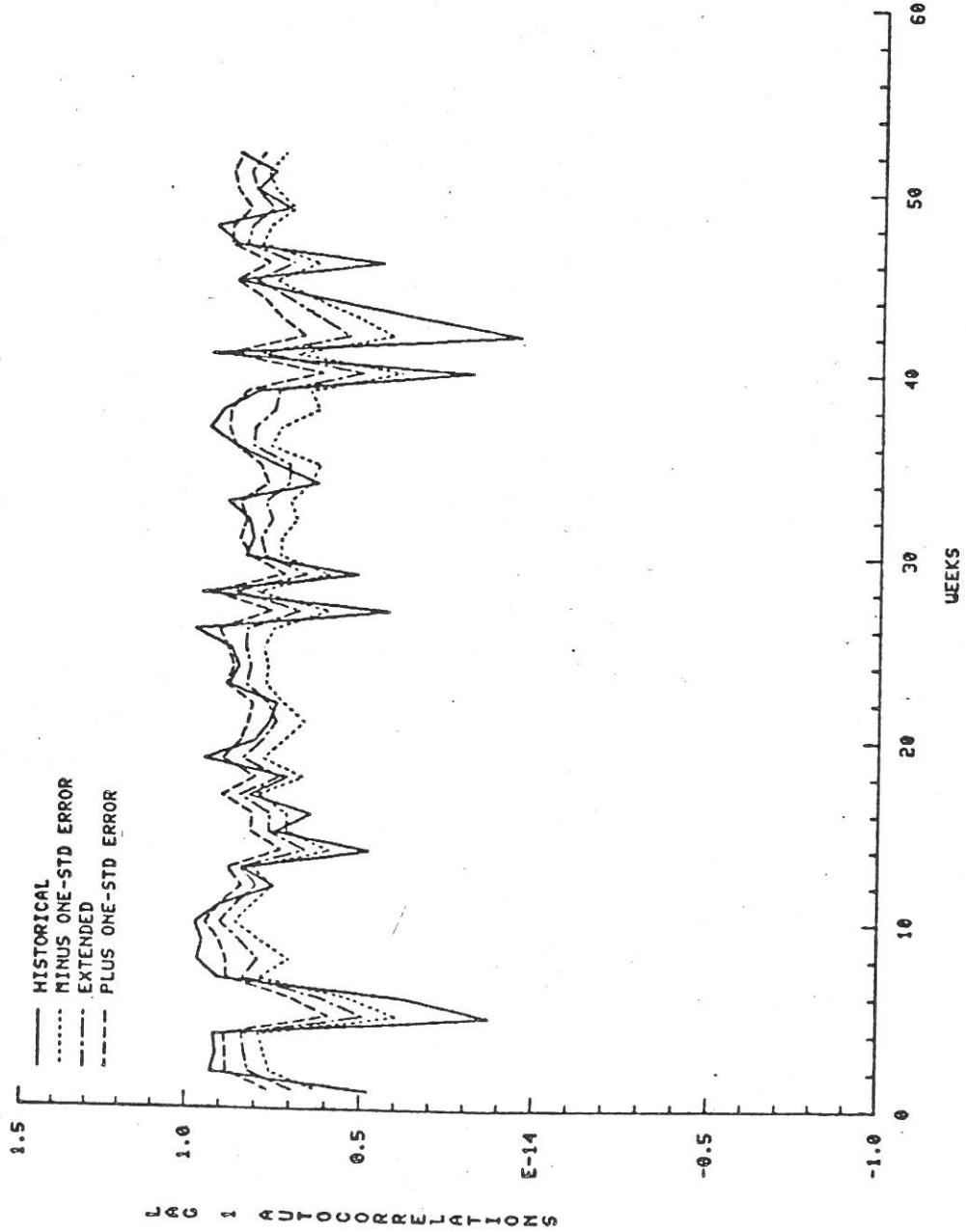


Figure 1.9.B.54. Historical and extended series weekly lag-1 autocorrelation coefficients of Rancho Arriba in logarithmic domain of flows.

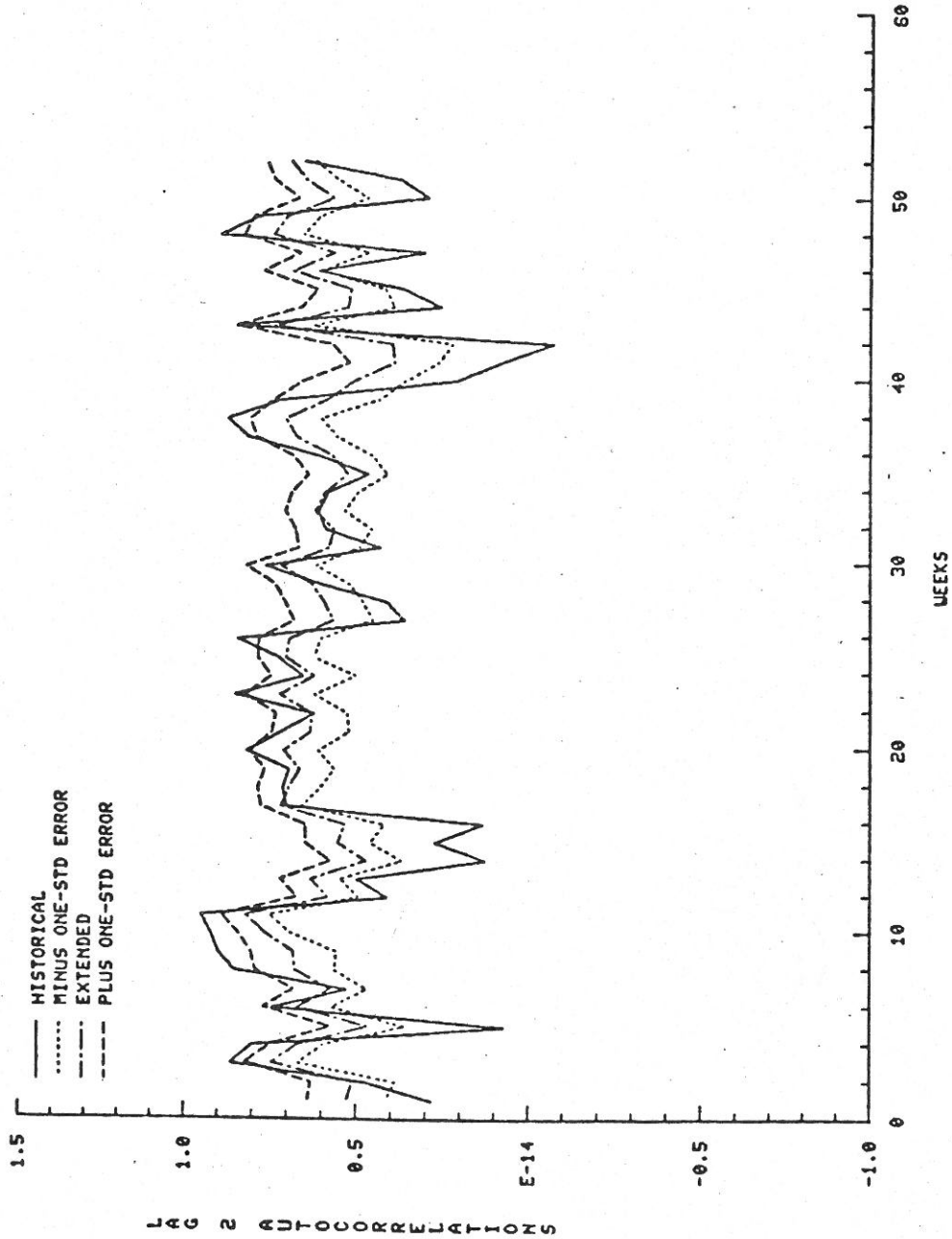


Figure 1.9.B.55. Historical and extended series weekly lag-2 autocorrelation coefficients of Rancho Arriba in logarithmic domain of flows.

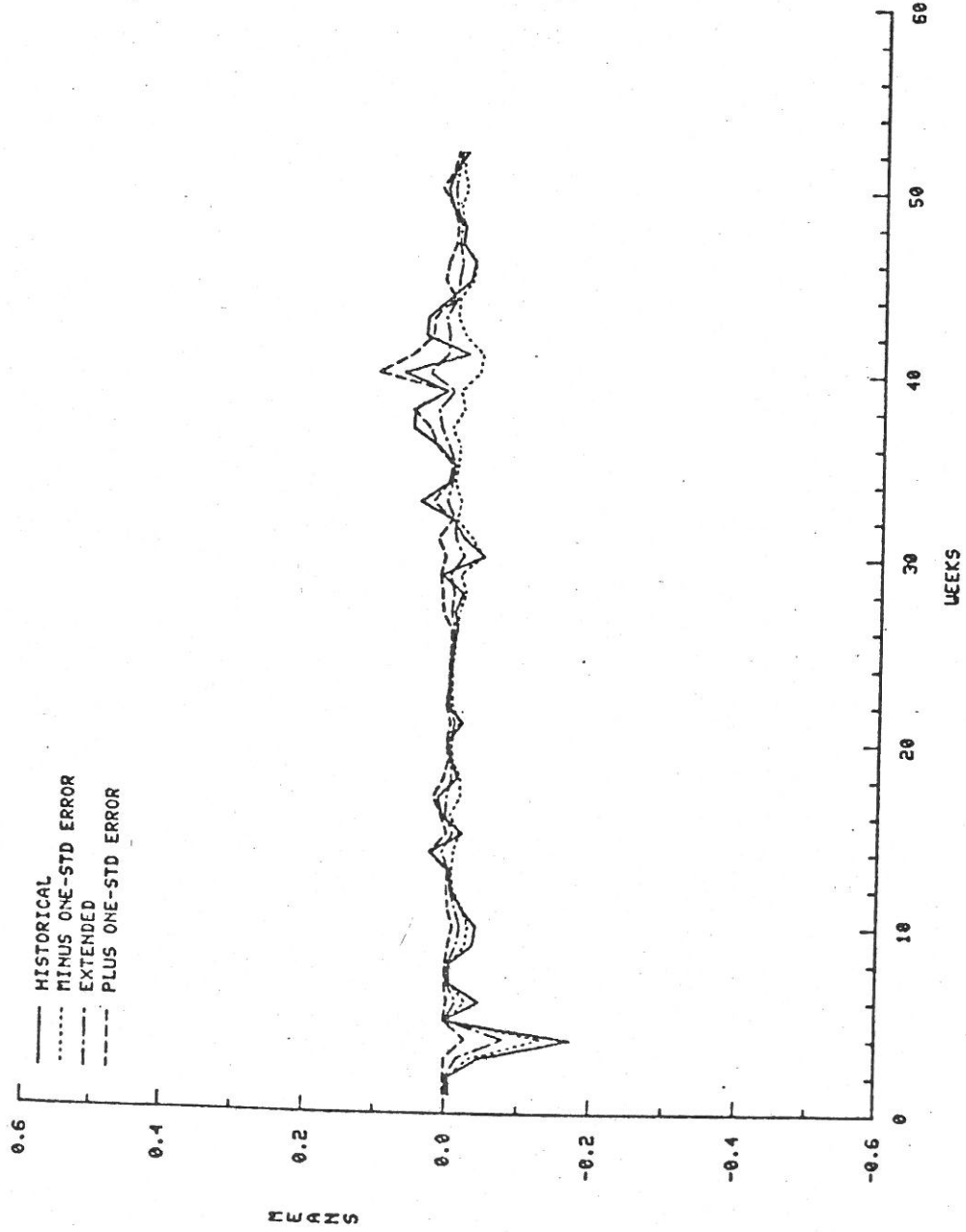


Figure 1.9.B.56. Historical and extended series weekly means of Rancho Arriba in log-Wilson-Hilferty domain of flows.

UNSATISFIED EXT P • 011666

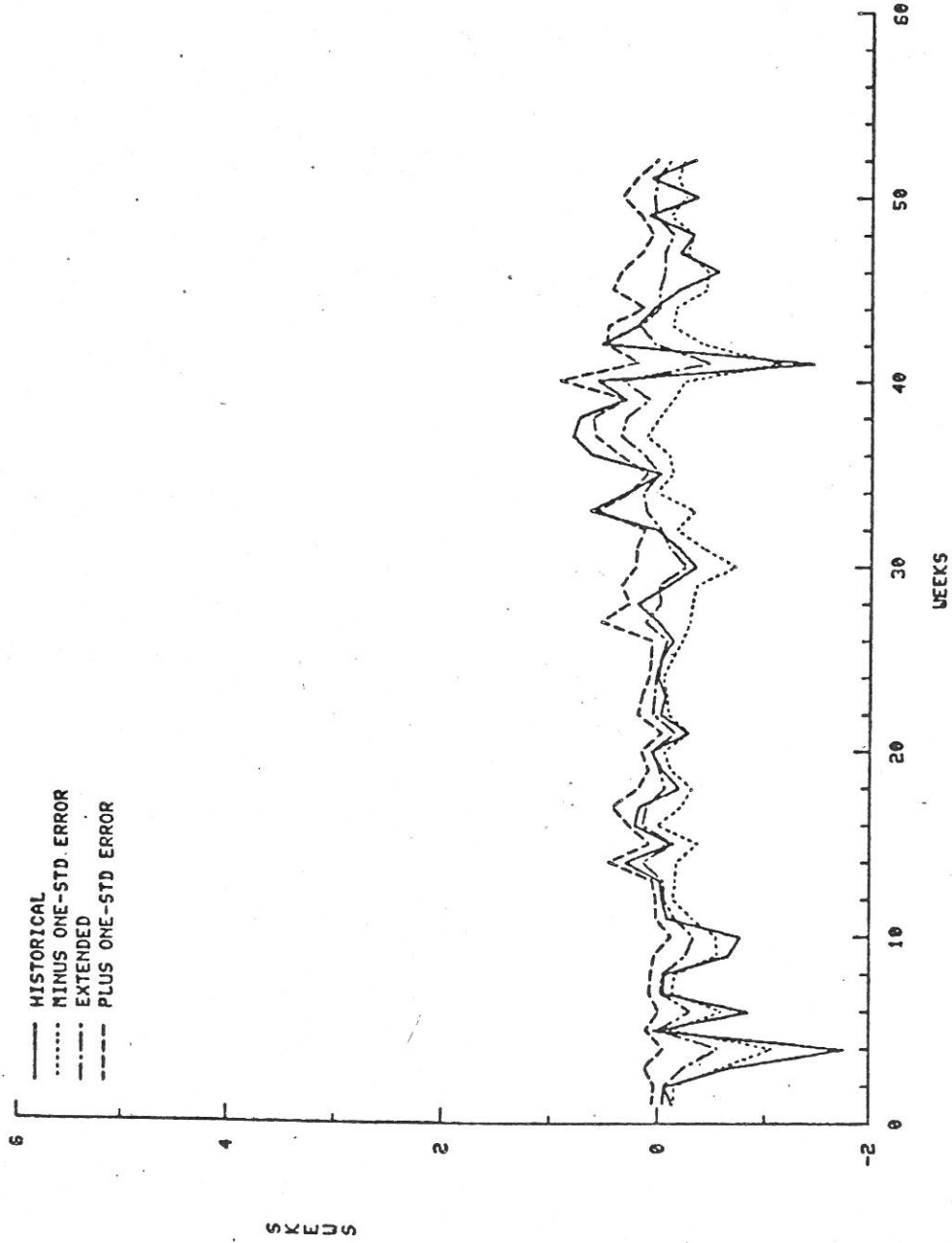


Figure 1.9.B.58. Historical and extended series weekly skewness coefficients of Rancho Arriba in log-Wilson-Hilferty domain of flows.

UNSAT

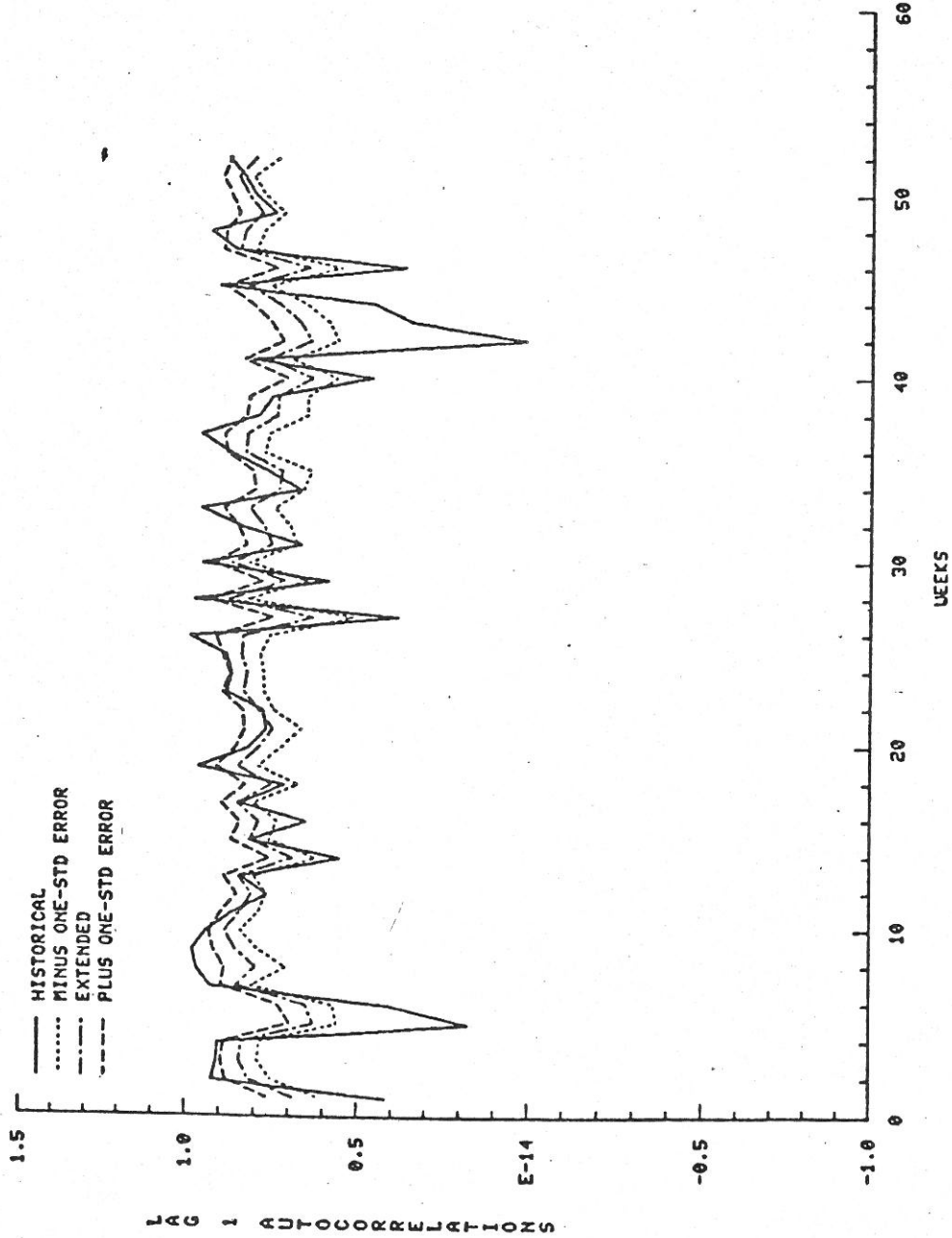


Figure 1.9.B.59. Historical and extended series weekly lag-1 autocorrelation coefficients of Rancho Arriba in log-Wilson-Hilferty domain of flows.

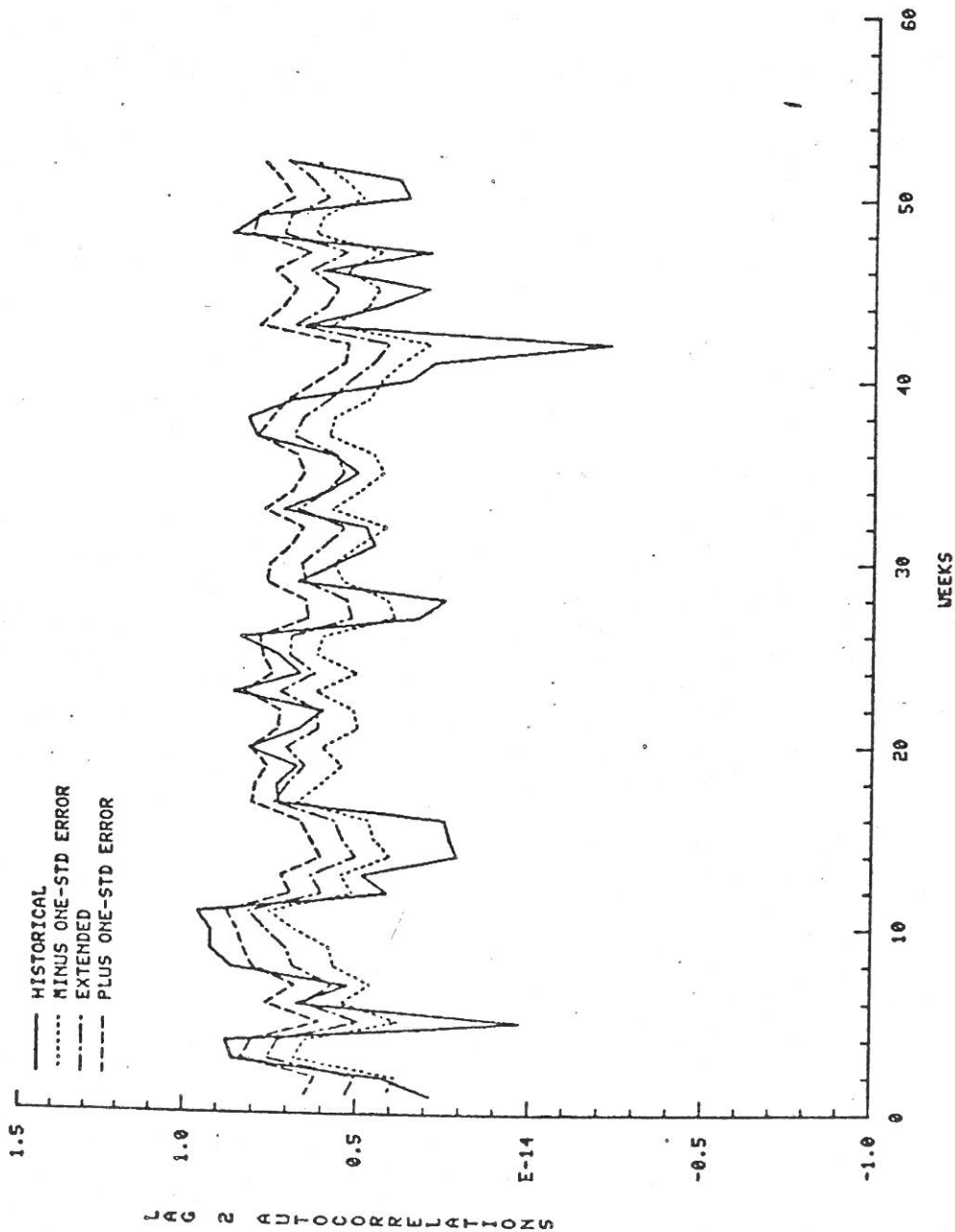


Figure 1.9.B.60. Historical and extended series weekly lag-2 autocorrelation coefficients of Rancho Arriba in log-Wilson-Hilferty domain of flows.

APPENDIX 1.9.C

FOURIER SERIES FITTING OF PERIODIC STATISTICAL PARAMETERS

Consider that U_τ represents the periodic statistical parameter such as skewness G_τ or the autocorrelation coefficient $R_\tau(k)$. The Fourier series representation of U_τ denoted by U_τ^* is obtained from (Salas, et al., 1980):

$$U_\tau^* = \bar{U} + \sum_{j=1}^h [A_j \cos(2\pi j\tau/\omega) + B_j \sin(2\pi j\tau/\omega)] \quad (1.9.C.1)$$

for $\tau = 1, \dots, \omega$ semi-months. The mean \bar{U} and Fourier coefficients A_j and B_j are determined by

$$\bar{U} = \frac{1}{\omega} \sum_{\tau=1}^{\omega} U_\tau \quad (1.9.C.2a)$$

$$A_j = \frac{2}{\omega} \sum_{\tau=1}^{\omega} U_\tau \cos\left(\frac{2\pi j\tau}{\omega}\right); \quad j = 1, \dots, h \quad (1.9.C.2b)$$

and

$$B_j = \frac{2}{\omega} \sum_{\tau=1}^{\omega} U_\tau \sin\left(\frac{2\pi j\tau}{\omega}\right); \quad j = 1, \dots, h \quad (1.9.C.2c)$$

The total number of harmonics h is theoretically equal to $\omega/2$ for ω an even number or equal to $(\omega-1)/2$ for ω an odd number. However, for purposes of removing sampling variabilities in the sample series U_τ , only a few harmonics are necessary. The selection of

harmonics may be decided based on the significance of explained variance of each harmonic component. The so-called explained variance for each harmonic is computed from

$$EV_j = \frac{(A_j^2 + B_j^2)}{S^2(u)} \times 100 \text{ percent} \quad (1.9.C.3)$$

where EV_j is the explained variance in percent and $S^2(u)$ is the variance of the series $\{U_r\}$. Further details of selection of significant harmonics are given by Salas, et al. (1980).

Results of the Fourier series fitting of the monthly and weekly skewness coefficients in the log-domain for Palo de Caja, Paso del Ermitaño and Rancho Arriba are given in Figures 1.9.C.1 through 1.9.C.6. For monthly skewness coefficients, the first 2, 3, and 4 harmonics were fitted while those for weekly, the first 4, 5, and 6 harmonics were fitted. In the figures, the skewness coefficients of the extended series are referred to as "historical" and has been corrected for bias prior to Fourier series fitting. The bias correction for the skewness is based on gamma distribution which implies that the extended series in the log-domain of flows are assumed to be gamma distributed. This assumption is made post-de-facto since the suitable normalizing transformation is found to be the Wilson-Hilferty transformation (which is a gamma-based transformation) which was performed after logarithmic transformation. In equation form, the skewness $(G_r)_c$ corrected for bias is given by (Yevjevich and Obeysekera, 1984):

$$(G_r)_c = G_r \left[\left(1 + \frac{6.51}{N} + \frac{20.20}{N^2} \right) + \left(\frac{1.48}{N} + \frac{6.77}{N^2} \right) G_r^2 \right]$$

where G_r is the extended series, average skewness of season r and N is the number of years of record. From results herein, it is decided that the Fourier fitted functions using the first two harmonics be used for monthly skews while the first four harmonics be used for weekly skews.

Figures 1.9.C.7 to 1.9.C.18 show the results of Fourier series fitting of lag-1 and lag-2, monthly and weekly autocorrelations for the three stations in the log-Wilson-Hilferty domain of flows. From results herein, it is likewise decided that the first two harmonic fitted functions be used for monthly autocorrelations while the first four harmonic functions be used for weekly autocorrelations.

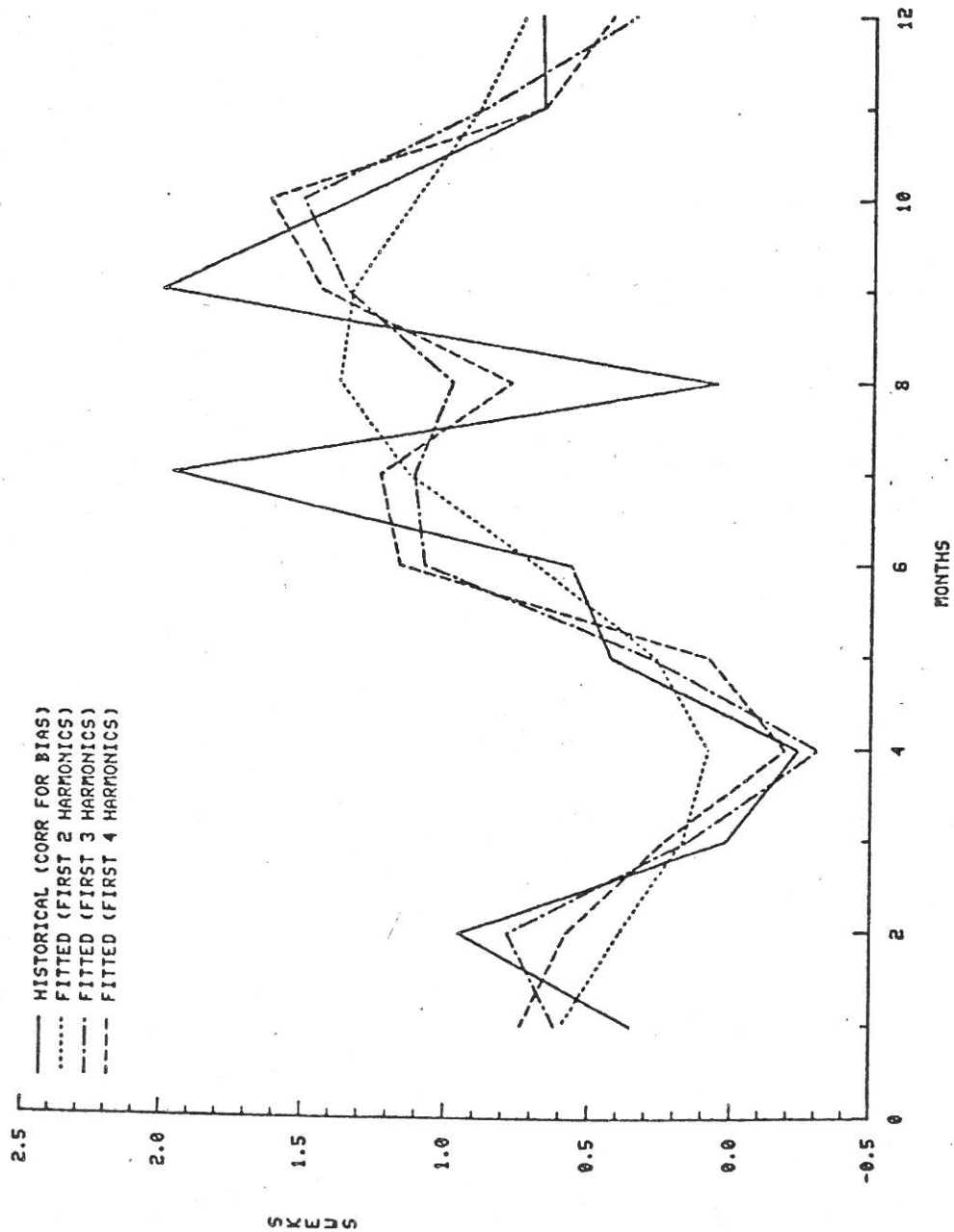


Figure 1.9.C.1 PALO DE CAJA - SKEWS IN LOG DOMAIN

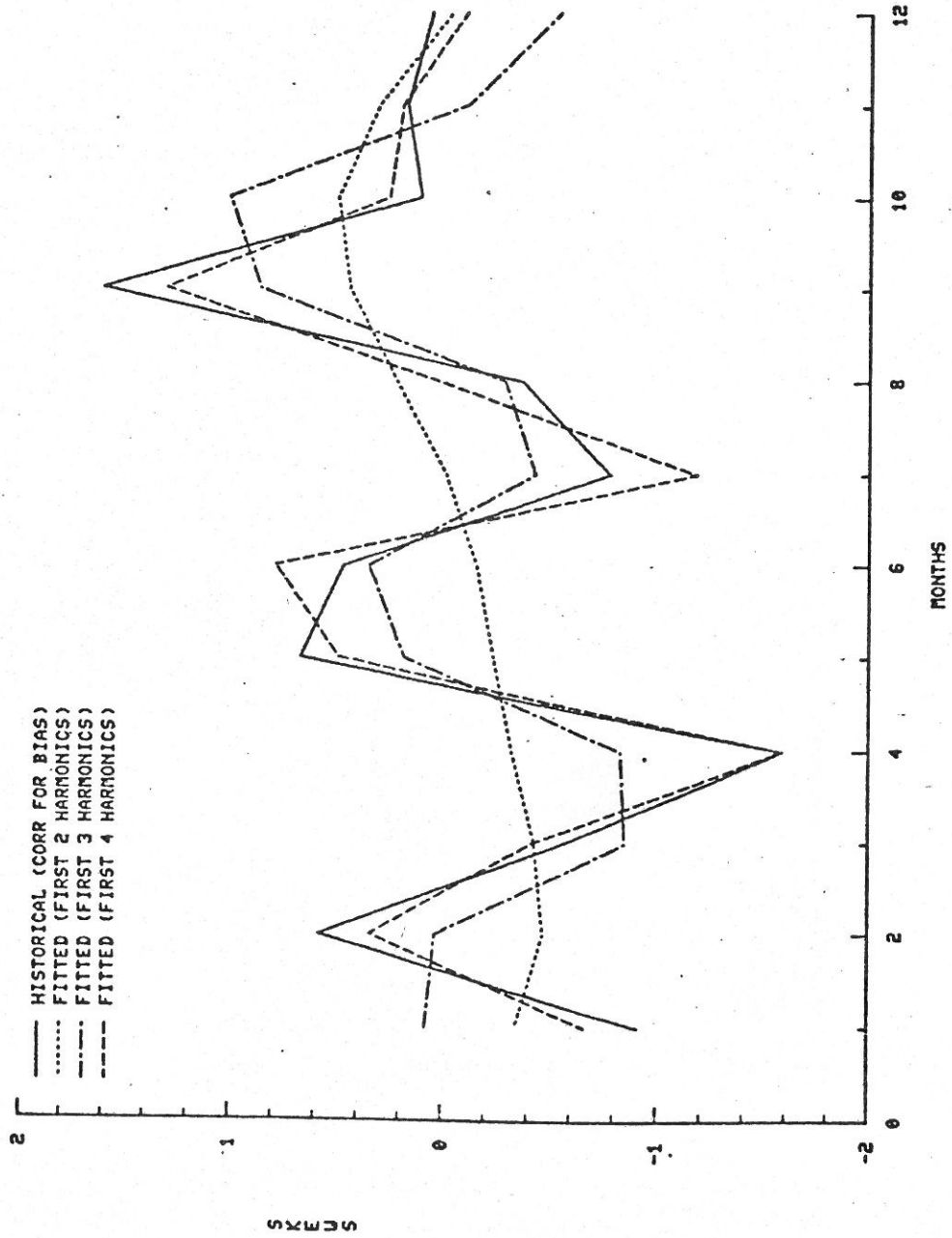


Figure 1.9.C.2 PASO DEL ERMITANO - SKEWS IN LOG DOMAIN

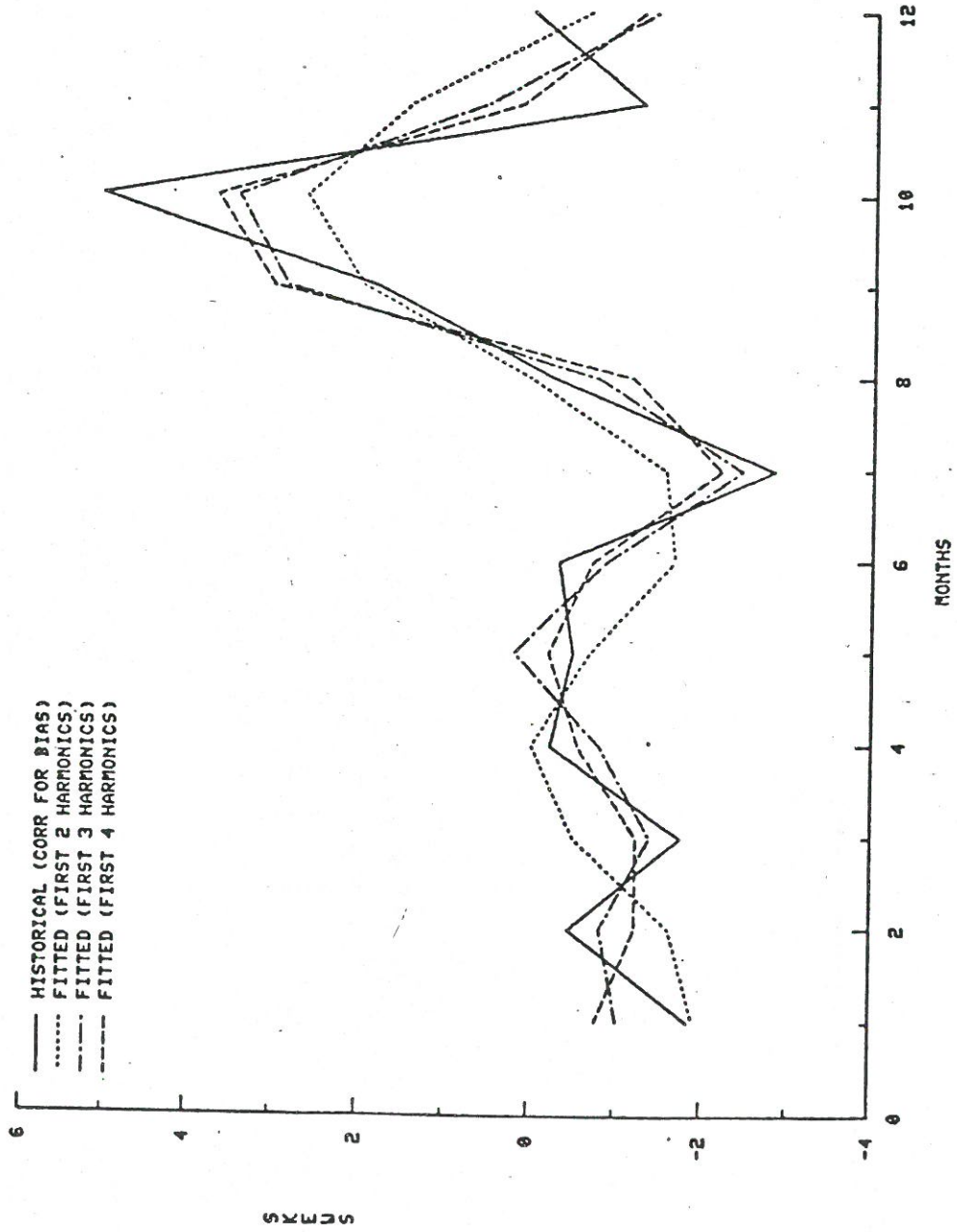


Figure 1.9.C.3 RANCHO ARRIBA - SKEWS IN LOG DOMAIN

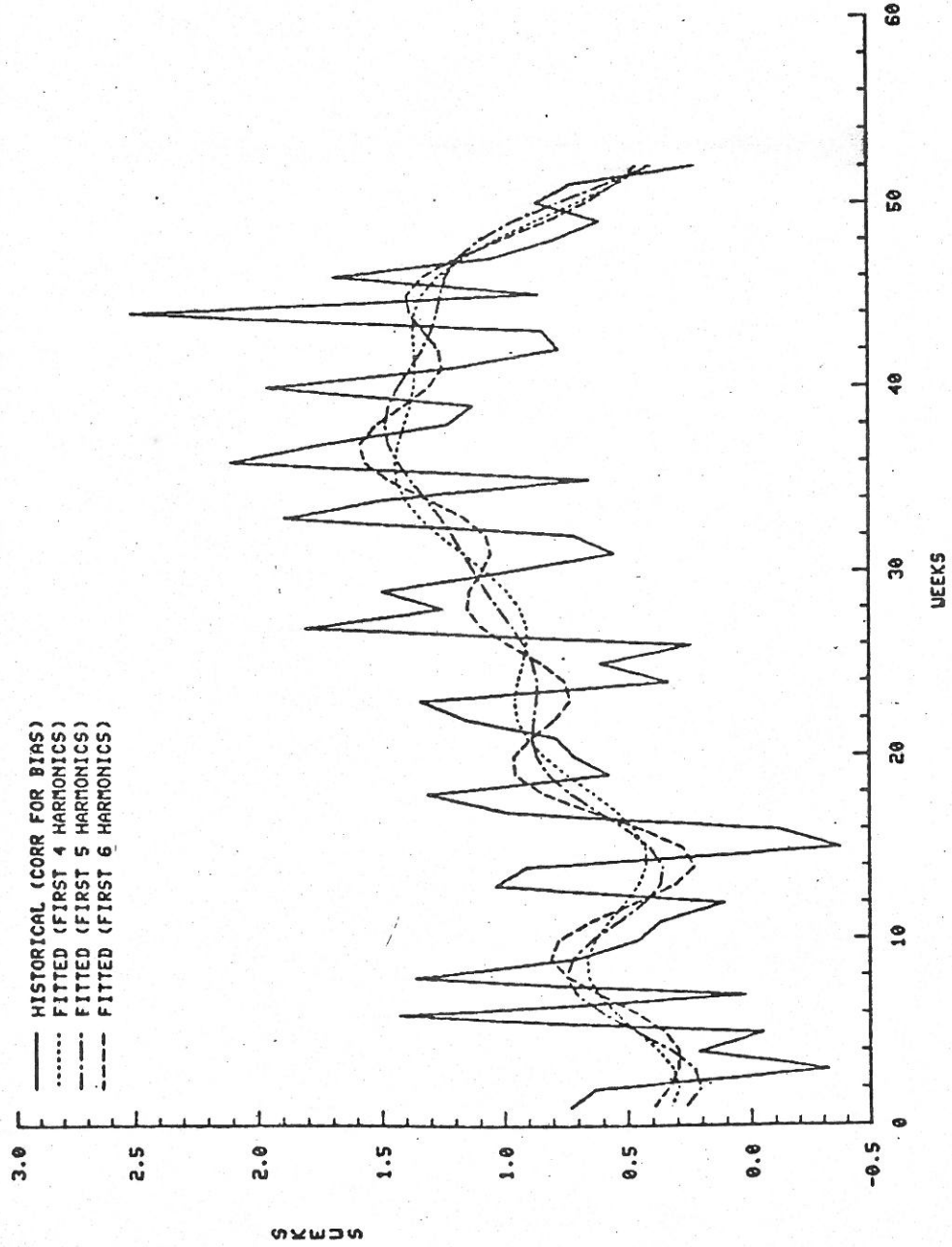


Figure 1.9.C.4 PALO DE CAJA - SKEWS IN LOG DOMAIN

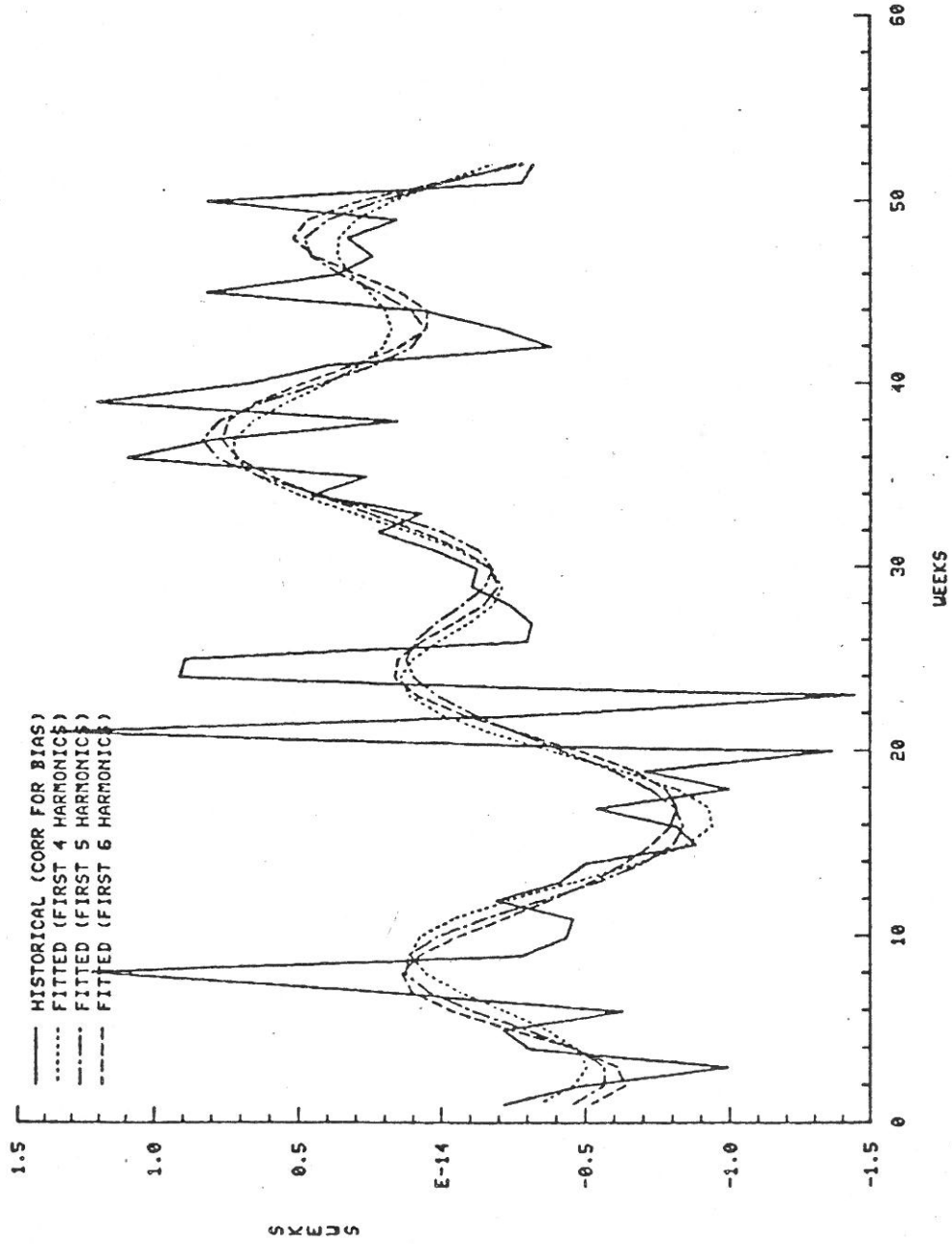


Figure 1.9.C.5 PASO DEL ERMITANO - SKEWS IN LOG DOMAIN

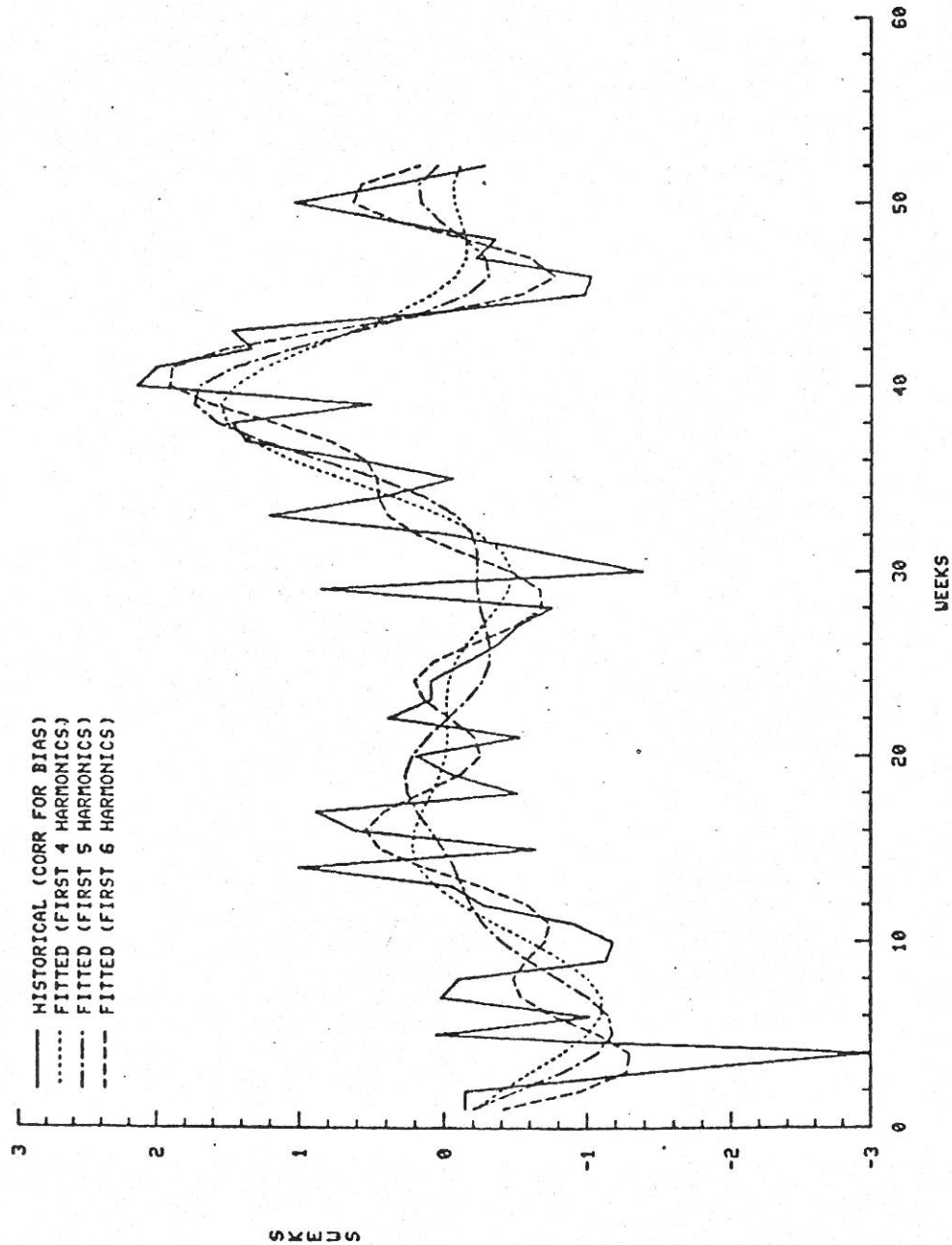


Figure 1.9.C.6 RANCHO ARRIBA - SKEWS IN LOG DOMAIN

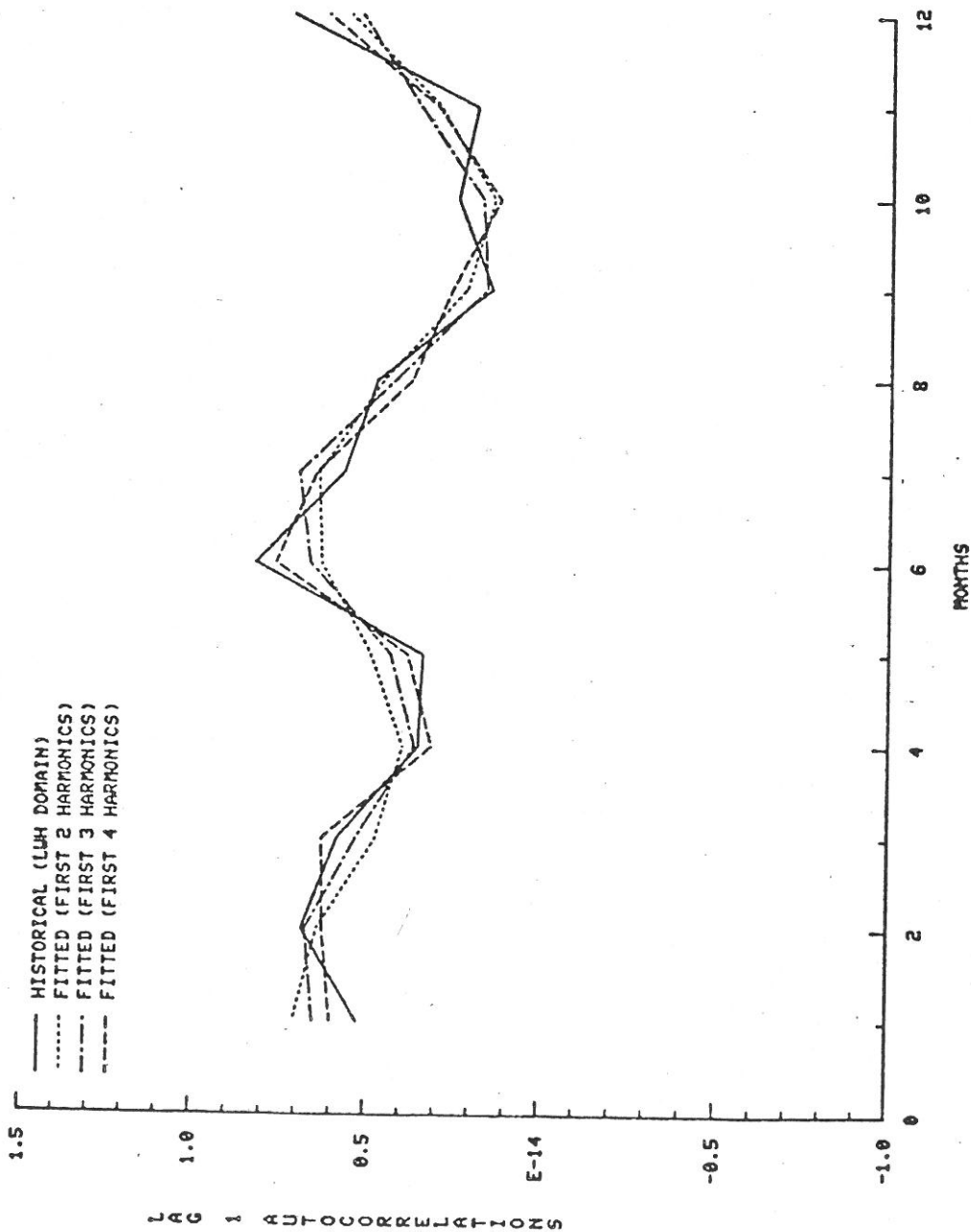


Figure 1.9.C.7 PALO DE CAJA - LAG 1 AUTOCORRELATIONS IN LWH DOMAIN

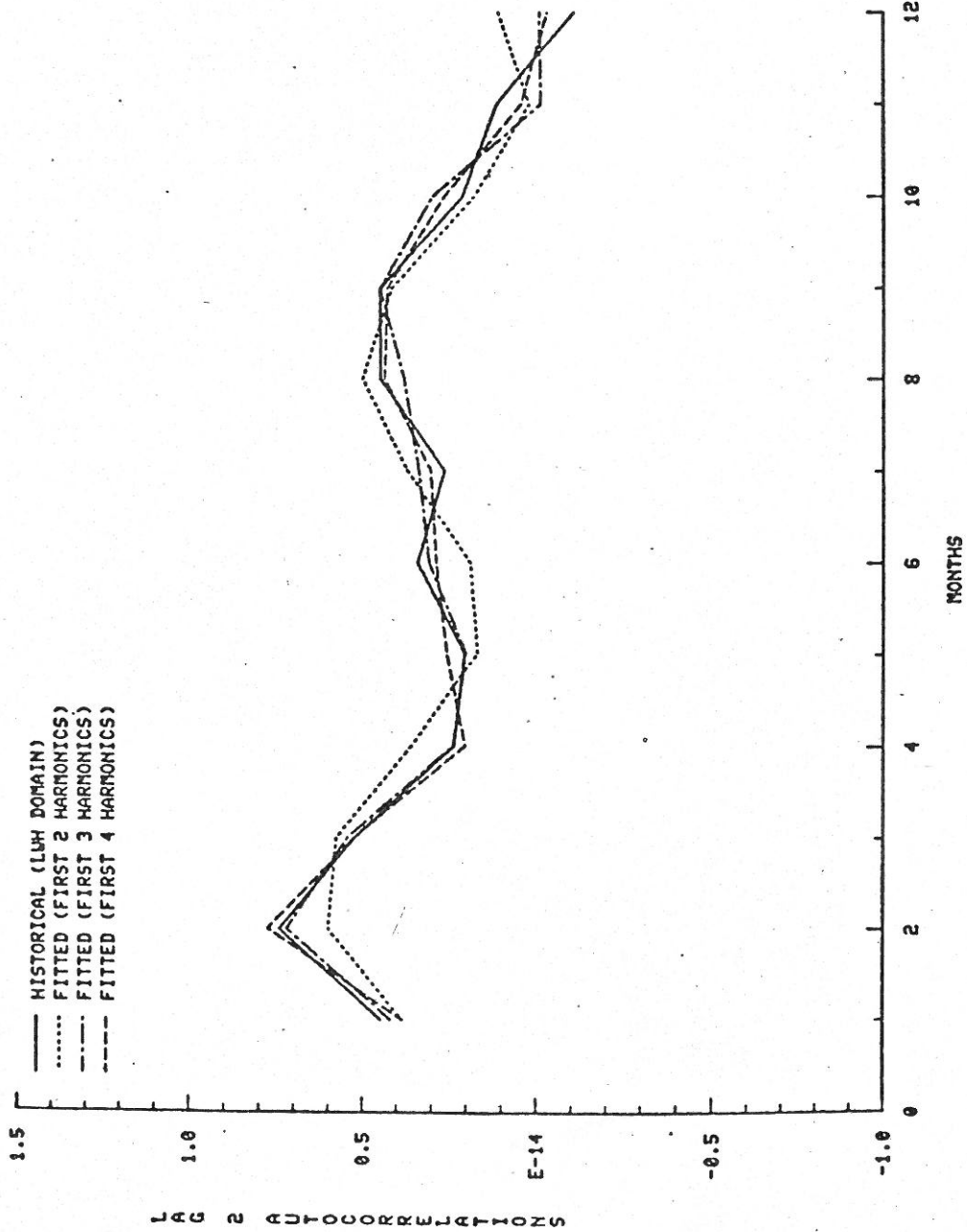


Figure 1.9.C.8 PALO DE CAJA - LAG 2 AUTOCORRELATIONS IN LWH DOMAIN

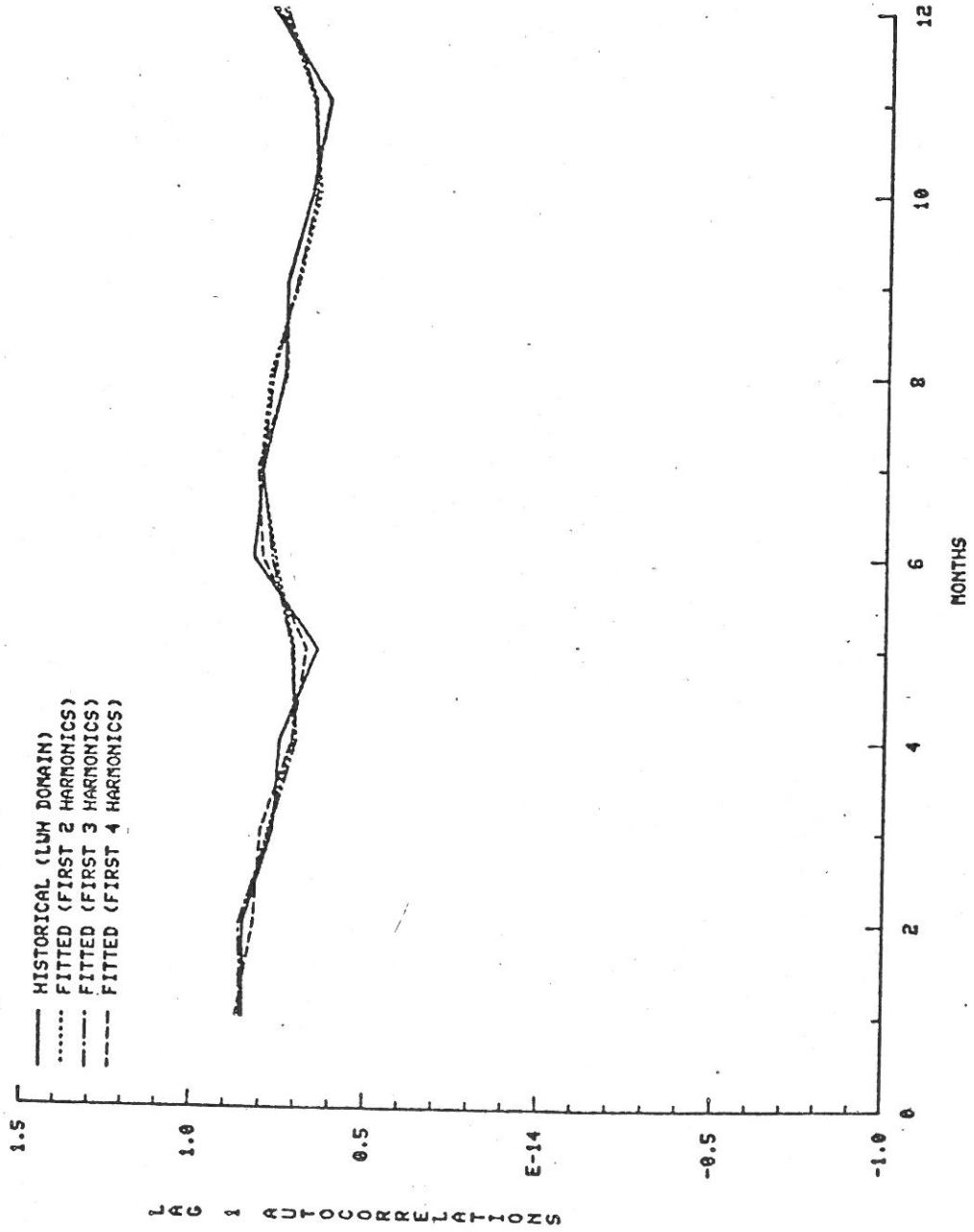


Figure 1.9.C.9 PASO DEL ERMITANO - LAG 1 AUTOCORRELATIONS IN LUH DOMAIN.

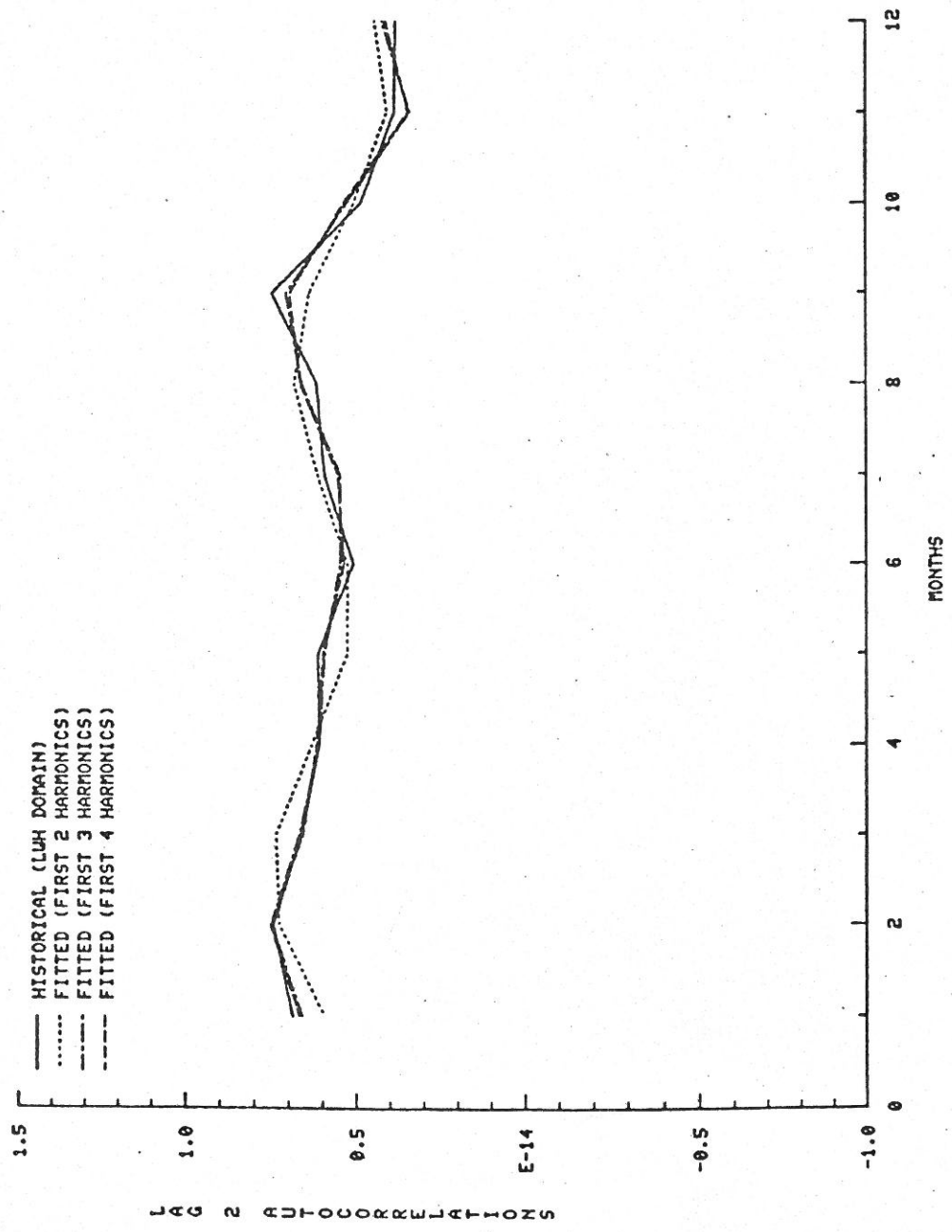


Figure 1.9.C.10 PASO DEL ERMITANO - LAG 2 AUTOCORRELATIONS IN LUH DOMAIN

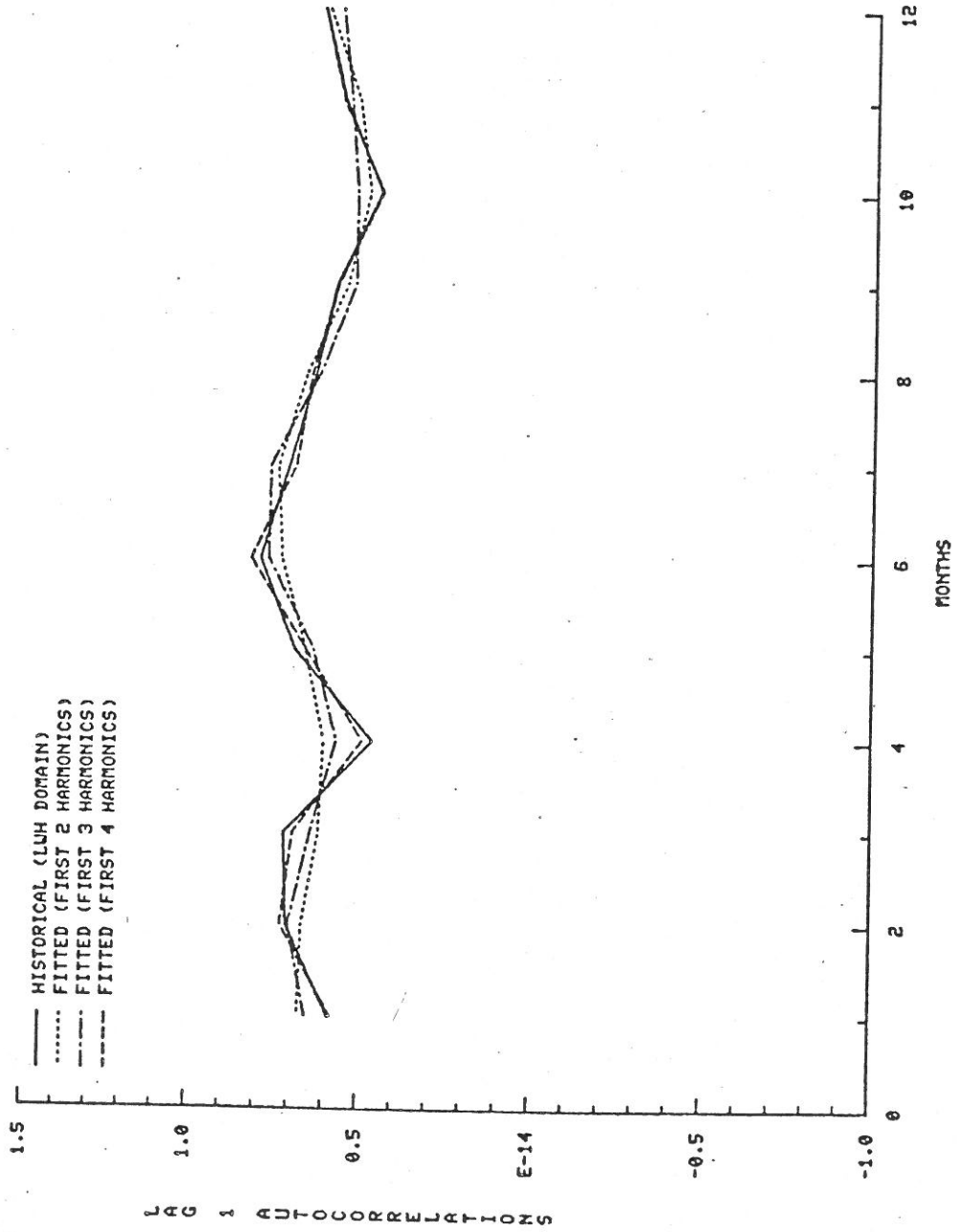


Figure 1.9.C.11 RANCHO ARRIBA - LAG 1 AUTOCORRELATIONS IN LUH DOMAIN

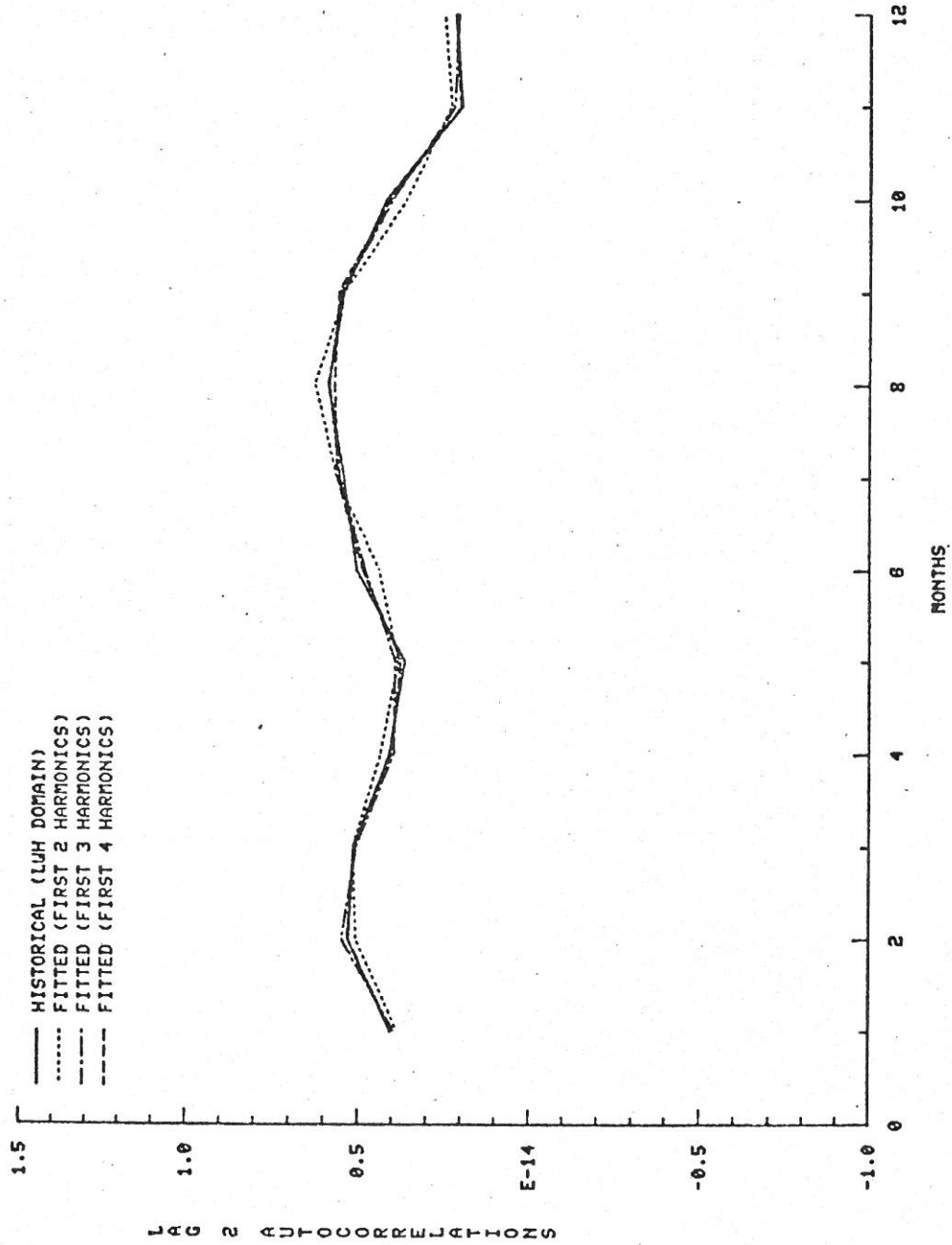


Figure 1.9.C.12 RANCHO ARRIBA - LAG 2 AUTOCORRELATIONS IN LWH DOMAIN

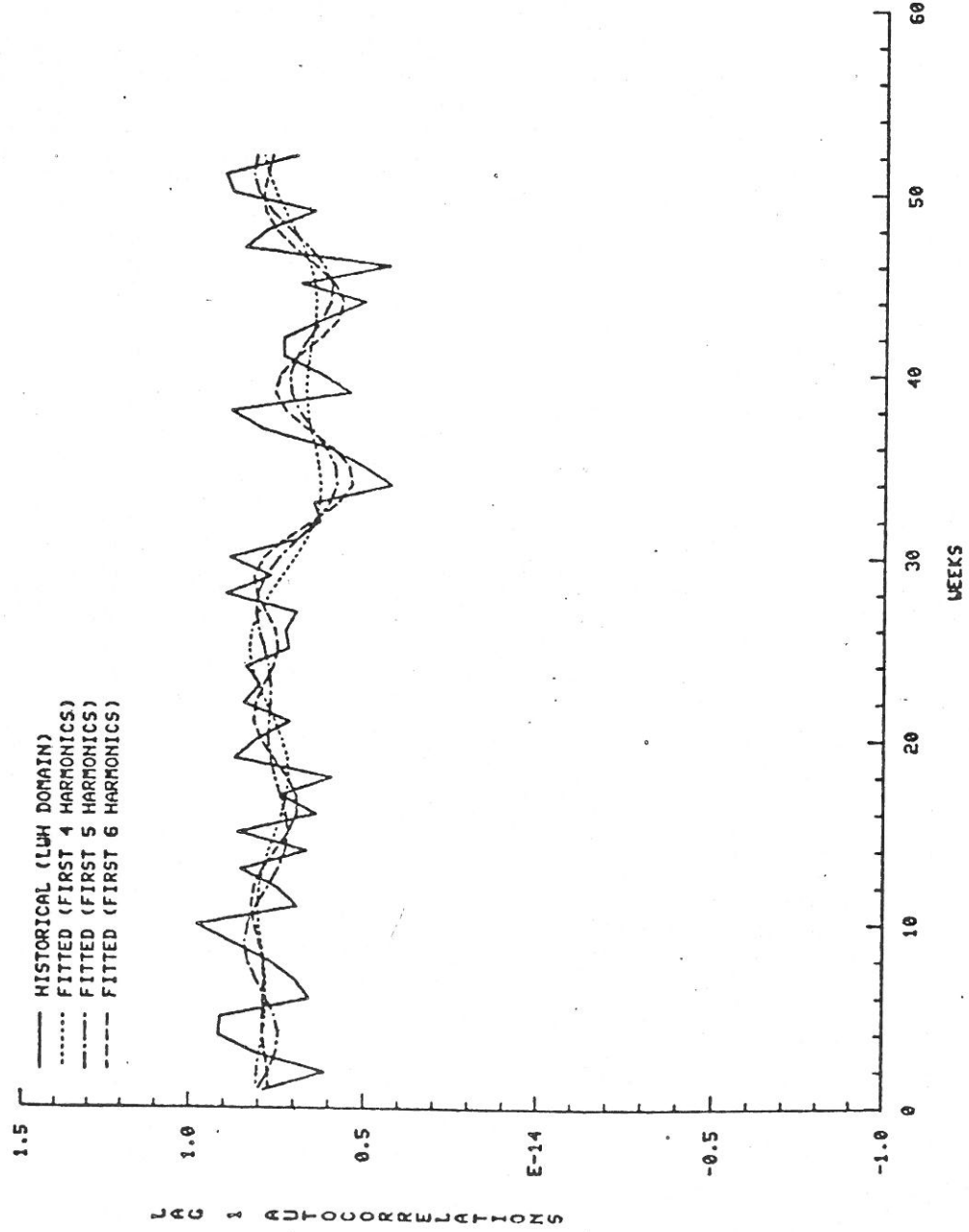


Figure 1.9.C.13 PALO DE CAJA - LAG 1 AUTOCORRELATIONS IN LUH DOMAIN

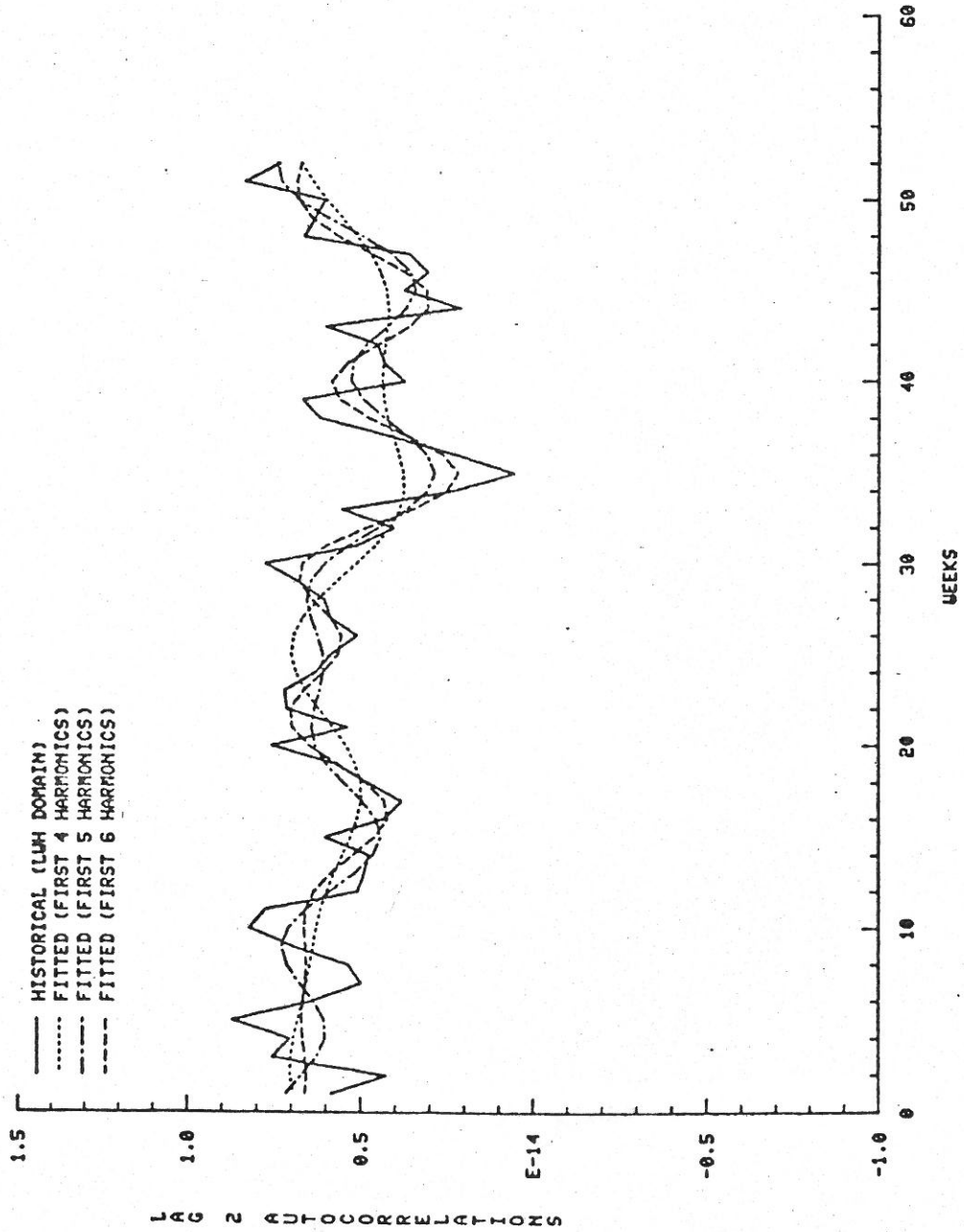


Figure 1.9.C.14 PALO DE CAJA - LAG 2 AUTOCORRELATIONS IN LUH DOMAIN

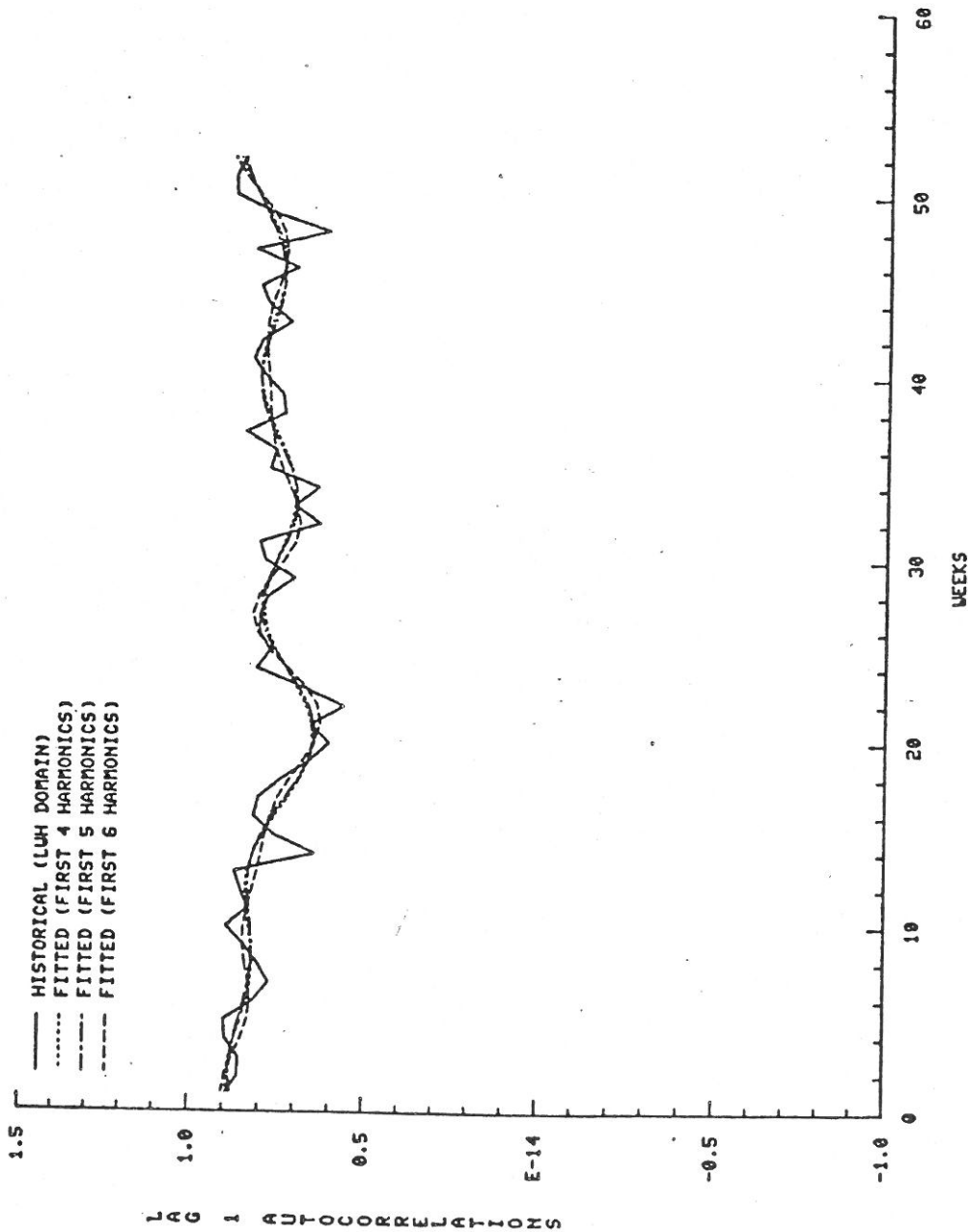


Figure 1.9.C.15 PASO DEL ERMITANO - LAG 1 AUTOCORRELATIONS IN LUH DOMAIN

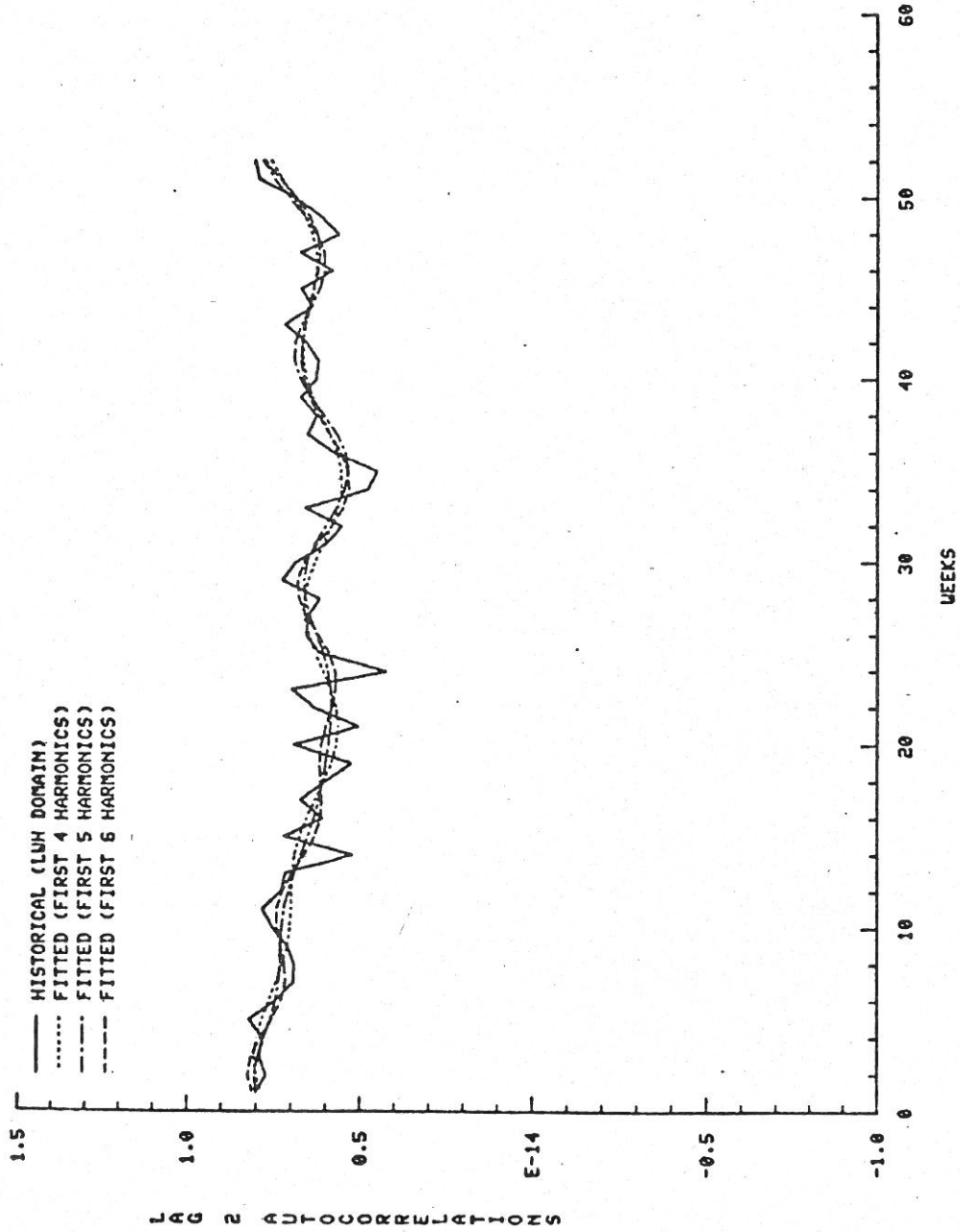


Figure 1.9.C.16 PASO DEL ERMITANO - LAG 2 AUTOCORRELATIONS IN LUH DOMAIN

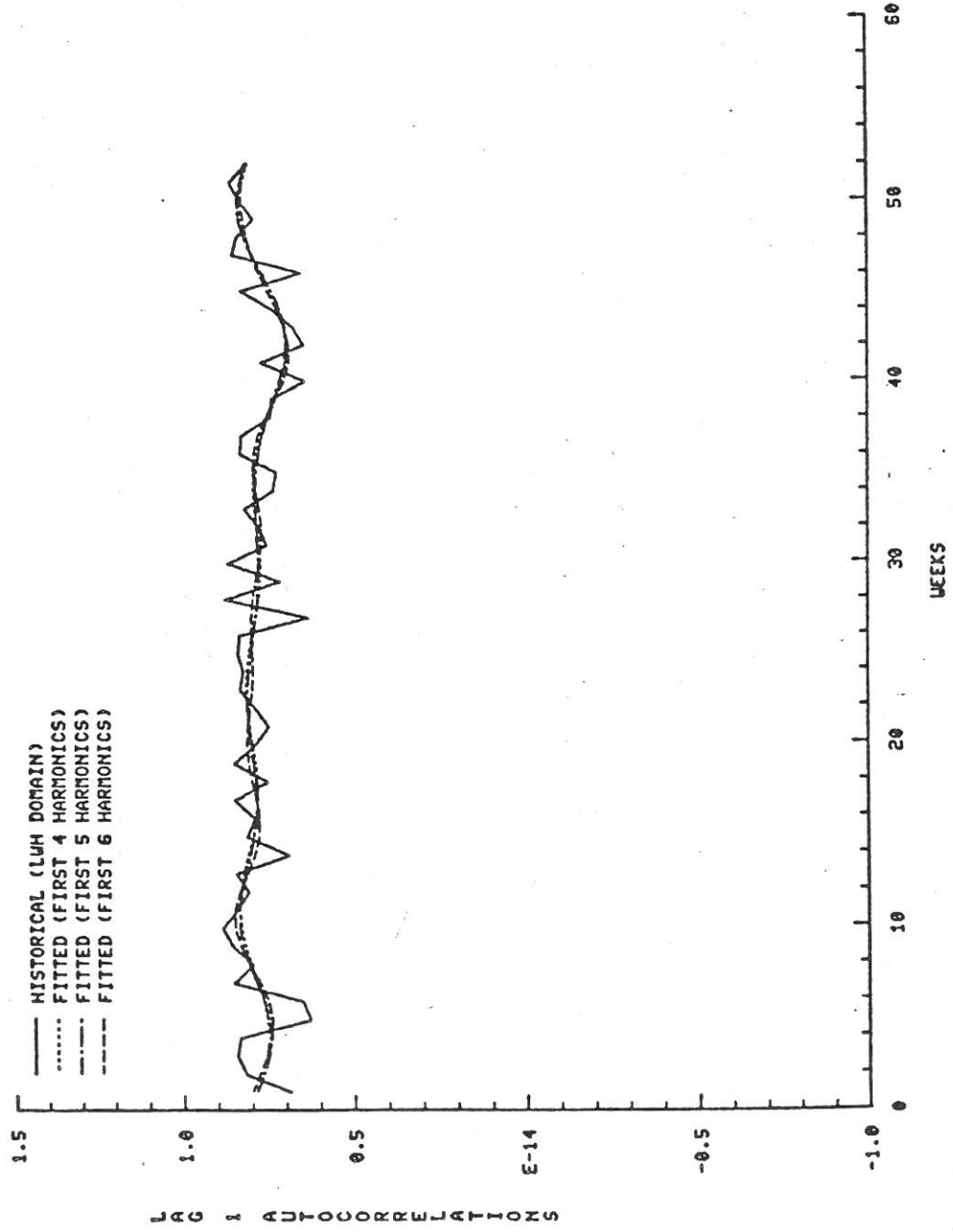


Figure 1.9.C.17 RANCHO ARRIBA - LAG 1 AUTOCORRELATIONS IN LUH DOMAIN

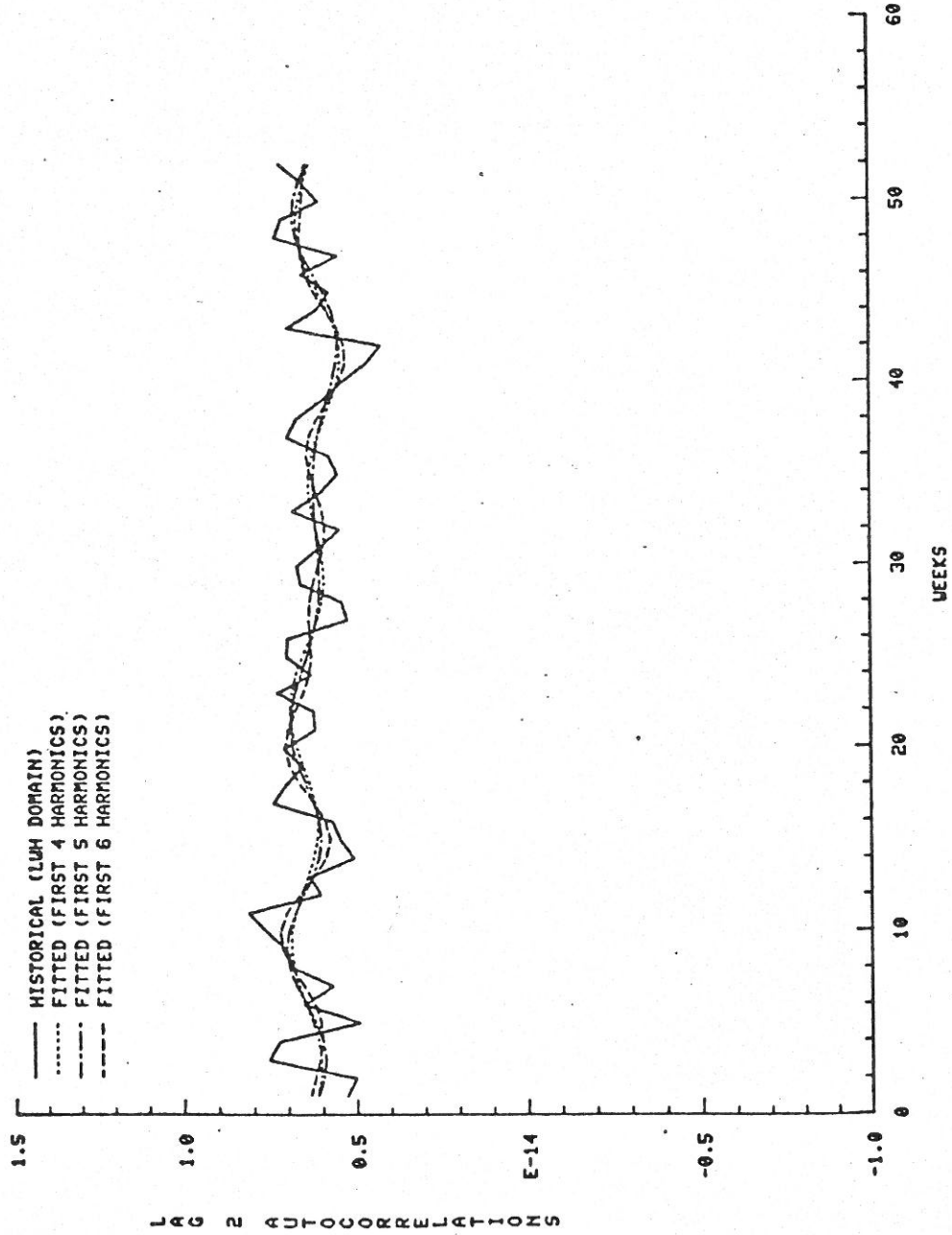


Figure 1.9.C.18 RANCHO ARRIBA - LAG 2 AUTOCORRELATIONS IN LUH DOMAIN

APPENDIX 1.9.D

HISTORICAL (EXTENDED SERIES) AND GENERATED MONTHLY AND WEEKLY
STATISTICS FOR PALO DE CAJA (PALODE), PASO DEL ERMITAÑO (PASODE)
AND RANCHO ARRIBA (RANCHO) USING MODELS A, B AND C IN THE
ORIGINAL DOMAIN OF FLOWS

Note: See Appendix 1.9.E for details of computing generated statistics.

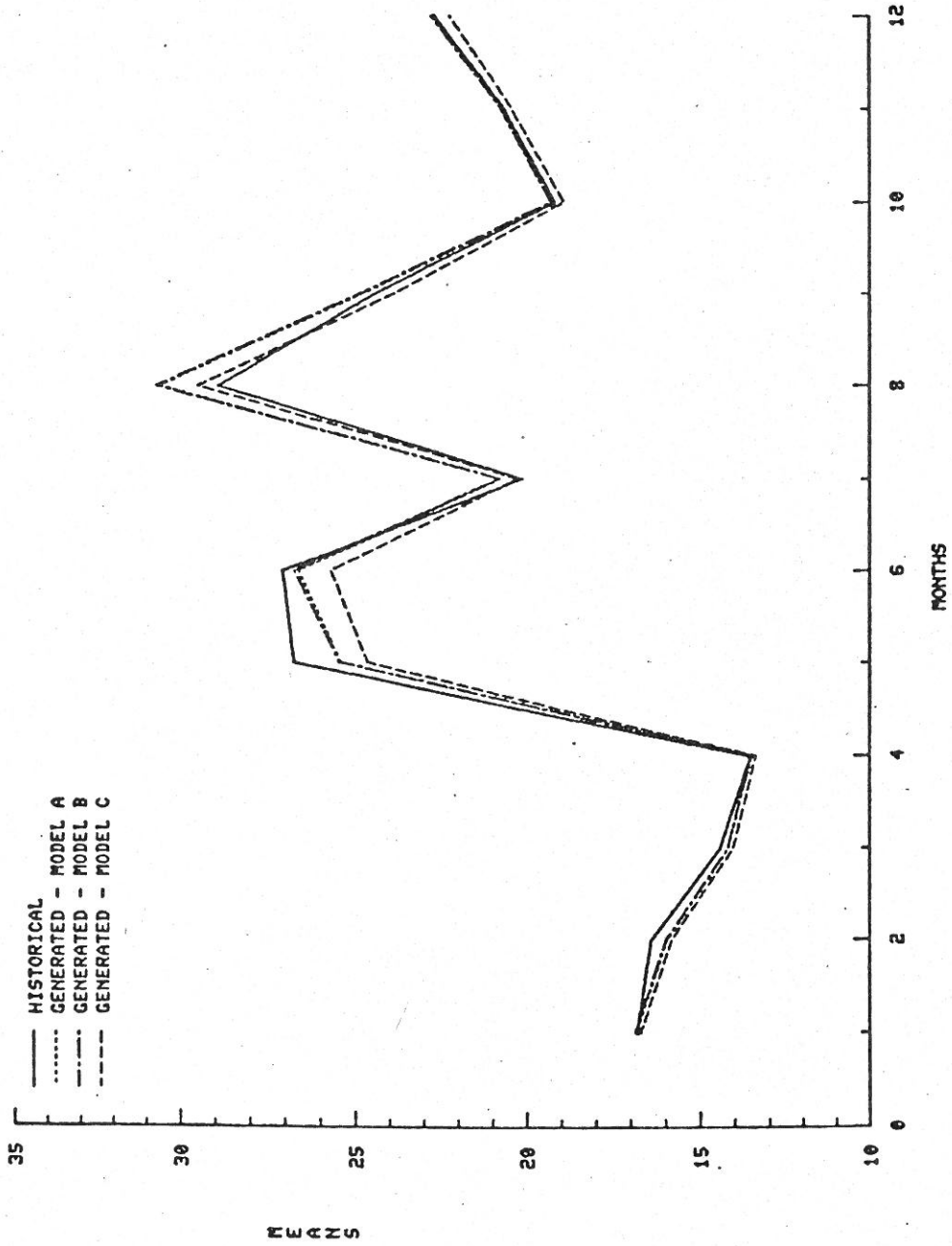


Figure 1.9.D.1 PASODE - MEANS (ORIGINAL DOMAIN)

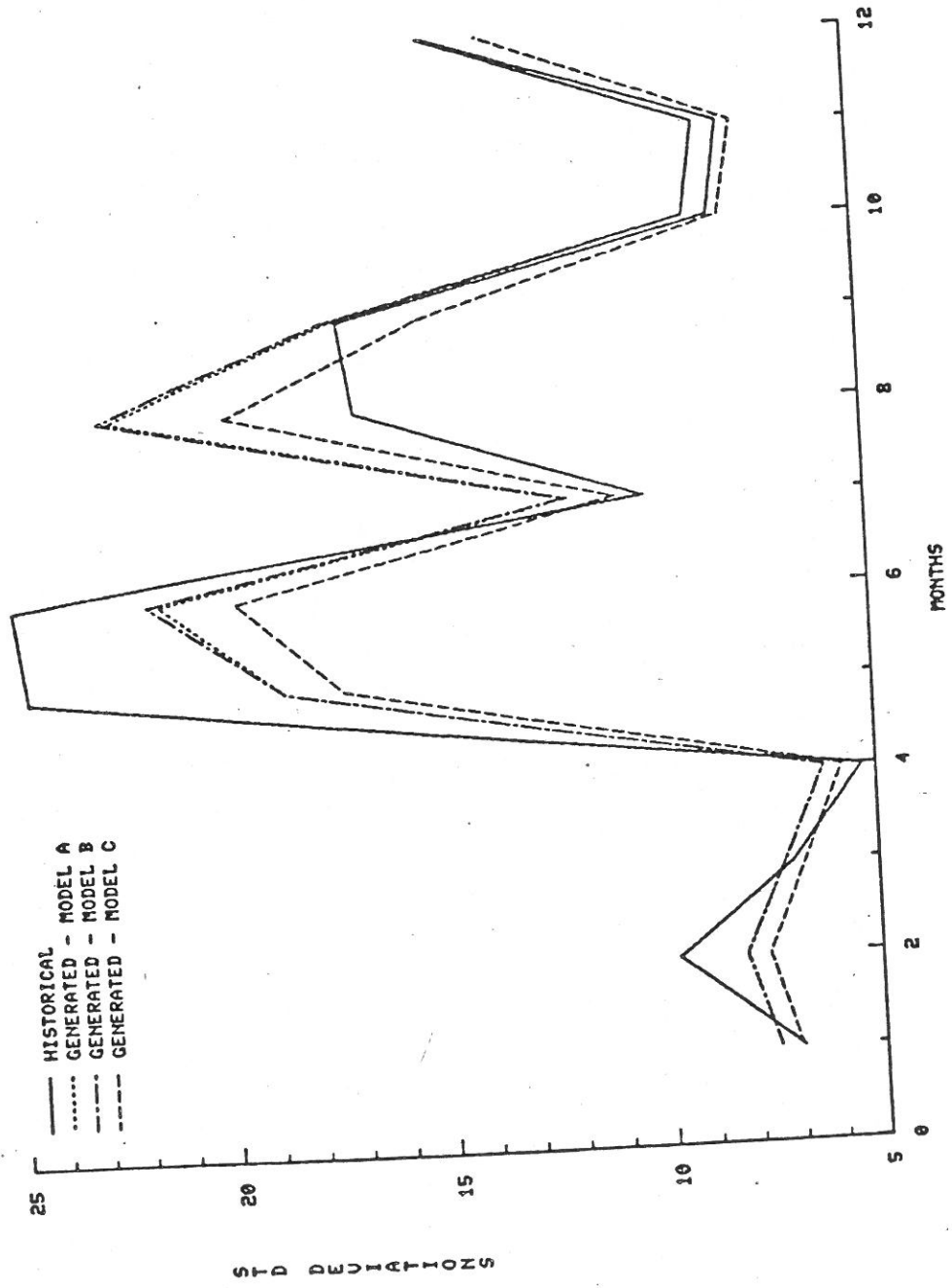


Figure 1.9.D.2 PASODE - STD DEVIATIONS (ORIGINAL DOMAIN)

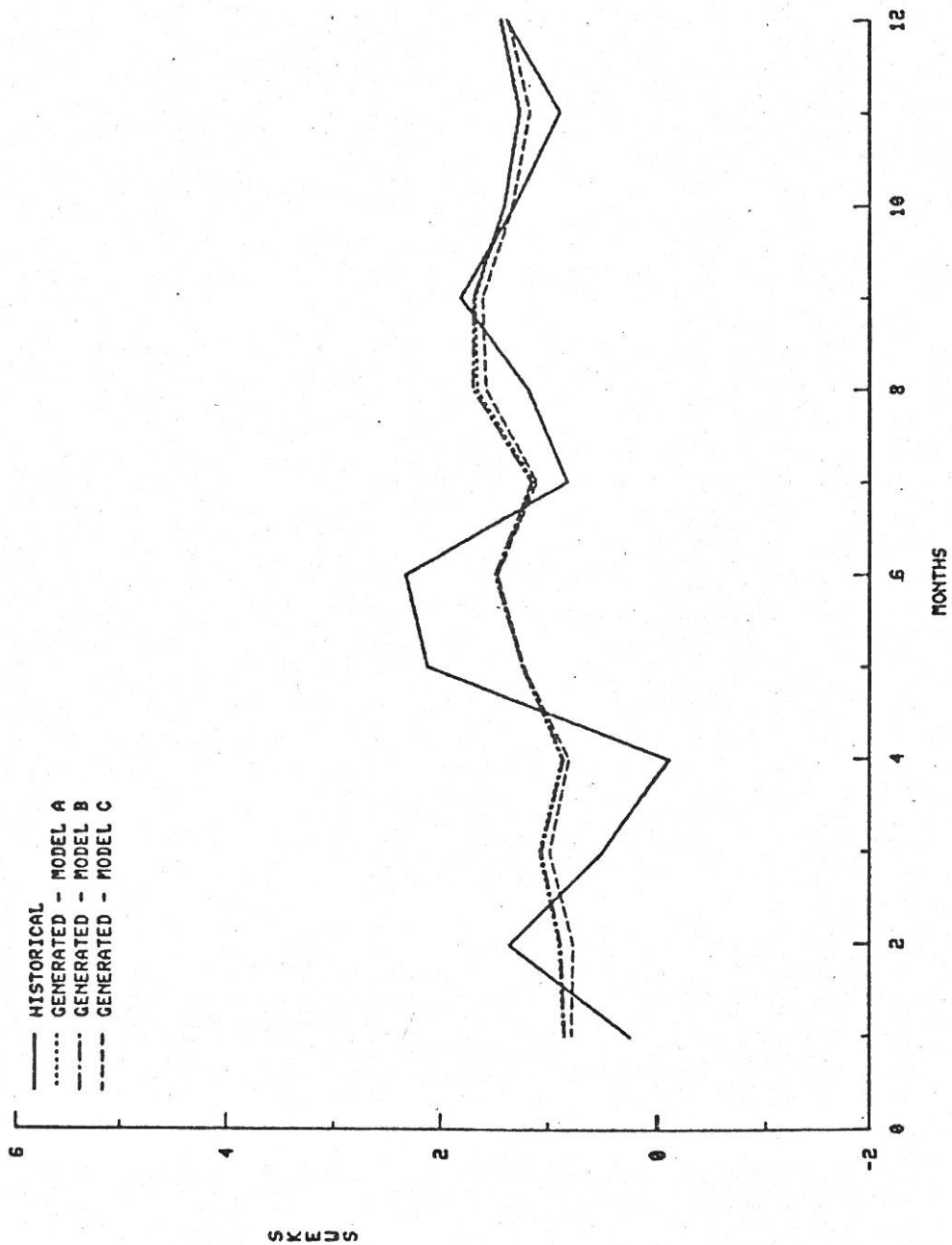


Figure 1.9.D.3 PASODE - SKEWS (ORIGINAL DOMAIN)

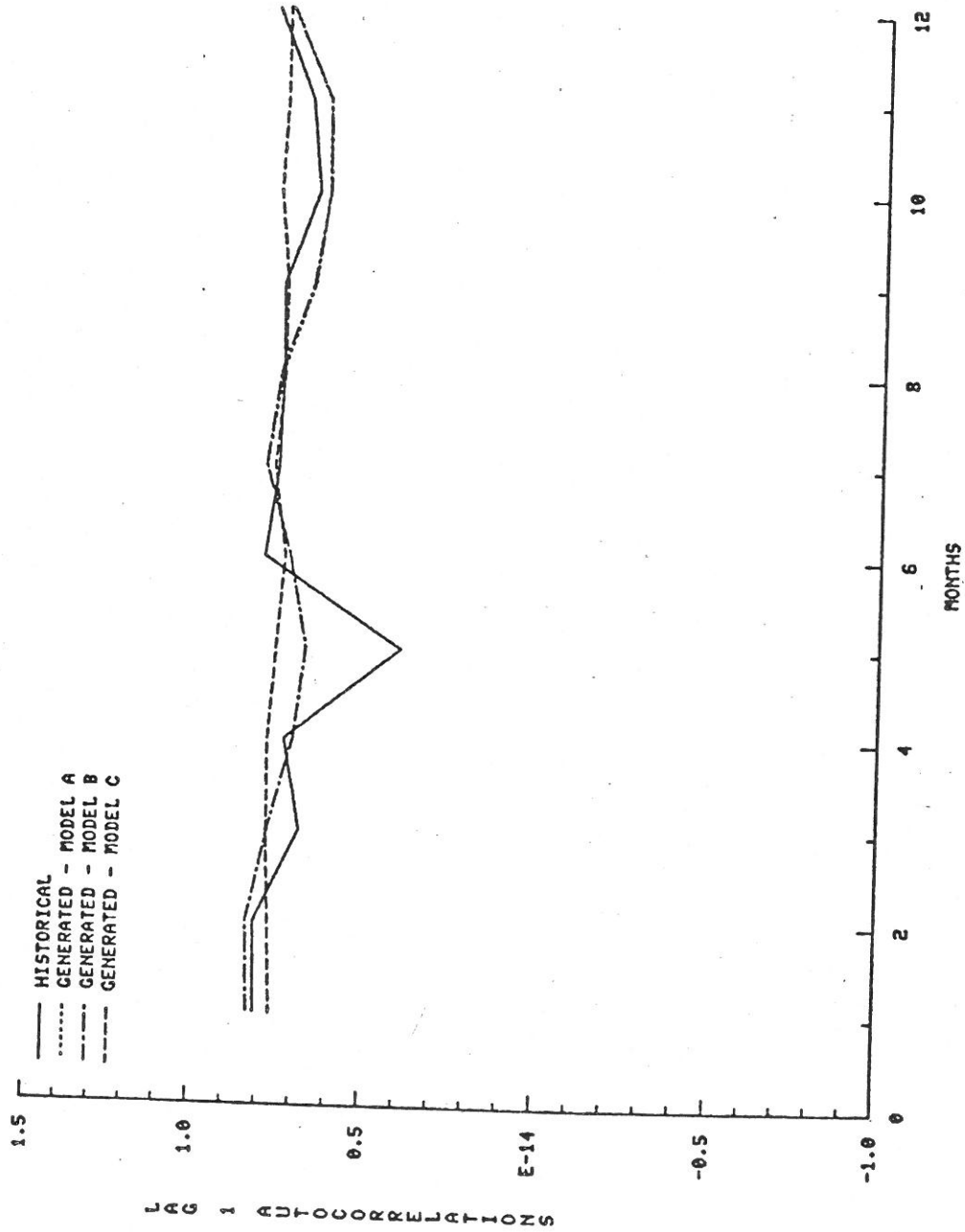


Figure 1.9.D.4 PASODE - LAG 1 AUTOCORRELATIONS (ORIGINAL DOMAIN)

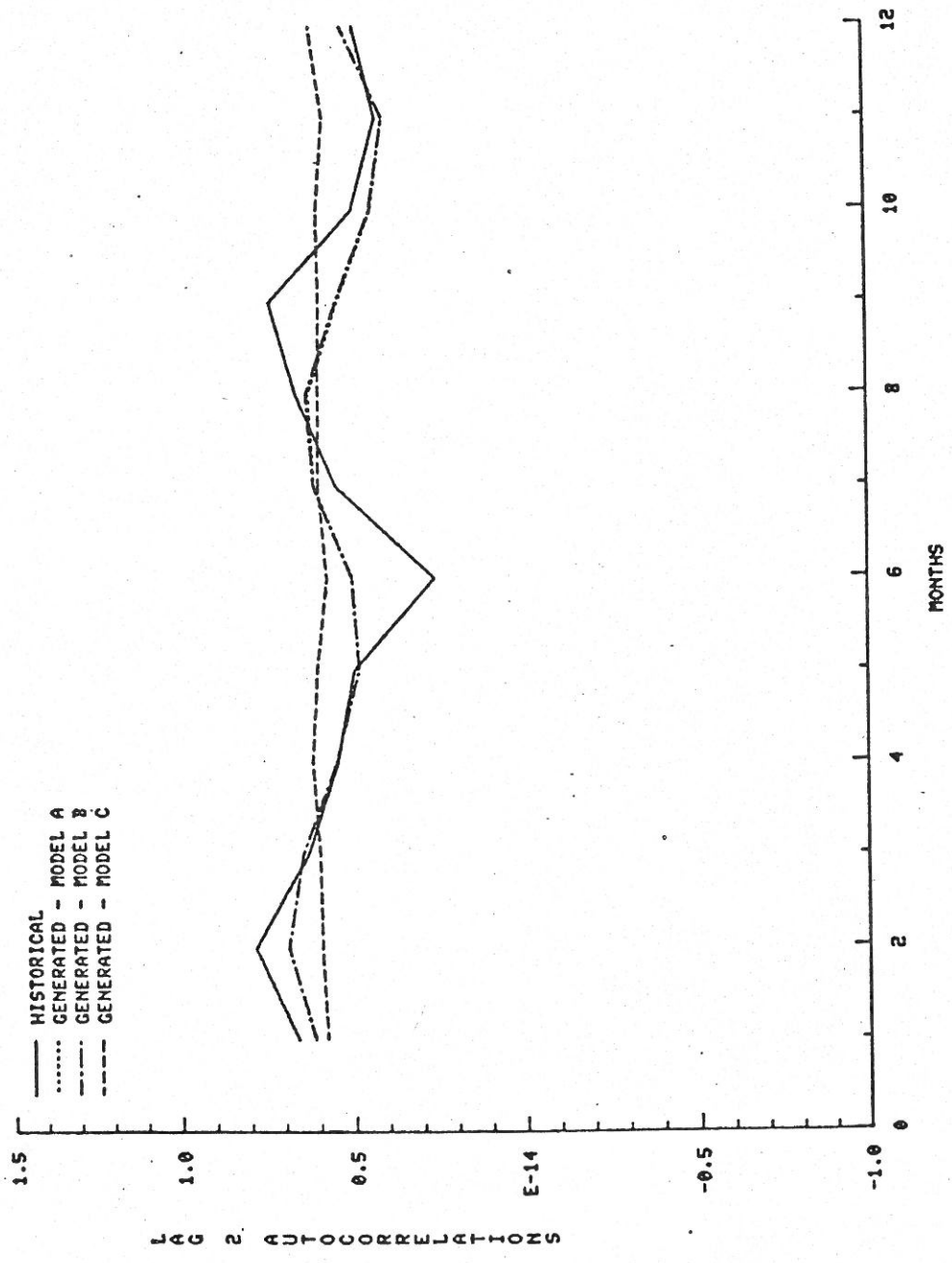


Figure 1.9.D.5 PASODE - LAG 2 AUTOCORRELATIONS (ORIGINAL DOMAIN)

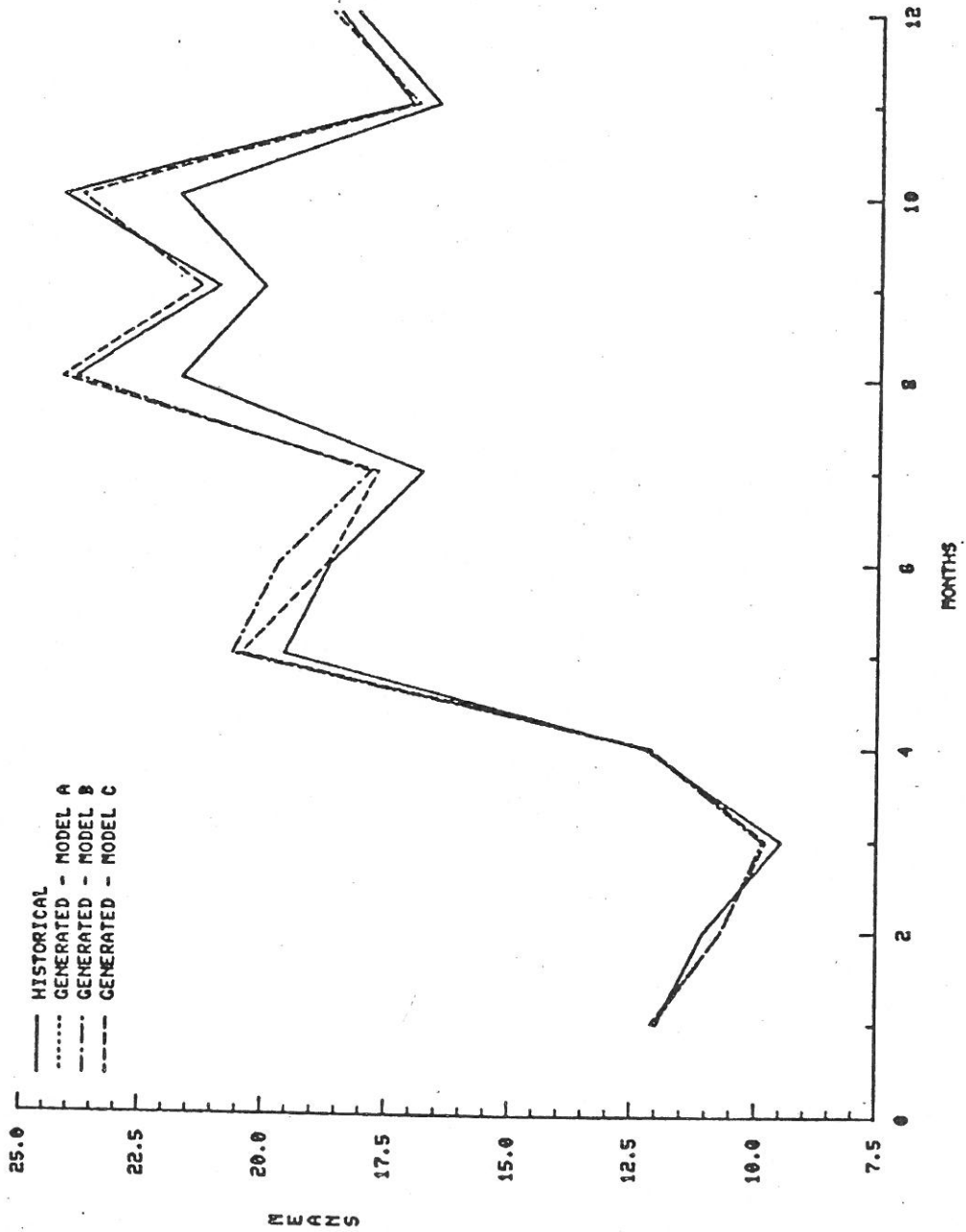


Figure 1.9.D.6 PALODE - MEANS (ORIGINAL DOMAIN)

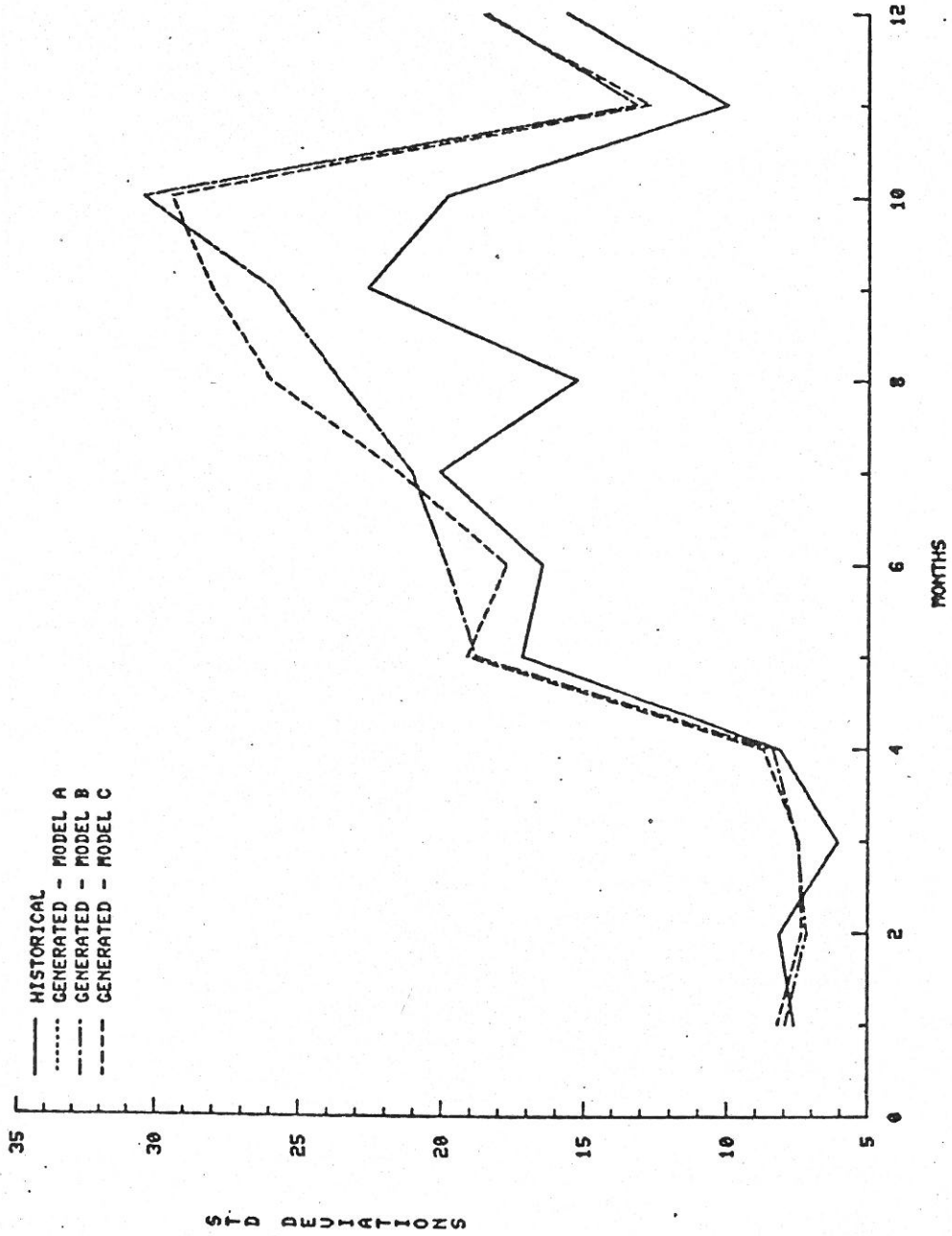


Figure 1.9.D.7 PALODE - STD DEVIATIONS (ORIGINAL DOMAIN)

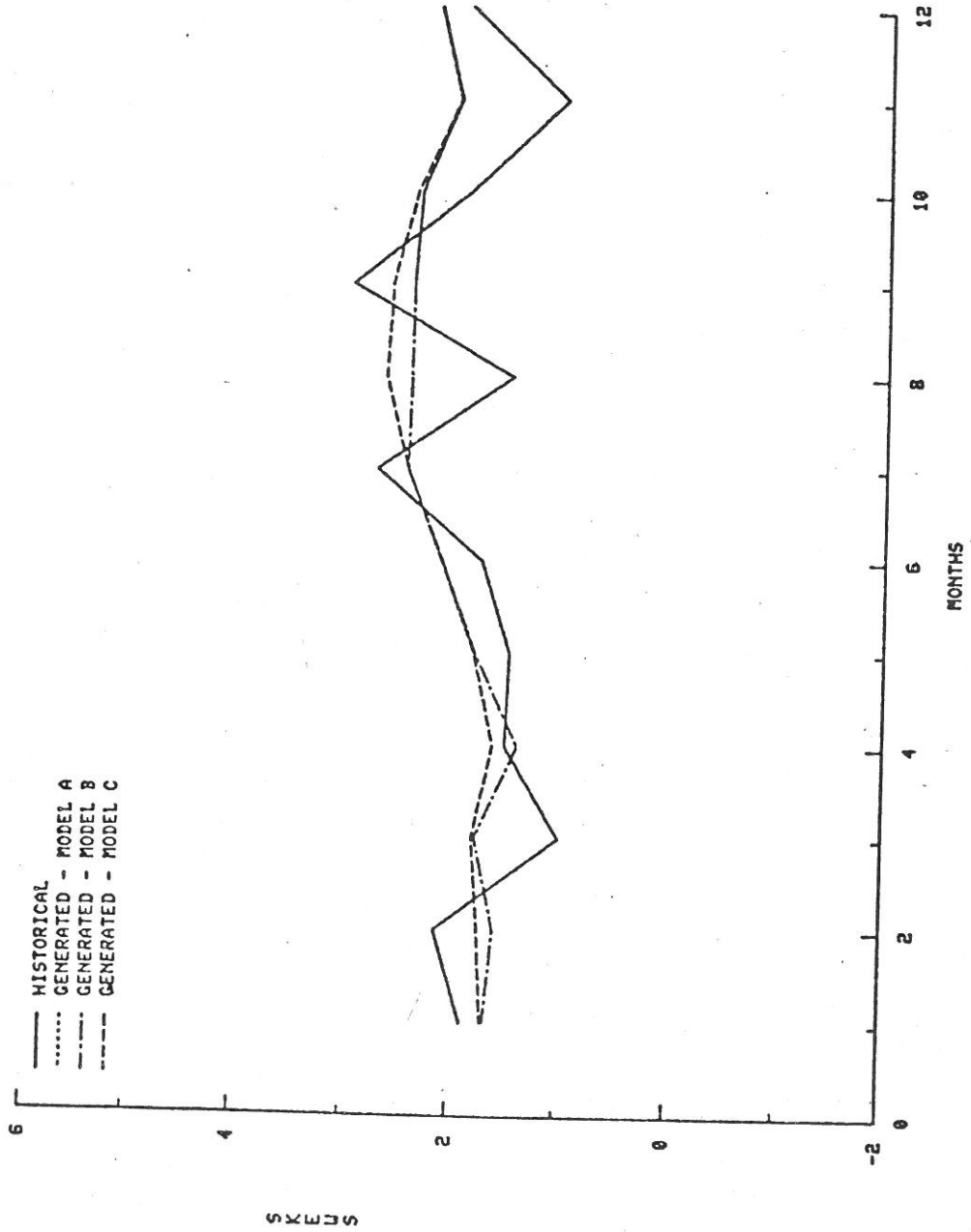


Figure 1.9.D.8 PALODE - SKEWS (ORIGINAL DOMAIN)

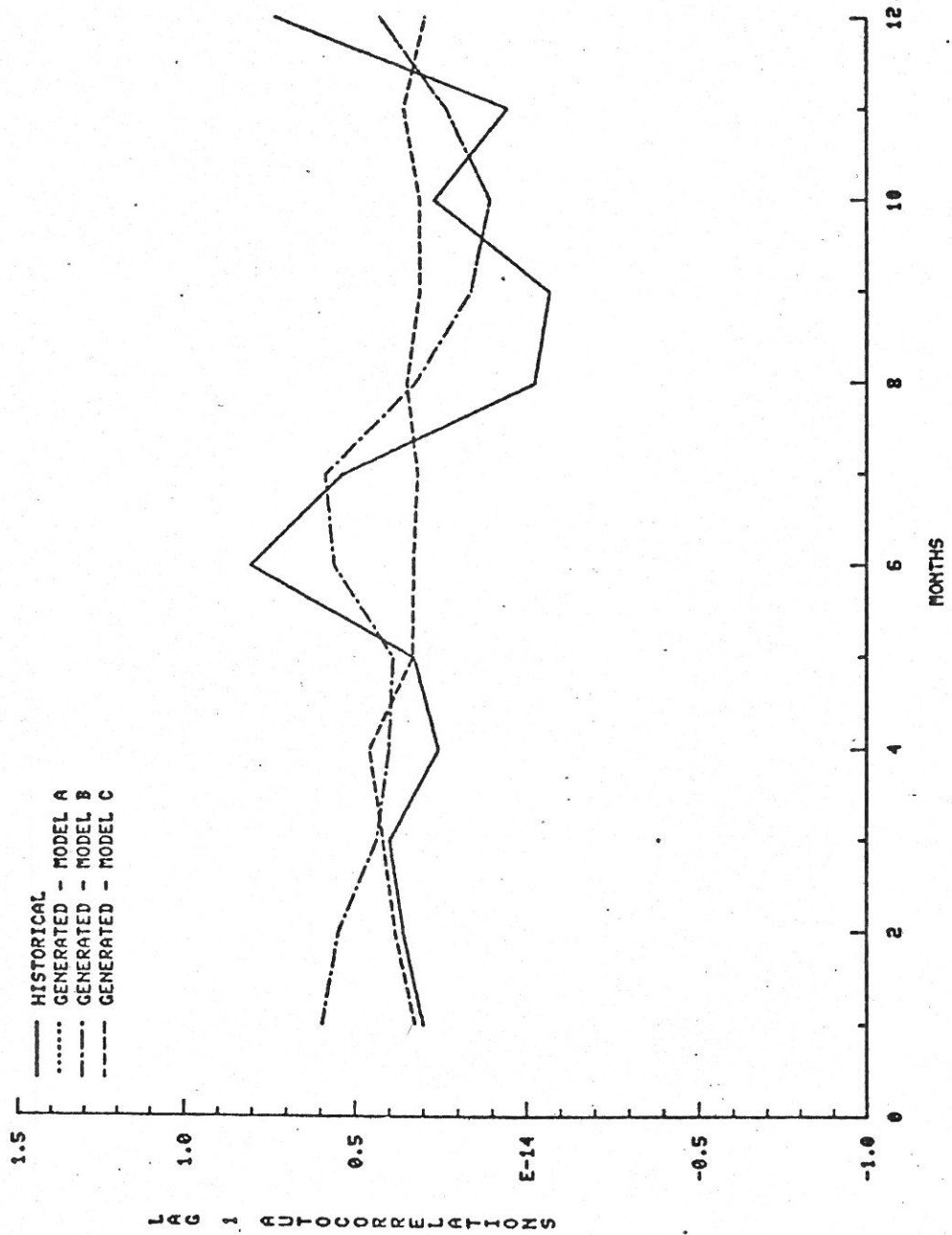


Figure I.9.D.9 PALODE - LAG 1 AUTOCORRELATIONS (ORIGINAL DOMAIN)

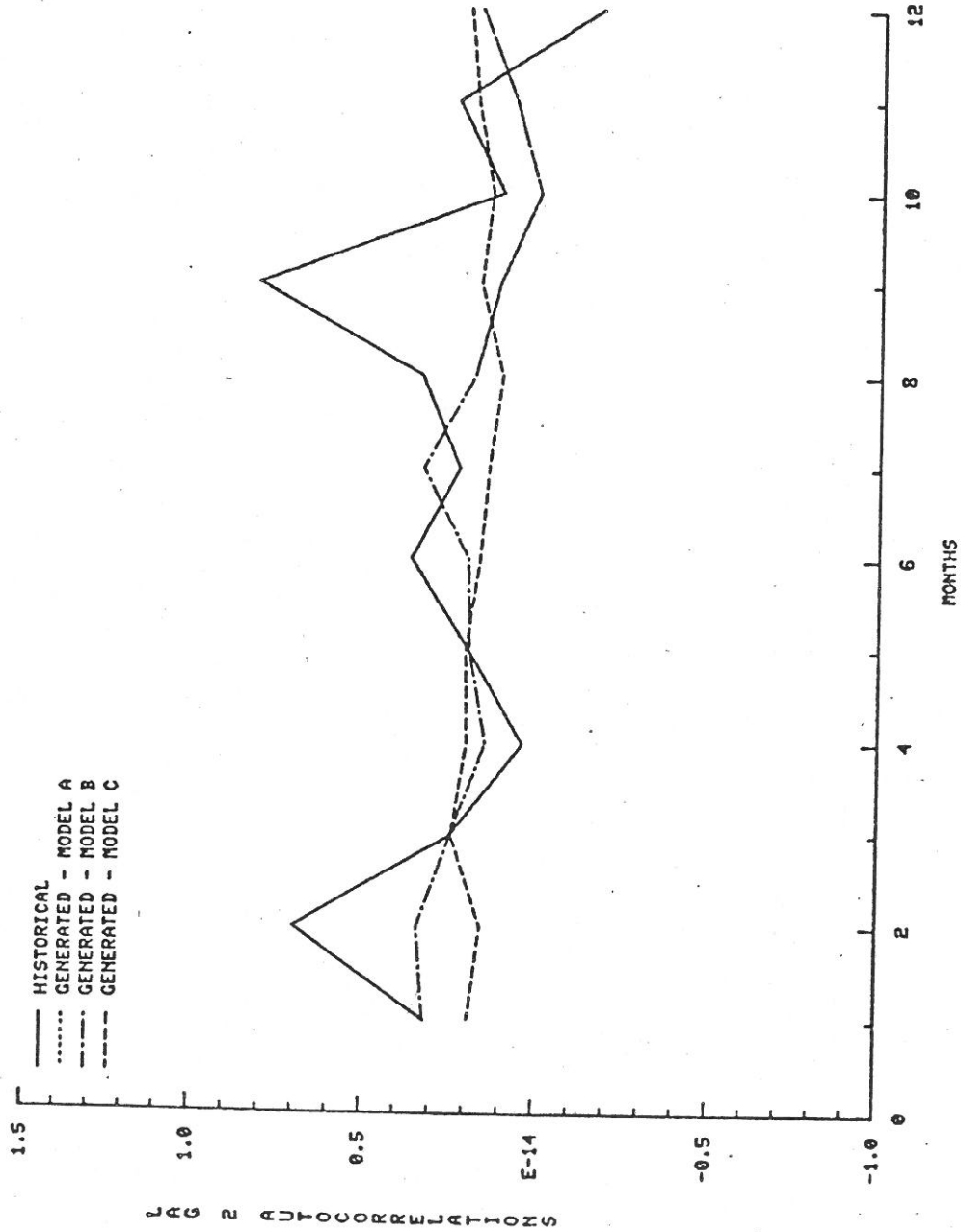


Figure 1.9.D.10 PALODE - LAG 2 AUTOCORRELATIONS (ORIGINAL DOMAIN)

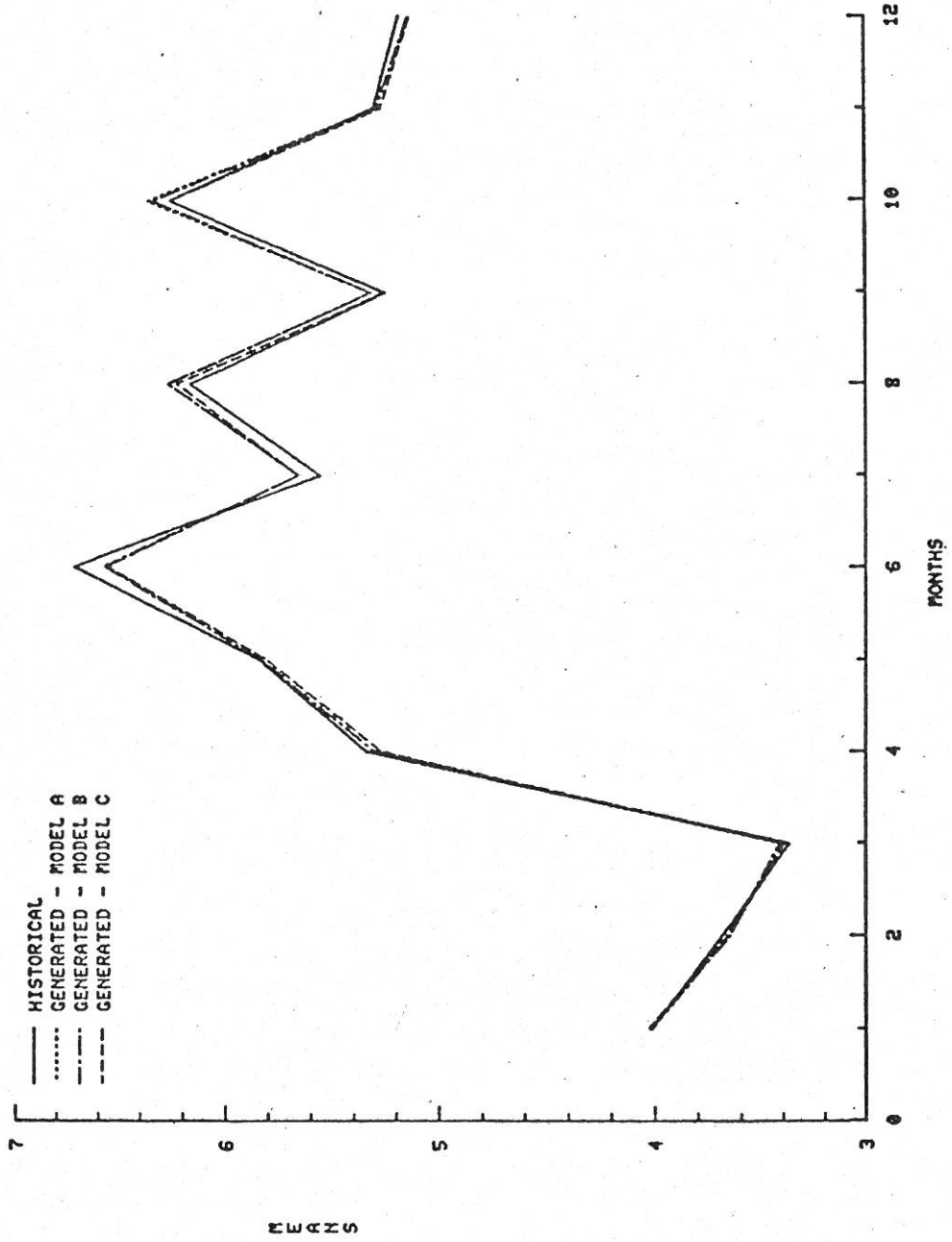


Figure 1.9.D.11 RANCHO - MEANS (ORIGINAL DOMAIN)

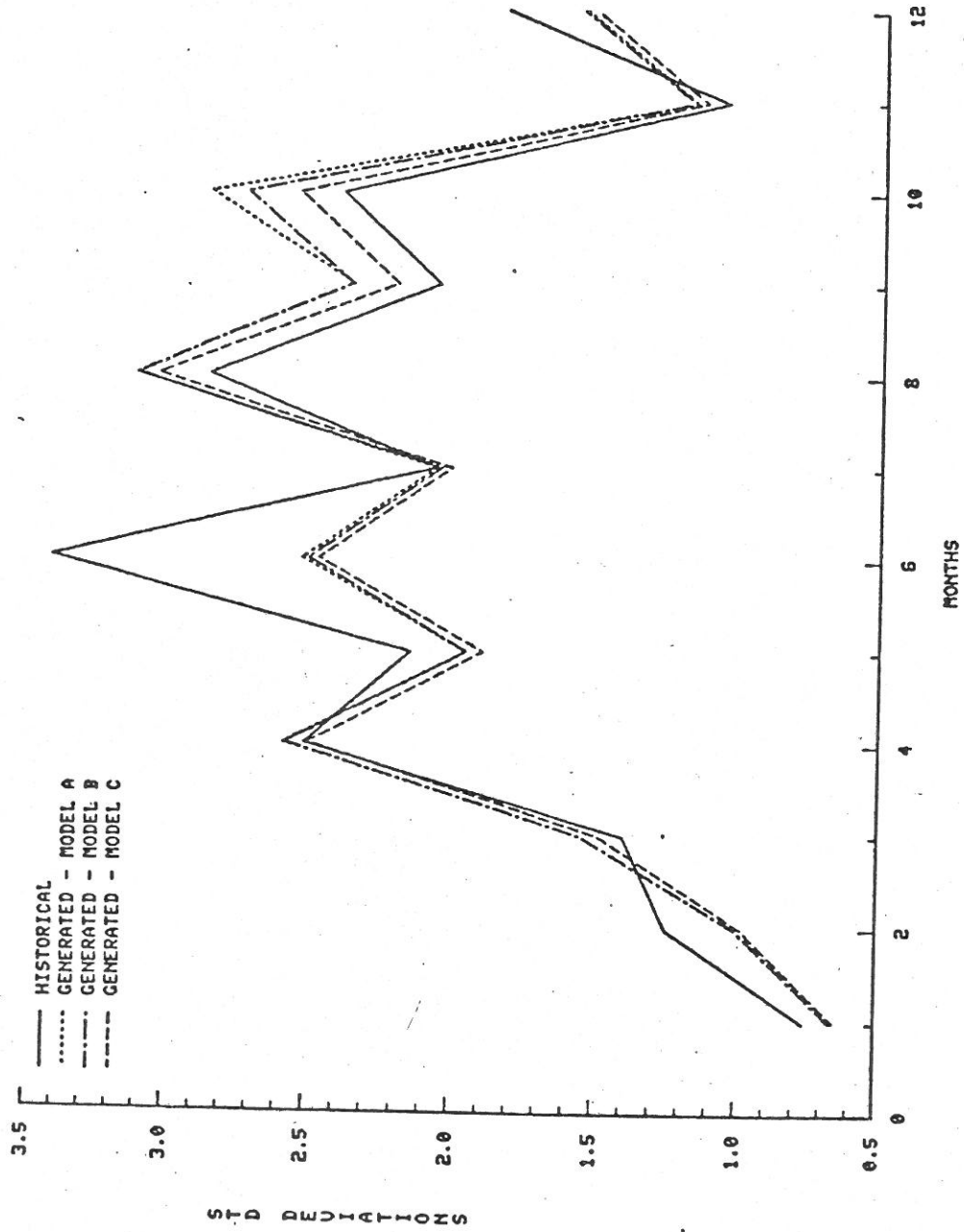


Figure 1.9.D.12 RANCHO - STD DEVIATIONS (ORIGINAL DOMAIN)

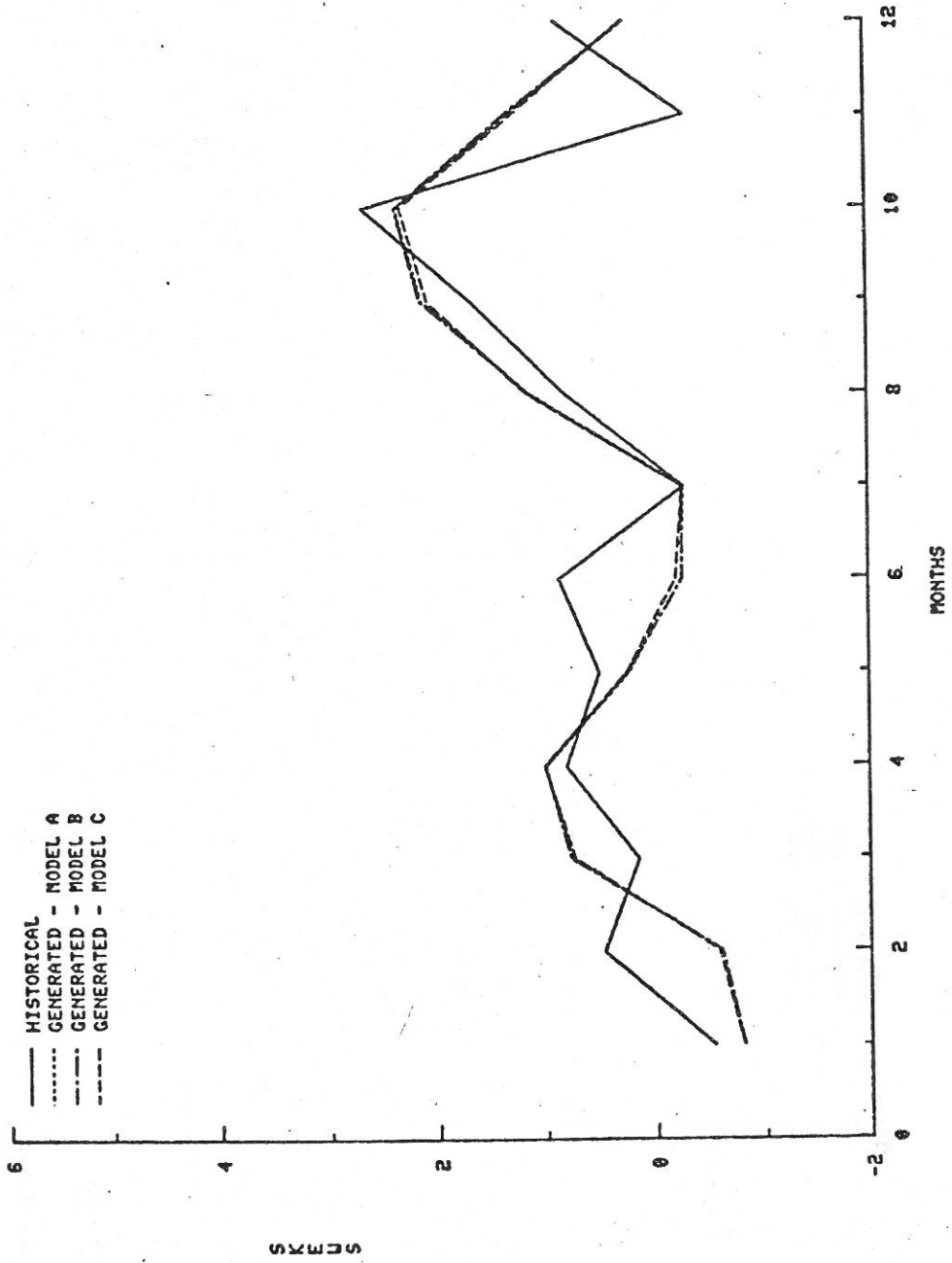


Figure 1.9.D.13 RANCHO - SKEWS (ORIGINAL DOMAIN)

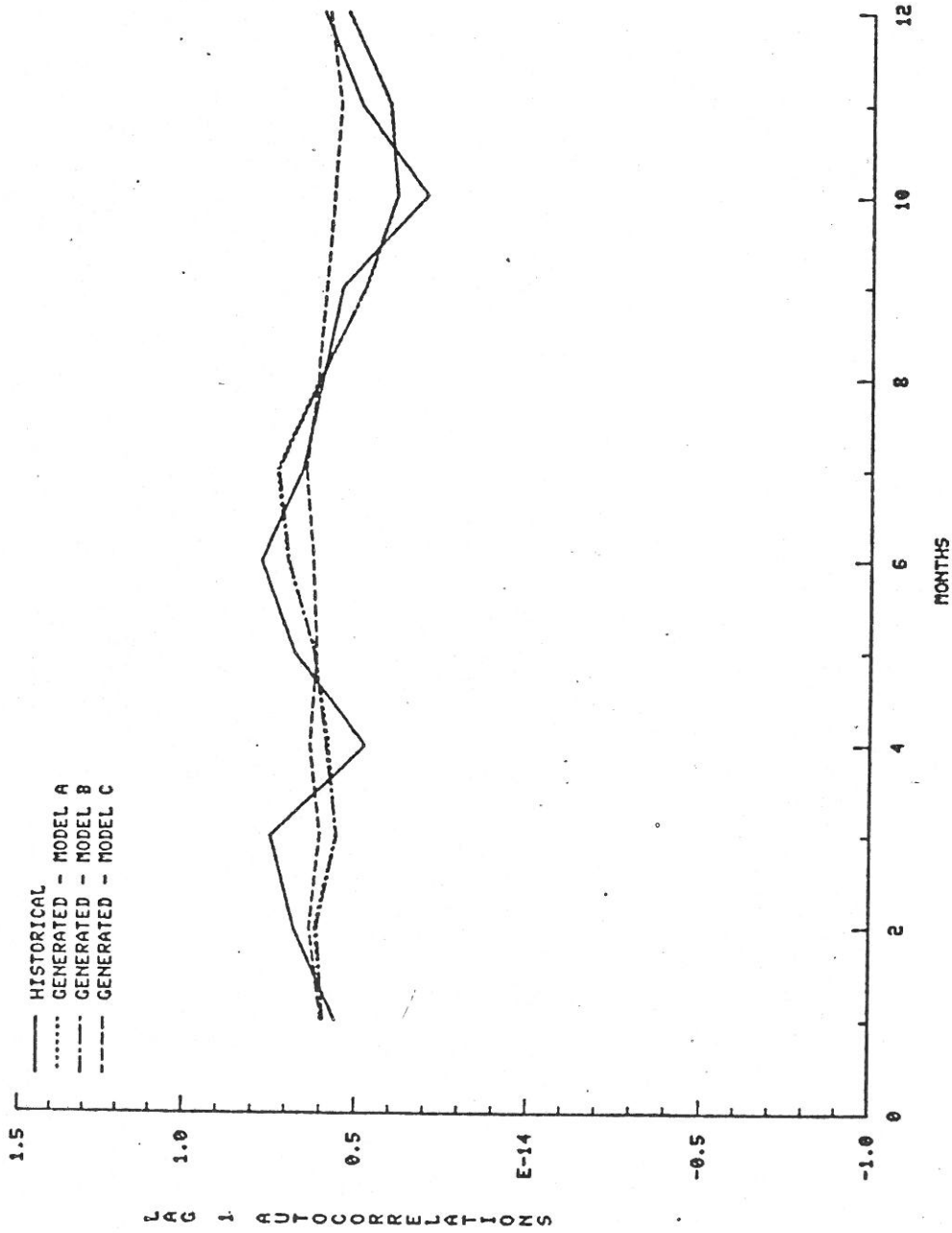


Figure 1.9.D.14 RANCHO - LAG 1 AUTOCORRELATIONS (ORIGINAL DOMAIN)

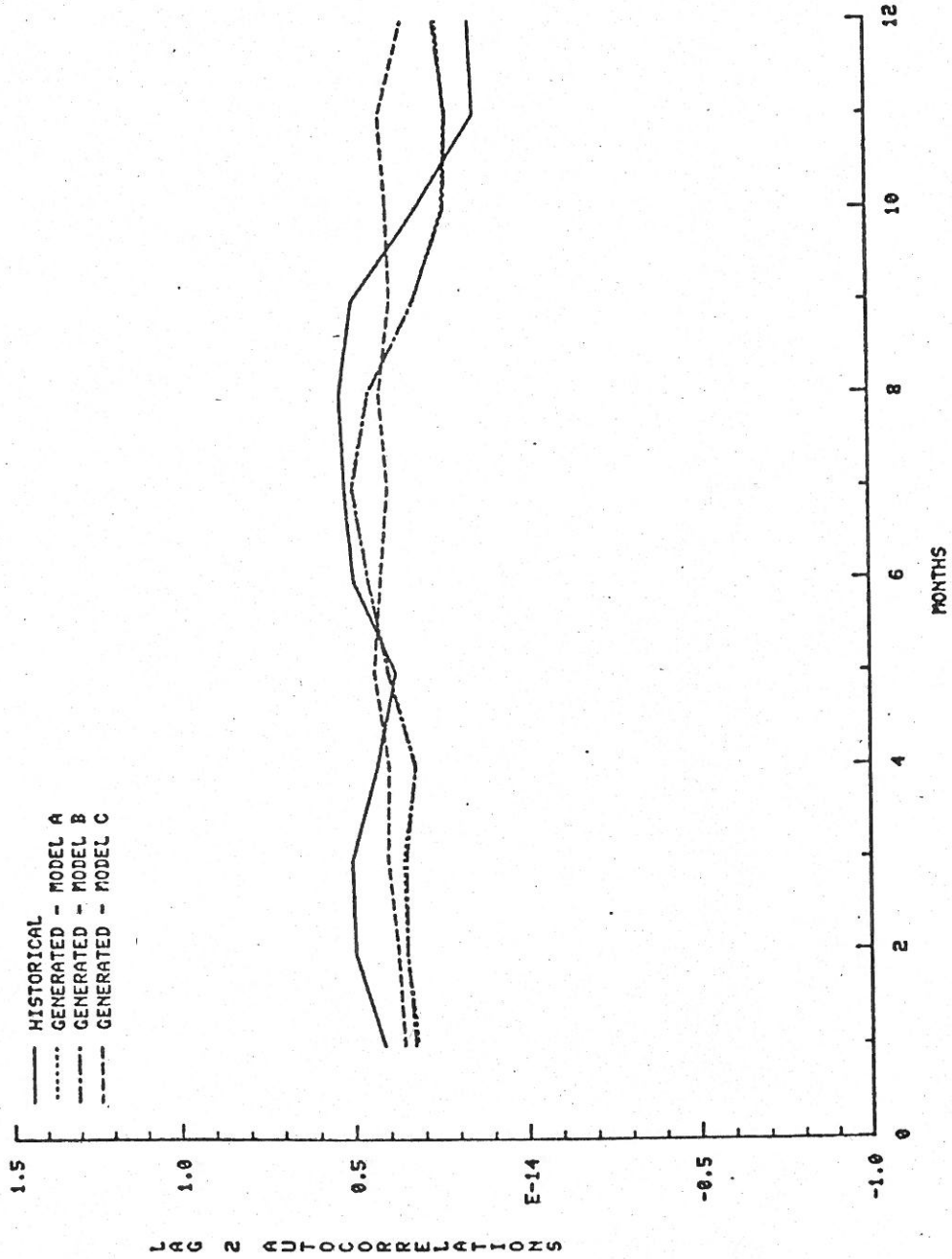


Figure 1.9.D.15 RANCHO - LAG 2 AUTOCORRELATIONS (ORIGINAL DOMAIN)

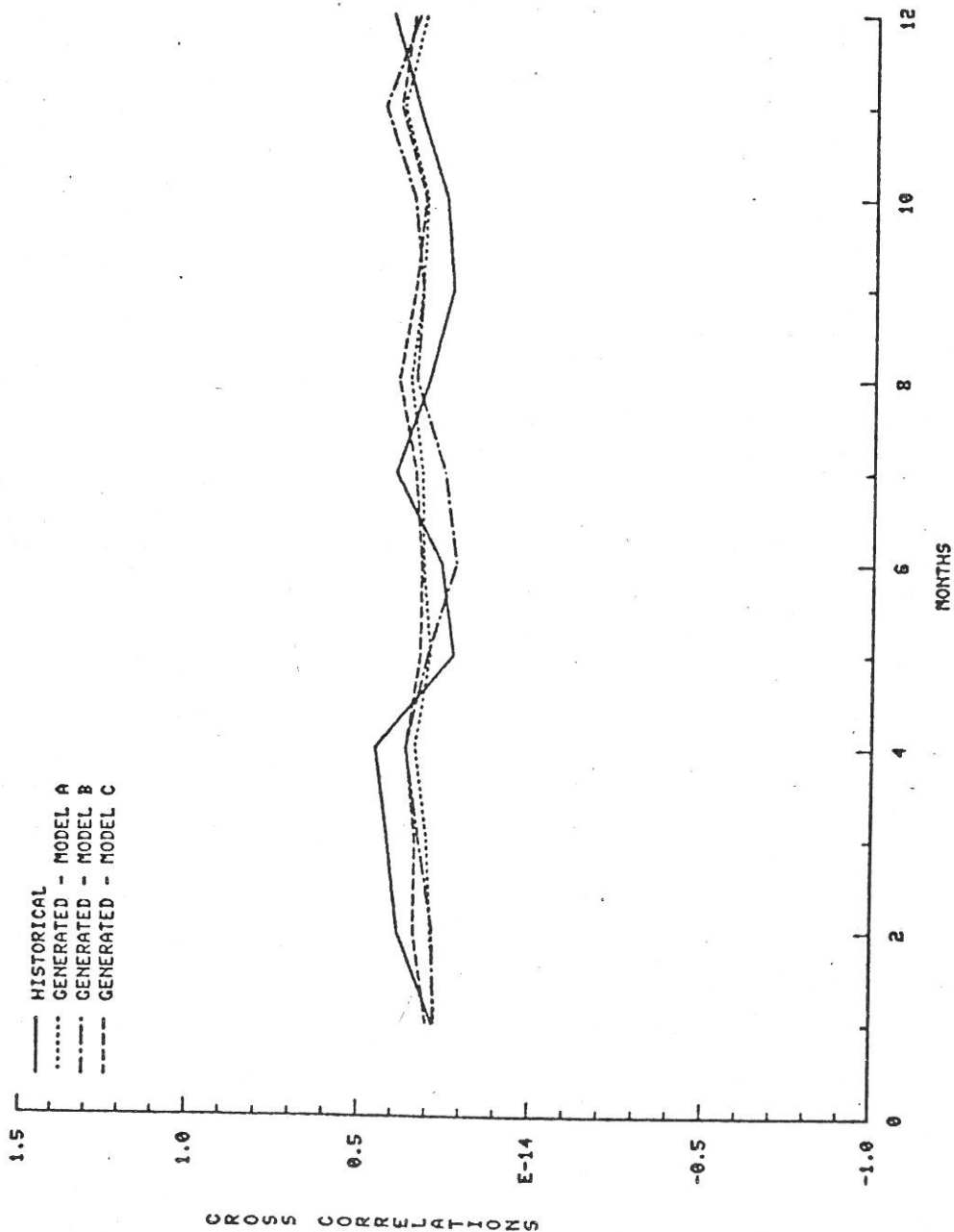


Figure I.9.D.16 CROSS-CORR PASODE AND RANCHO (ORIGINAL DOMAIN)

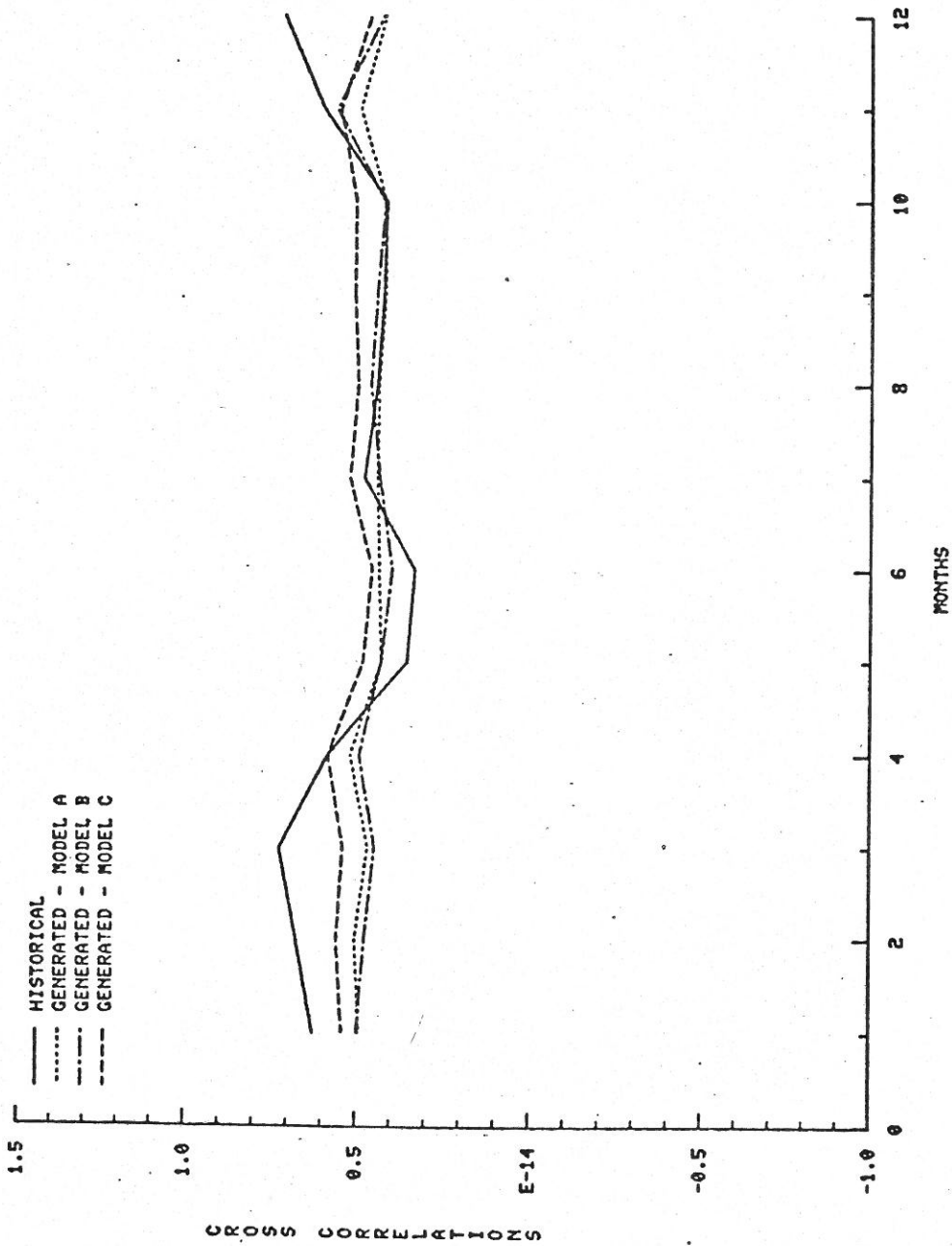


Figure 1.9.D.17 CROSS-CORR PALODE AND PASODE (ORIGINAL DOMAIN)

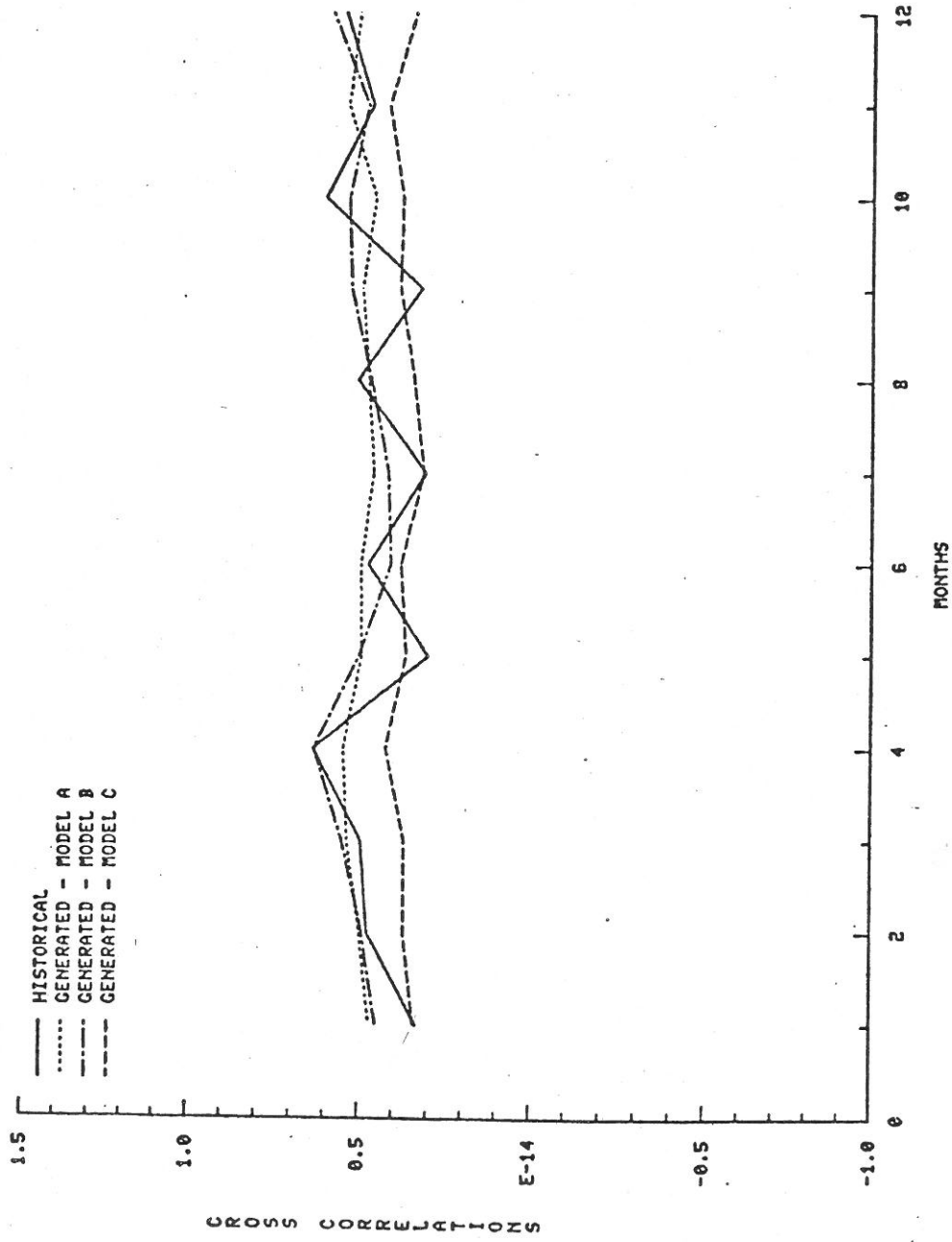


Figure 1.9.D.18 CROSS-CORR PALODE AND RANCHO (ORIGINAL DOMAIN)

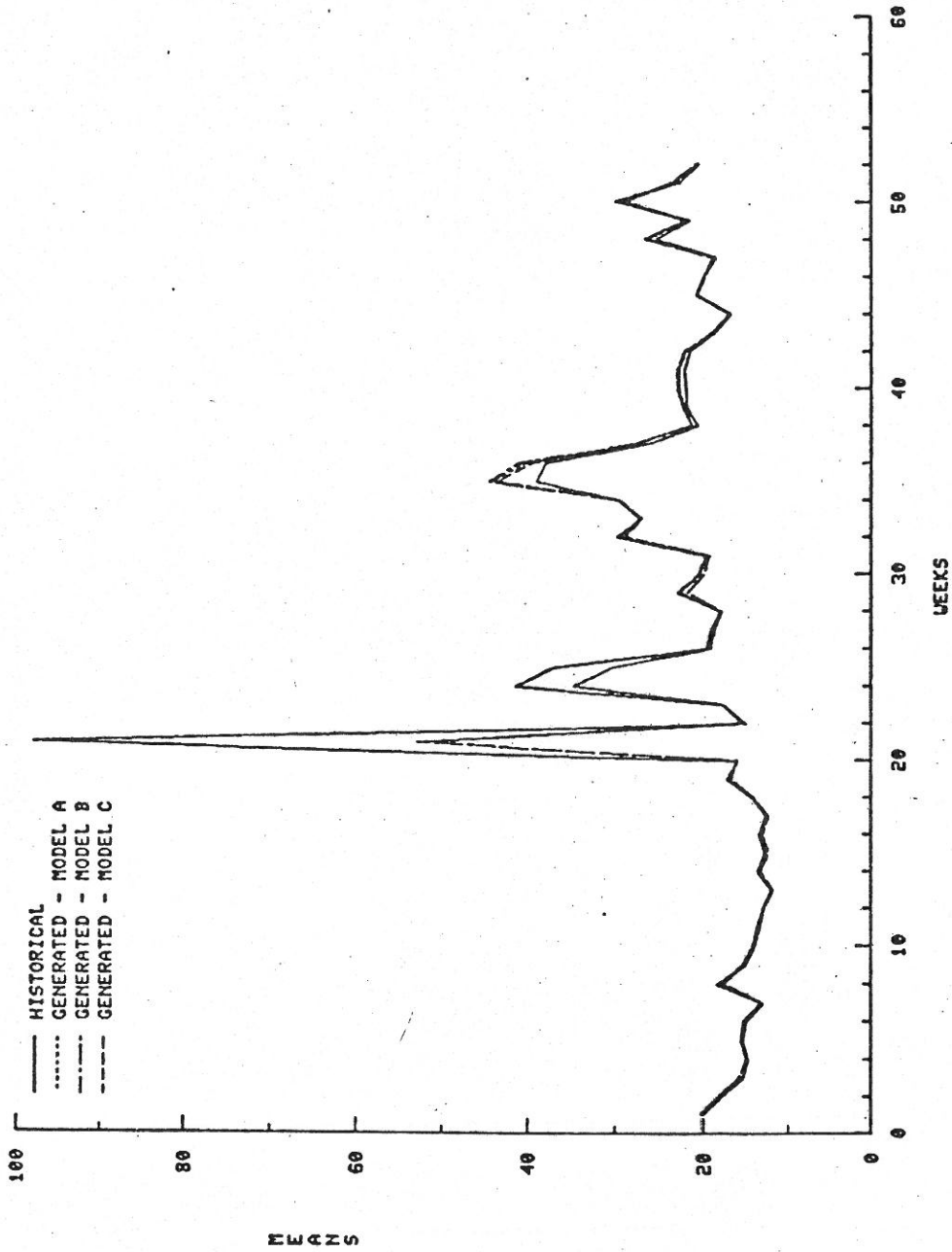


Figure 1.9.D.19 PASODE - MEANS (ORIGINAL DOMAIN)

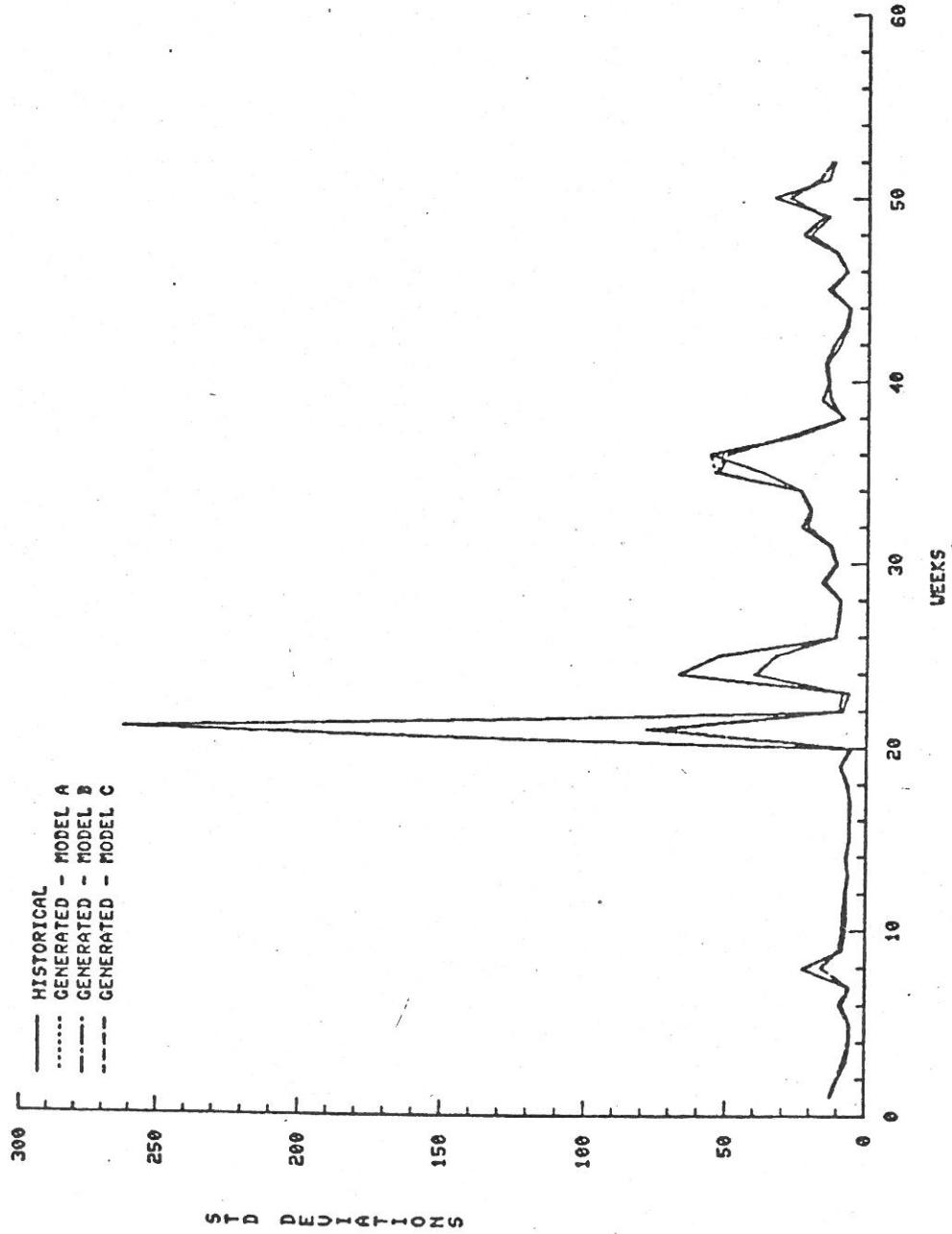


Figure 1.9.D.20 PASODE - STD DEVIATIONS (ORIGINAL DOMAIN)

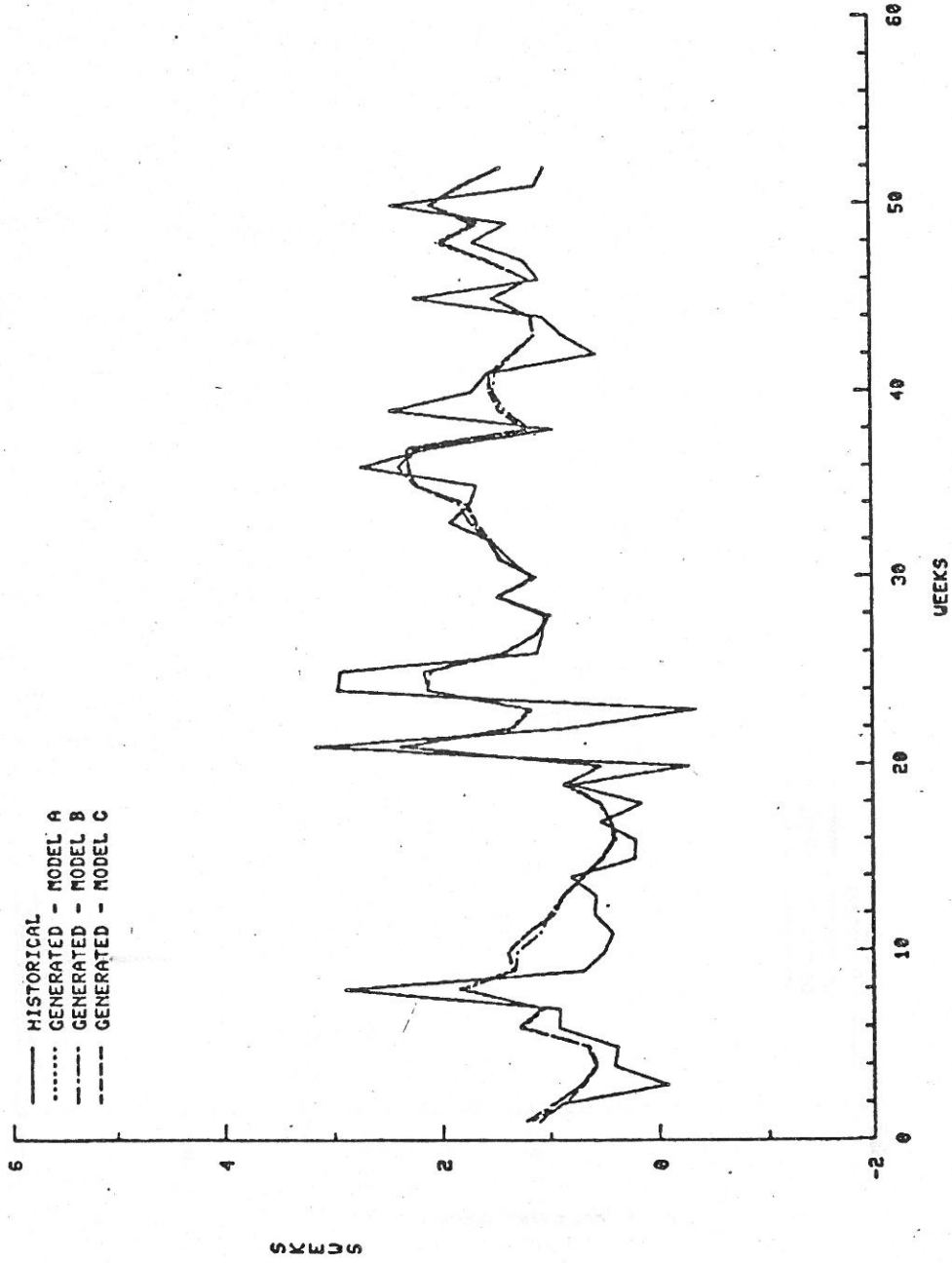


Figure 1.9.D.21 PASODE - SKEWS (ORIGINAL DOMAIN)

UNSATISFIE

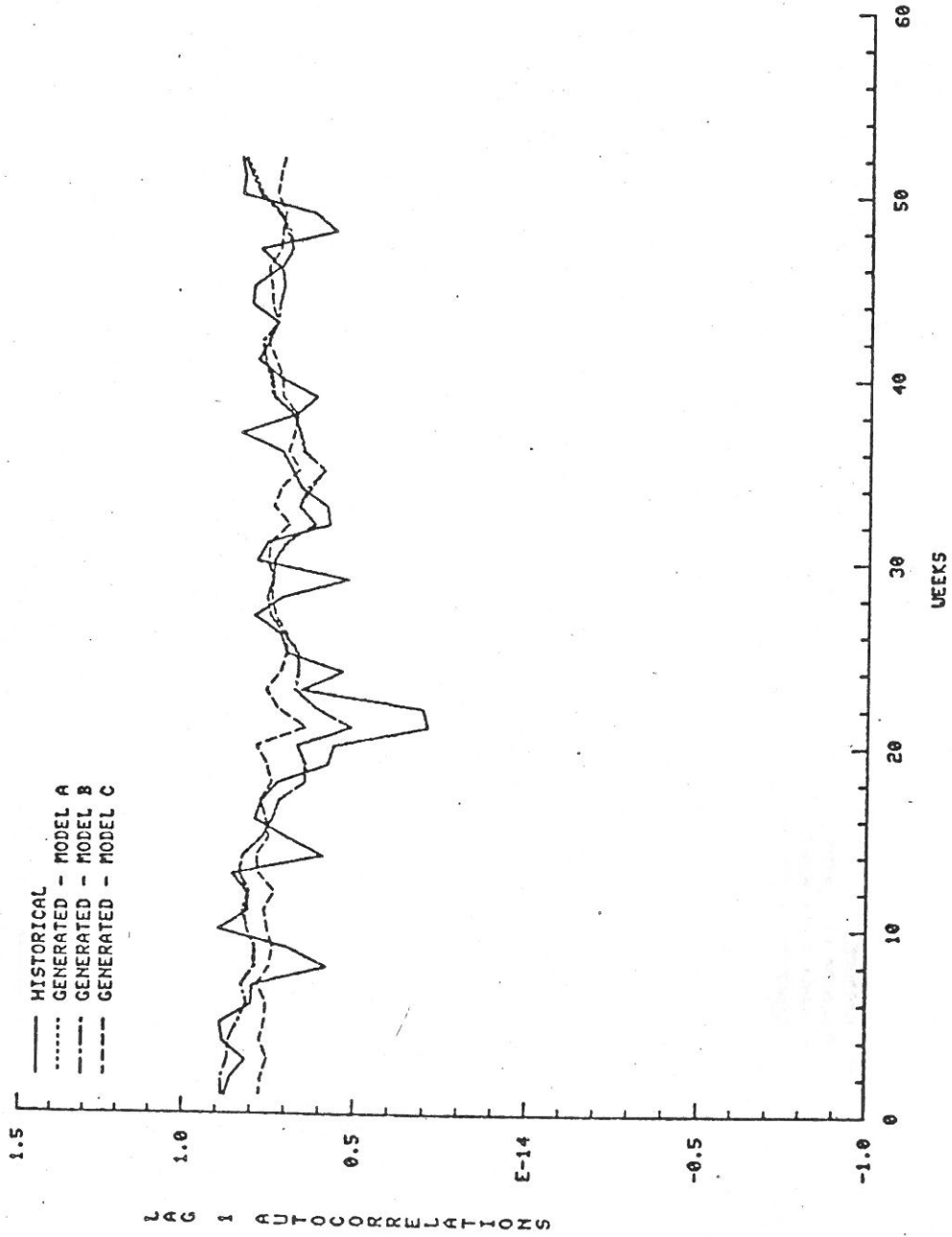


Figure 1.9.D.22 PASODE - LAG 1 AUTOCORRELATIONS (ORIGINAL DOMAIN)

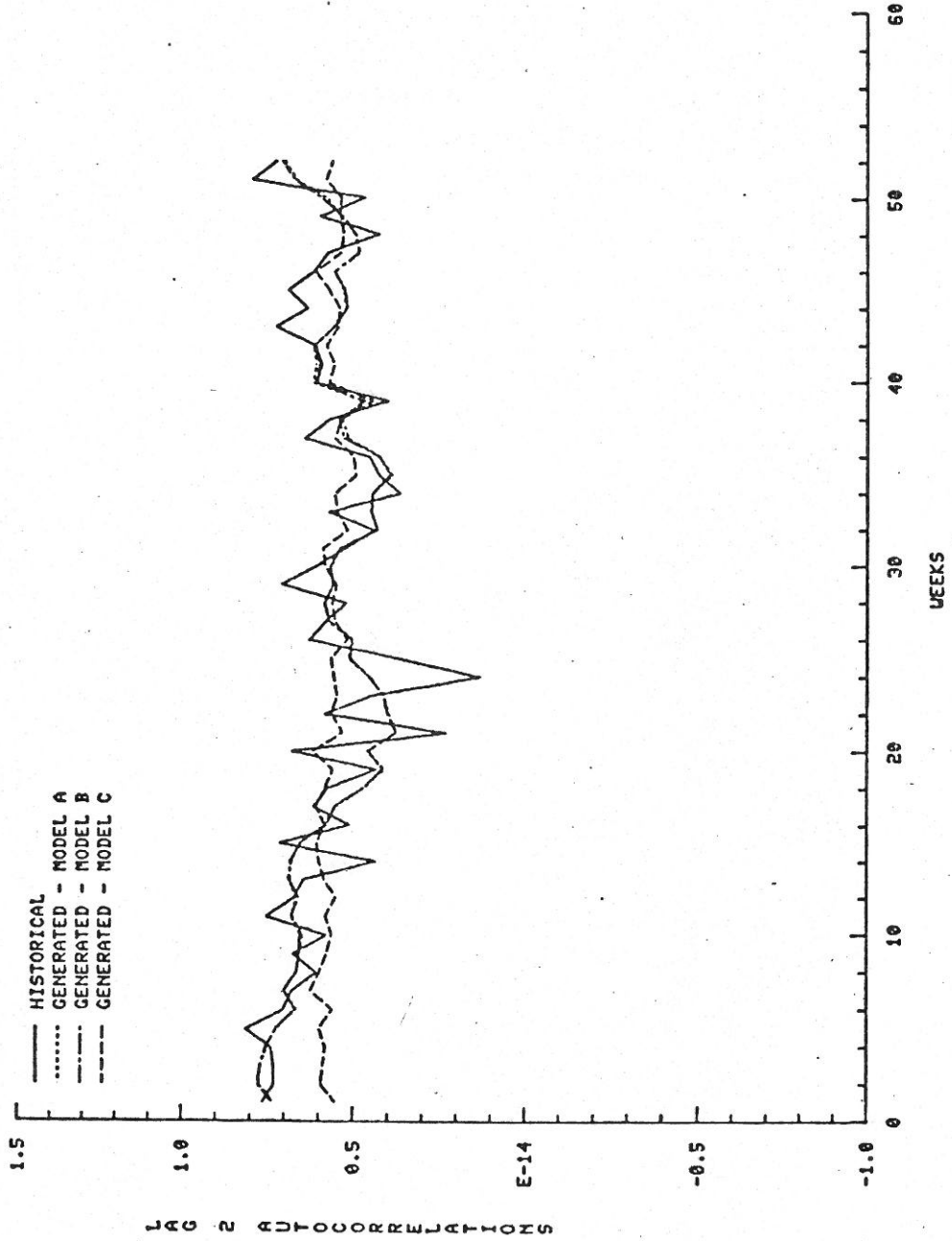


Figure 1.9.D.23 PASODE - LAG 2 AUTOCORRELATIONS (ORIGINAL DOMAIN)

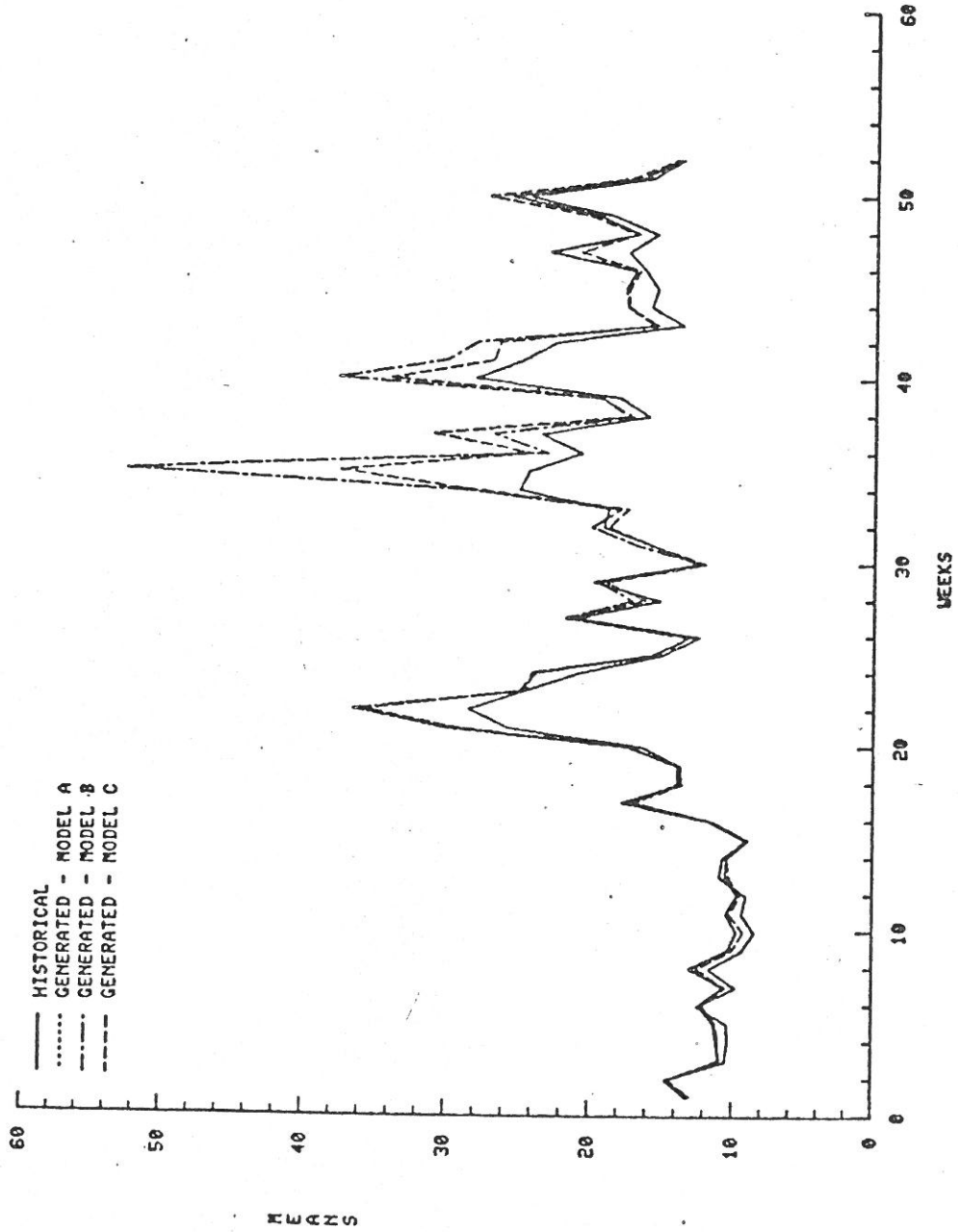


Figure 1.9.D.24 PALODE - MEANS (ORIGINAL DOMAIN)

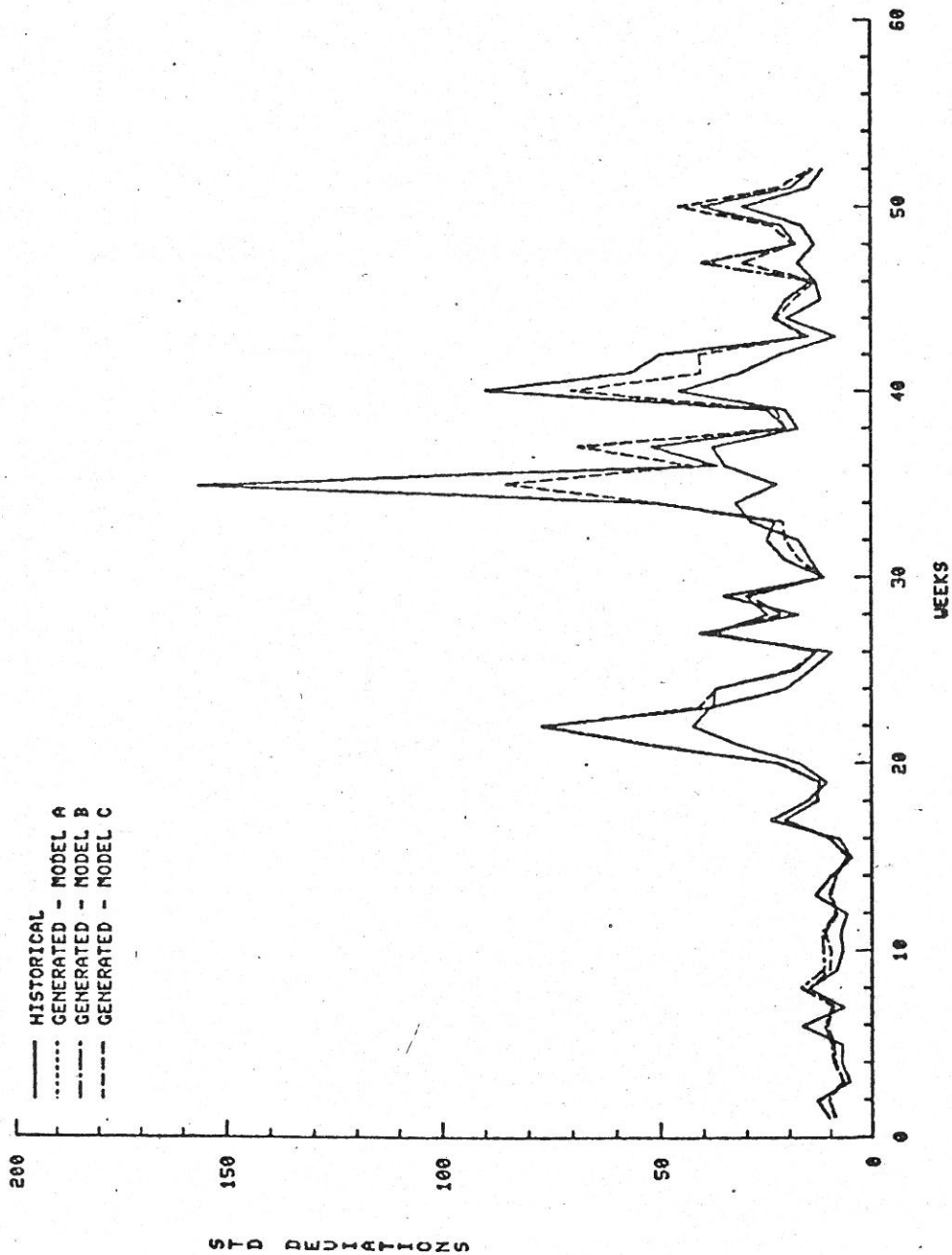


Figure 1.9.D.25 PALODE - STD DEVIATIONS (ORIGINAL DOMAIN)

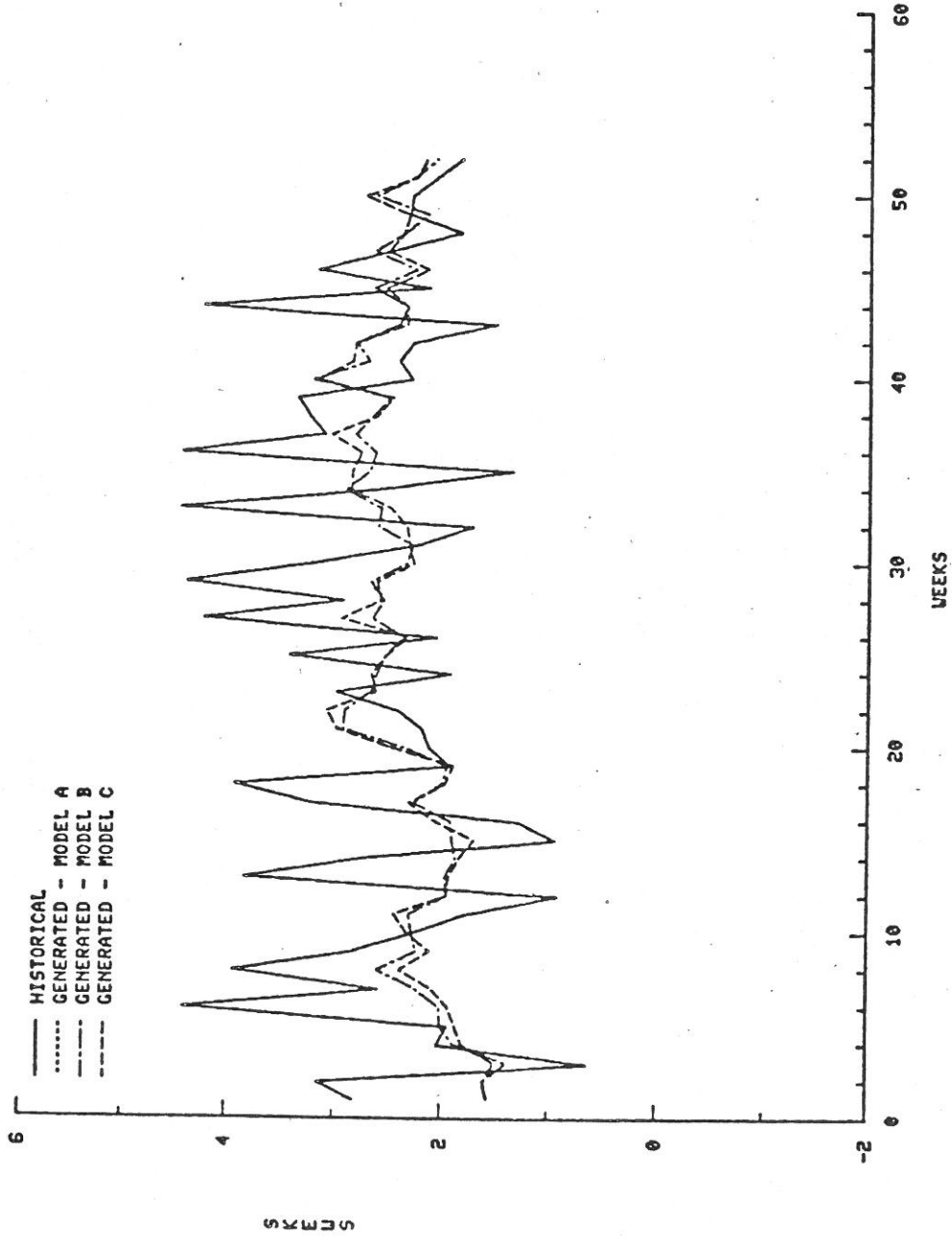


Figure 1.9.D.26 PALODE - SKEWS (ORIGINAL DOMAIN)

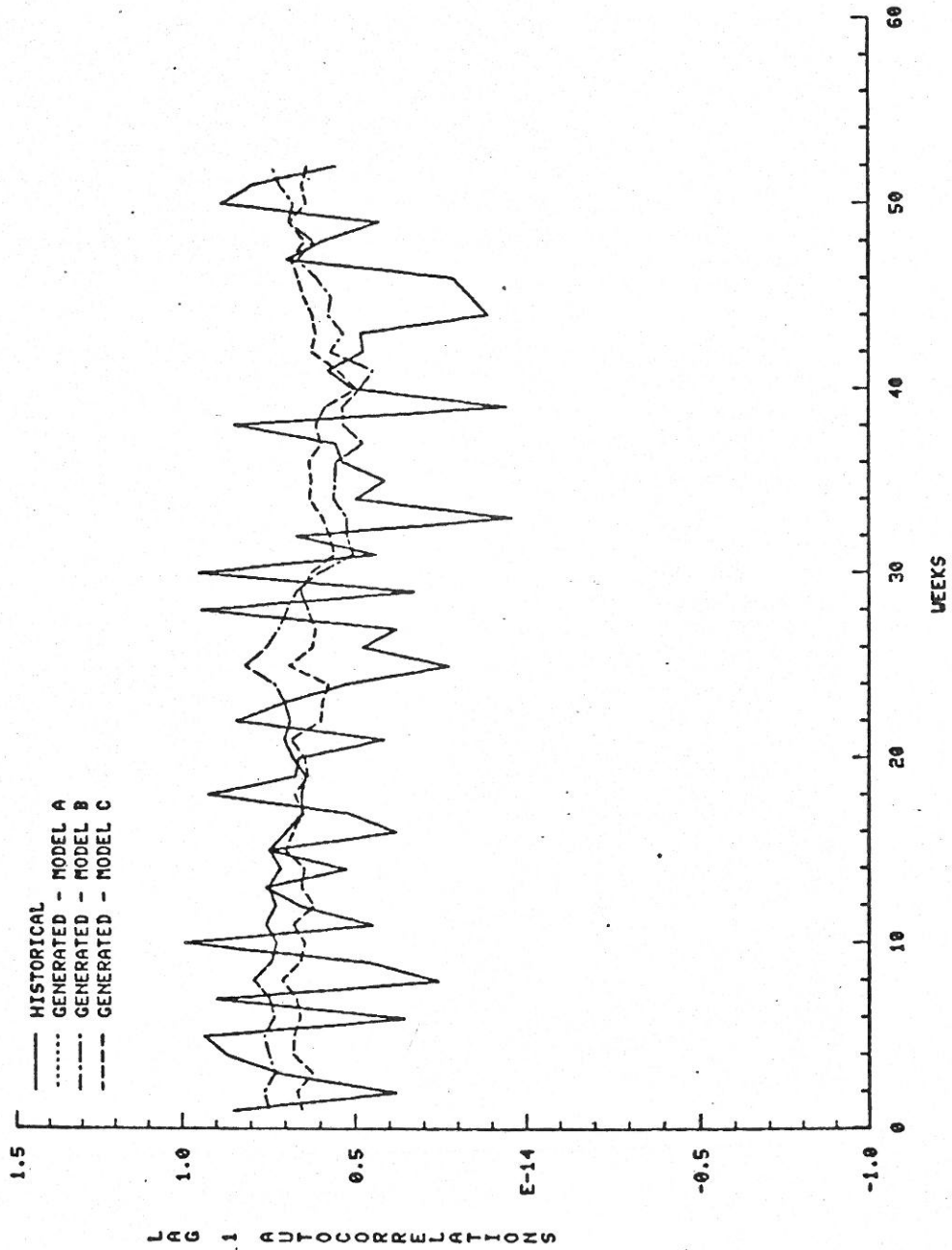


Figure 1.9.D.27 PALODE - LAG 1 AUTOCORRELATIONS (ORIGINAL DOMAIN)

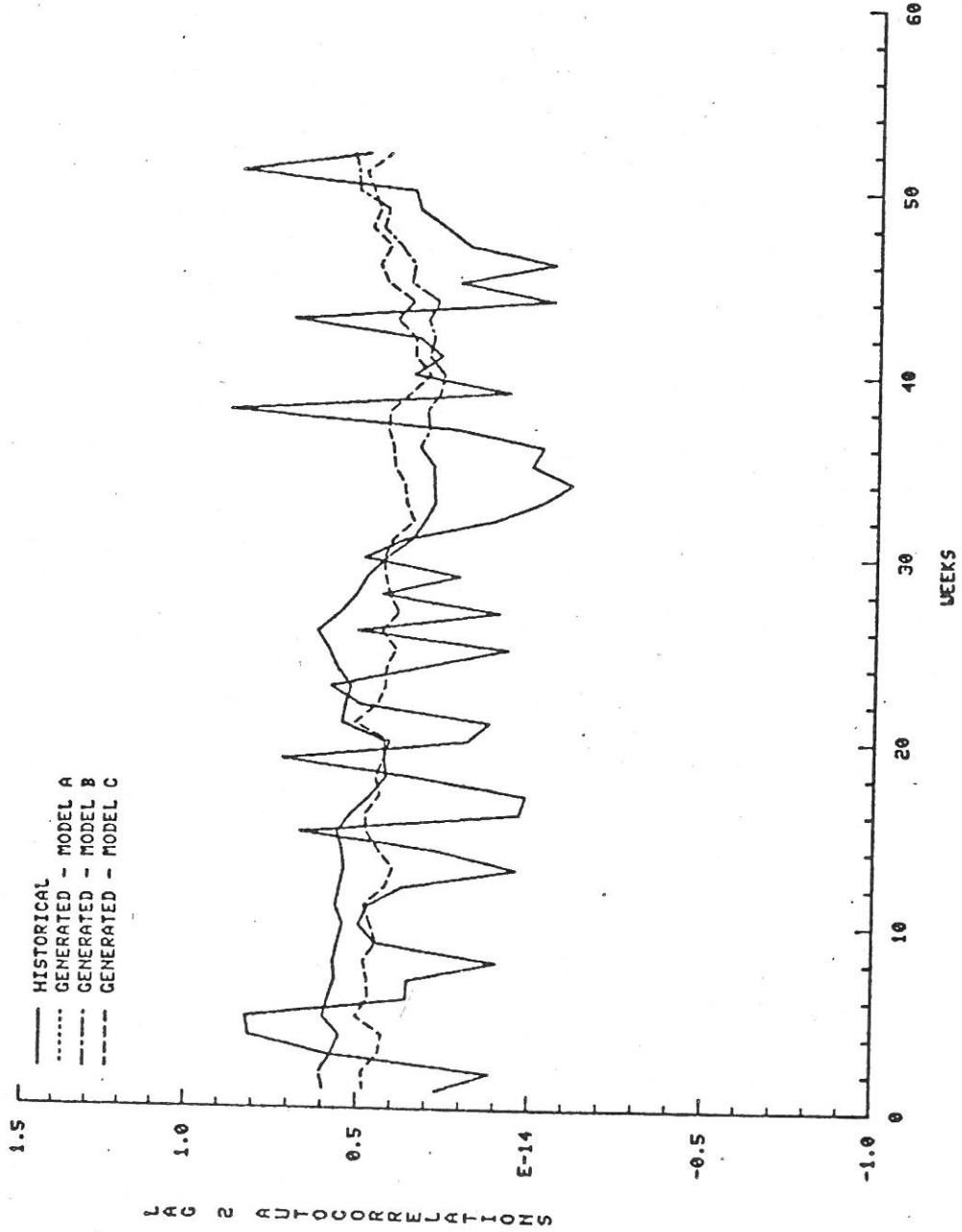


Figure 1.9.D.28 PALODE - LAG 2 AUTOCORRELATIONS (ORIGINAL DOMAIN)

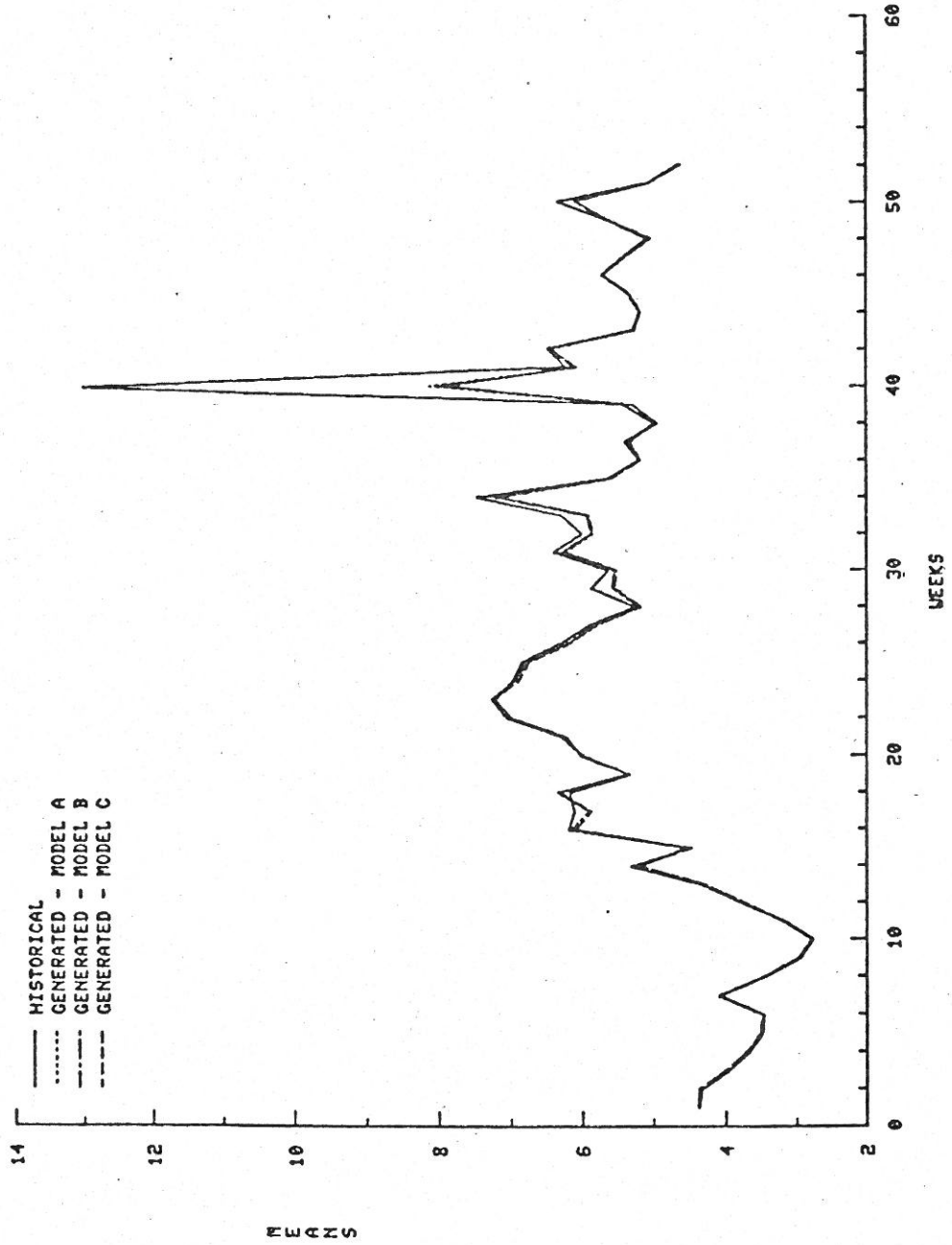


Figure 1.9.D.29 RANCHO - MEANS (ORIGINAL DOMAIN)

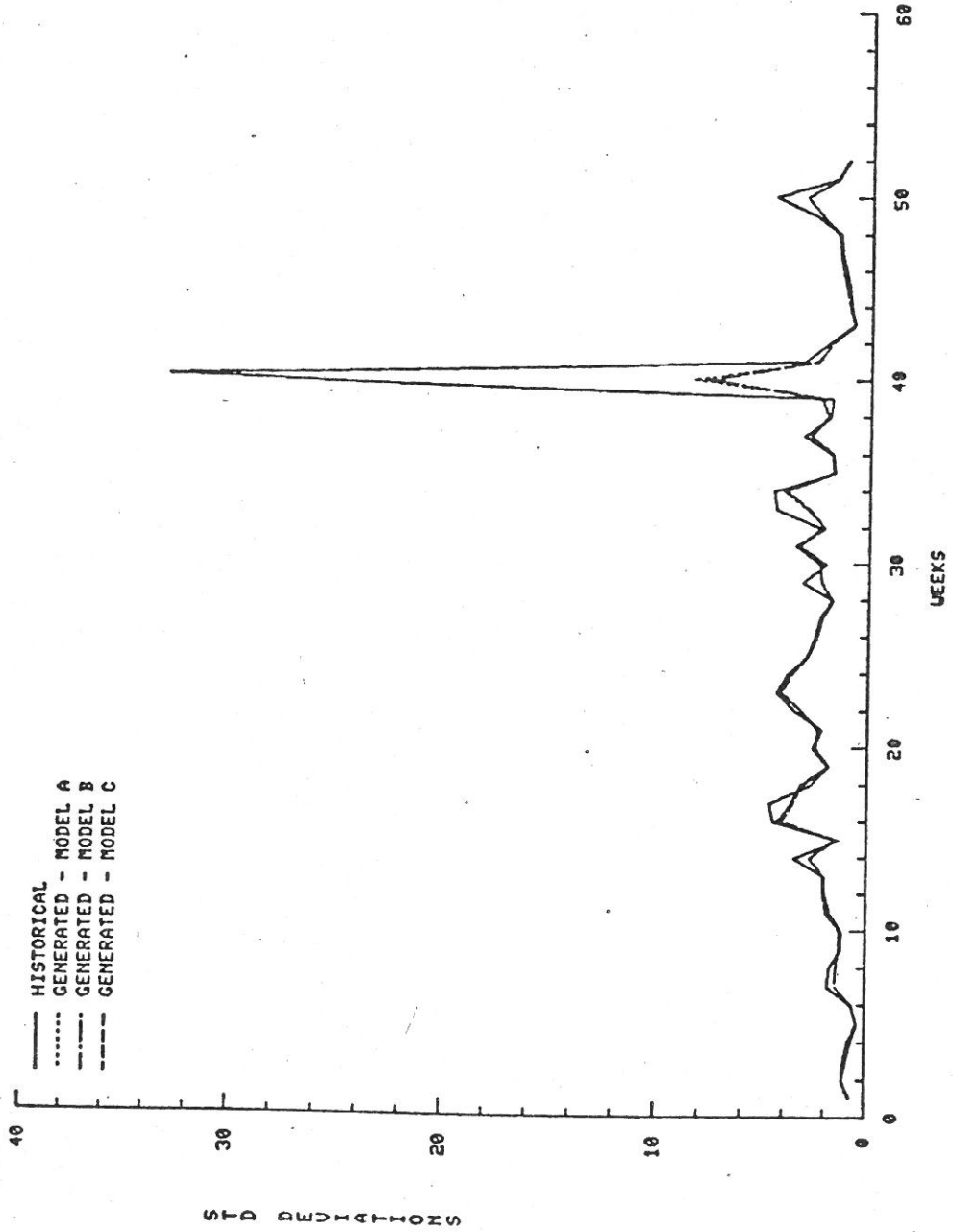


Figure 1.9.D.30 RANCHO - STD DEVIATIONS (ORIGINAL DOMAIN)

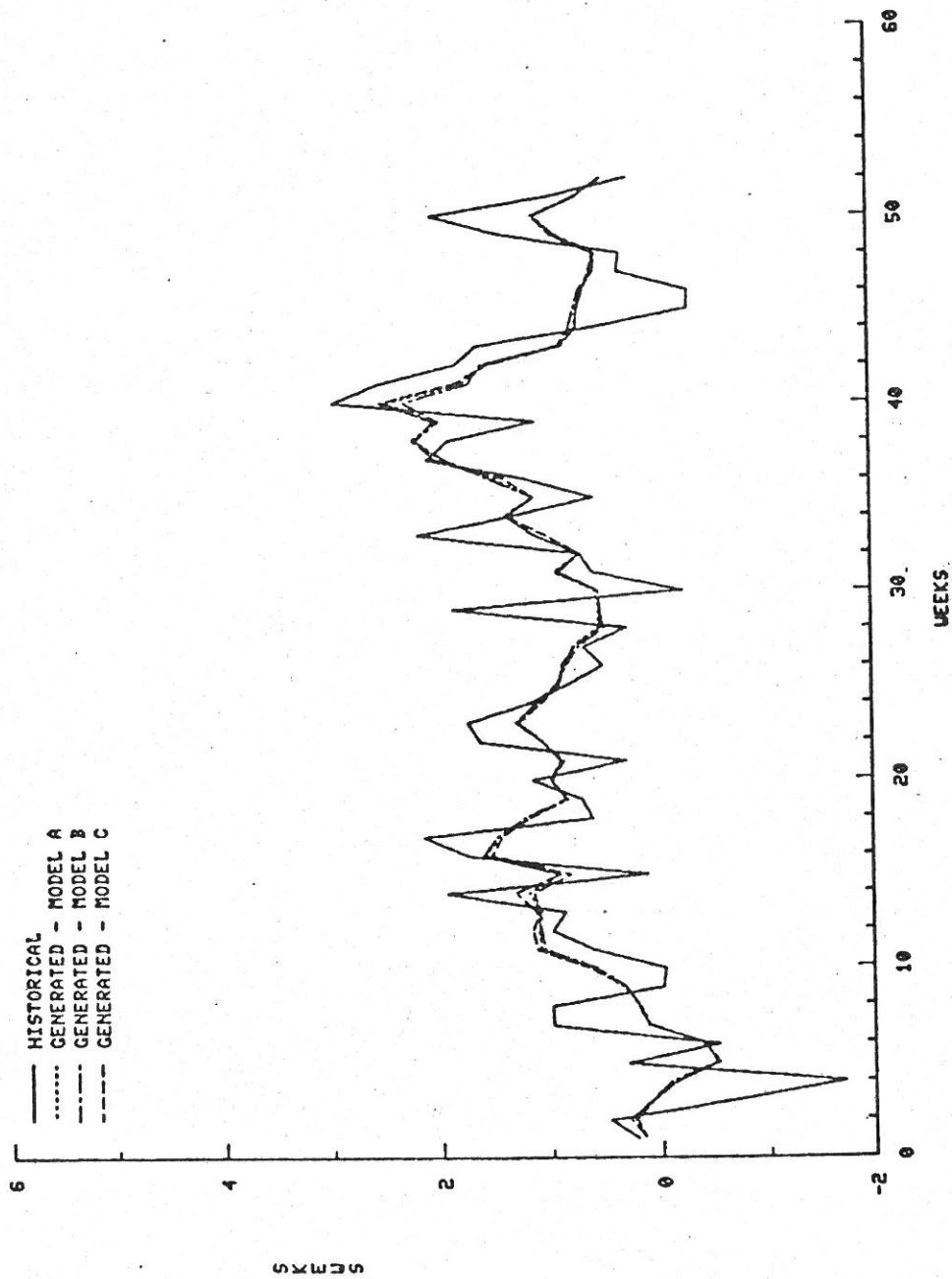


Figure 1.9.D.31 RANCHO - SKEWS (ORIGINAL DOMAIN)

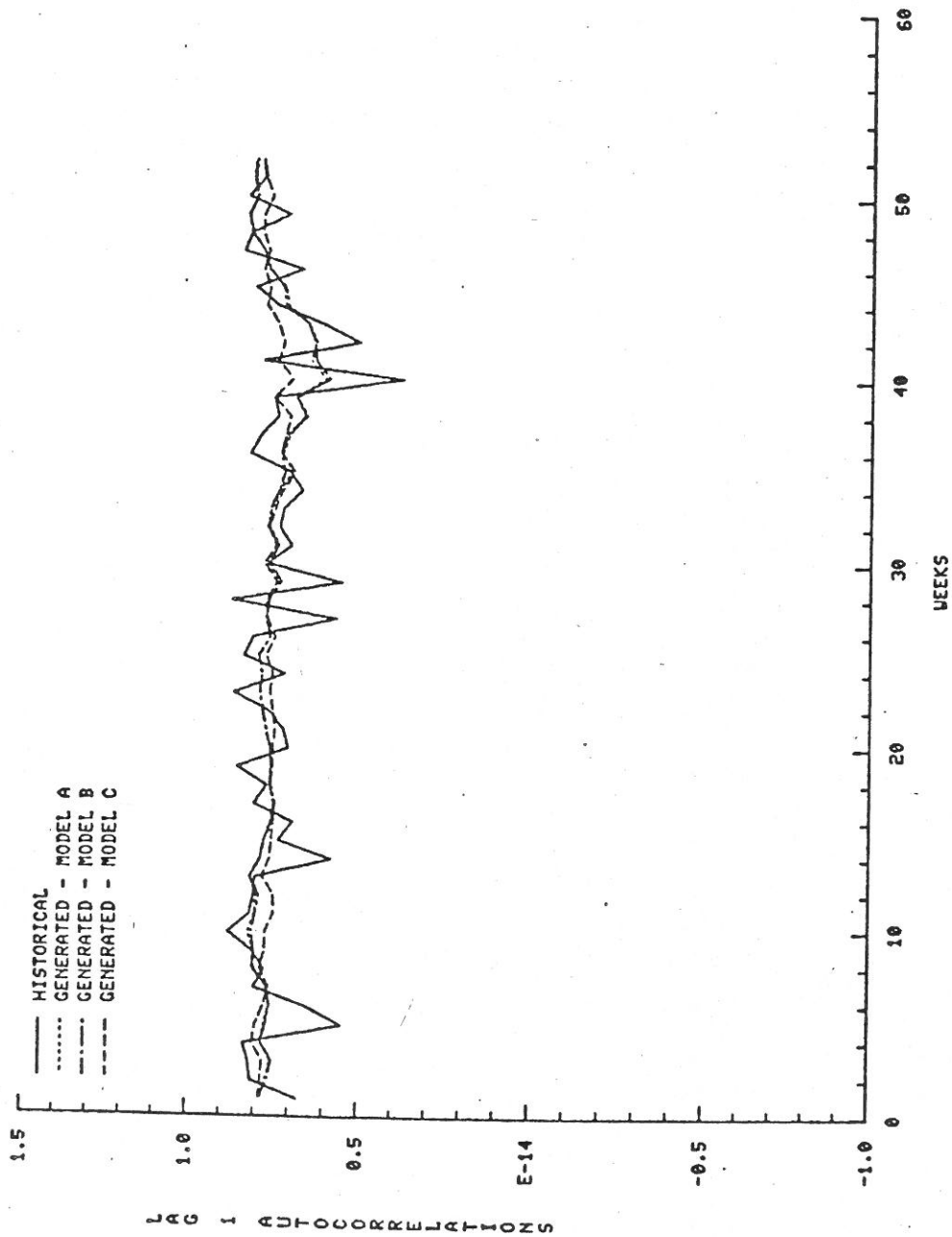


Figure I.9.D.32 RANCHO - LAG 1 AUTOCORRELATIONS (ORIGINAL DOMAIN)

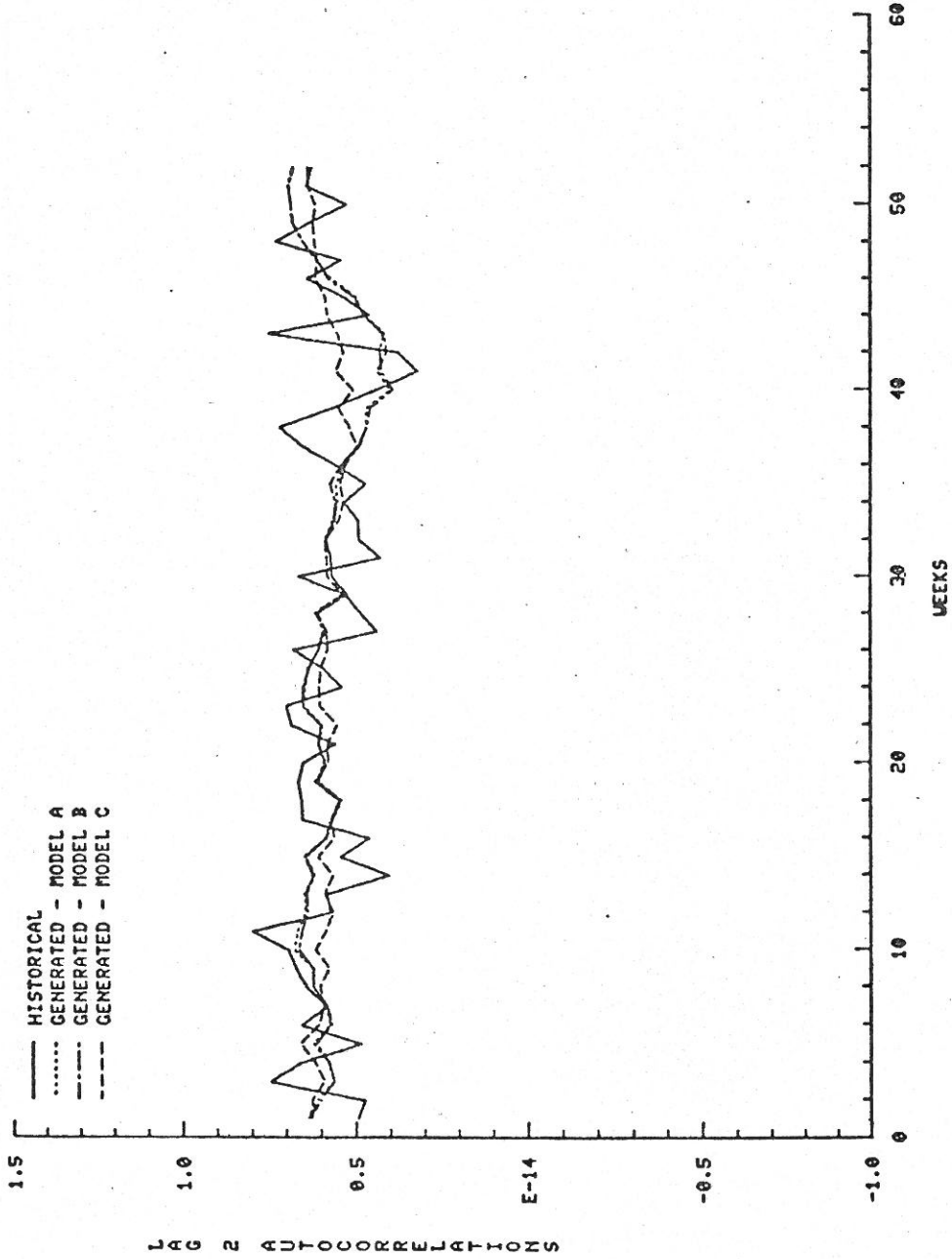


Figure 1.9.D.33 RANCHO - LAG 2 AUTOCORRELATIONS (ORIGINAL DOMAIN)

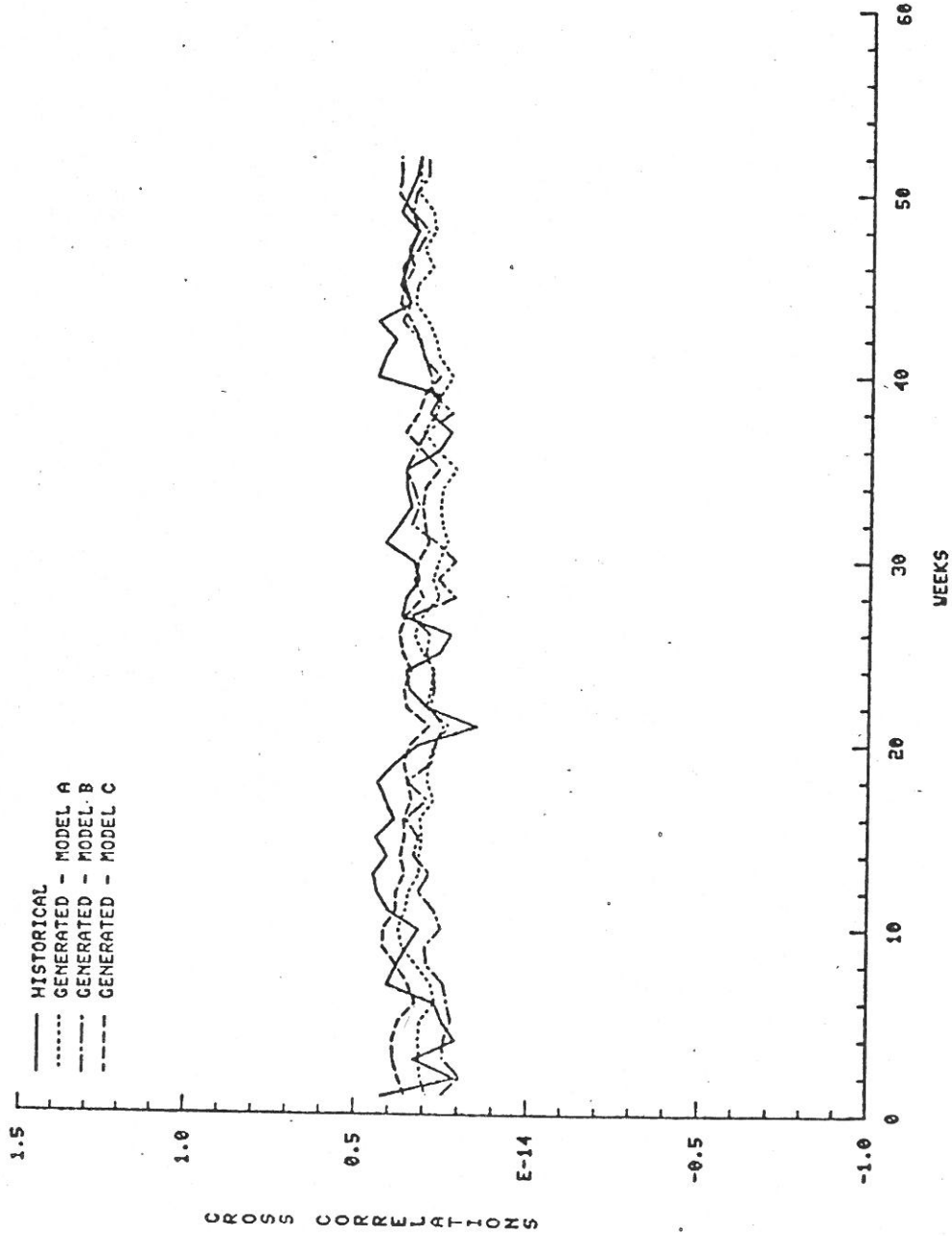


Figure 1.9.D.34 CROSS-CORR PASODE AND RANCHO (ORIGINAL DOMAIN)

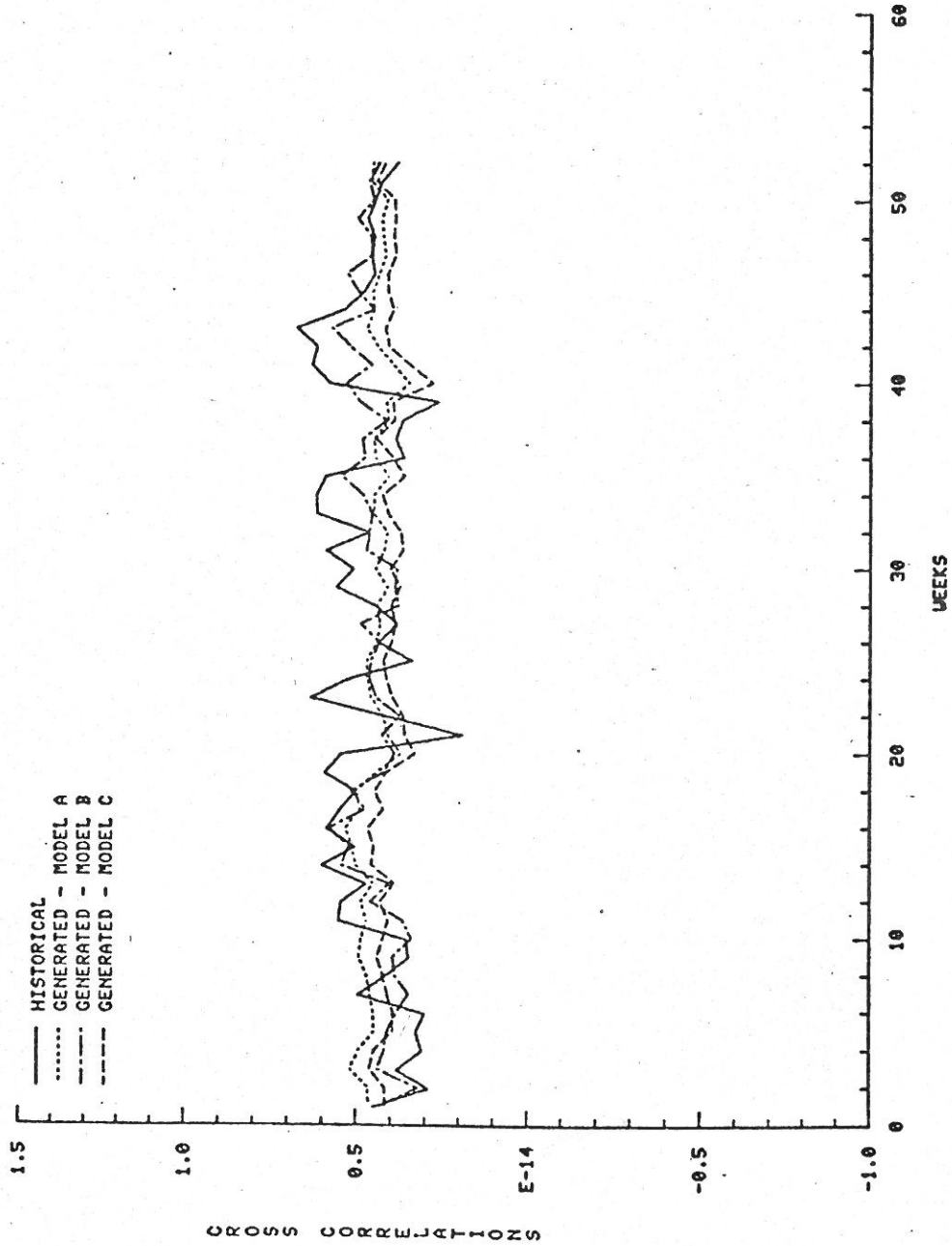


Figure 1.9.D.35 CROSS-CORR PALODE AND RANCHO (ORIGINAL DOMAIN)

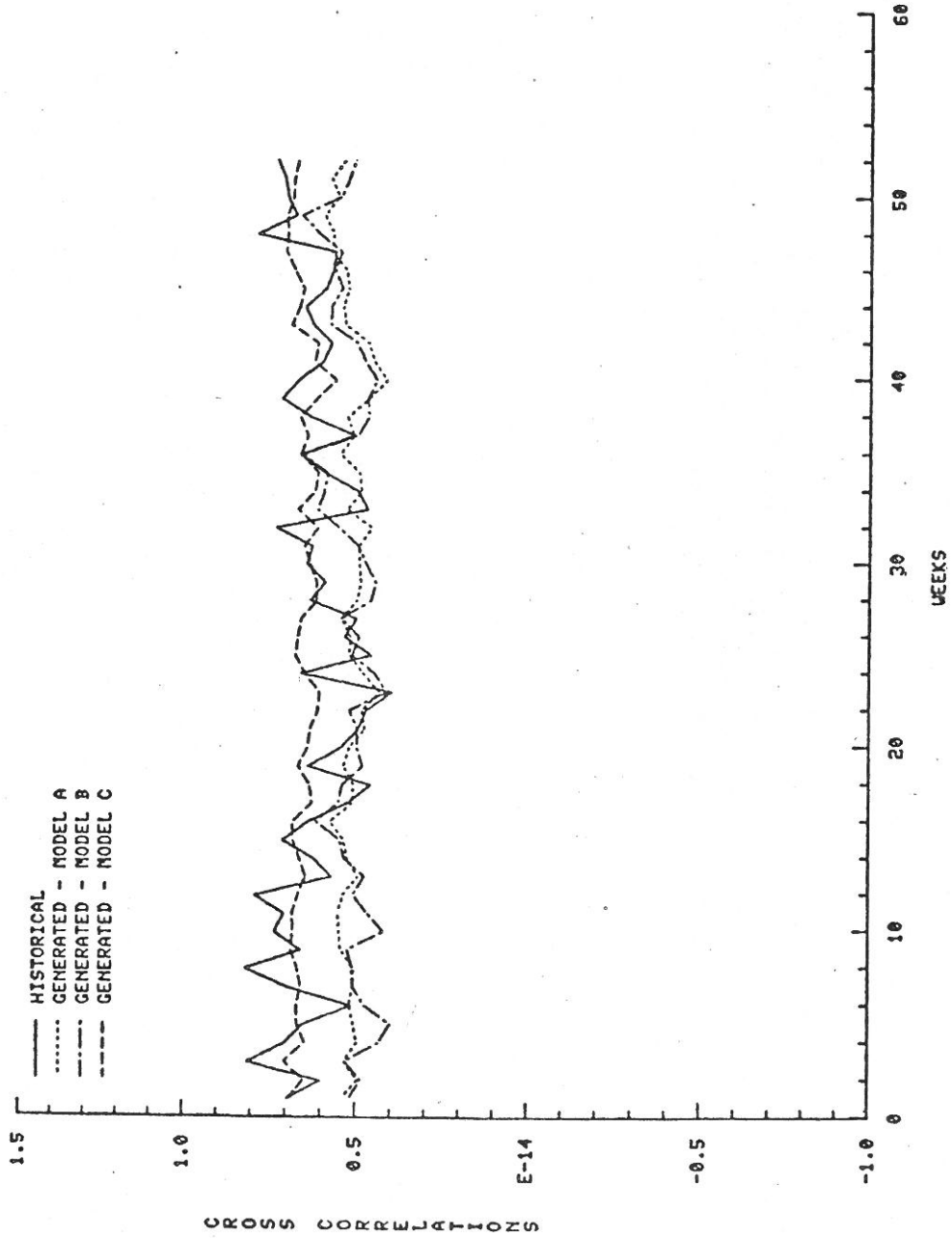


Figure 1.9.D.36 CROSS-CORR PALODE AND PASODE (ORIGINAL DOMAIN)

APPENDIX 1.9.E

HISTORICAL (EXTENDED SERIES) AND GENERATED MONTHLY AND WEEKLY
STATISTICS OF PALO DE CAJA, PASO DEL ERMITAÑO AND RANCHO ARRIBA
FOR MODEL B IN THE ORIGINAL, LOGARITHMIC AND LOG-WILSON-HILFERTY
DOMAIN OF FLOWS

In computing the generated statistics of each sample, Stedinger and Taylor (1982) suggested using the "theoretical" means and standard deviations (i.e., means and standard deviations of the extended data) for computing the generated statistics. This is done in order to reduce any statistical biases introduced from small-sample estimates of the standard deviations, skew coefficients and autocorrelations when computed based on the means and standard deviations of the generated flows. Thus, the unbiased estimates of the generated statistics are:

$$\bar{Y}_\tau = \frac{1}{n} \sum_{\nu=1}^n Y_{\nu,\tau} \quad (1.9.E.1)$$

$$S_\tau(Y) = \left\{ \frac{1}{n} \sum_{\nu=1}^n [Y_{\nu,\tau} - \hat{\mu}_\tau(Y)]^2 \right\}^{1/2} \quad (1.9.E.2)$$

$$G_\tau(Y) = \frac{1}{n\hat{\sigma}_\tau^3(Y)} \sum_{\nu=1}^n [Y_{\nu,\tau} - \hat{\mu}_\tau(Y)]^3 \quad (1.9.E.3)$$

$$R_\tau(k) = \frac{1}{n\hat{\sigma}_\tau(Y)\hat{\sigma}_{\tau-k}(Y)} \sum_{\nu=1}^n [Y_{\nu,\tau} - \hat{\mu}_\tau(Y)][Y_{\nu,\tau-k} - \hat{\mu}_{\tau-k}(Y)] \quad (1.9.E.4)$$

where $\hat{\mu}_\tau(Y)$ and $\hat{\sigma}_\tau(Y)$ are the seasonal means and standard deviations, n is the sample size, and τ is the seasonal index where $\tau = 1, \dots, \omega$ seasons. Equations (1.9.E.1) through (1.9.E.4) are used to compute the generated statistics of each sample not only for the

original domain of flows (represented by Y) but for the other domain of flows by replacing the notation Y by Z for the log-Wilson-Hilferty domain, and by X for the log-domain (see Appendix 1.9.A).

After computing the generated statistics for each sample (a total of 50 for each statistic), the average and standard error are determined. Denoting the mth sample generated statistic by $V_r(m)$, the average \bar{V}_r and standard error $S_r(V)$ are computed from

$$\bar{V}_r = \frac{1}{M} \sum_{m=1}^M V_r(m) \quad (1.9.E.5)$$

and

$$S_r(V) = \left\{ \frac{1}{M} \sum_{m=1}^M [V_r(m) - \bar{V}_r]^2 \right\}^{1/2} \quad (1.9.E.6)$$

where M is equal to 50 samples.

The computed averages and standard errors of the monthly and weekly generated statistics are given in the figures below for the three stations in the log-Wilson-Hilferty domain, logarithmic domain, and original domain. Also plotted in these figures are the historical statistics and the upper and lower confidence bands of the generated statistics. The confidence band B_r is computed as positive and negative one-standard error relative to the average as

$$B_r = \bar{V}_r \pm S_r(V) \quad (1.9.E.7)$$

where \bar{V}_r and $S_r(V)$ are as defined above. Note that for the plots of historical lag-1 autocorrelation coefficients in the log-Wilson-Hilferty domain, and the historical skew coefficients in the log-domain are the fitted Fourier functions (since, these are the parameters used in data generation).

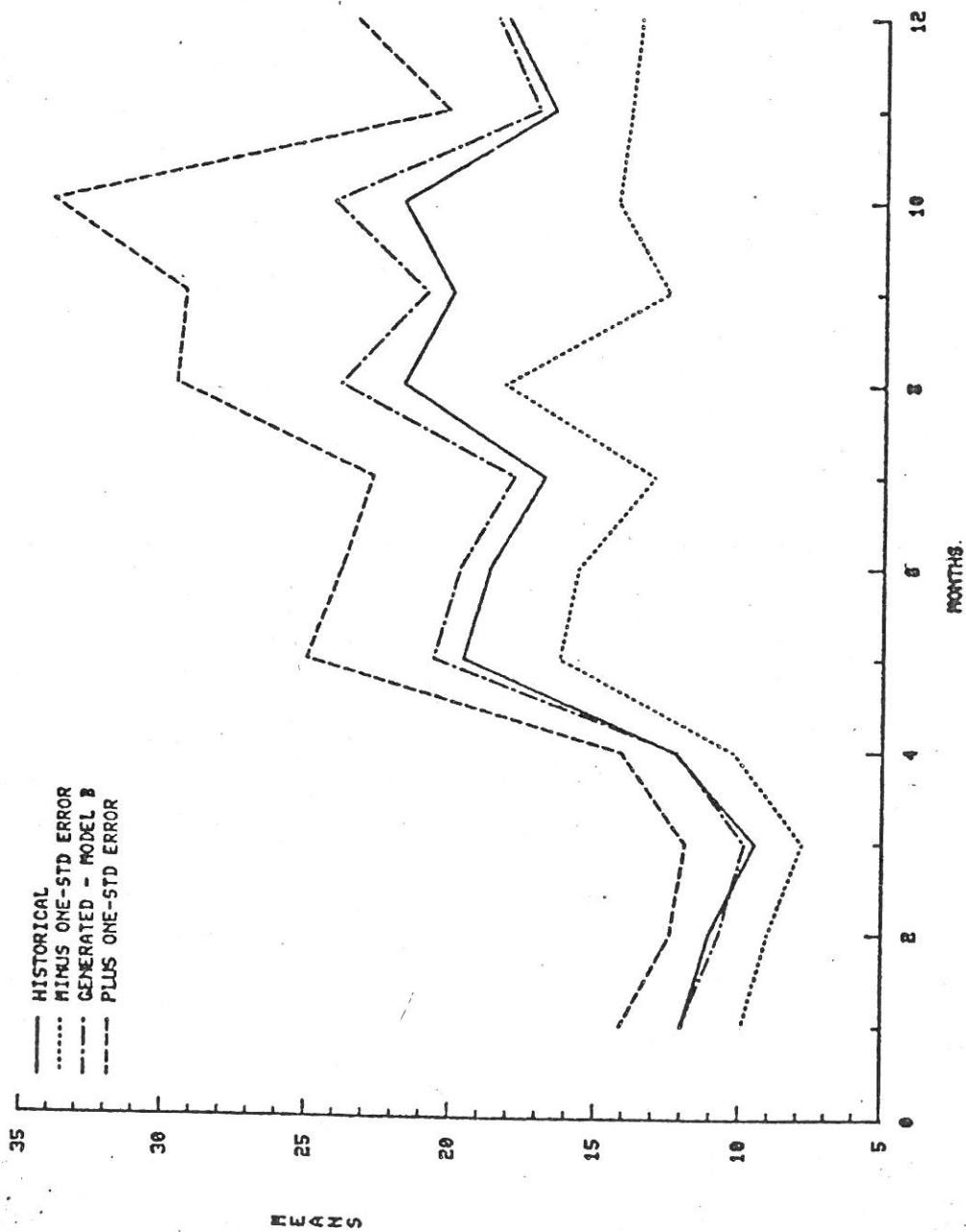


Figure 1.9.E.1. PALODE - MEANS (ORIGINAL DOMAIN)

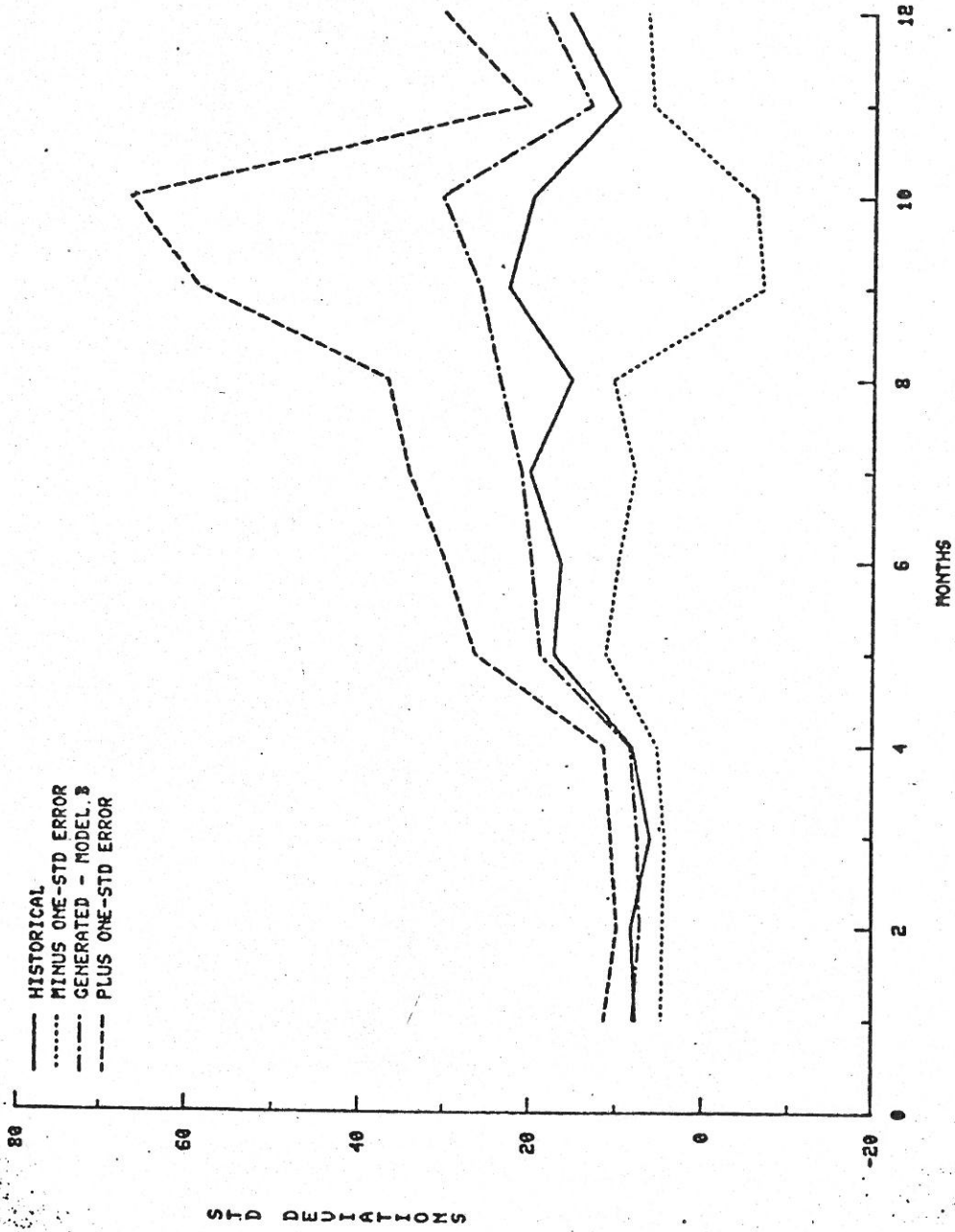


Figure 1.9.E.2. PALODE - STD DEVIATIONS (ORIGINAL DOMAIN)

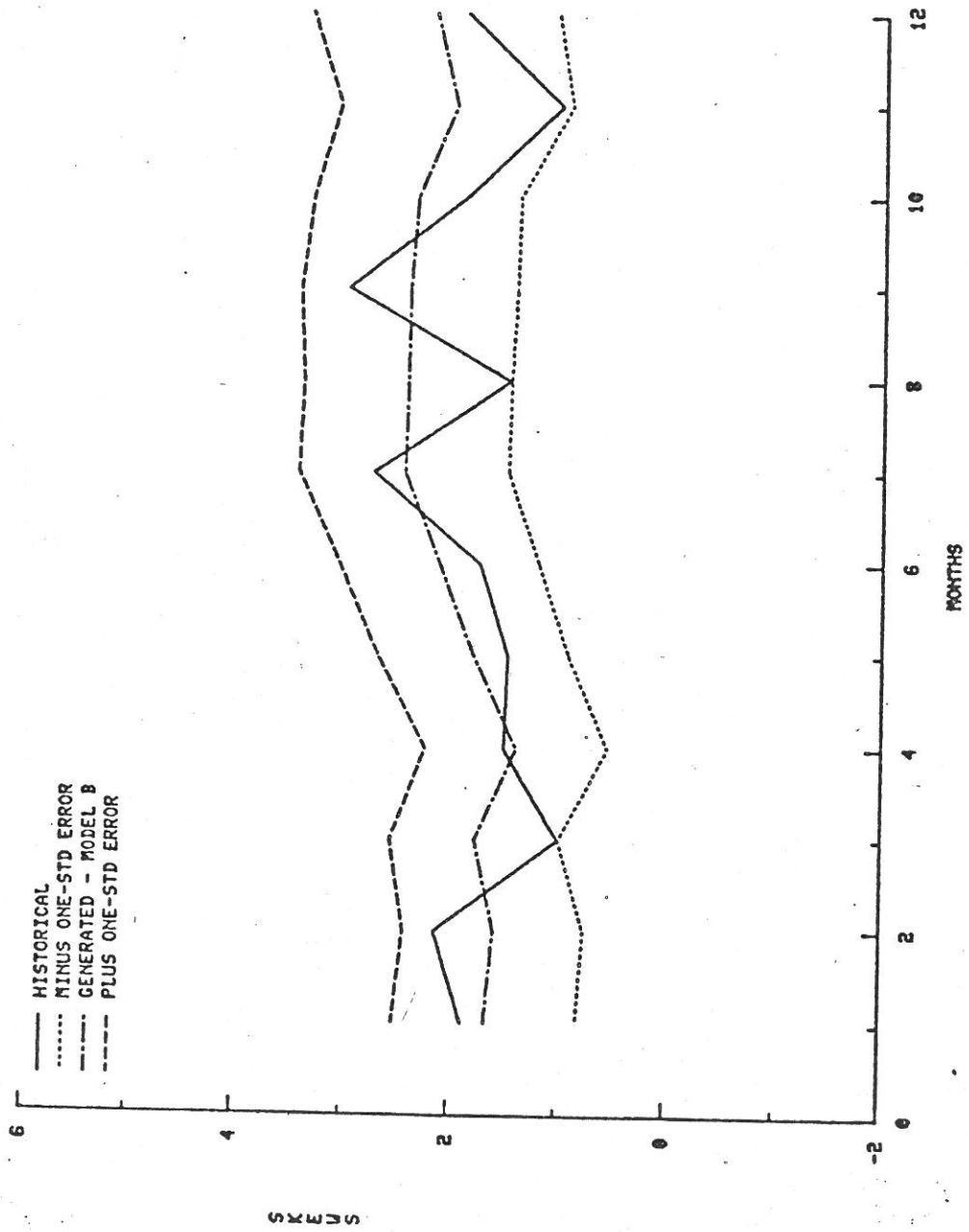


Figure 1.9.E.3. PALODE - SKEWS (ORIGINAL DOMAIN)

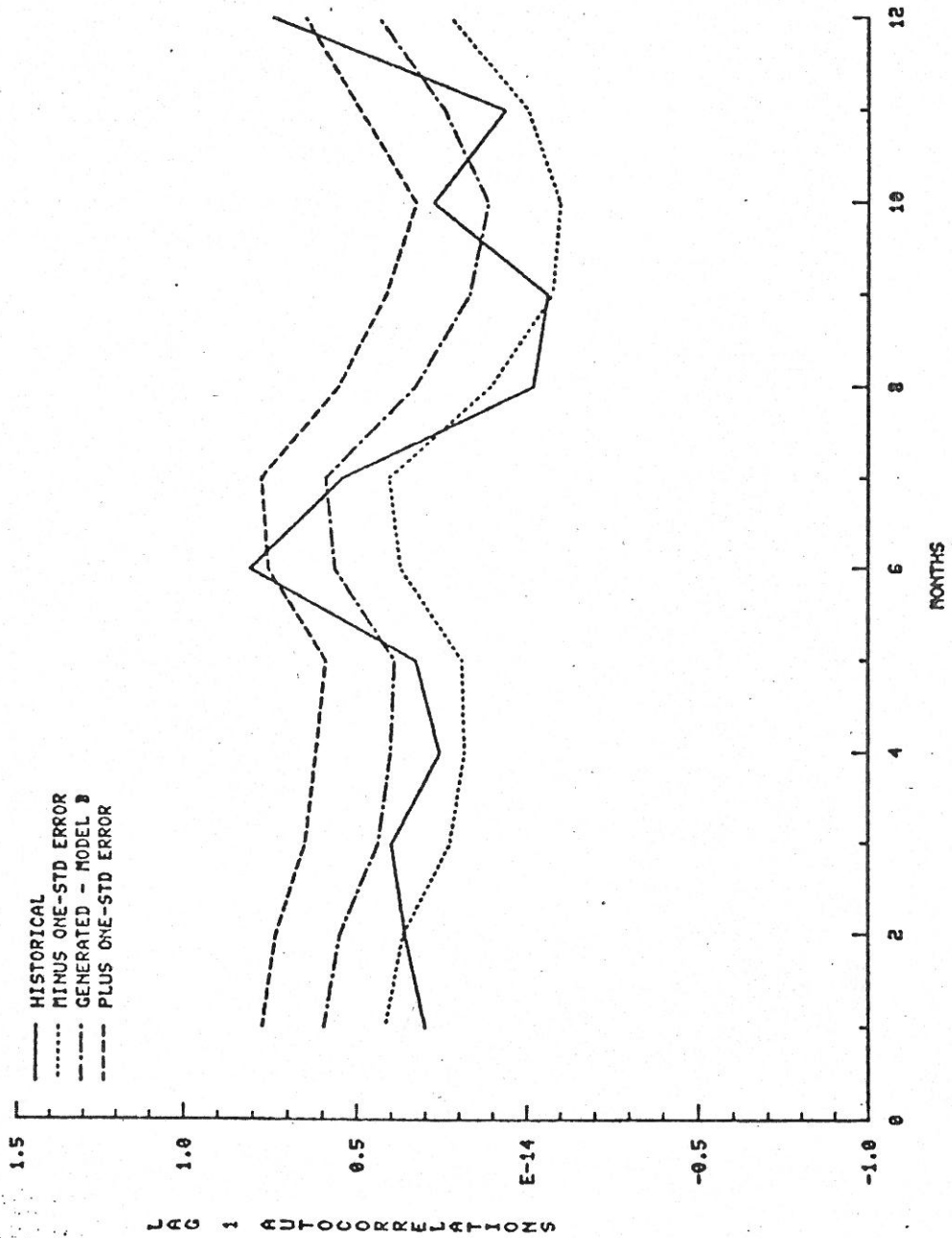


Figure 1:9.E.4. PALODE - LAG 1 AUTOCORRELATIONS (ORIGINAL DOMAIN)

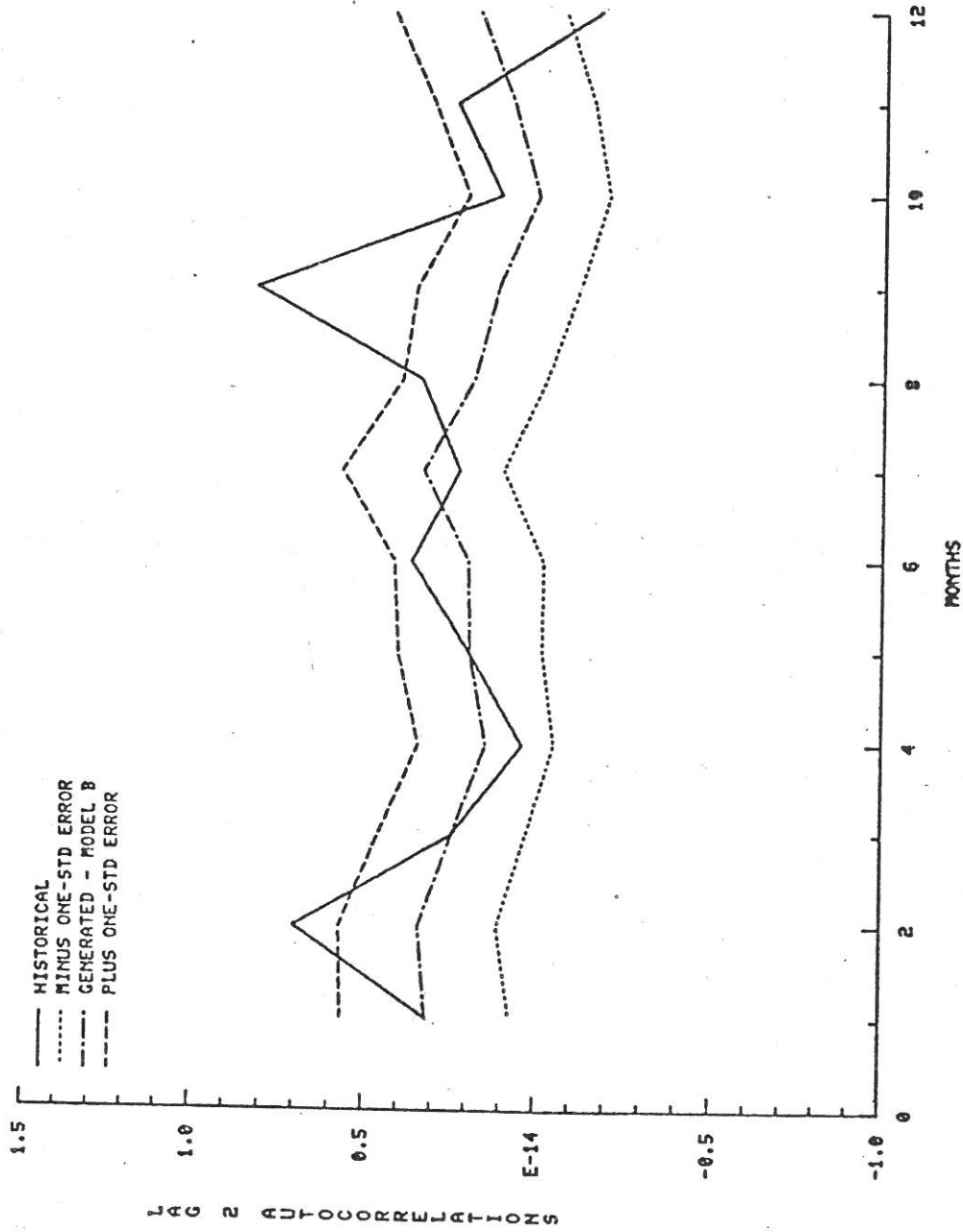


Figure 1.9.E.5. PALODE - LAG 2 AUTOCORRELATIONS (ORIGINAL DOMAIN)

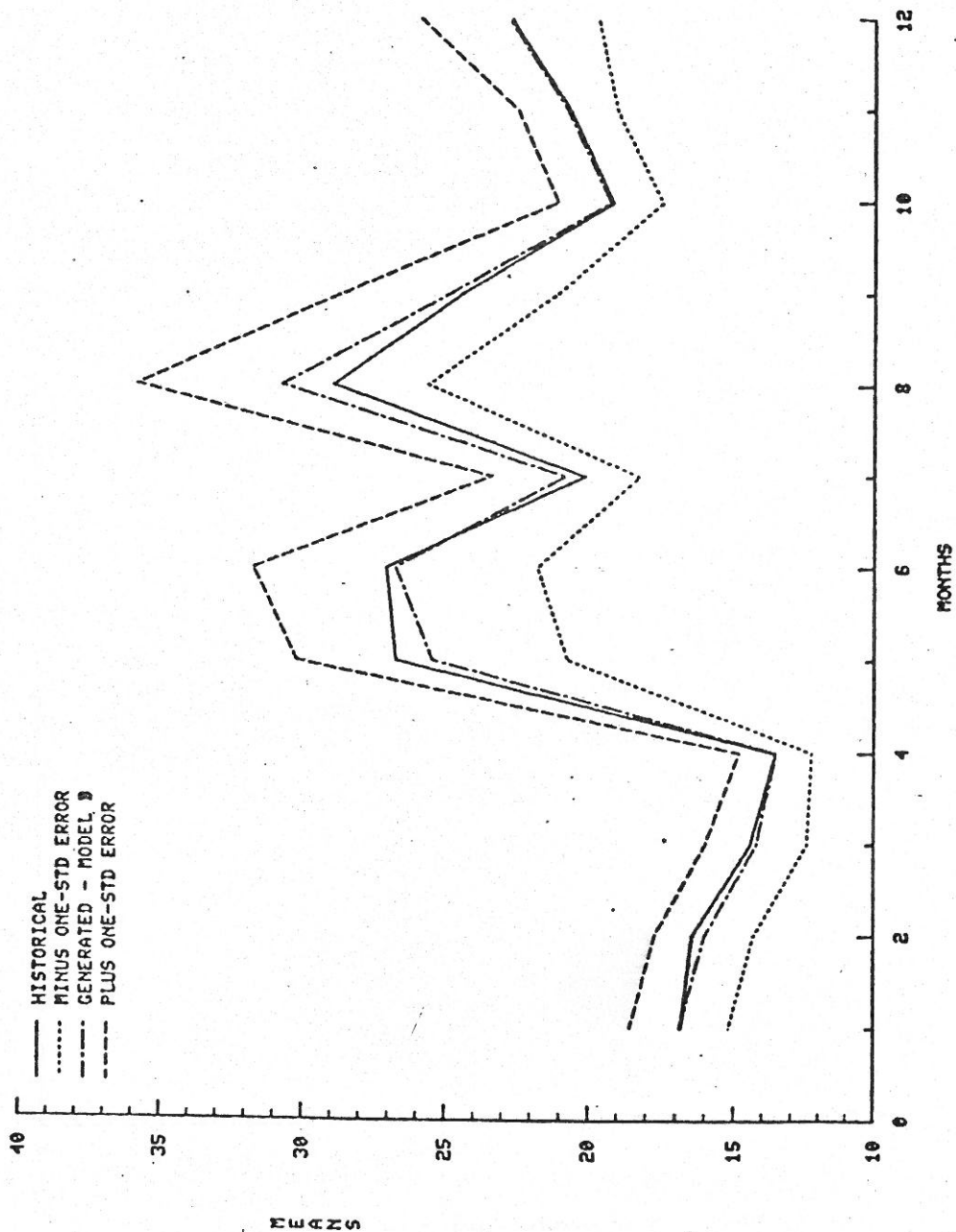


Figure 1.9.E.6. PASODE - MEANS (ORIGINAL DOMAIN)

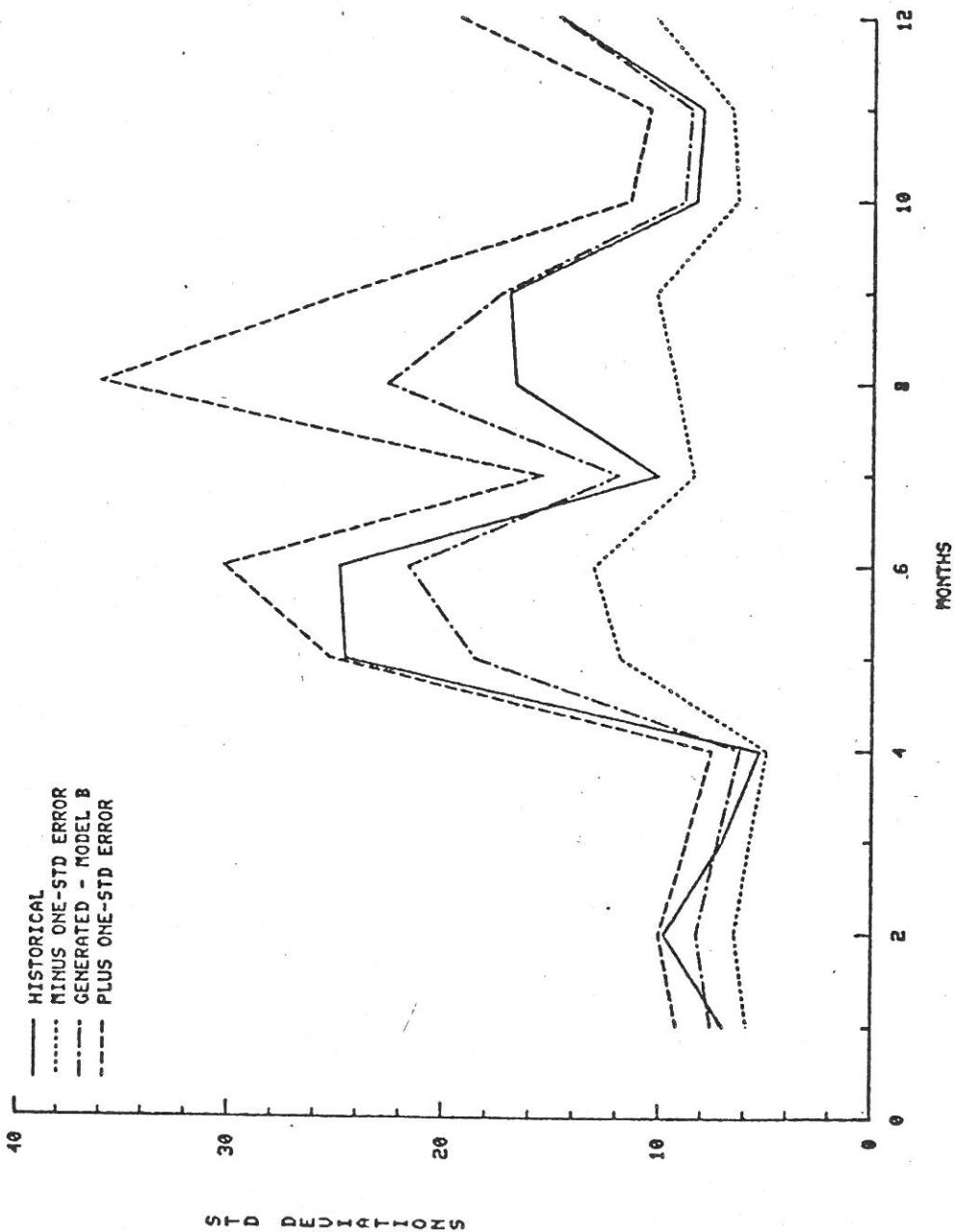


Figure 1.9.E.7. PASODE - STD DEVIATIONS (ORIGINAL DOMAIN)

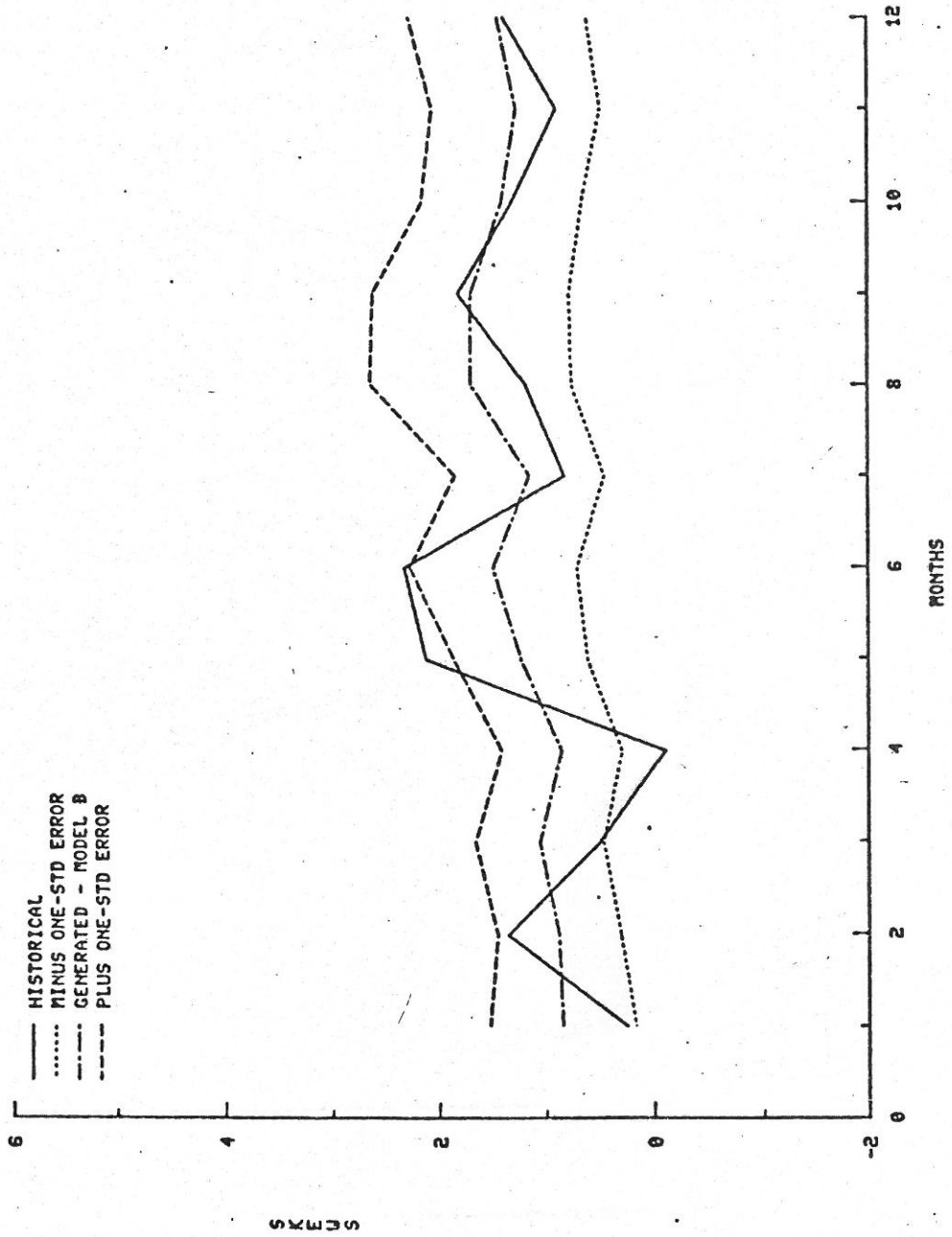


Figure 1.9.E.8. PASODE - SKEWS (ORIGINAL DOMAIN)

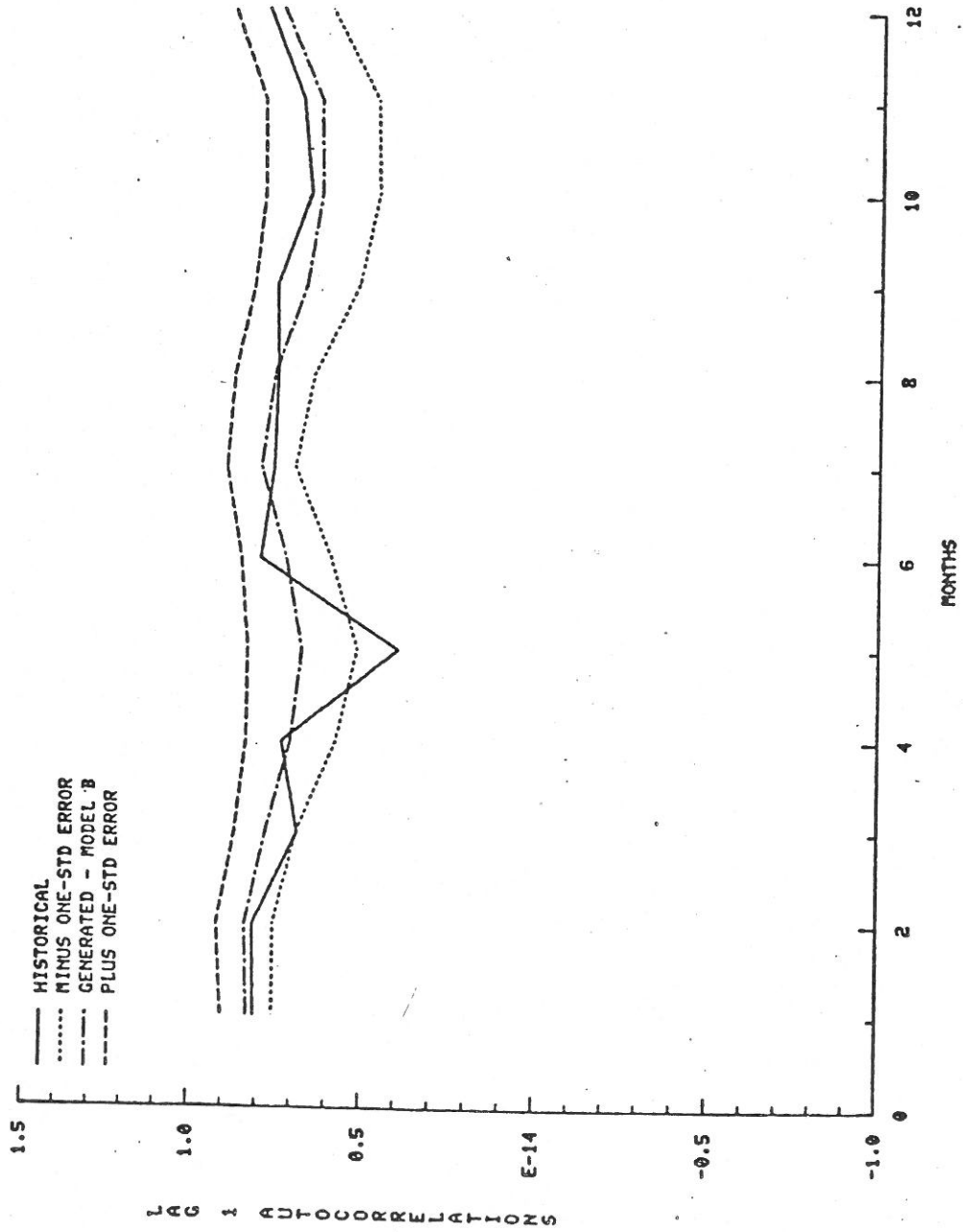


Figure 1.9.E.9. PASODE - LAG 1 AUTOCORRELATIONS (ORIGINAL DOMAIN)

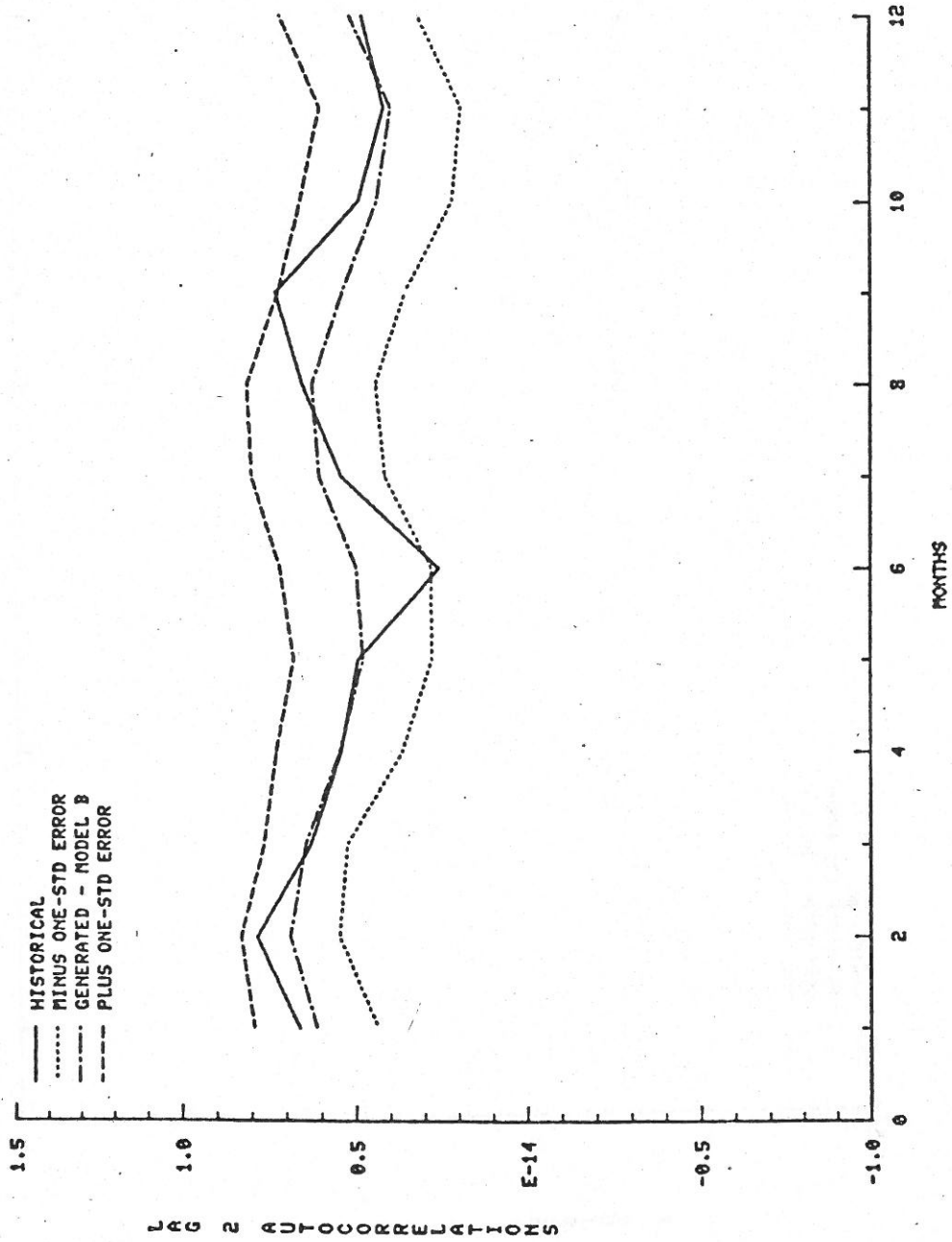


Figure 1.9.E.10. PASODE - LAG 2 AUTOCORRELATIONS (ORIGINAL DOMAIN)

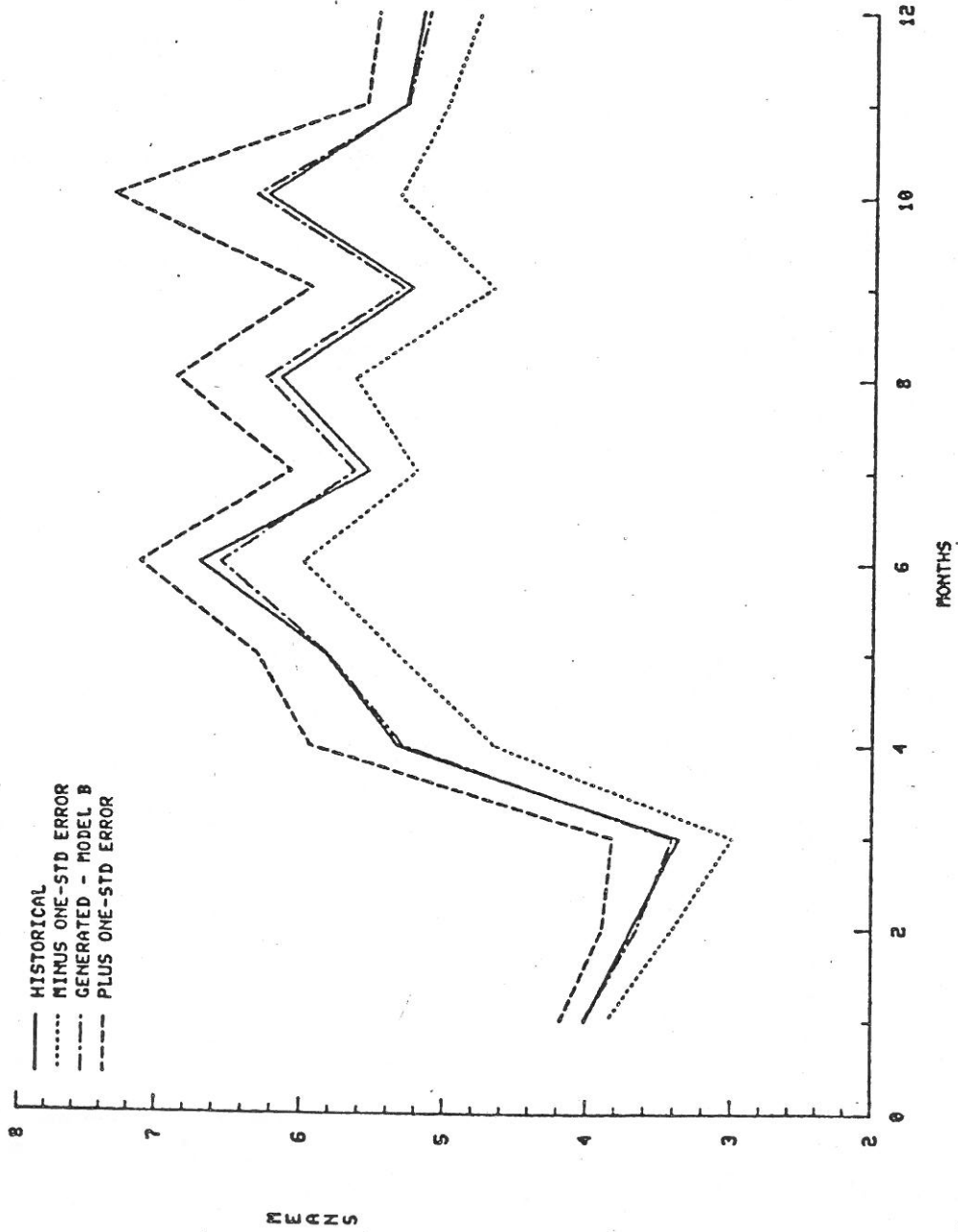


Figure 1.9.E.11. RANCHO - MEANS (ORIGINAL DOMAIN)

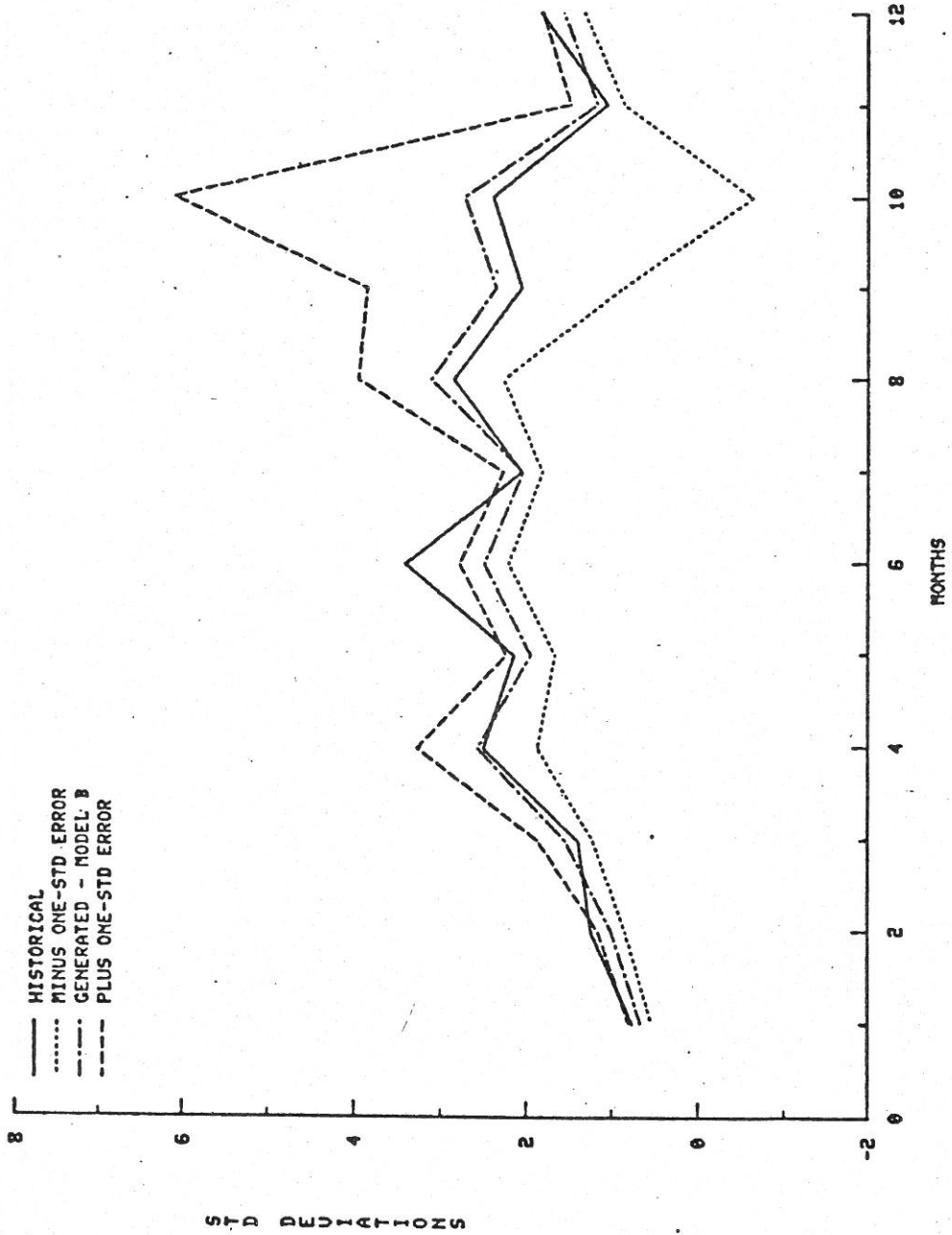


Figure 1.9.E.12. RANCHO - STD DEVIATIONS (ORIGINAL DOMAIN)

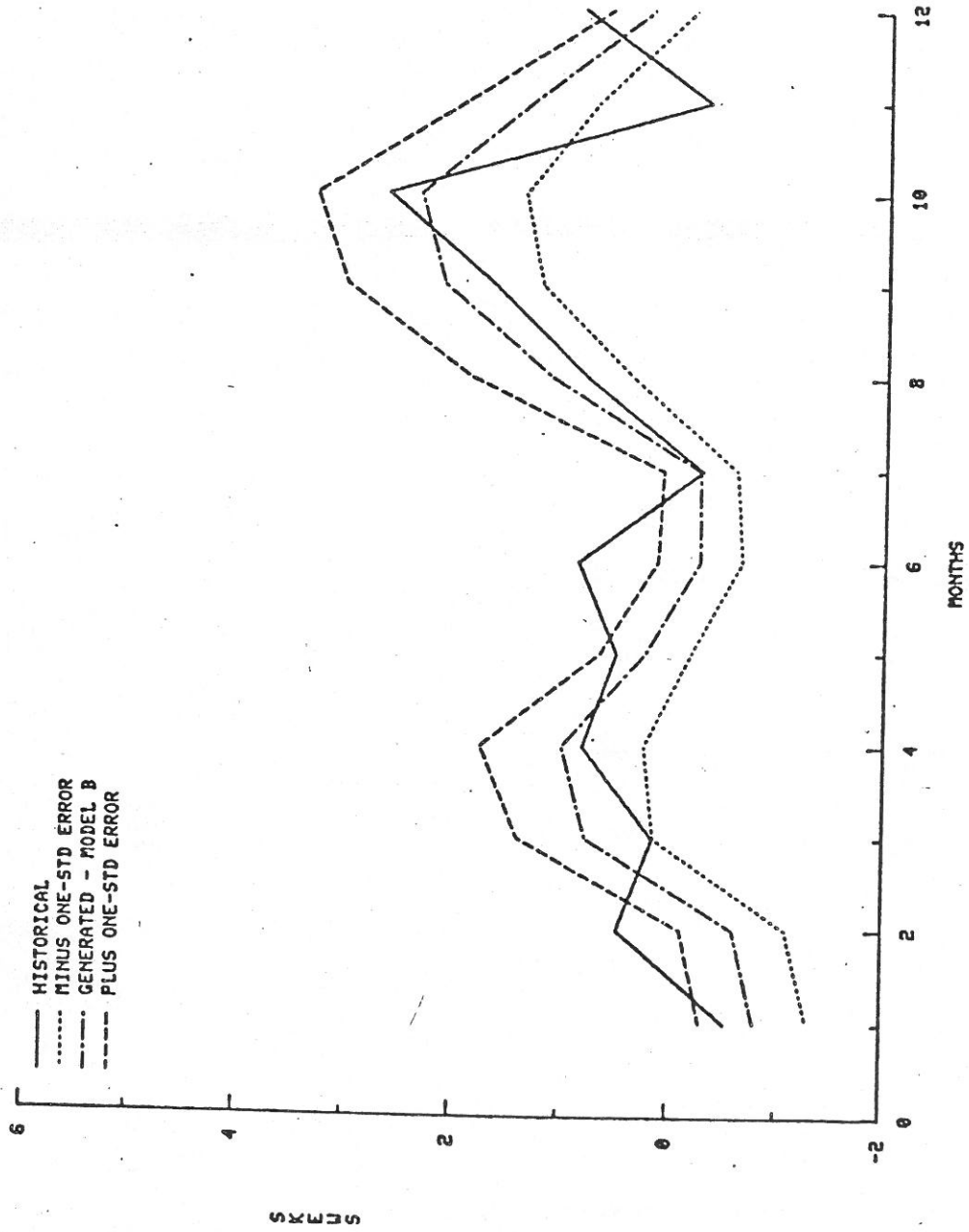


Figure 1.9.E.13. RANCHO - SKEWS (ORIGINAL DOMAIN)

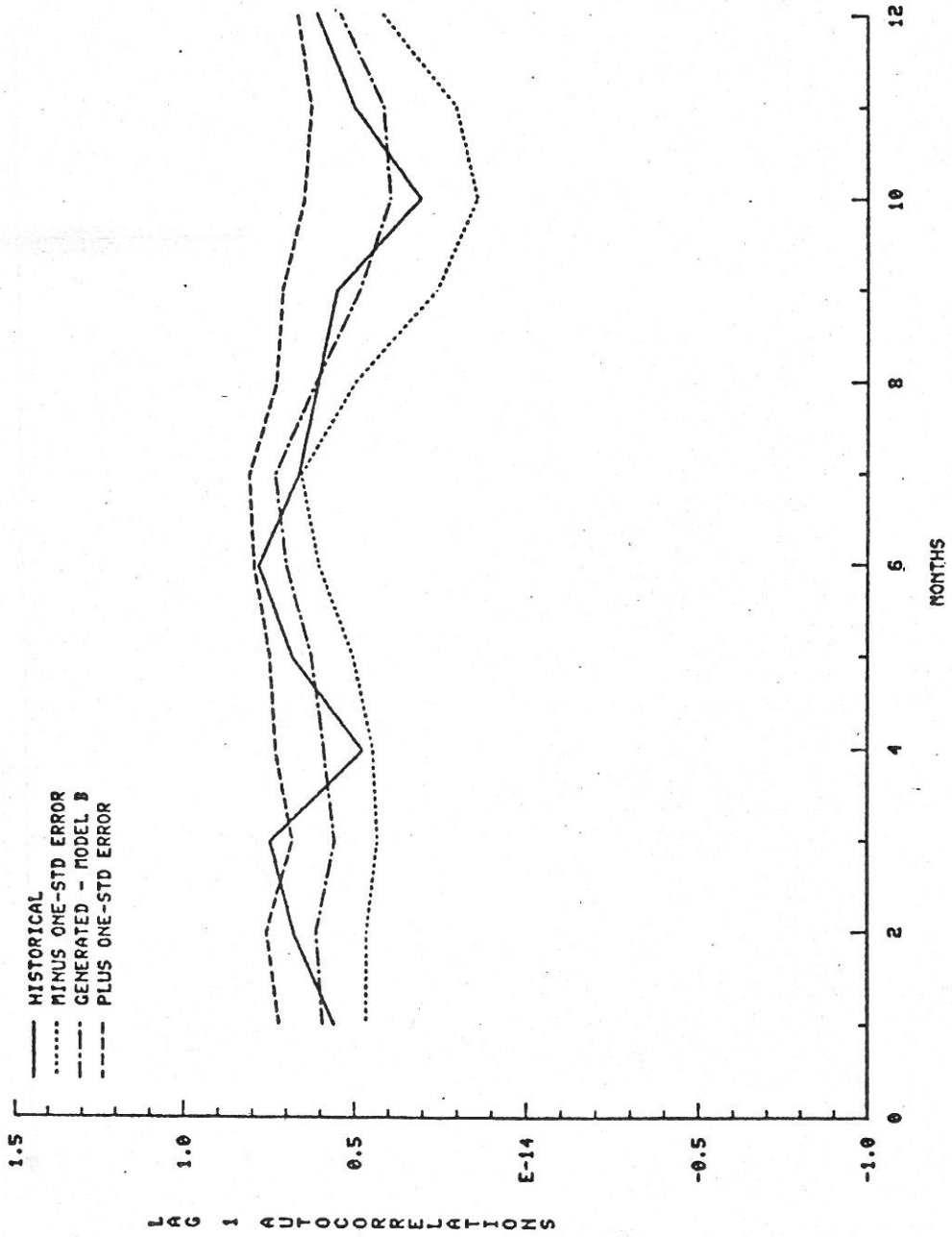


Figure 1.9.E.14. RANCHO - LAG 1 AUTOCORRELATIONS (ORIGINAL DOMAIN)

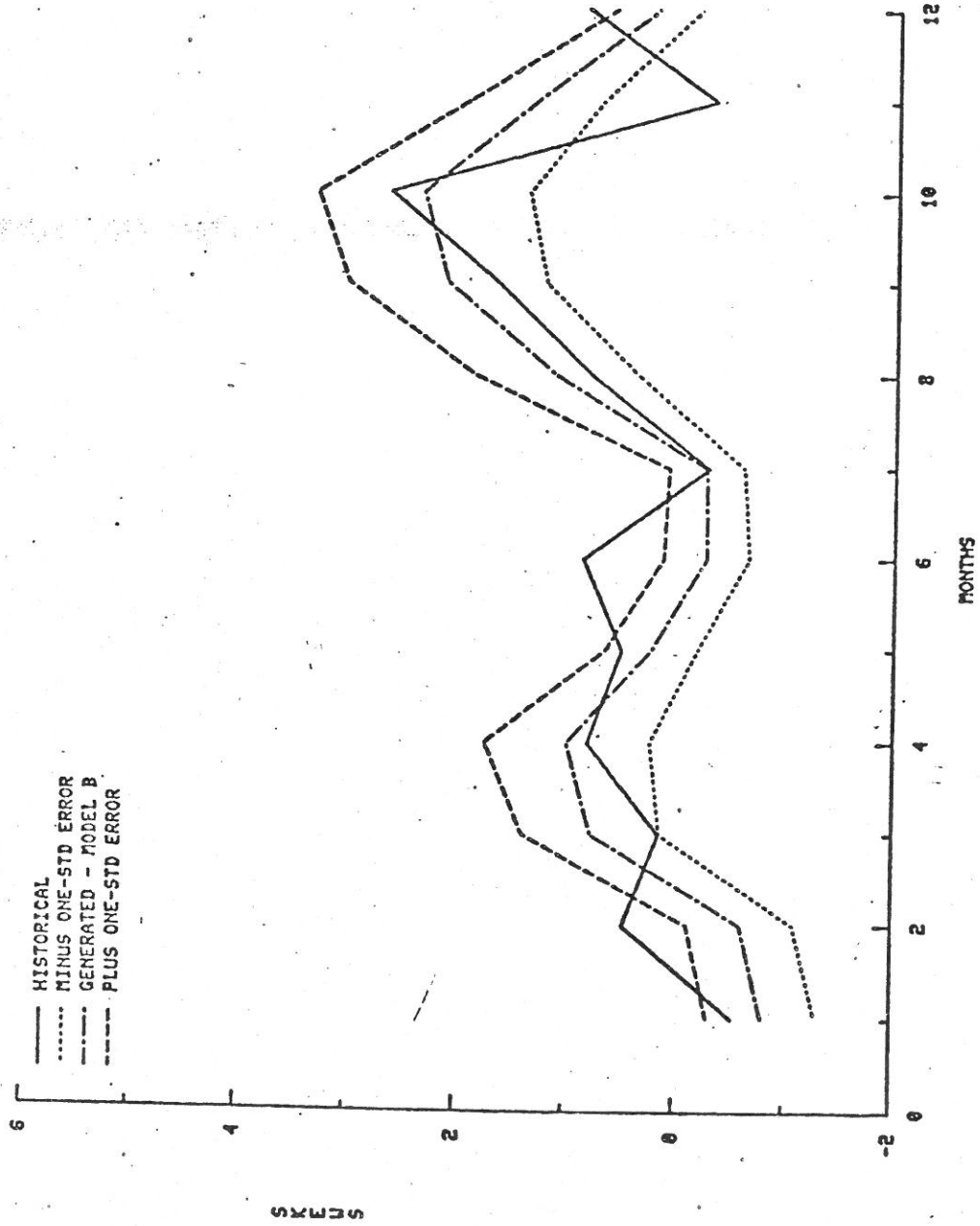


Figure 1.9.E.13. RANCHO - SKEWS (ORIGINAL DOMAIN)

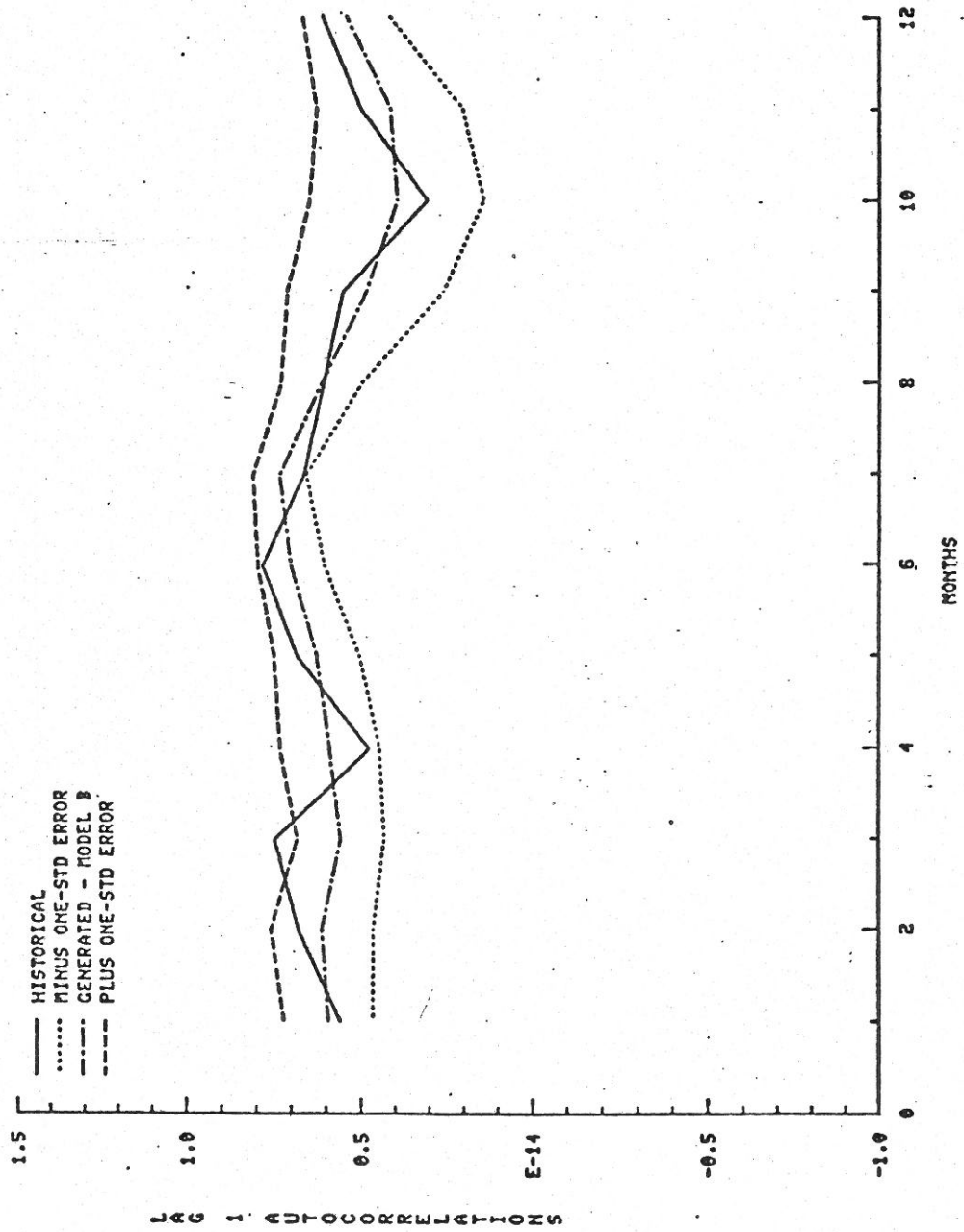


Figure 1.9.E.14. RANCHO - LAG 1 AUTOCORRELATIONS (ORIGINAL DOMAIN)

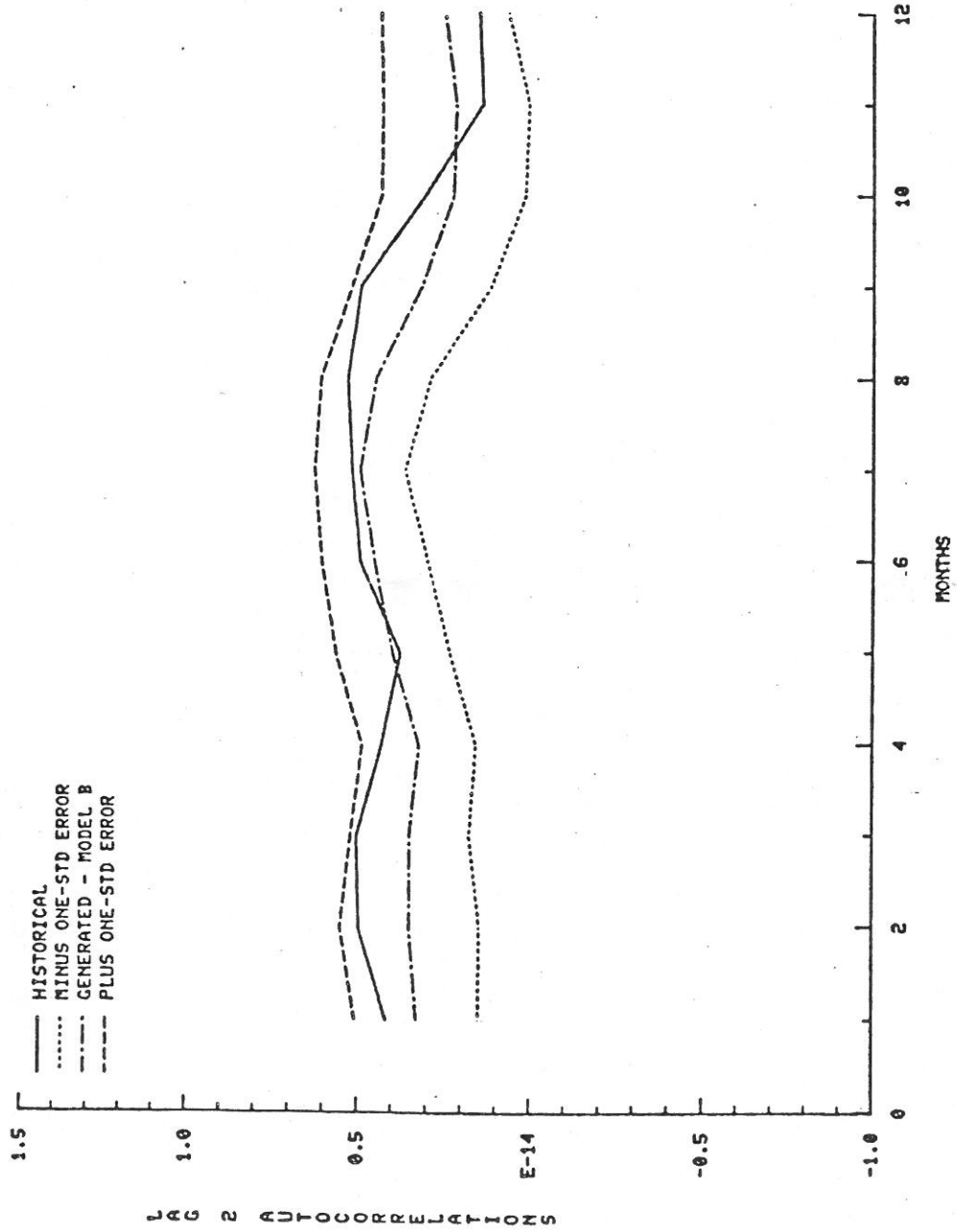


Figure 1.9.E.15. RANCHO - LAG 2 AUTOCORRELATIONS (ORIGINAL DOMAIN)

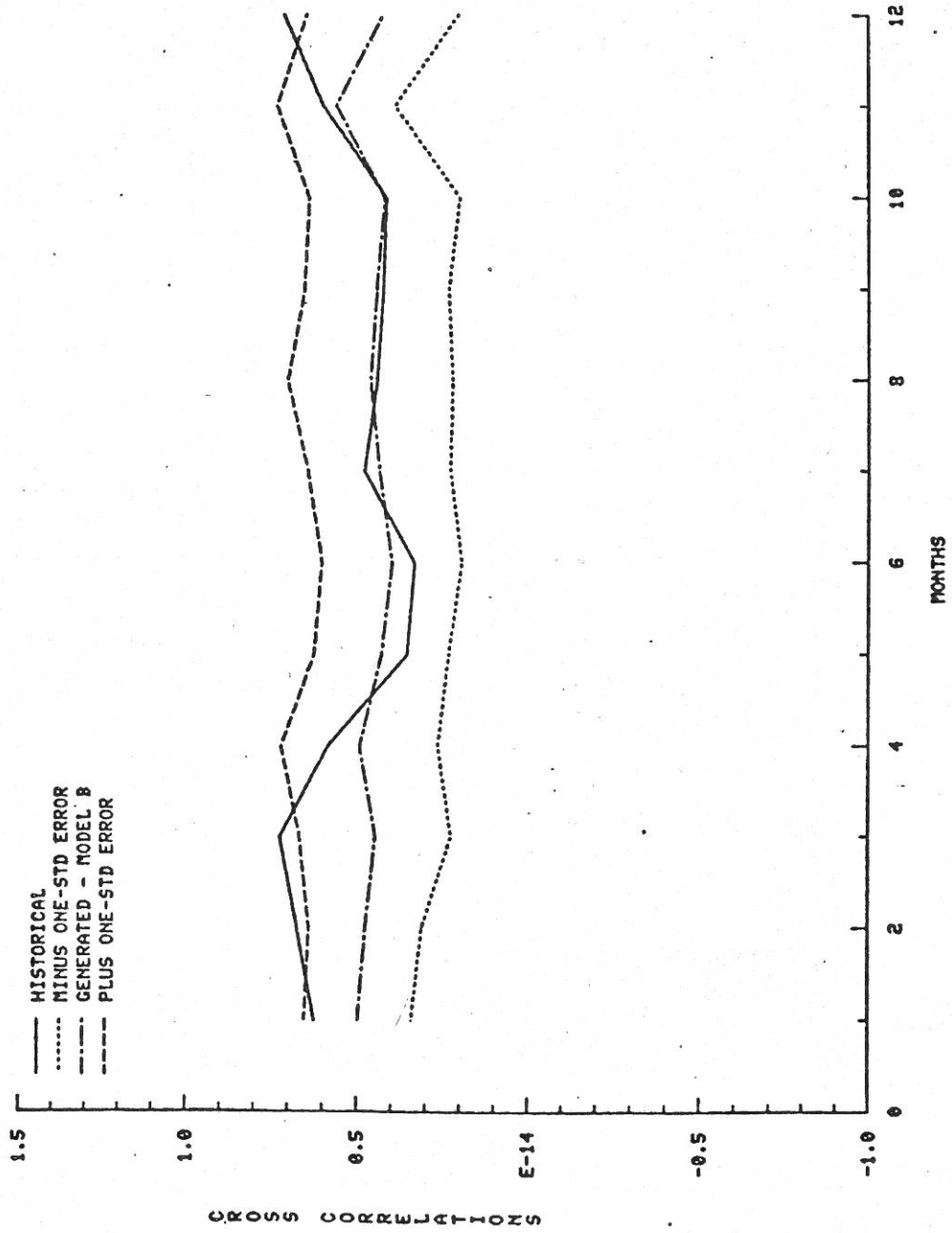


Figure 1.9.E.16. PALODE AND PASODE - CROSS CORRELATIONS (ORIGINAL DOMAIN)

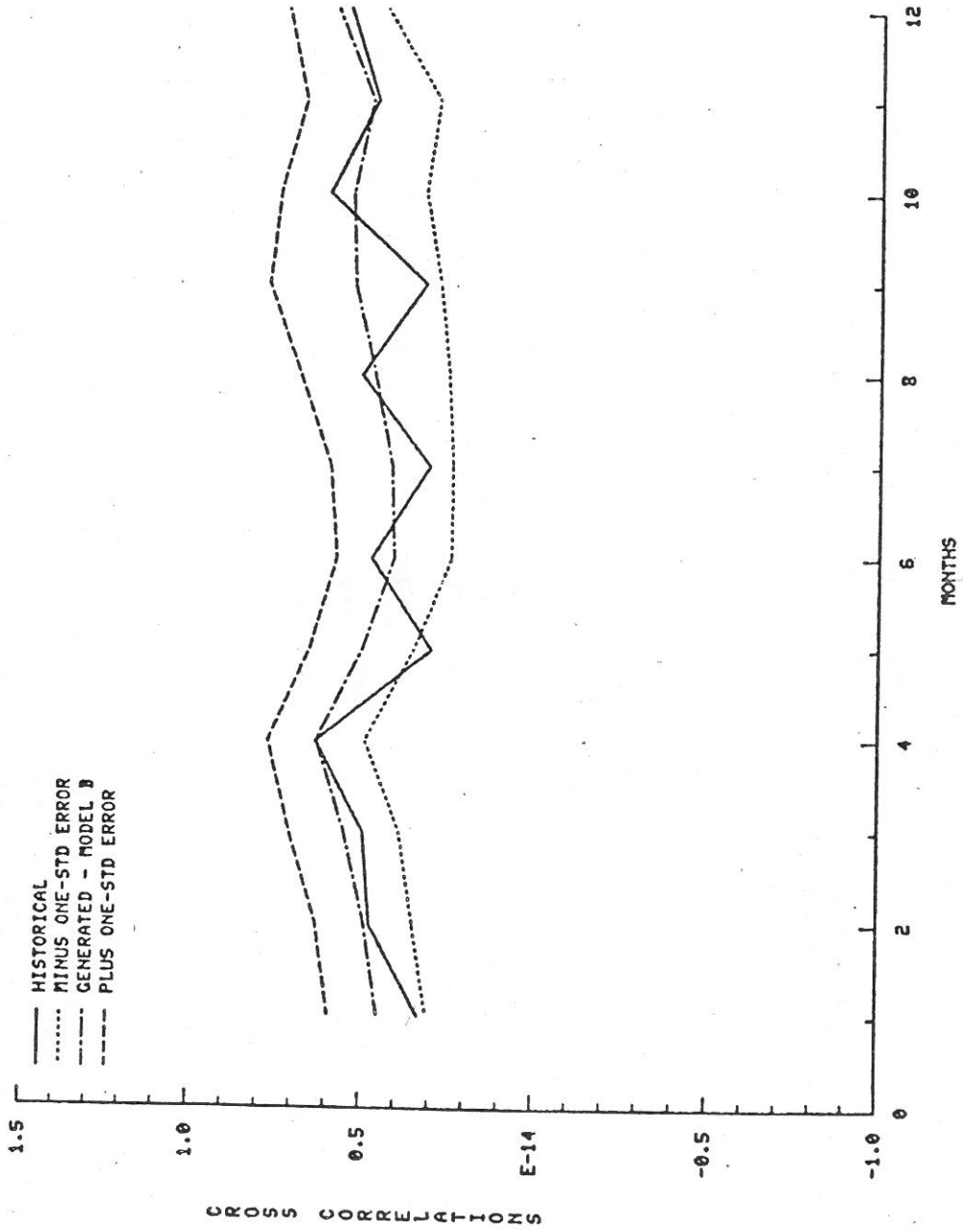


Figure 1.9.E.17. PALODE AND RANCHO - CROSS CORRELATIONS (ORIGINAL DOMAIN)

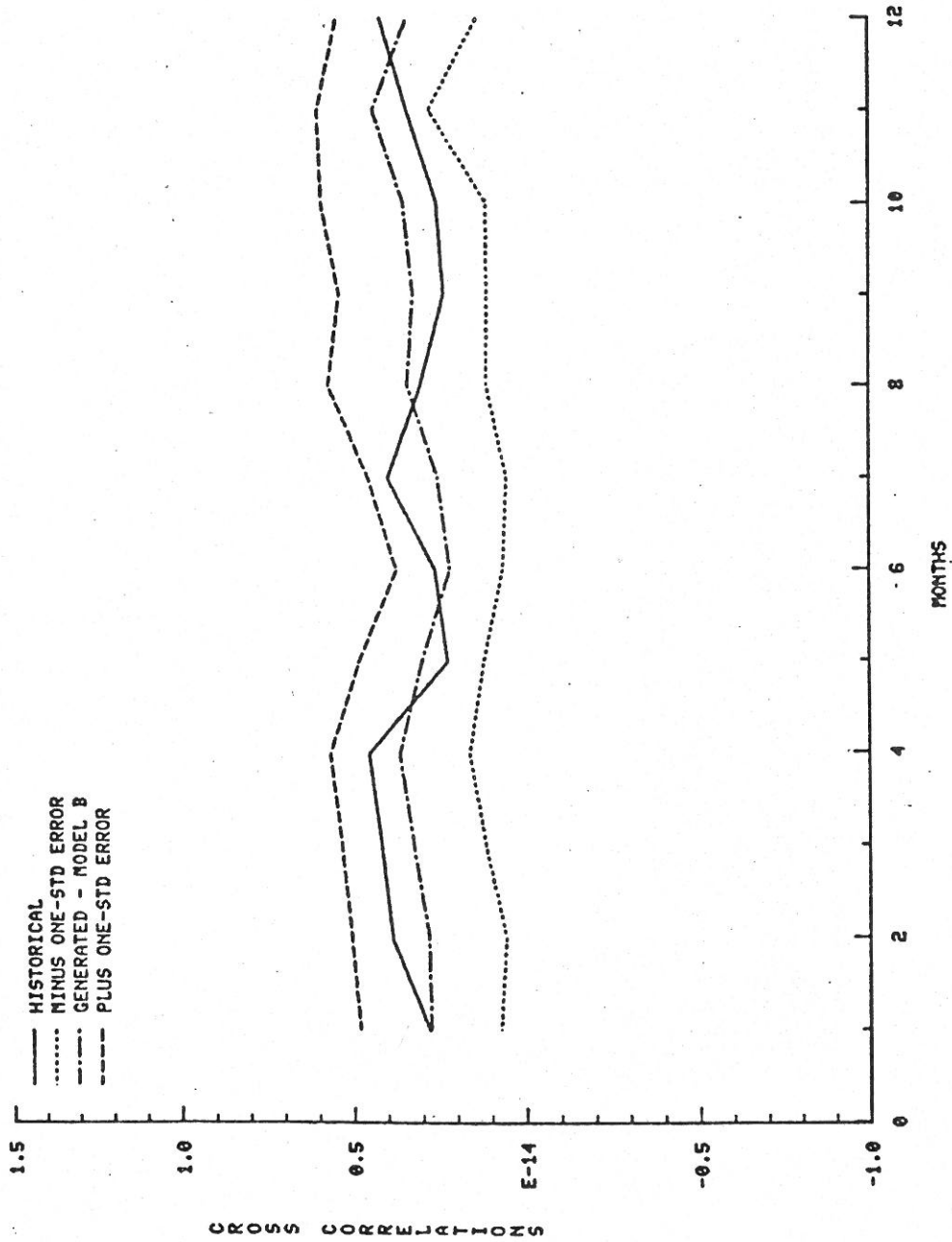


Figure 1.9.E.18. PASODE AND RANCHO - CROSS CORRELATIONS (ORIGINAL DOMAIN)

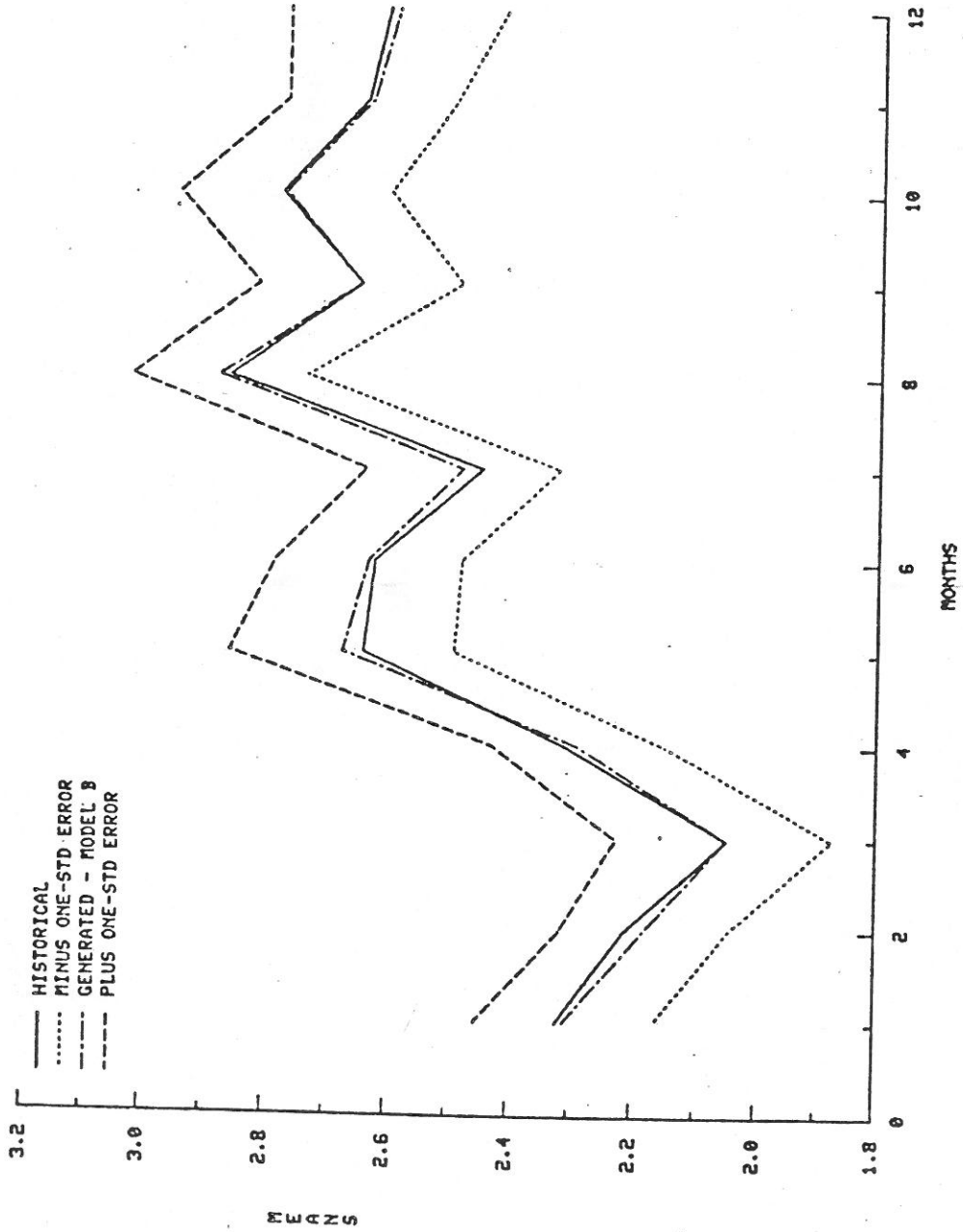


Figure 1.9.E.19. PALODE - MEANS (LOG DOMAIN)

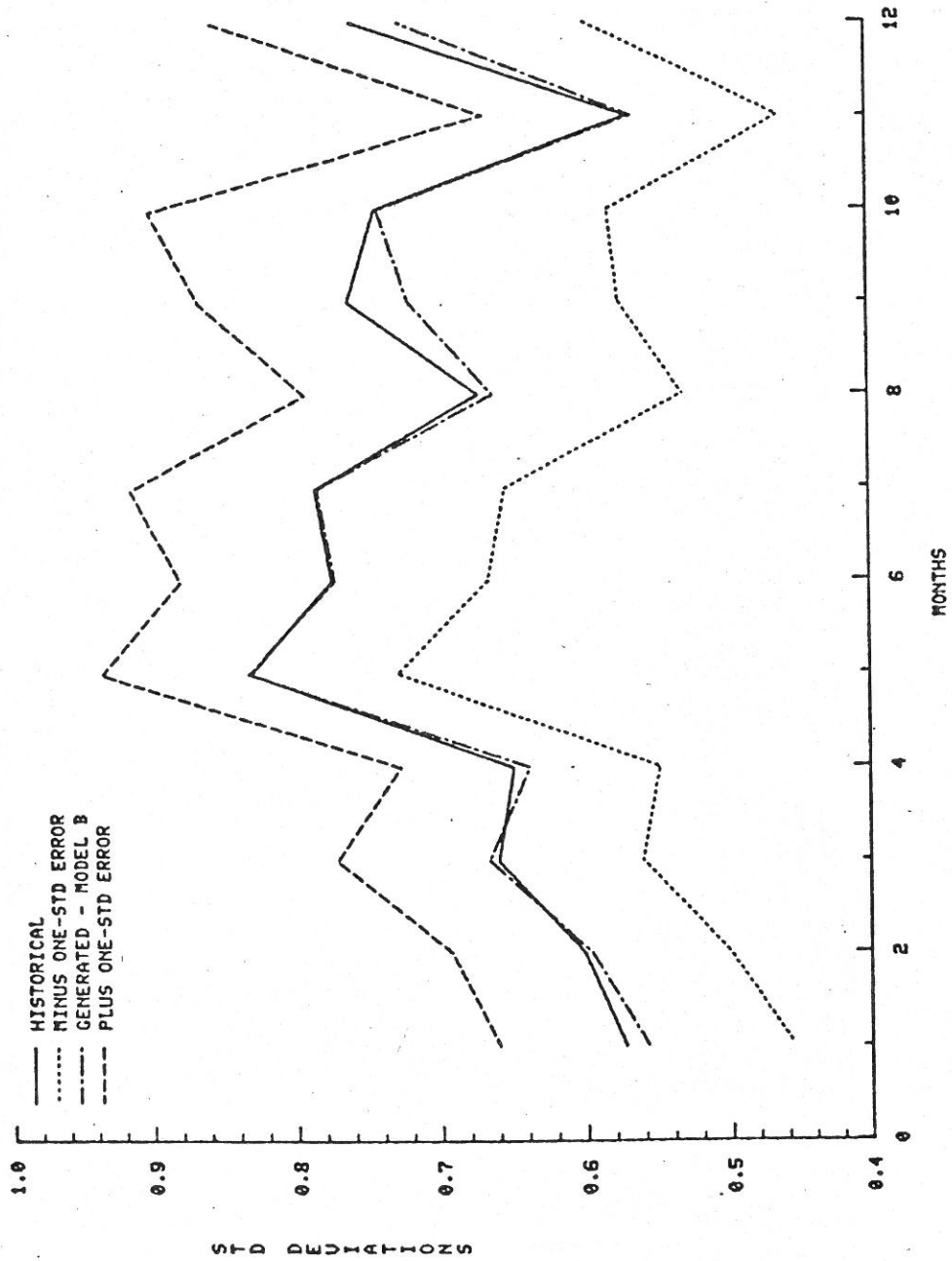


Figure 1.9.E.20. PALODE - STD DEVIATIONS (LOG DOMAIN)

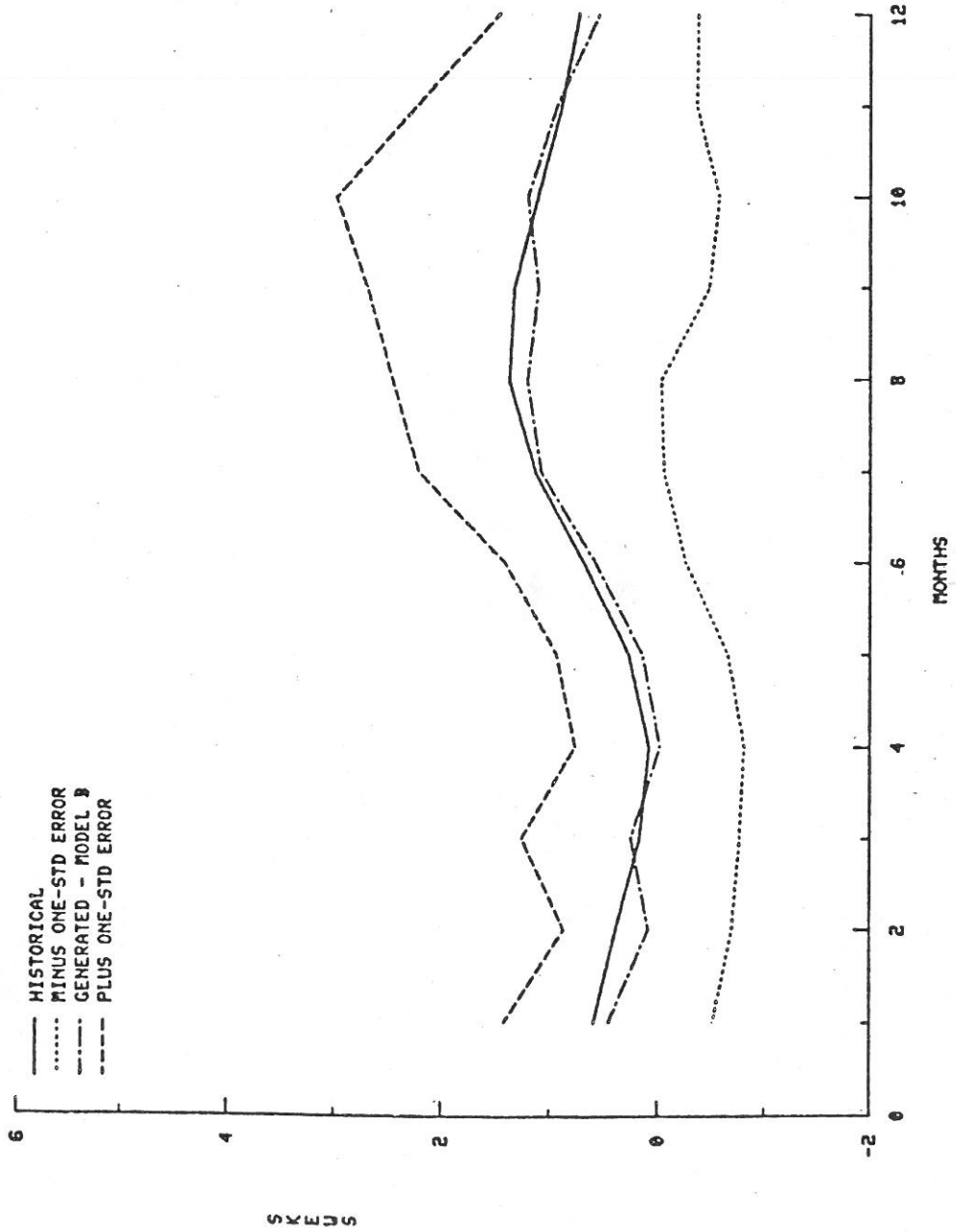


Figure 1.9.E.21. PALODE - SKEWS (LOG DOMAIN)

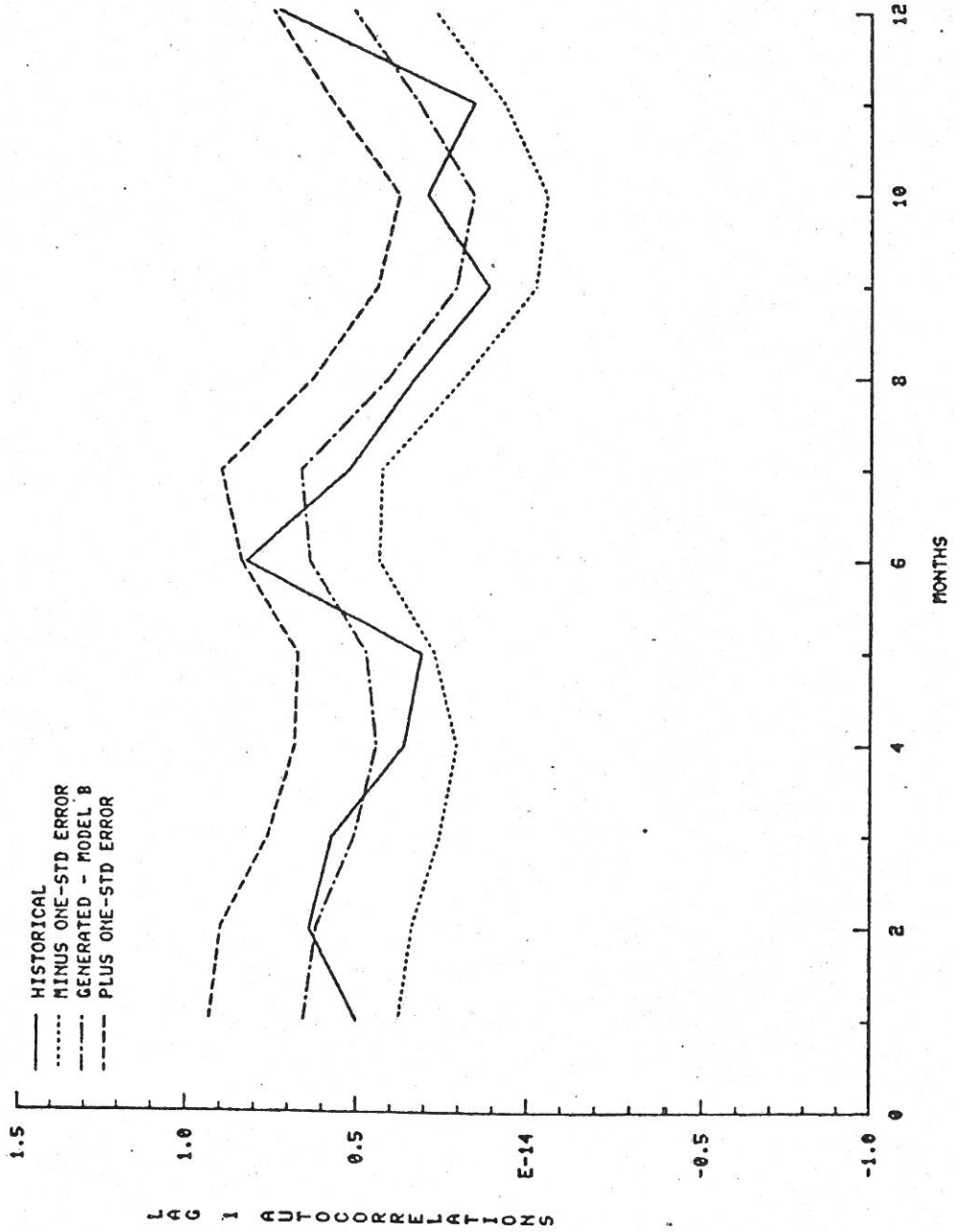


Figure 1.9.E.22. PALODE - LAG 1 AUTOCORRELATIONS (LOG DOMAIN)

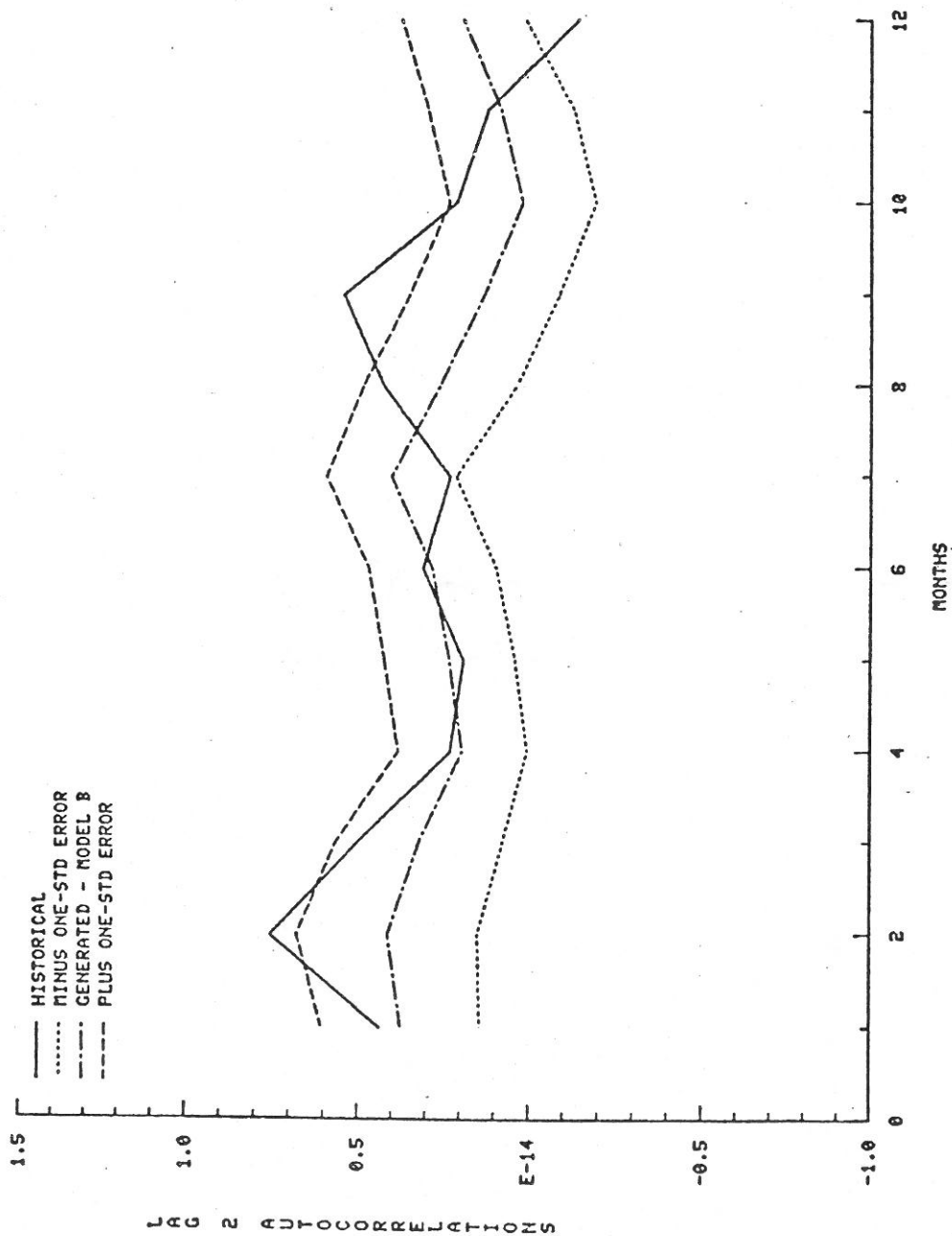


Figure 1.9.E.23. PALODE - LAG 2 AUTOCORRELATIONS (LOG DOMAIN)

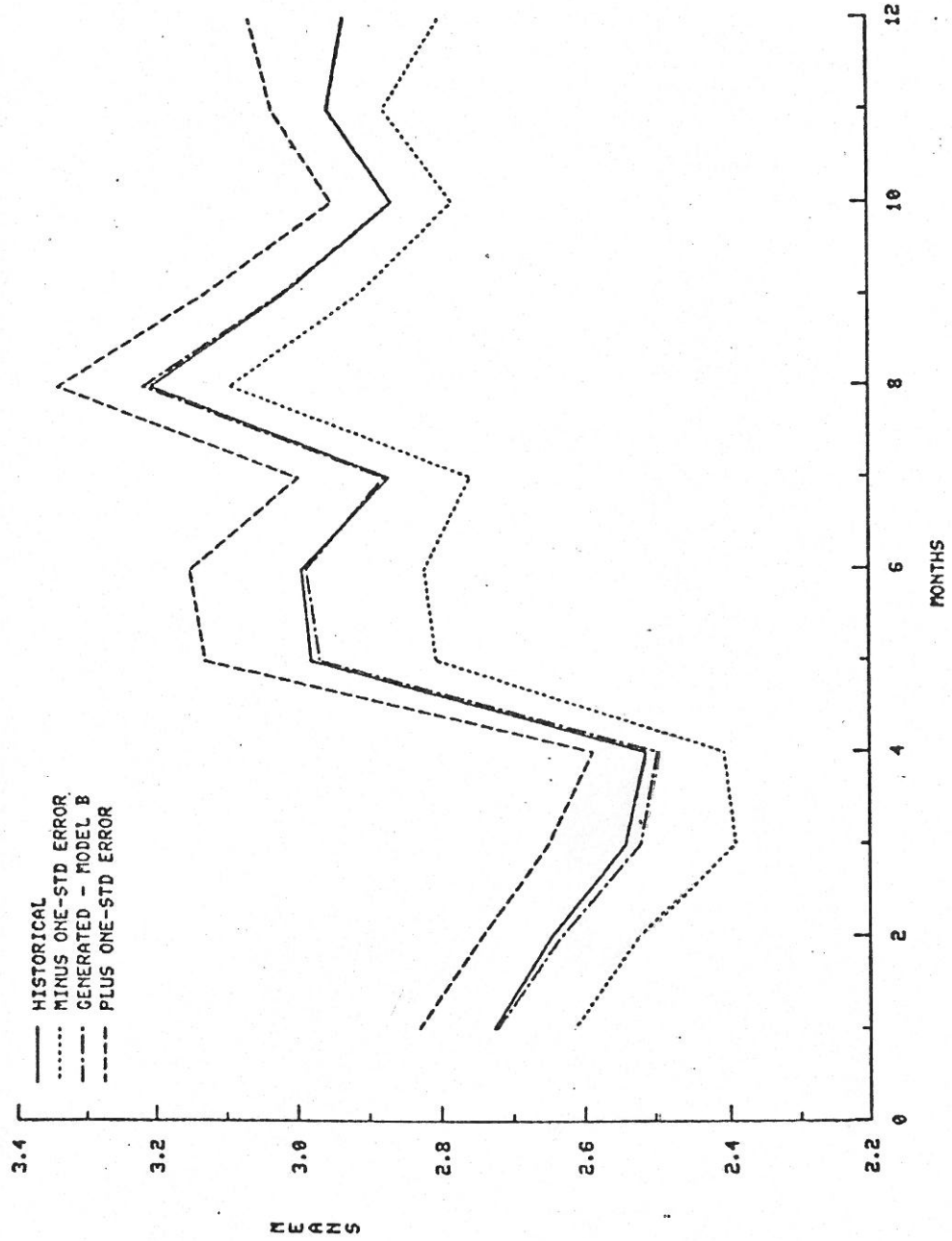


Figure 1.9.E.24. PASODE - MEANS (LOG DOMAIN)

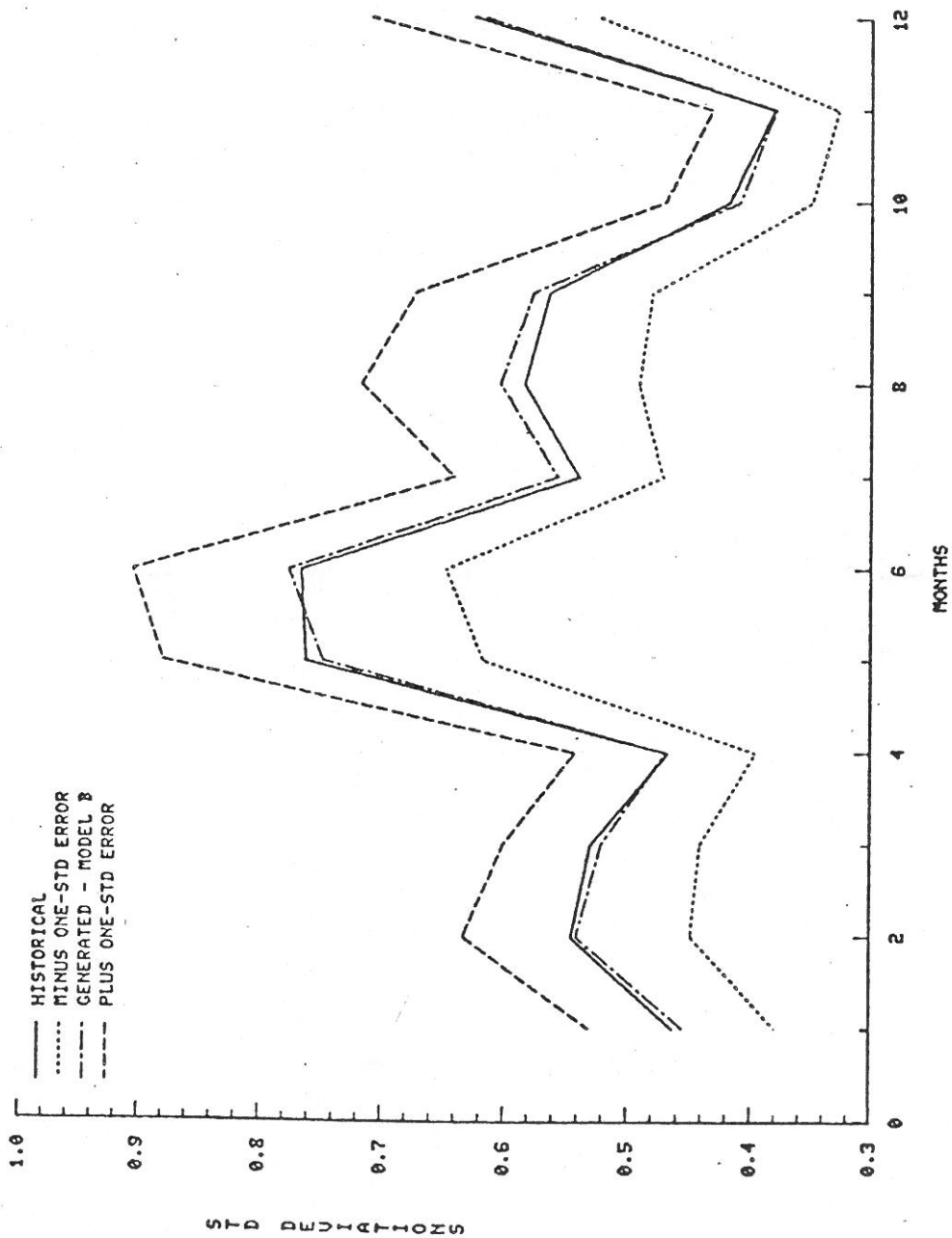


Figure 1.9.E.25. PASODE - STD DEVIATIONS (LOG DOMAIN)

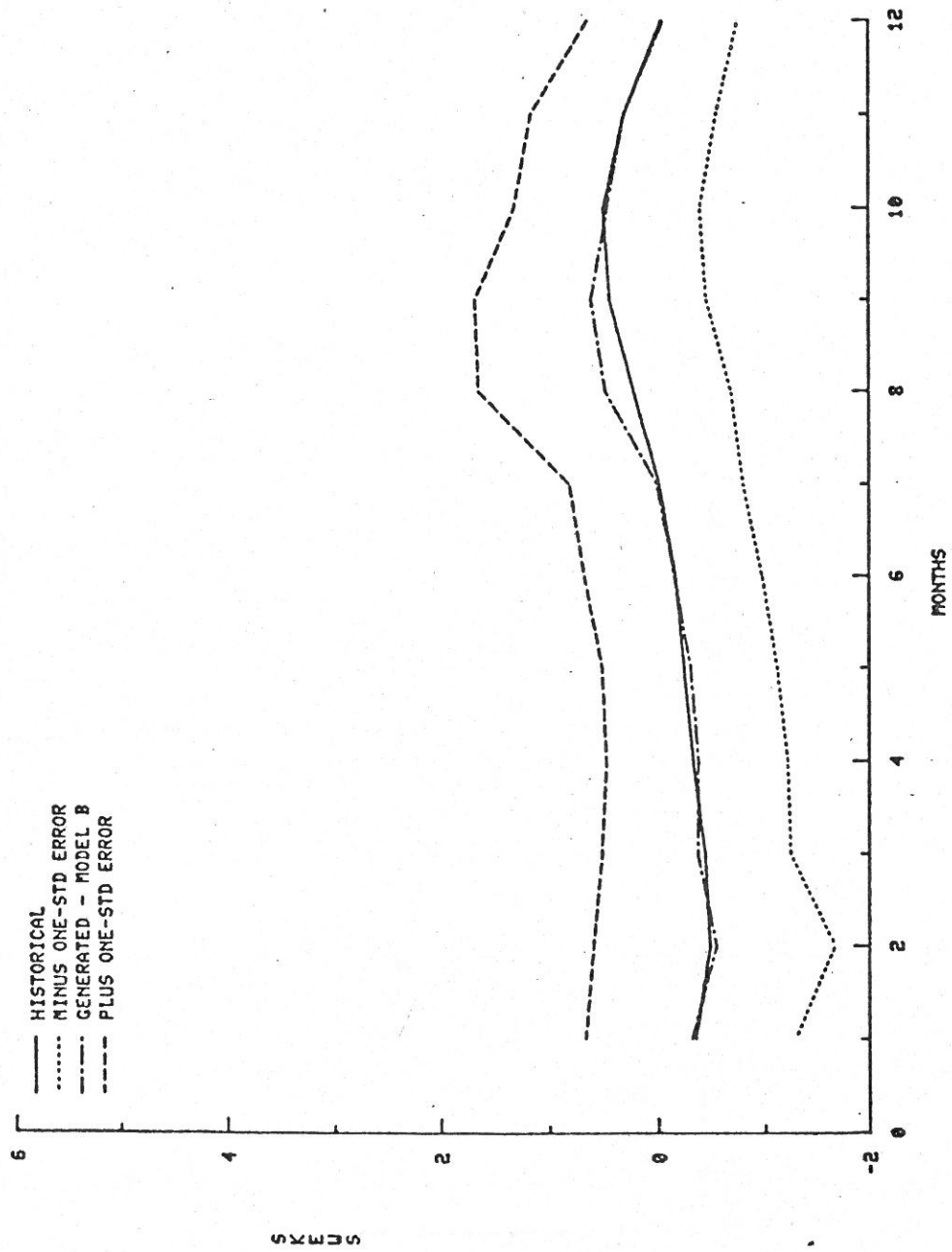


Figure 1.9.E.26. PASODE - SKEWS (LOG DOMAIN)

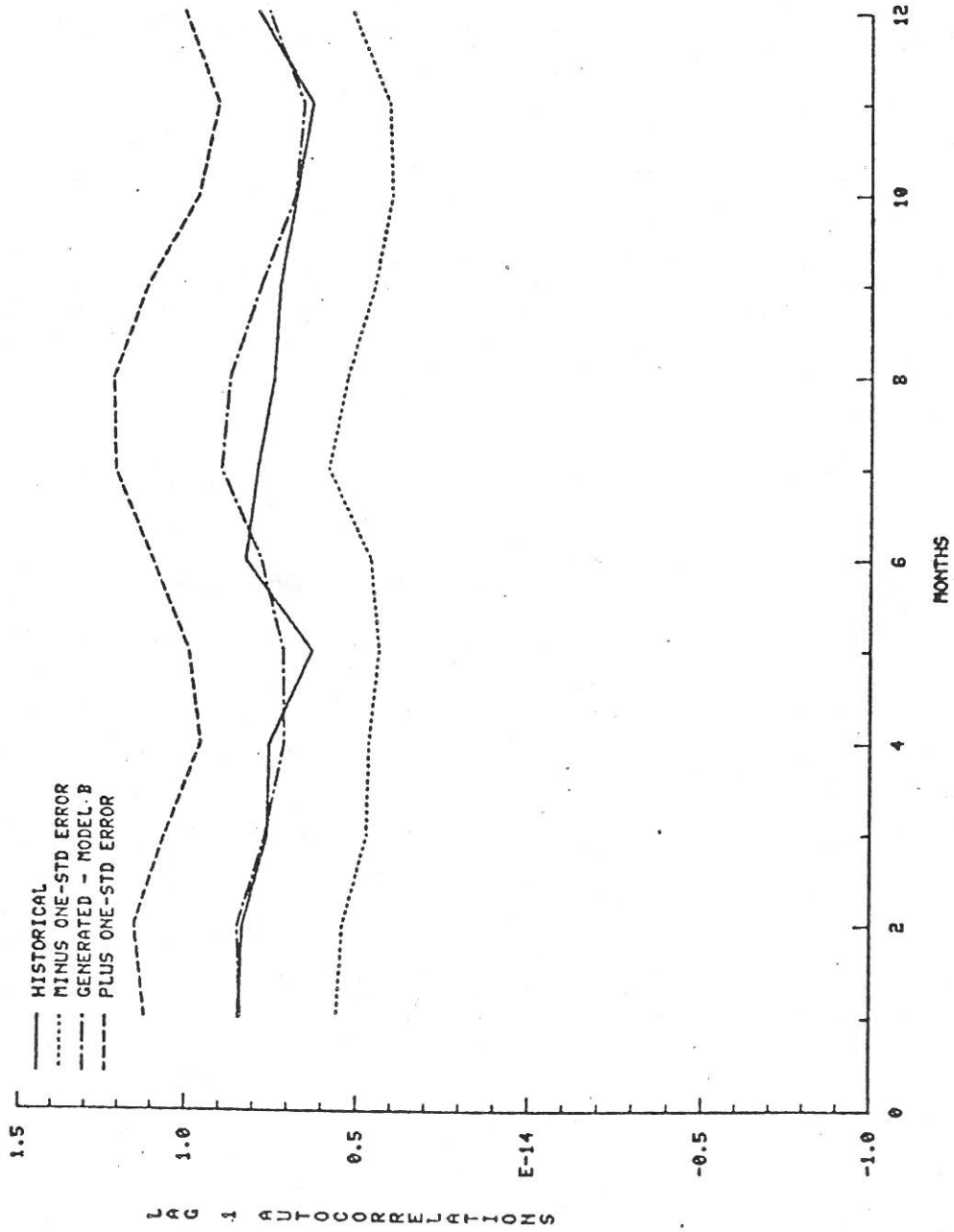


Figure 1.9.E.27. PASODE - LAG 1 AUTOCORRELATIONS (LOG DOMAIN)

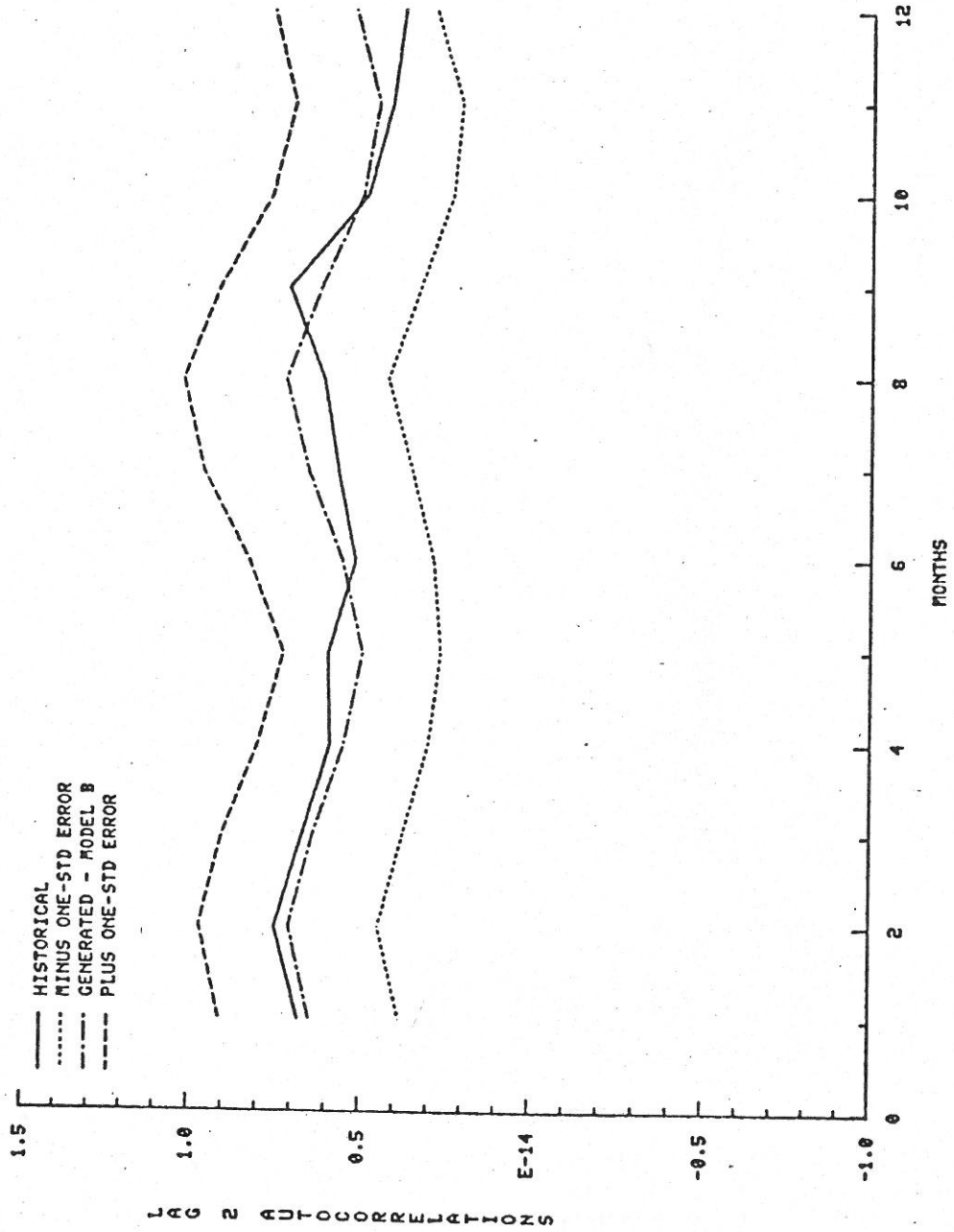


Figure 1.9.E.28. PASODE - LAG 2 AUTOCORRELATIONS (LOG DOMAIN)

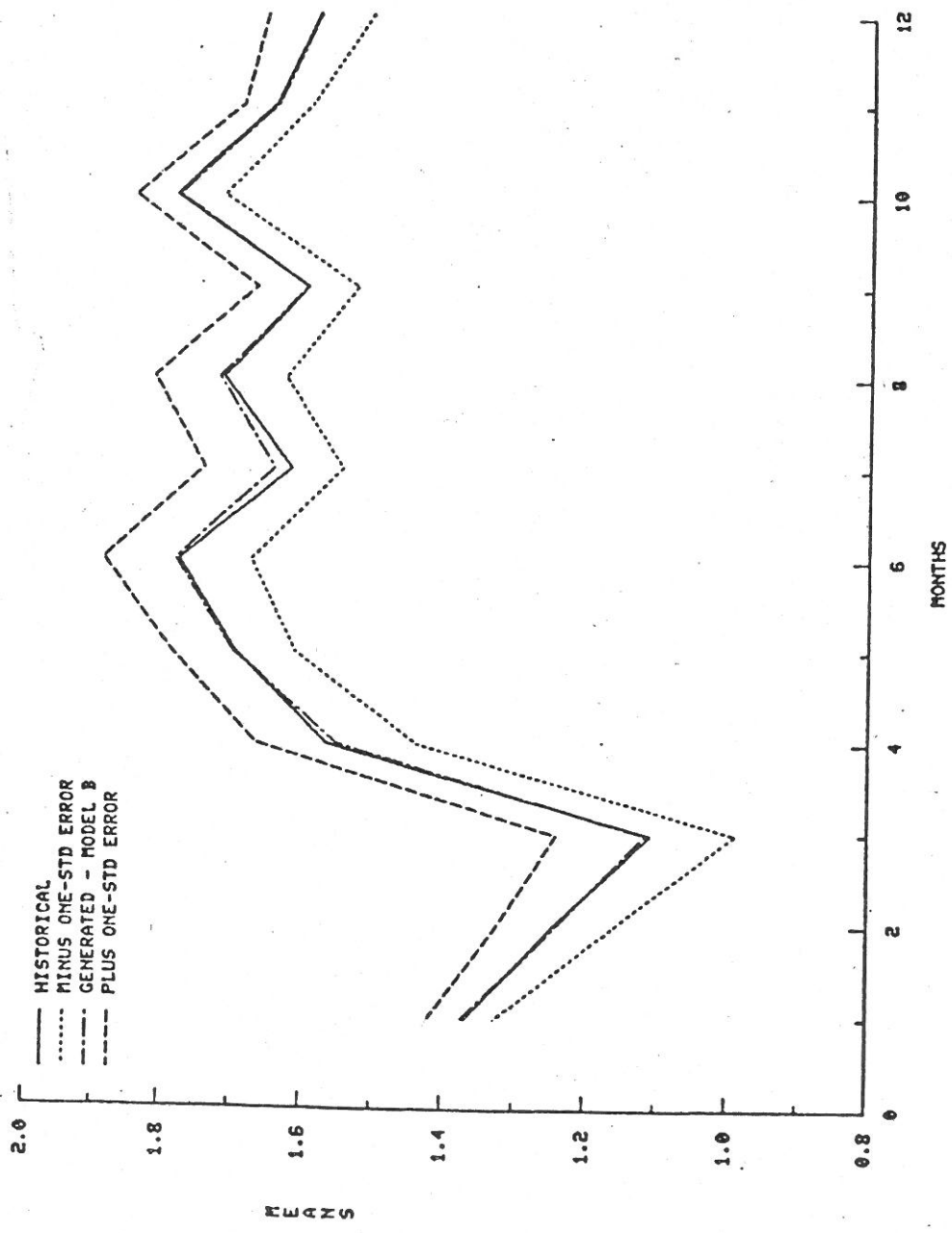


Figure 1.9.E.29. RANCHO - MEANS (LOG DOMAIN)

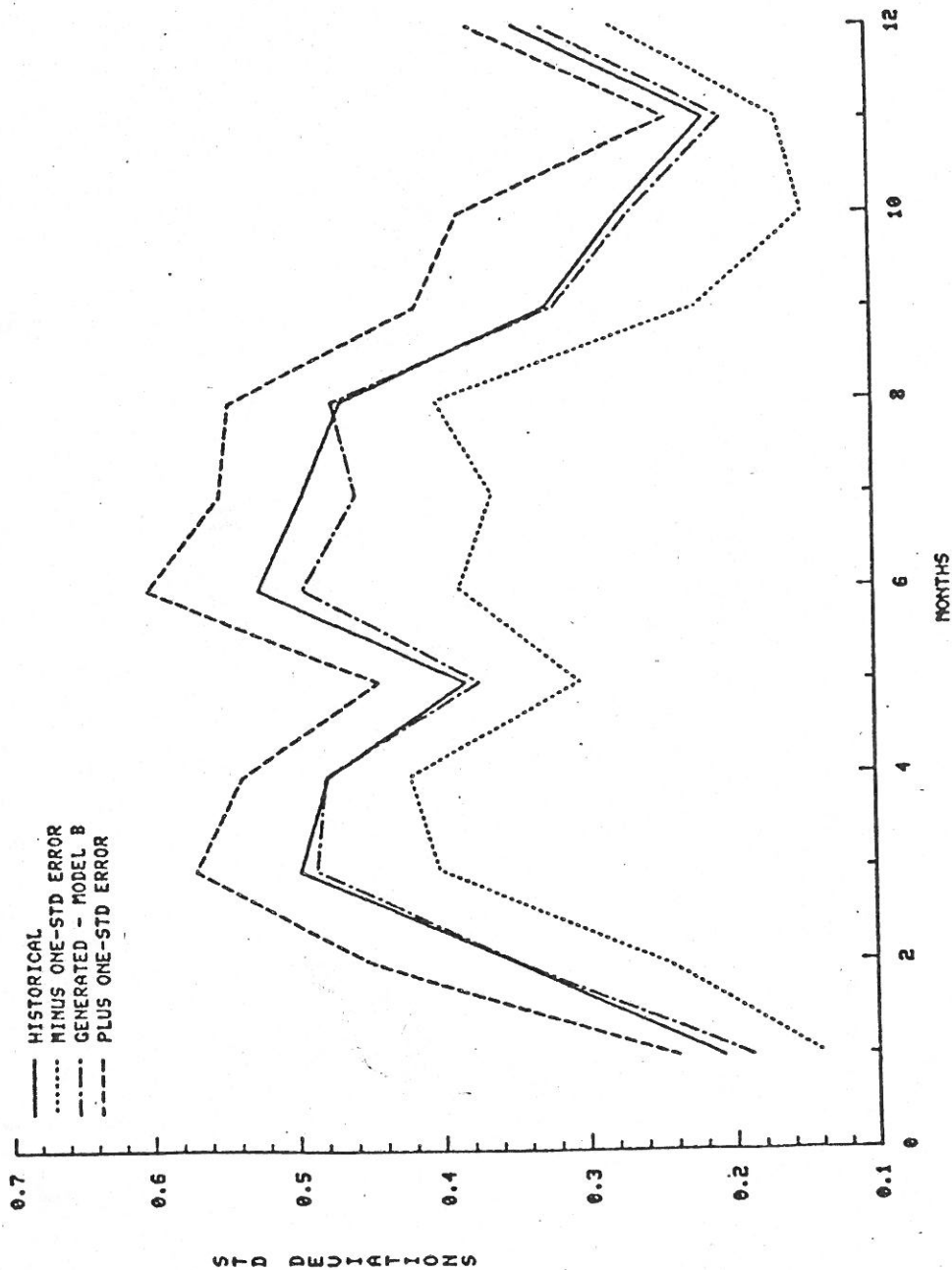


Figure 1.9.E.30. RANCHO - STD DEVIATIONS (LOG DOMAIN)

UNSATISFIE

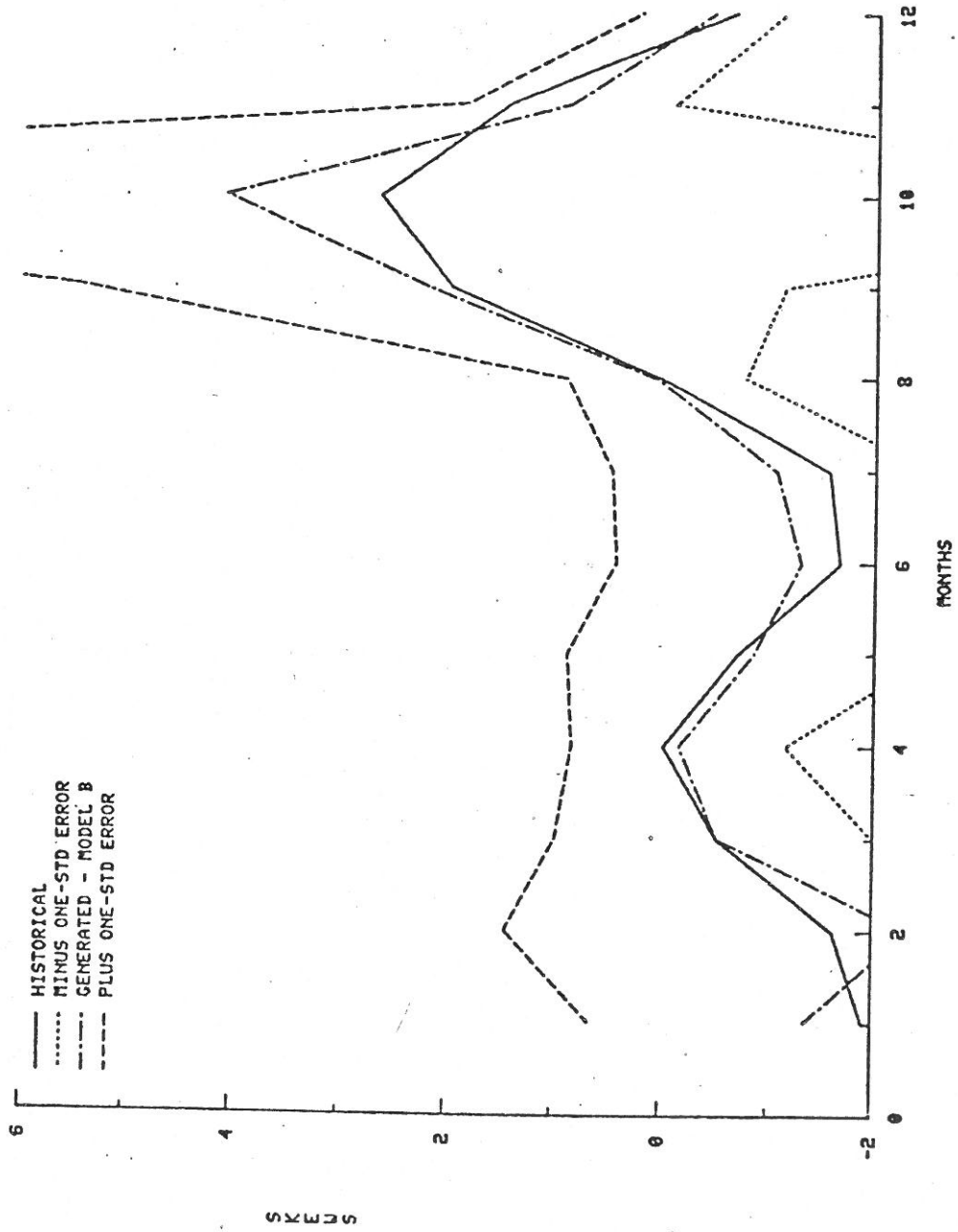


Figure 1.9.E.31. RANCHO - SKEWS (LOG DOMAIN)

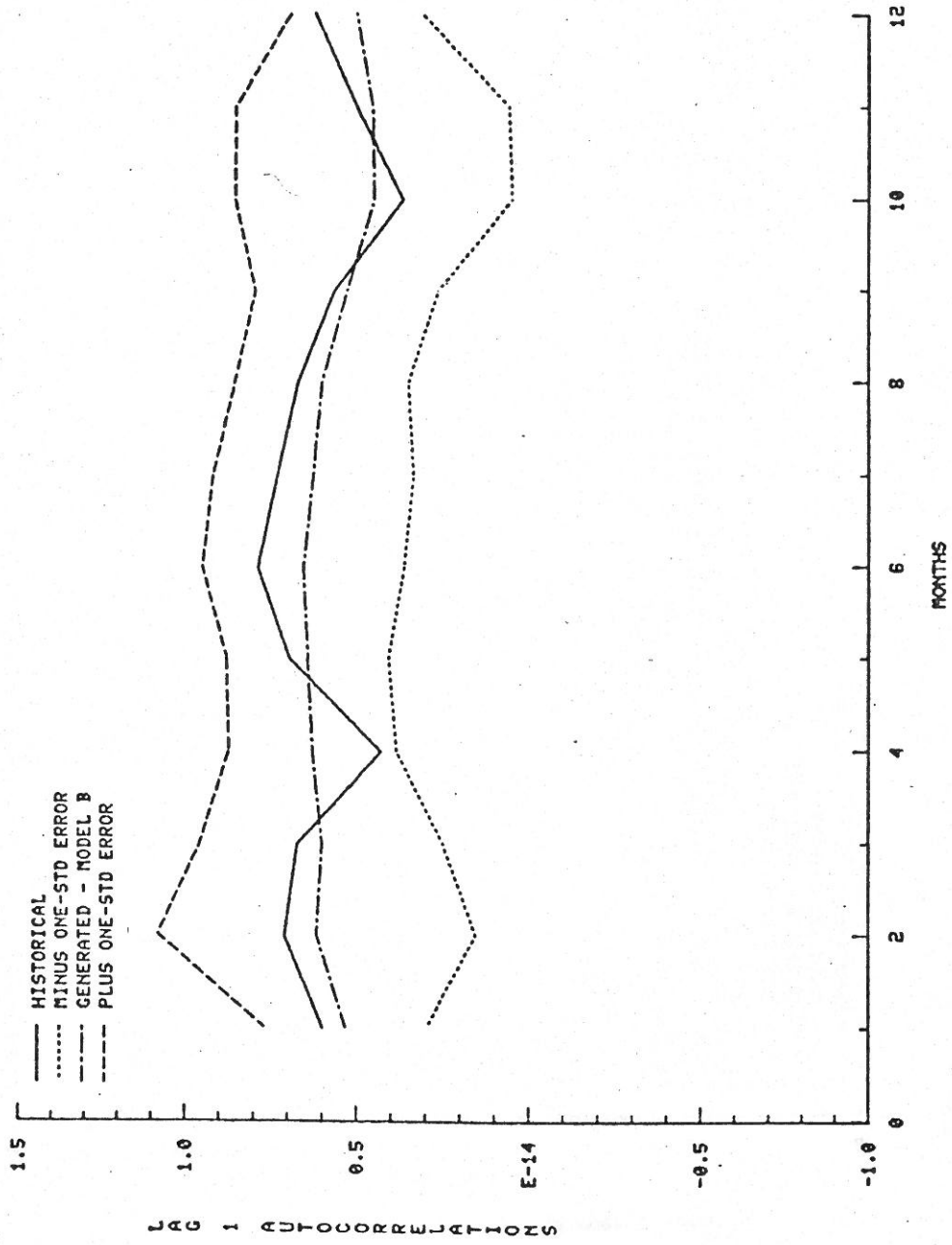


Figure 1.9.E.32. RANCHO - LAG 1 AUTOCORRELATIONS (LOG DOMAIN)

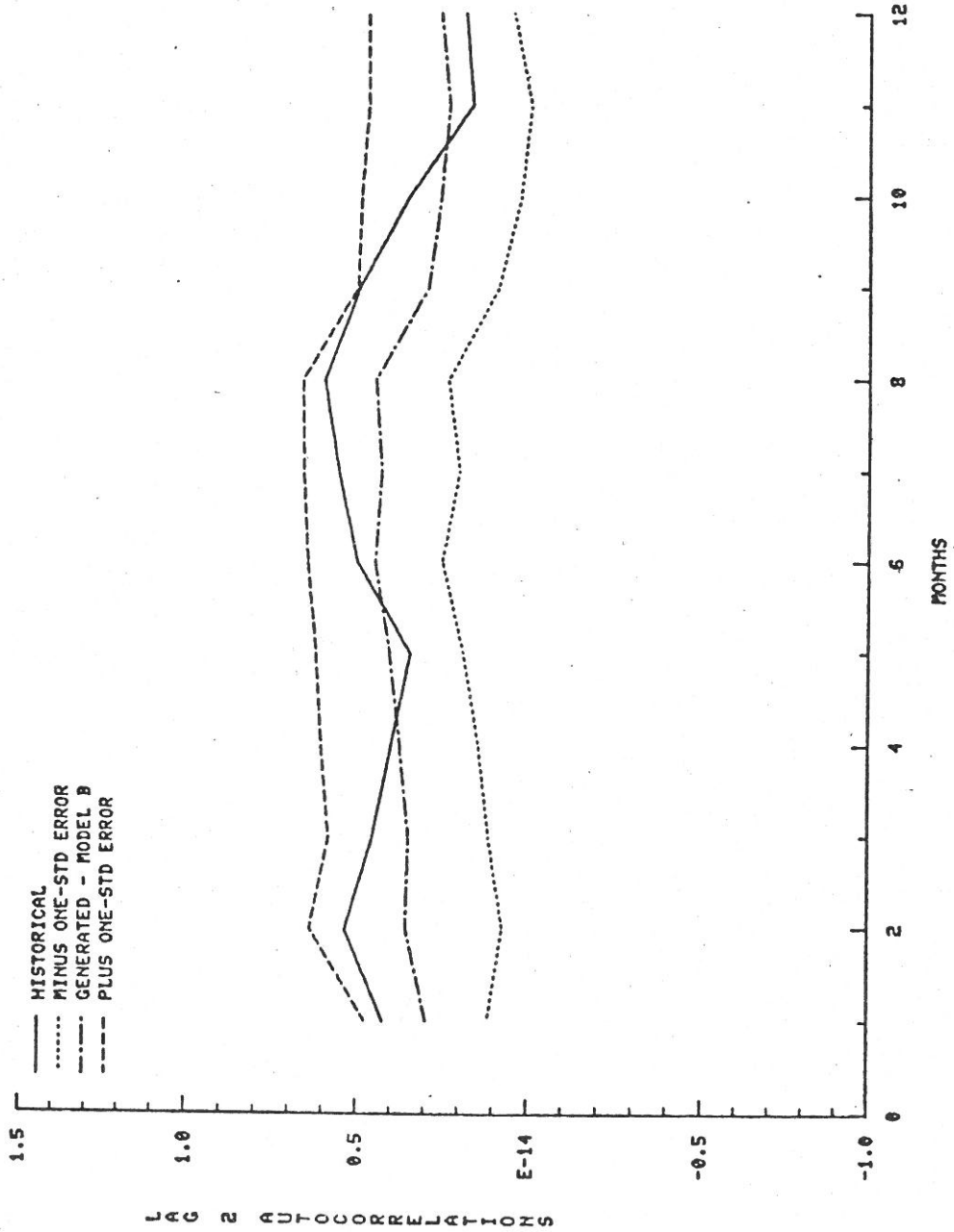


Figure 1.9.E.33. RANCHO - LAG 2 AUTOCORRELATIONS (LOG DOMAIN)

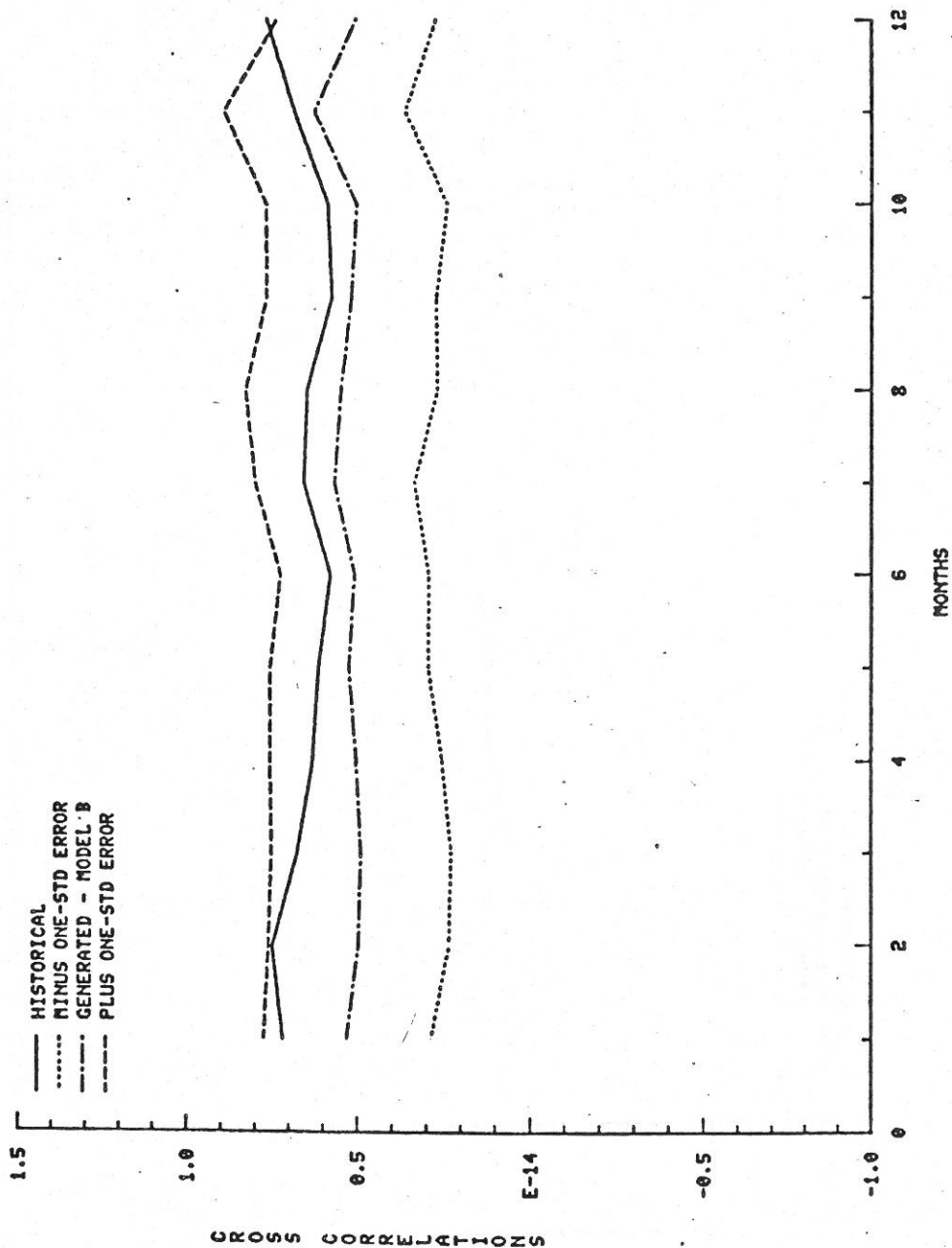


Figure 1.9.E.34. PALODE AND PASODE - CROSS CORRELATIONS (LOG DOMAIN)

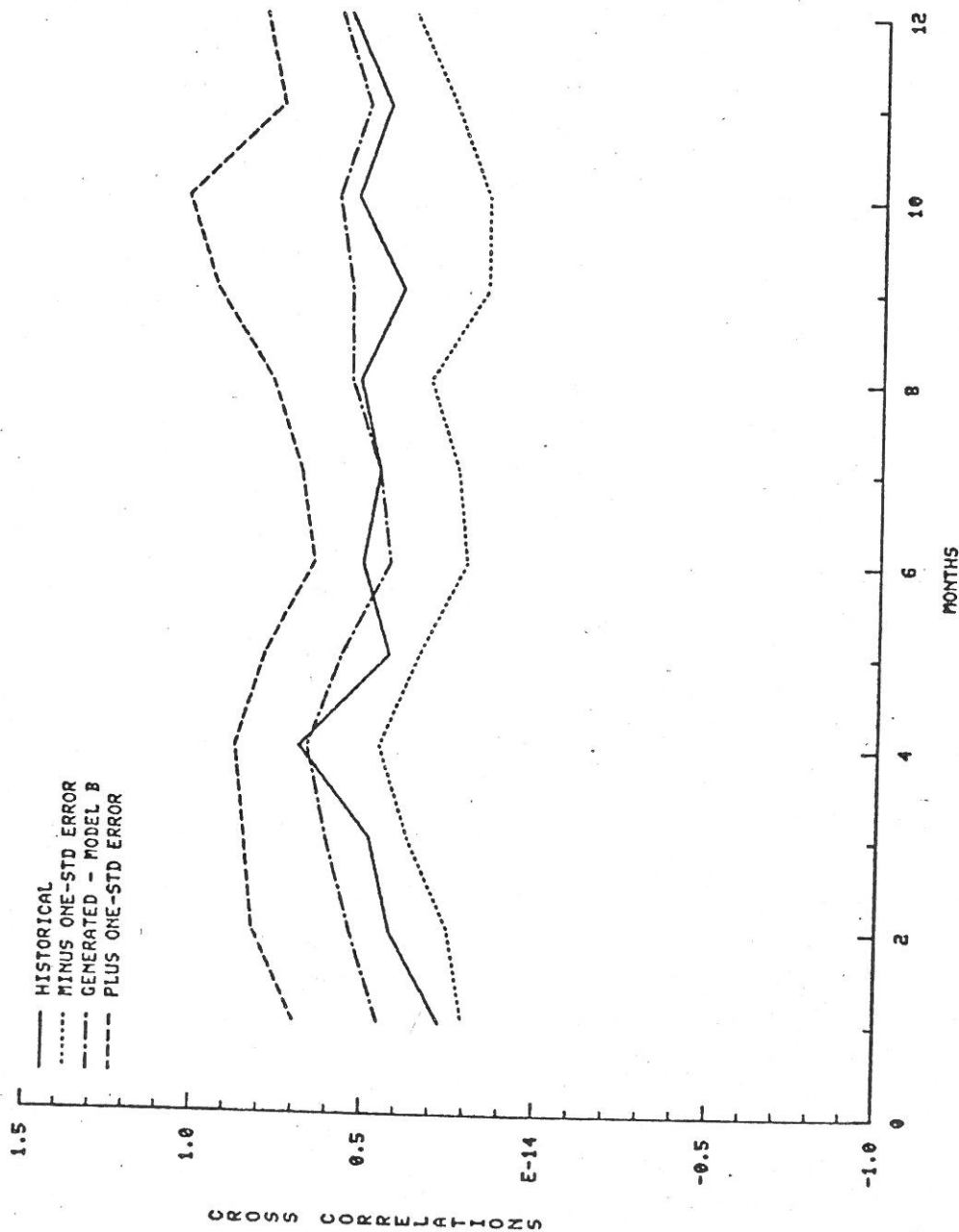


Figure 1.9.E.35. PALODE AND RANCHO - CROSS CORRELATIONS (LOG DOMAIN)

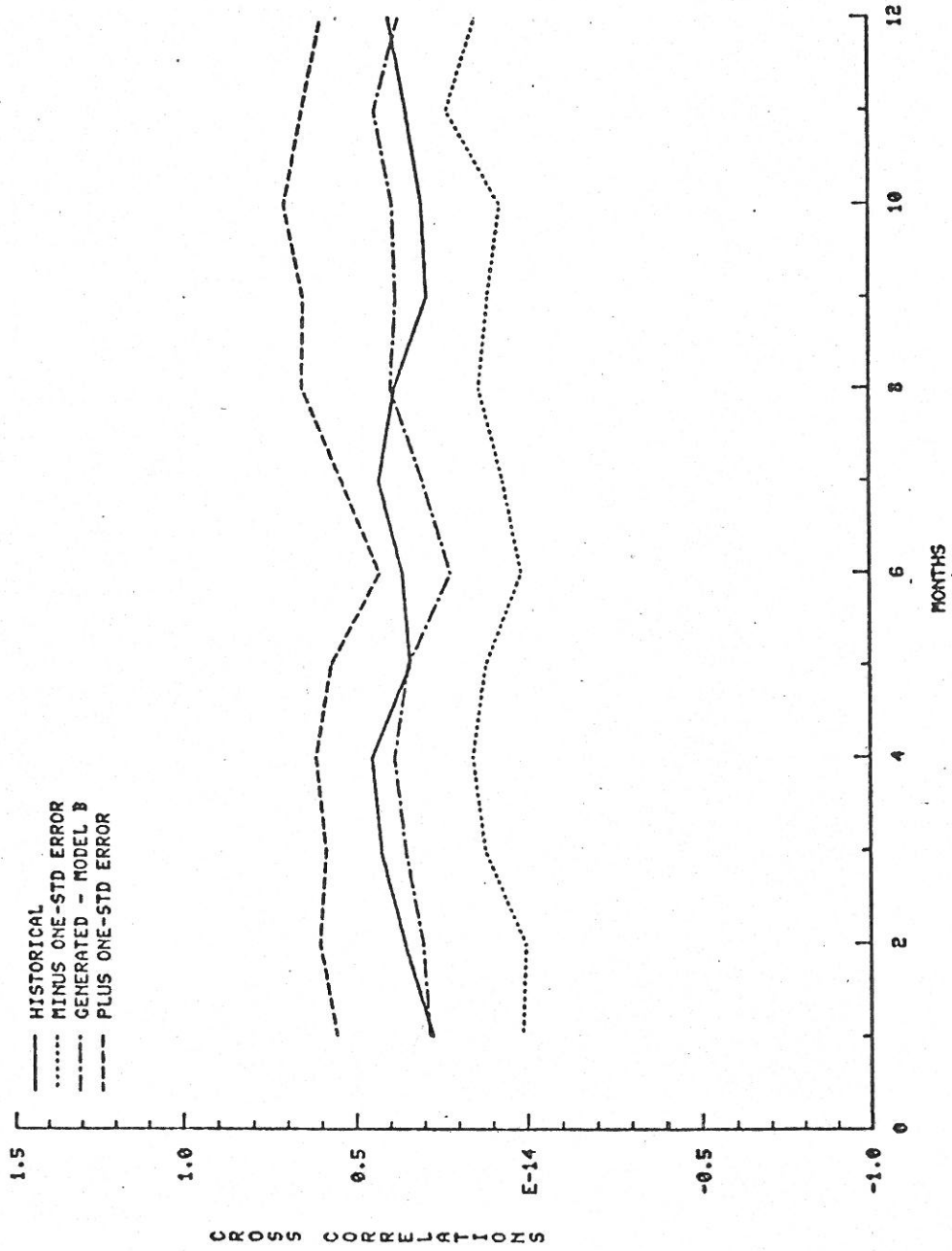


Figure 1.9.E.36. PASODE AND RANCHO - CROSS CORRELATIONS (LOG DOMAIN)

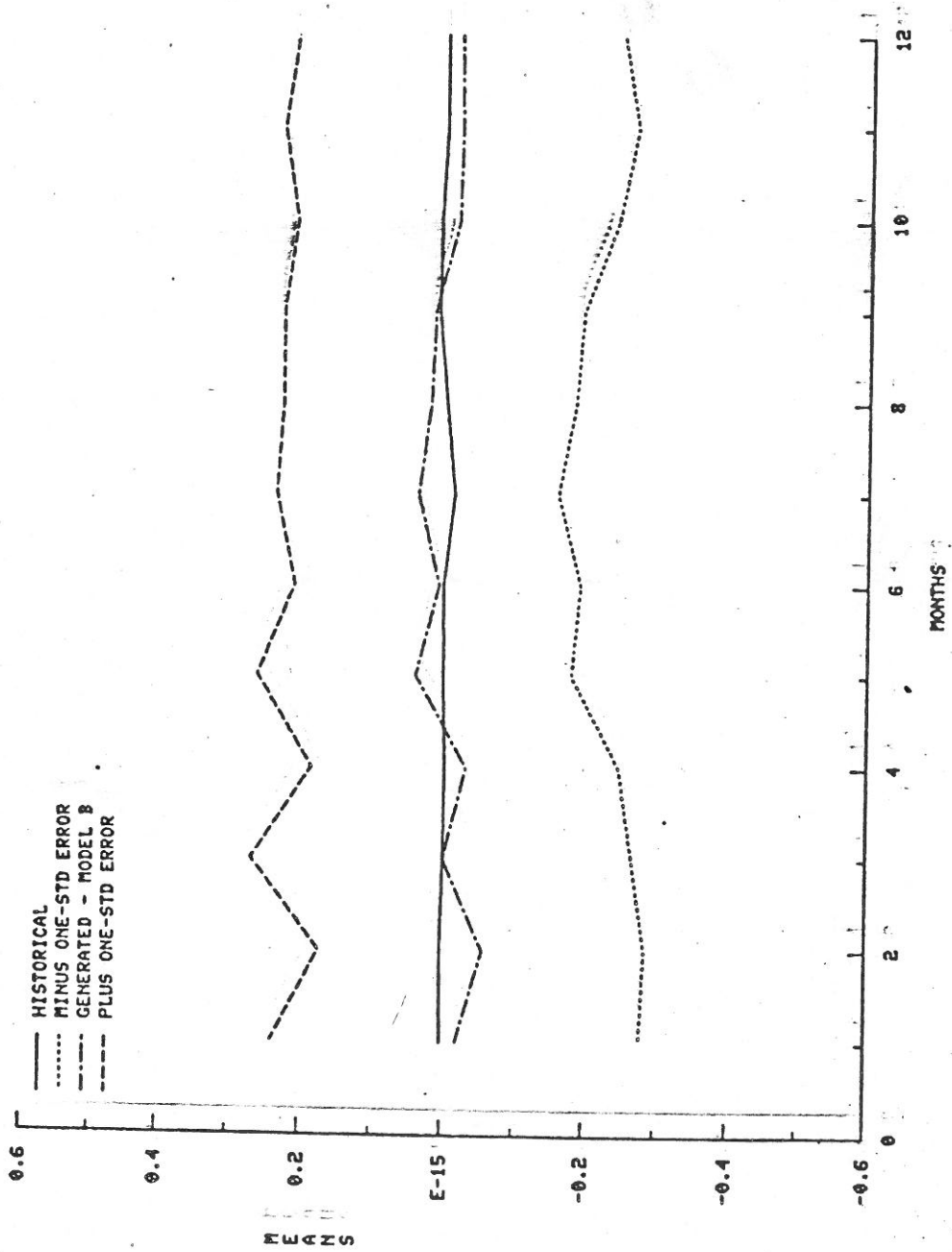


Figure 1.9.E.37. PALODE - MEANS (LUH DOMAIN)

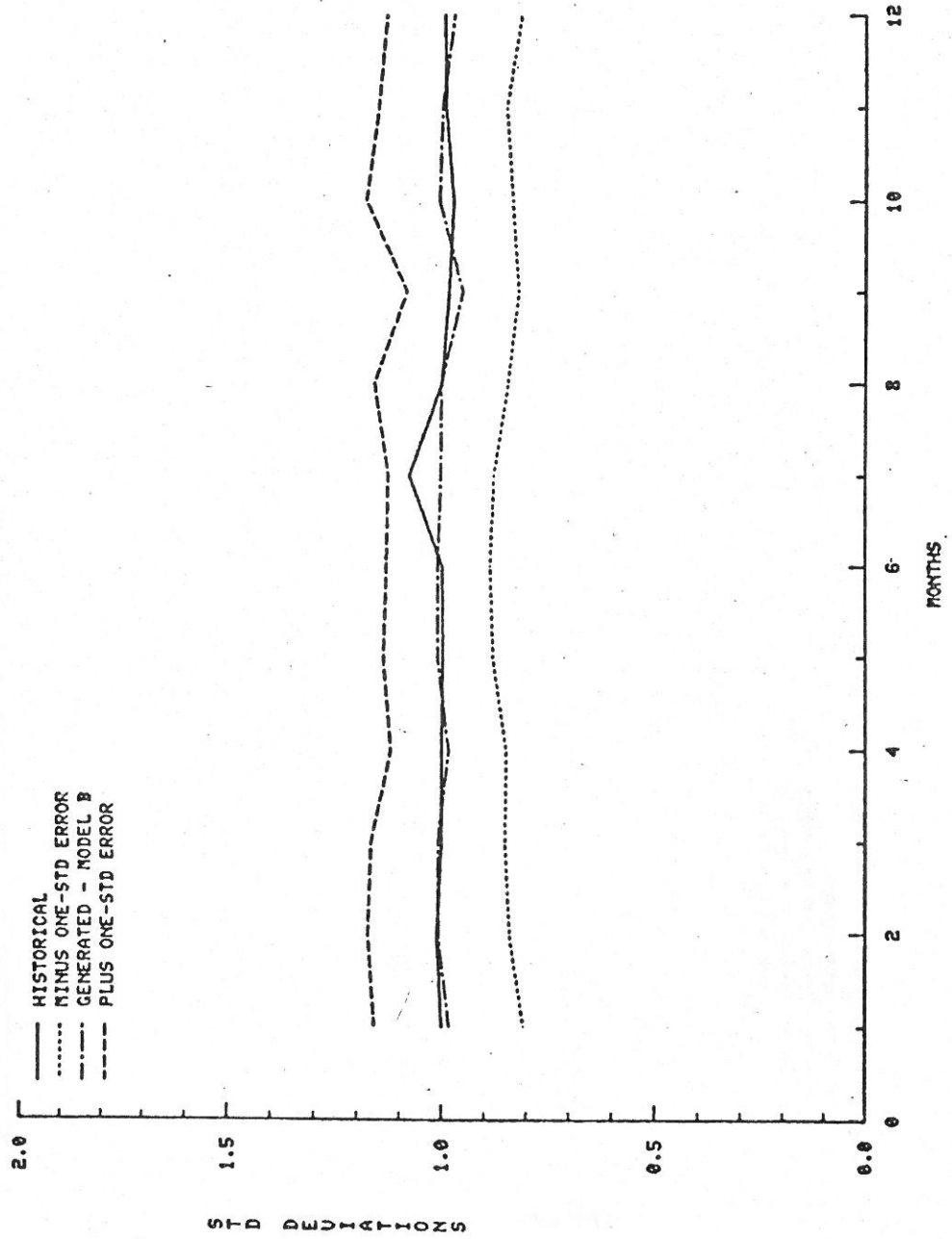


Figure 1.9.E.38. PALODE - STD DEVIATIONS (LWH DOMAIN)

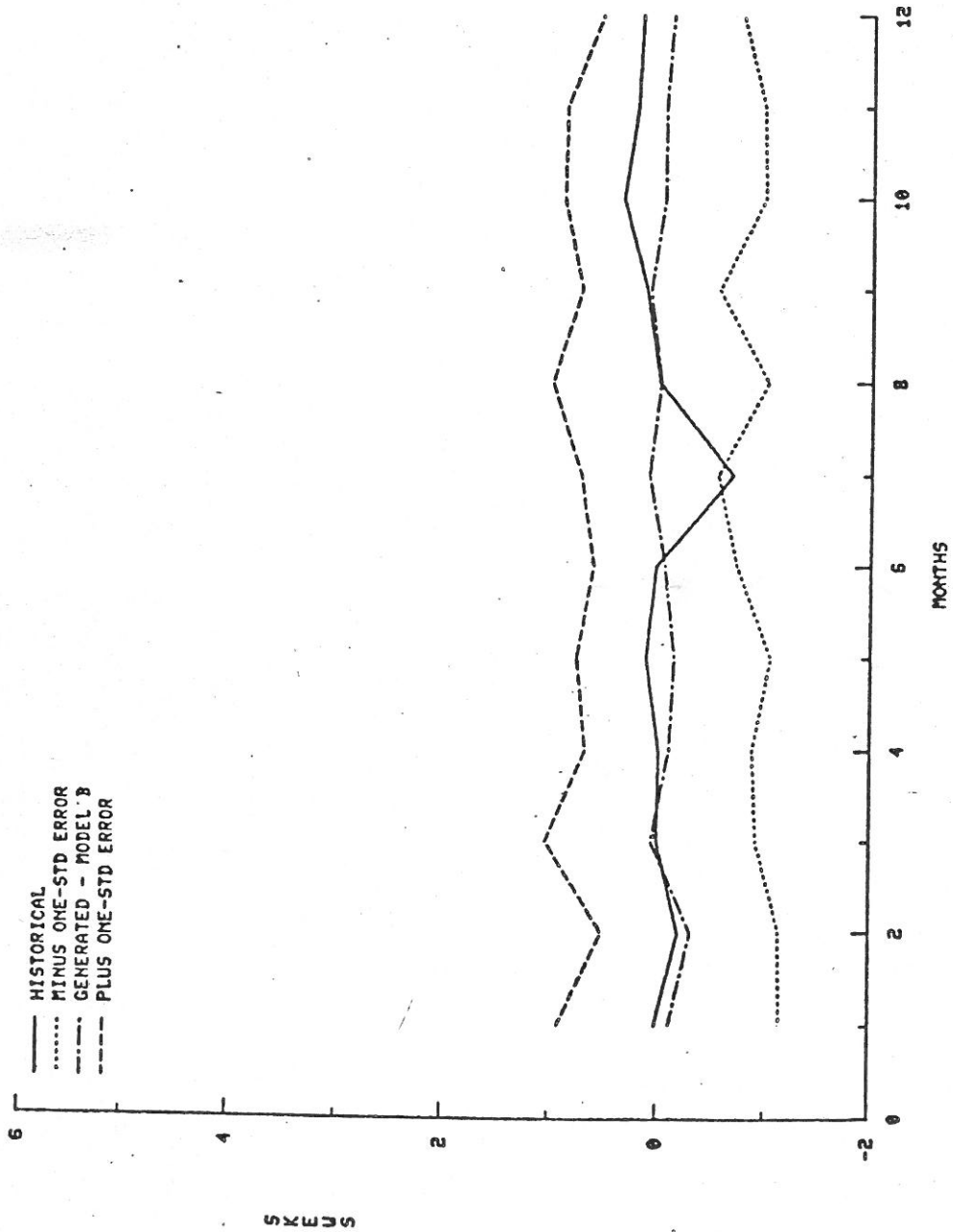


Figure 1.9.E.39. PALODE - SKEWS (LWH DOMAIN)

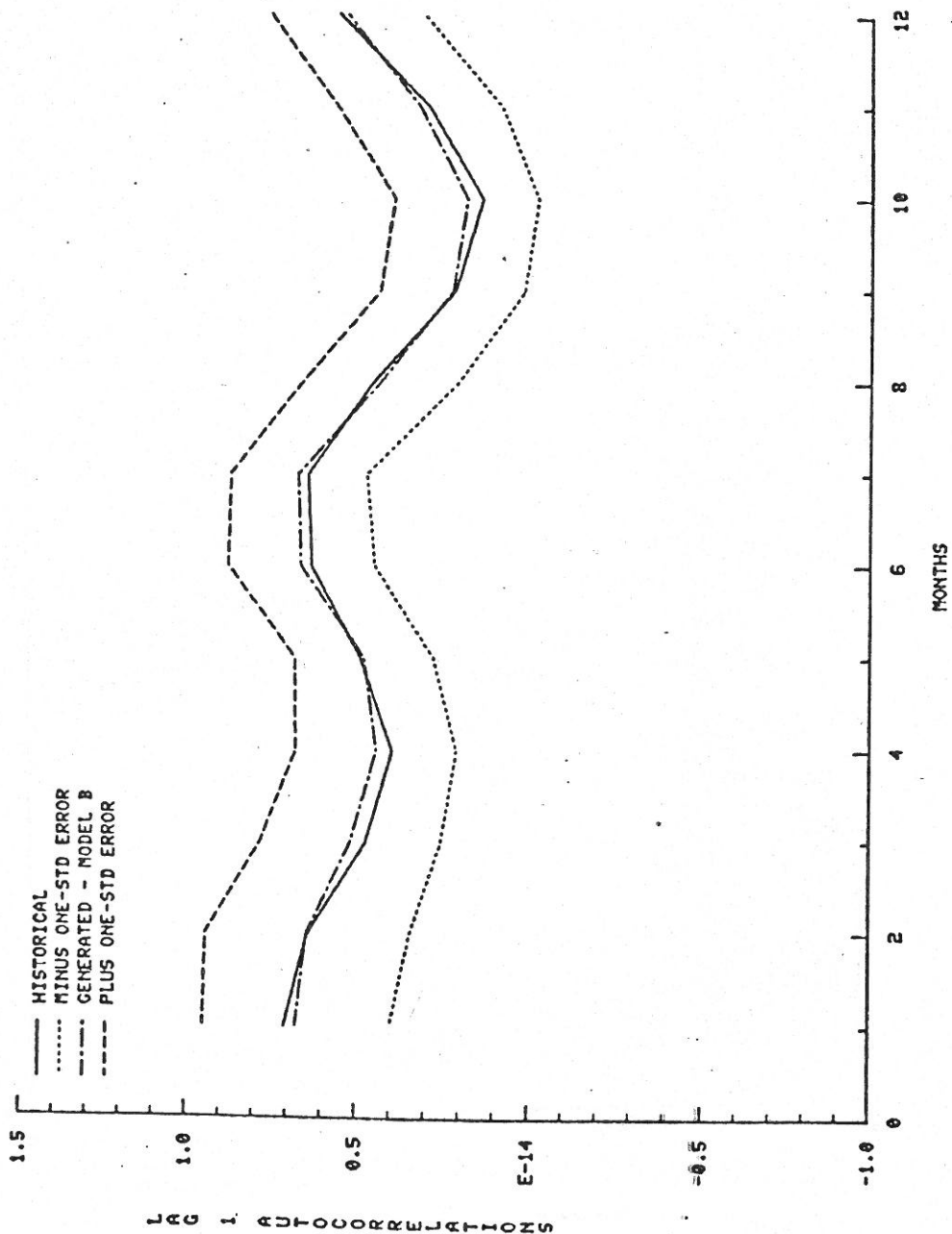


Figure 1.9.E.40. PALODE - LAG 1 AUTOCORRELATIONS (LWLH DOMAIN)

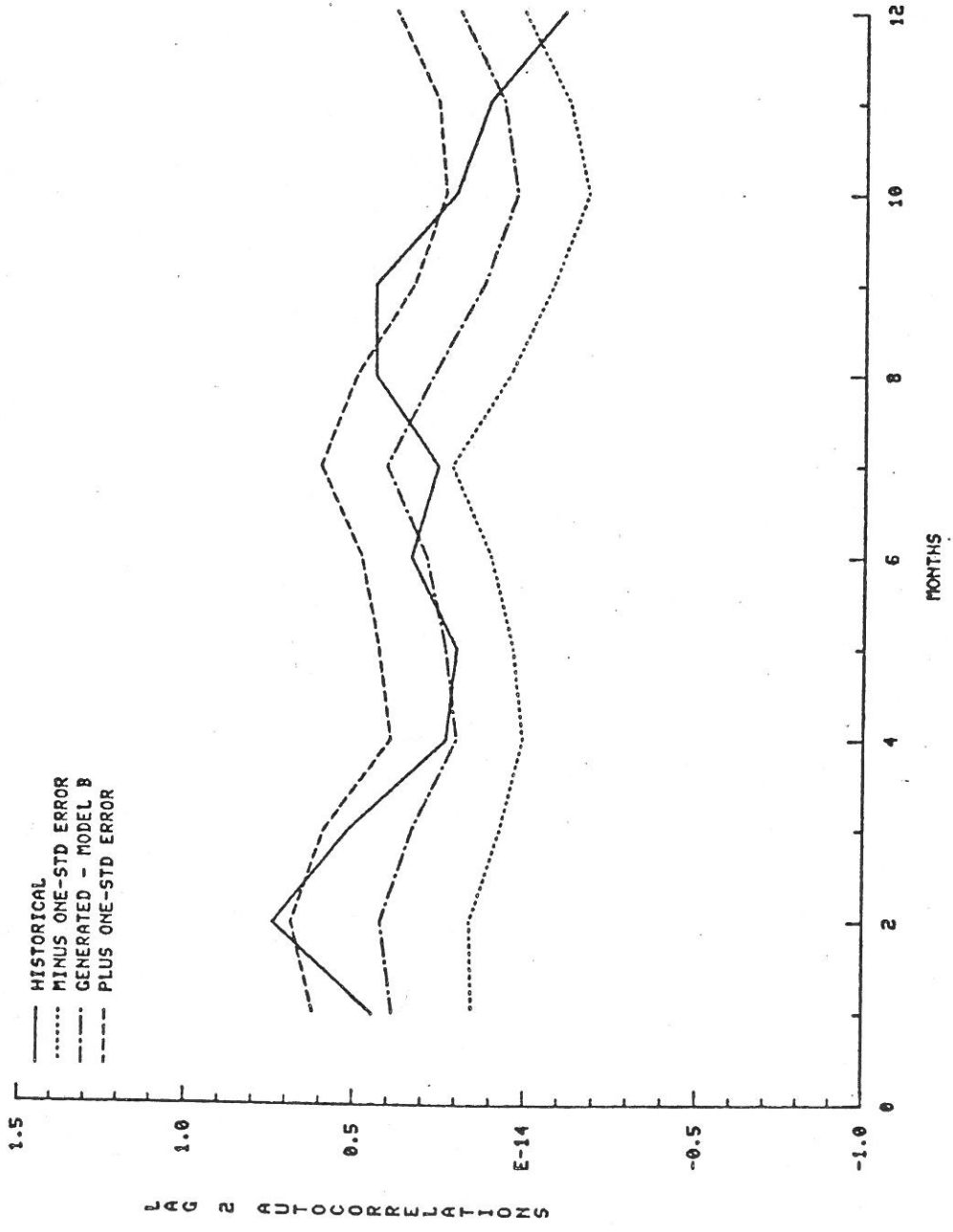


Figure 1.9.E.41. PALODE - LAG 2 AUTOCORRELATIONS (LWH DOMAIN)

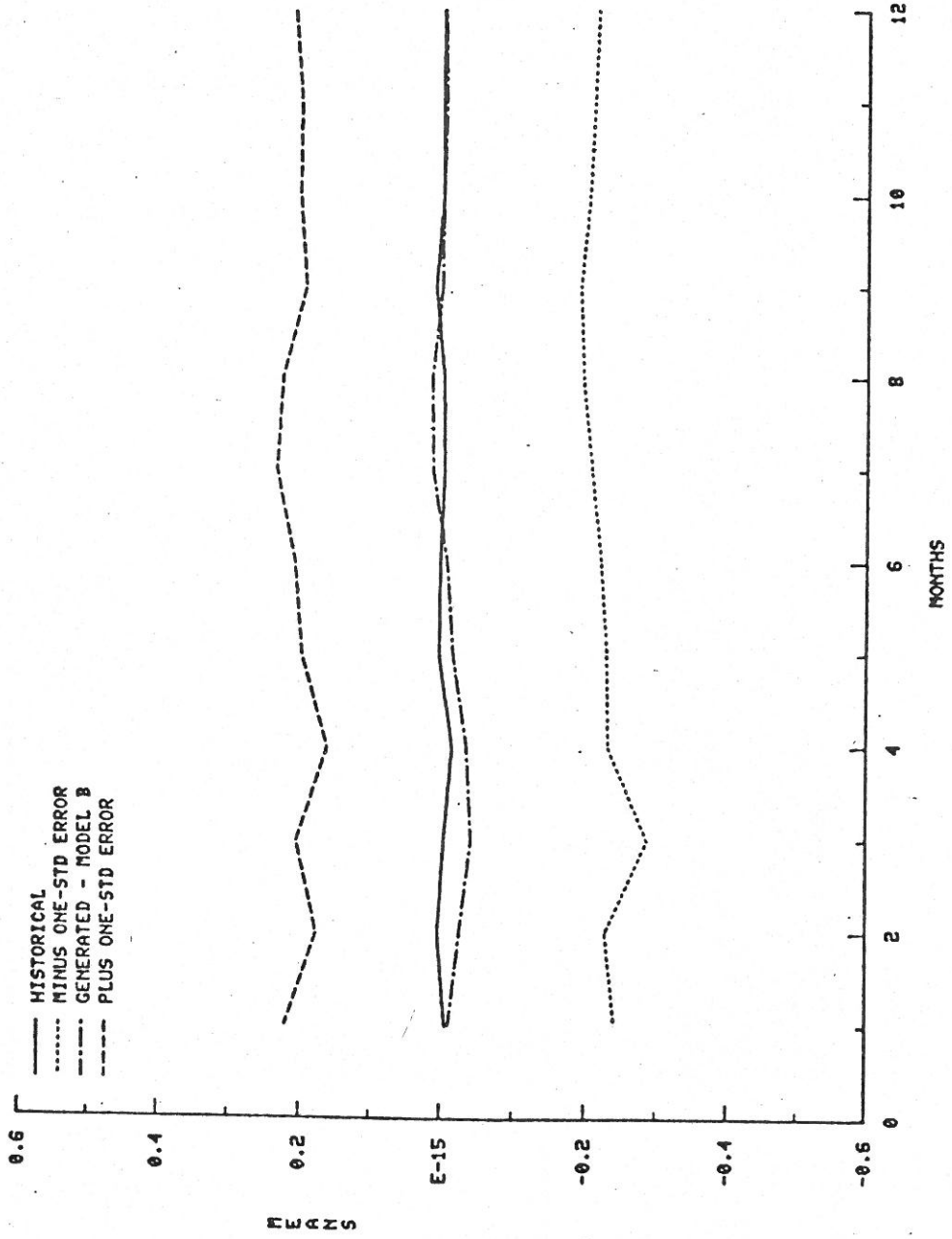


Figure 1.9.E.42. PASODE - MEANS (LWH DOMAIN)

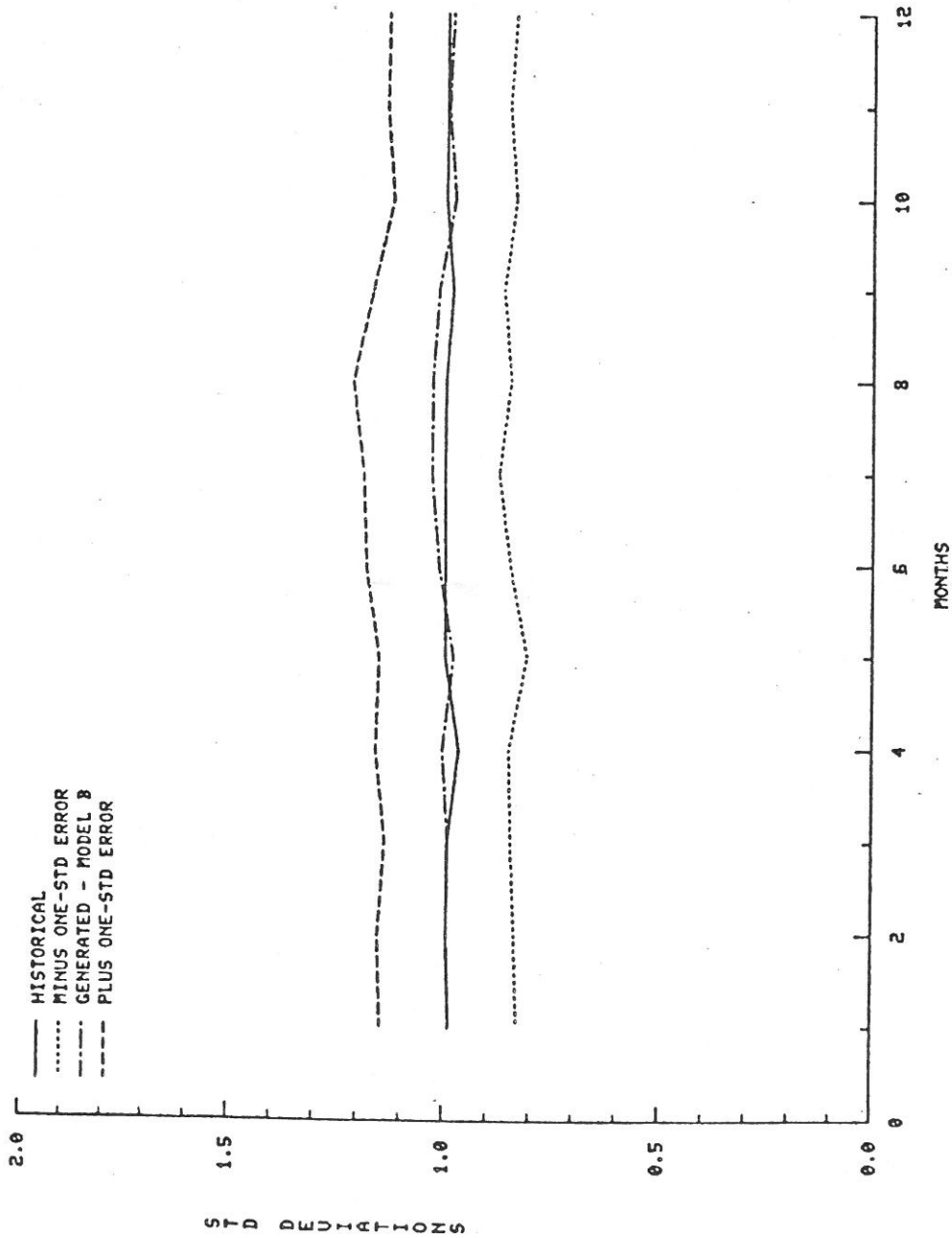


Figure 1.9.E.43. PASODE - STD DEVIATIONS (LJH DOMAIN)

UNSATISFIED EXT P • 011666

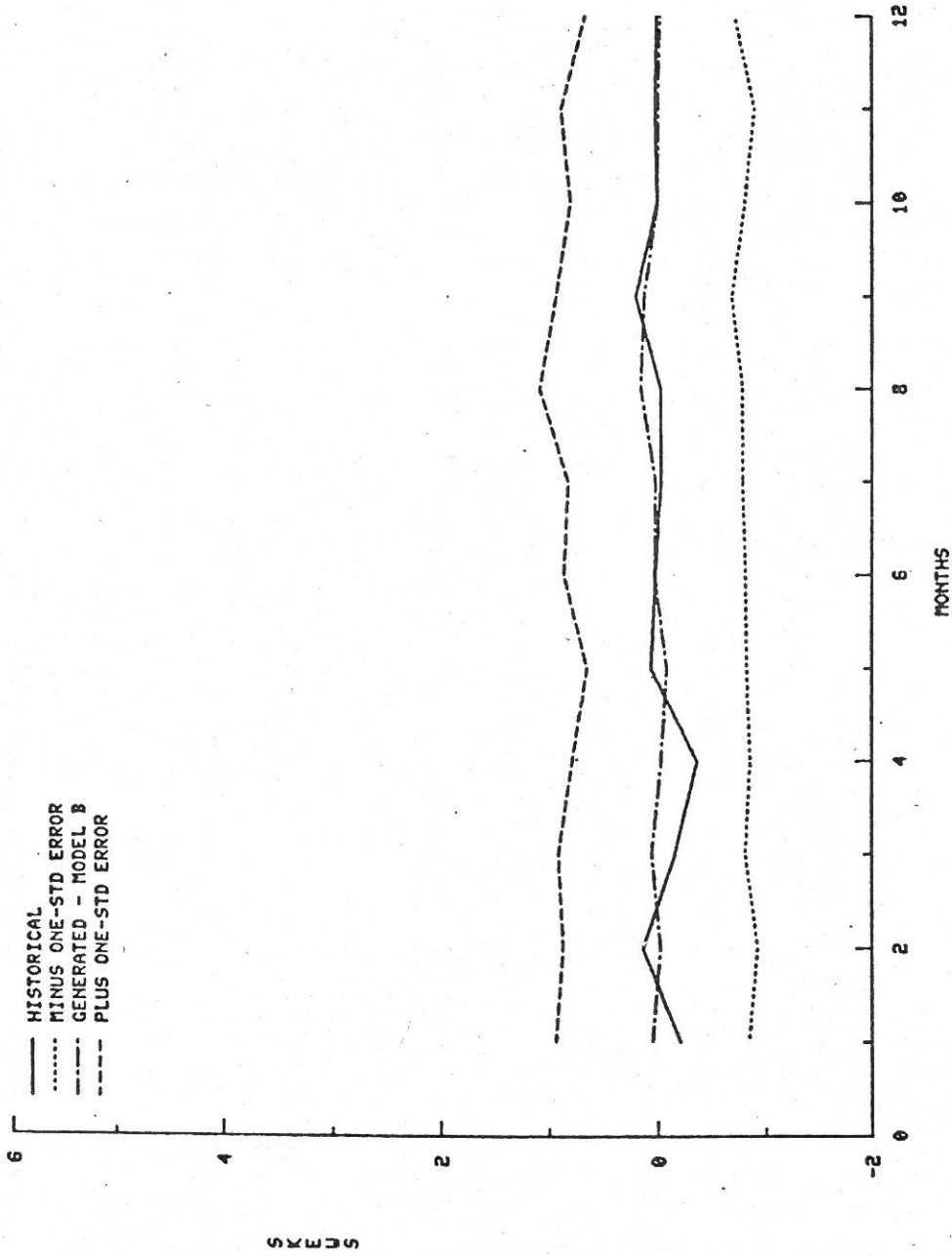


Figure 1.9.E.44. PASODE - SKEWS (LWH DOMAIN)

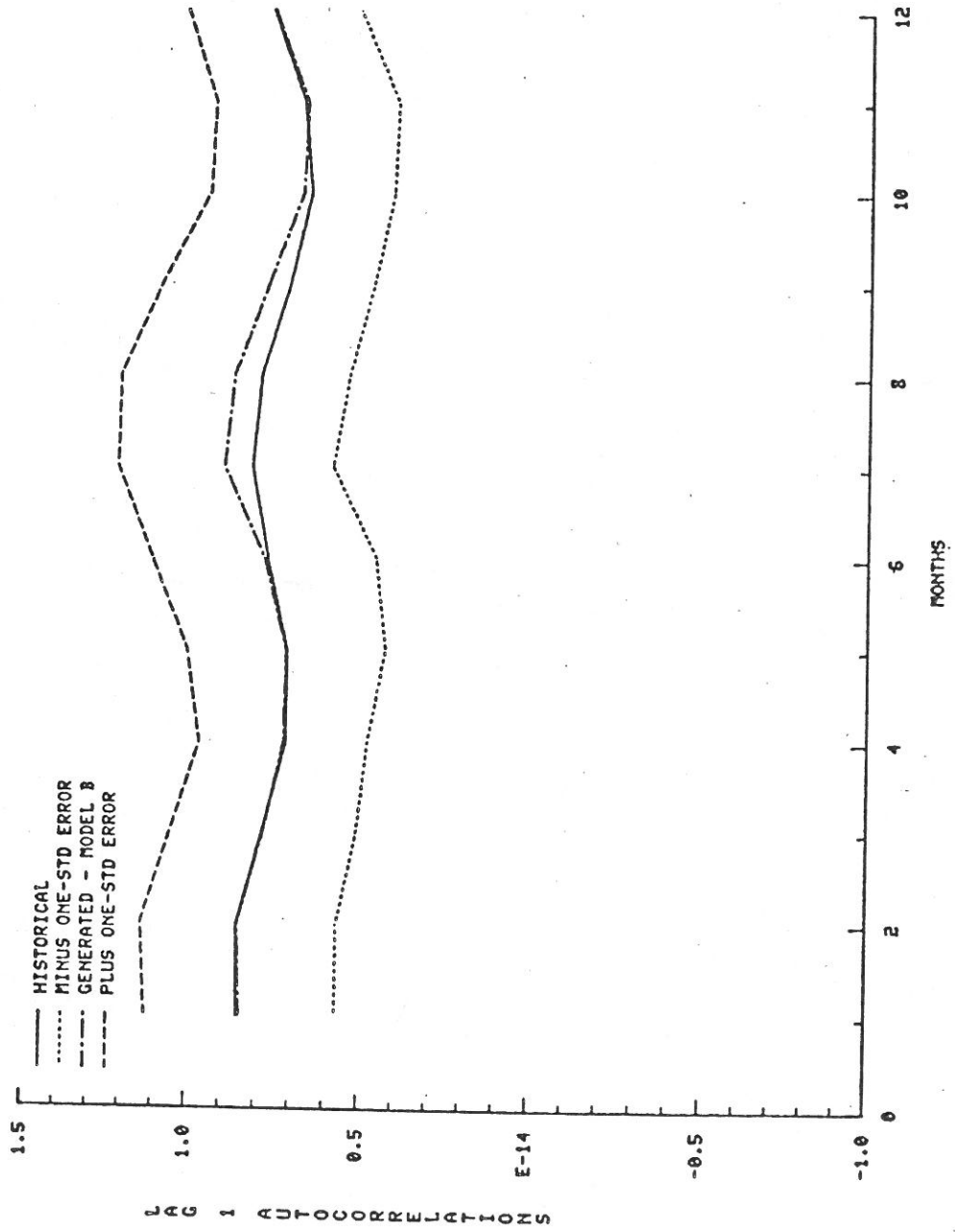


Figure 1.9.E.45. PASODE - LAG 1 AUTOCORRELATIONS (LJH DOMAIN)

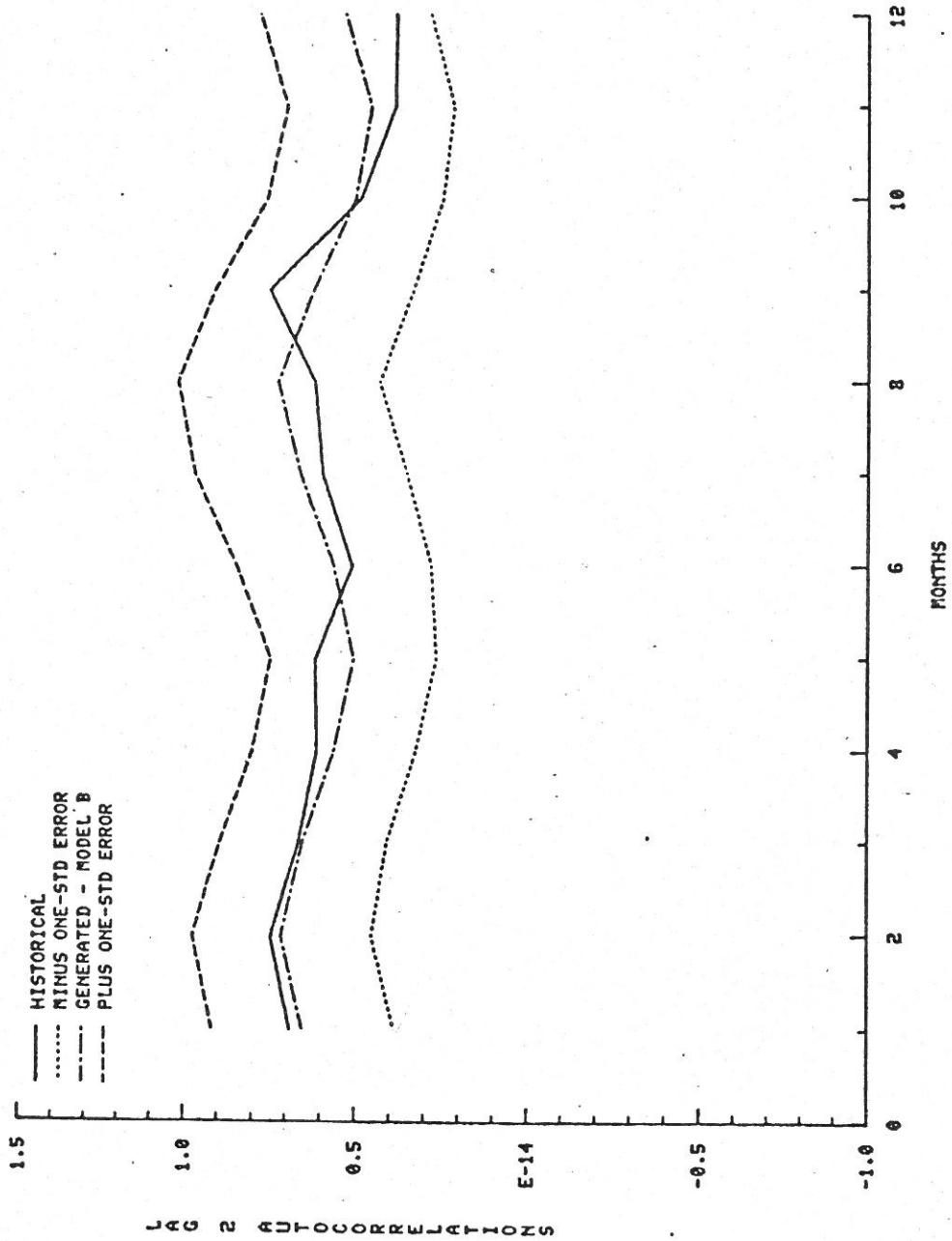


Figure 1.9.E.46. PASODE - LAG 2 AUTOCORRELATIONS (LWH DOMAIN)

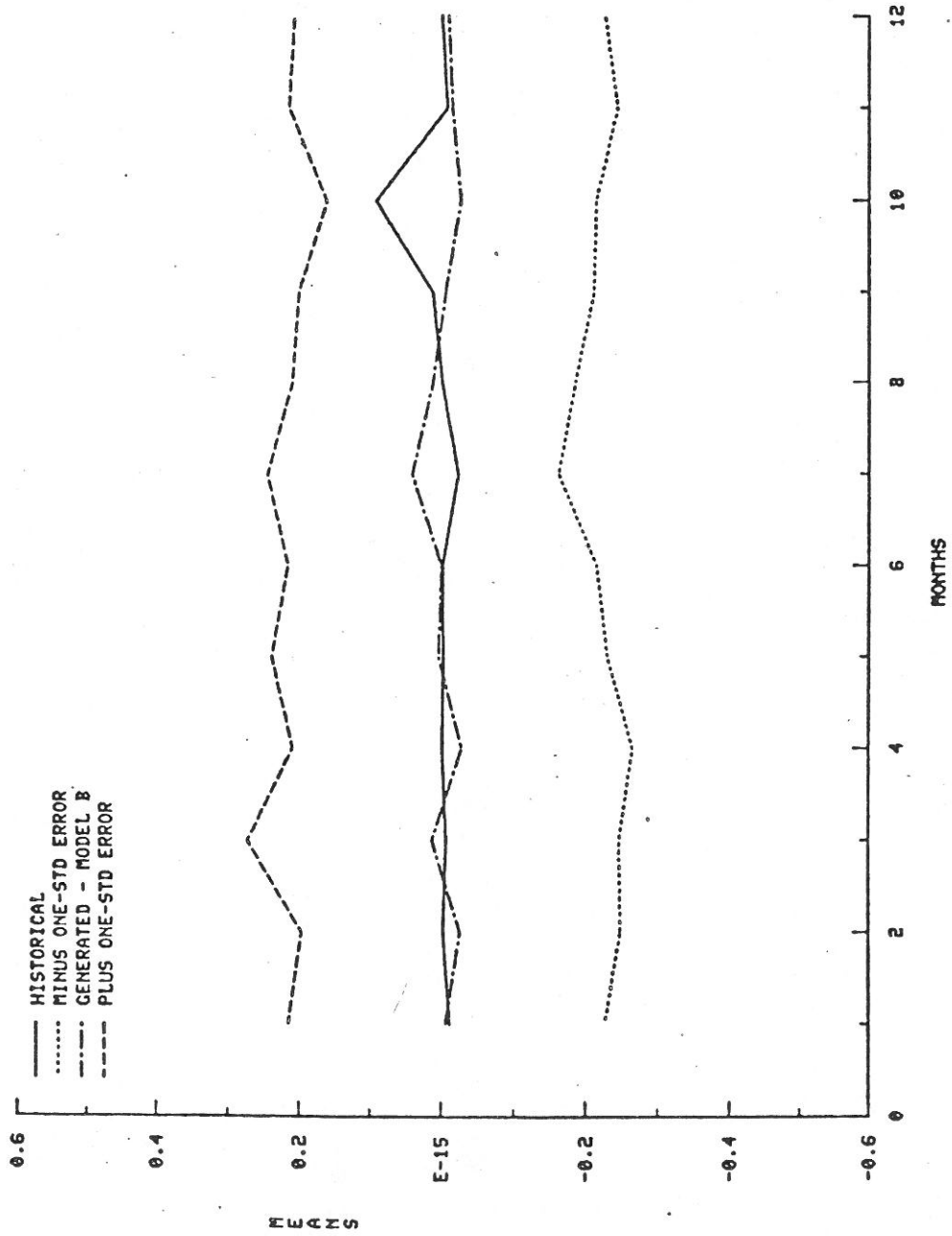


Figure 1.9.E.47. RANCHO - MEANS (LJH DOMAIN)

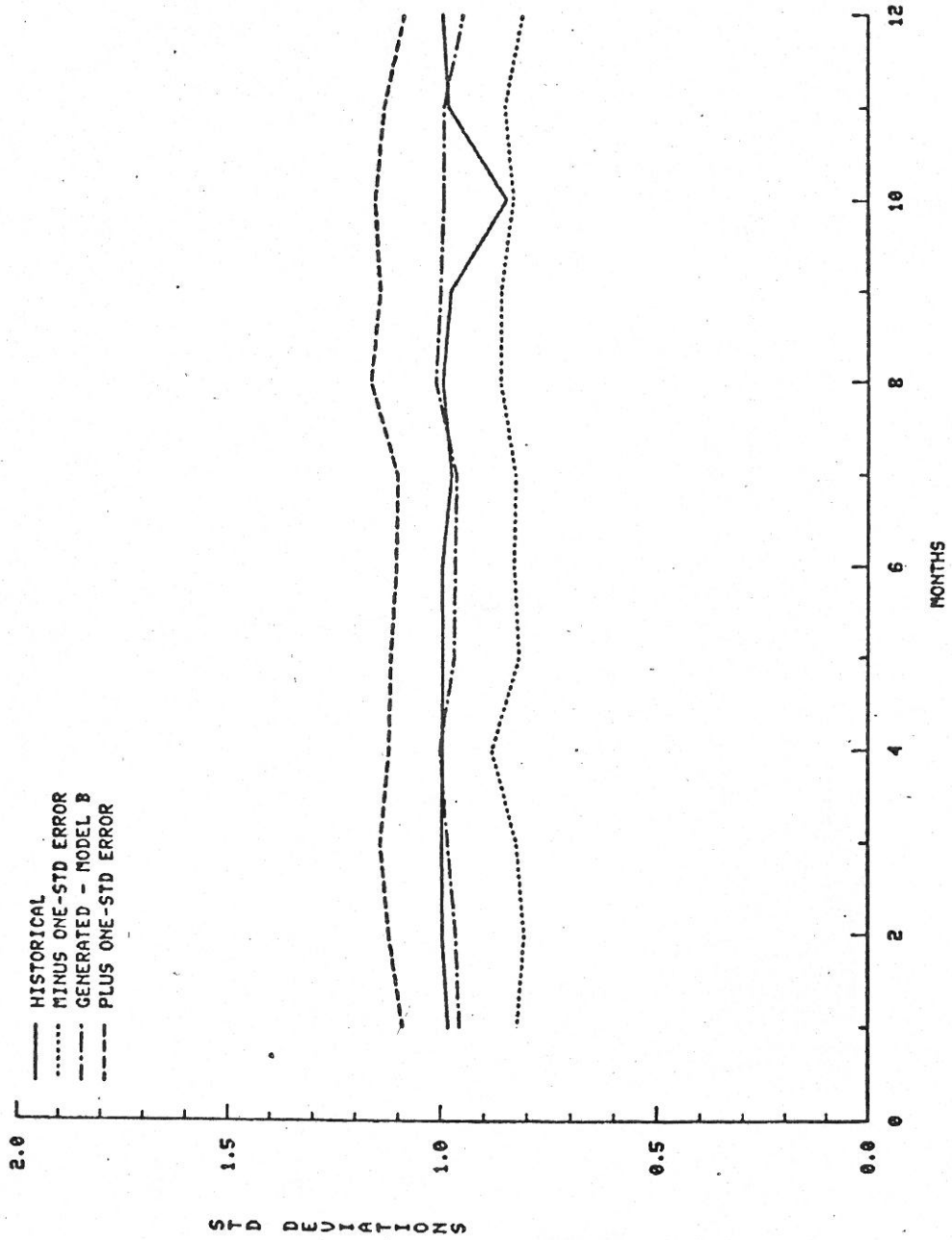


Figure 1.9.E.48. RANCHO - STD DEVIATIONS (LWA Domain)

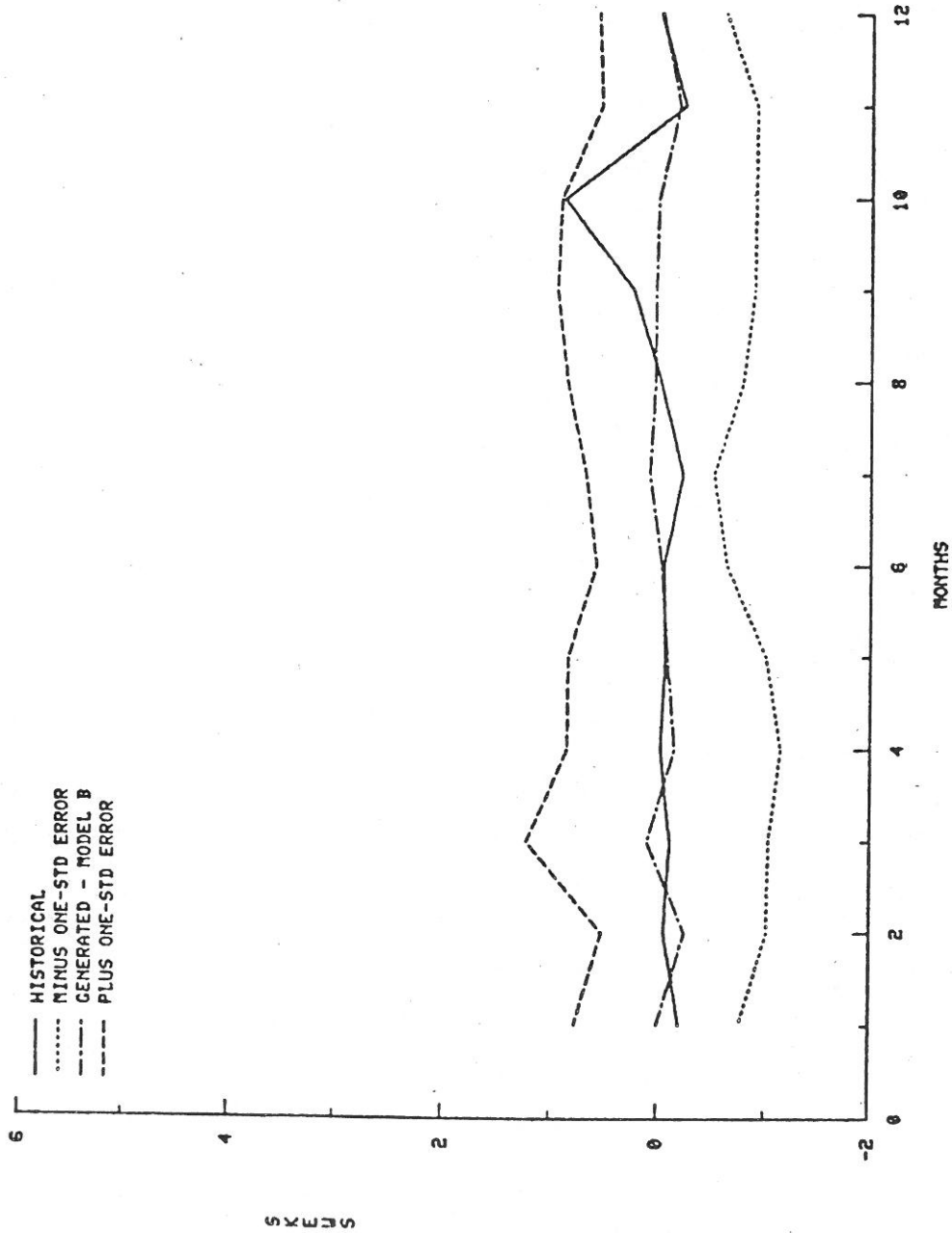


Figure 1.9.E.49. RANCHO - SKEWS (LWH DOMAIN)

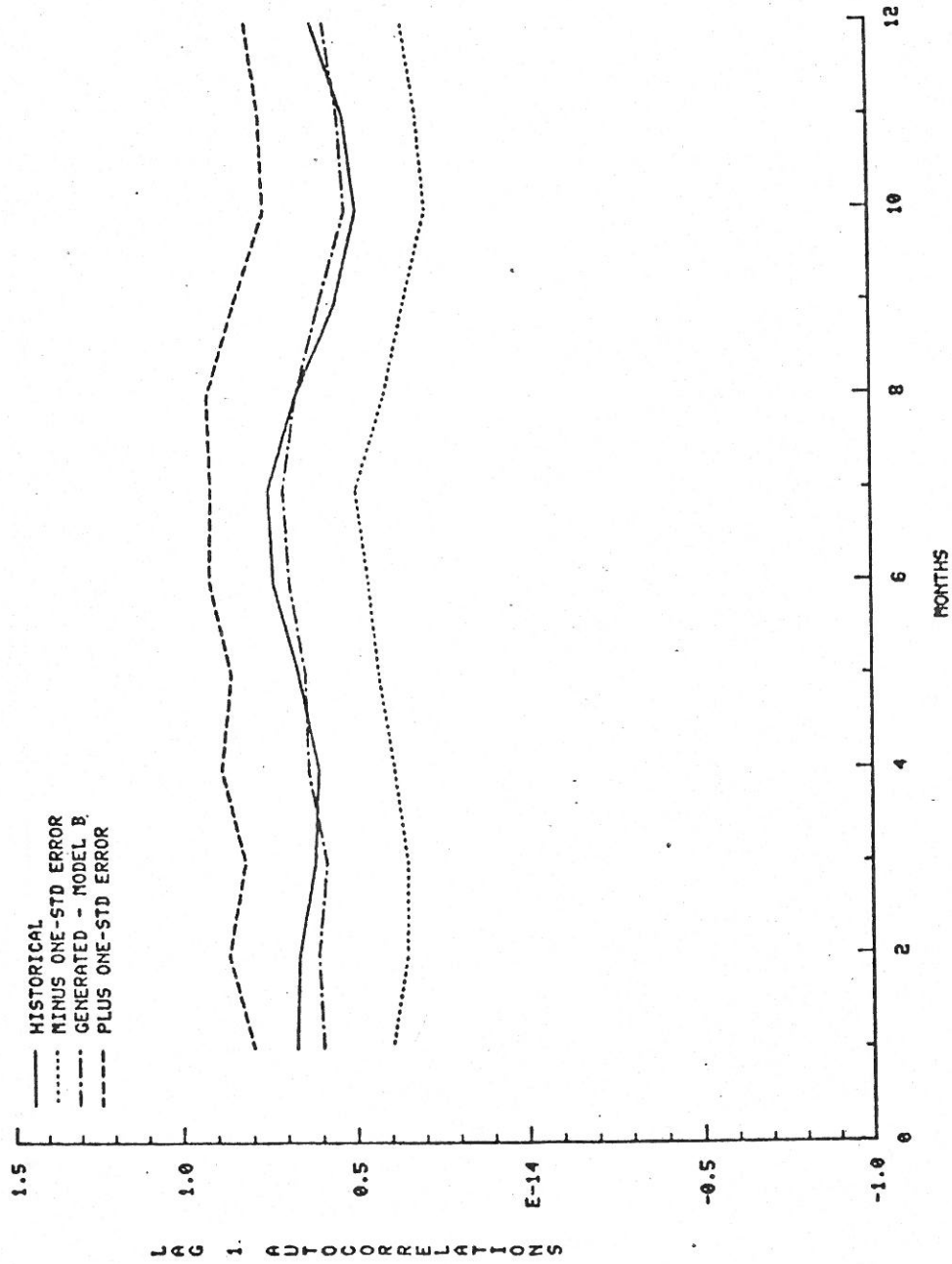


Figure 1.9.E.50. RANCHO - LAG 1 AUTOCORRELATIONS (LWH DOMAIN)

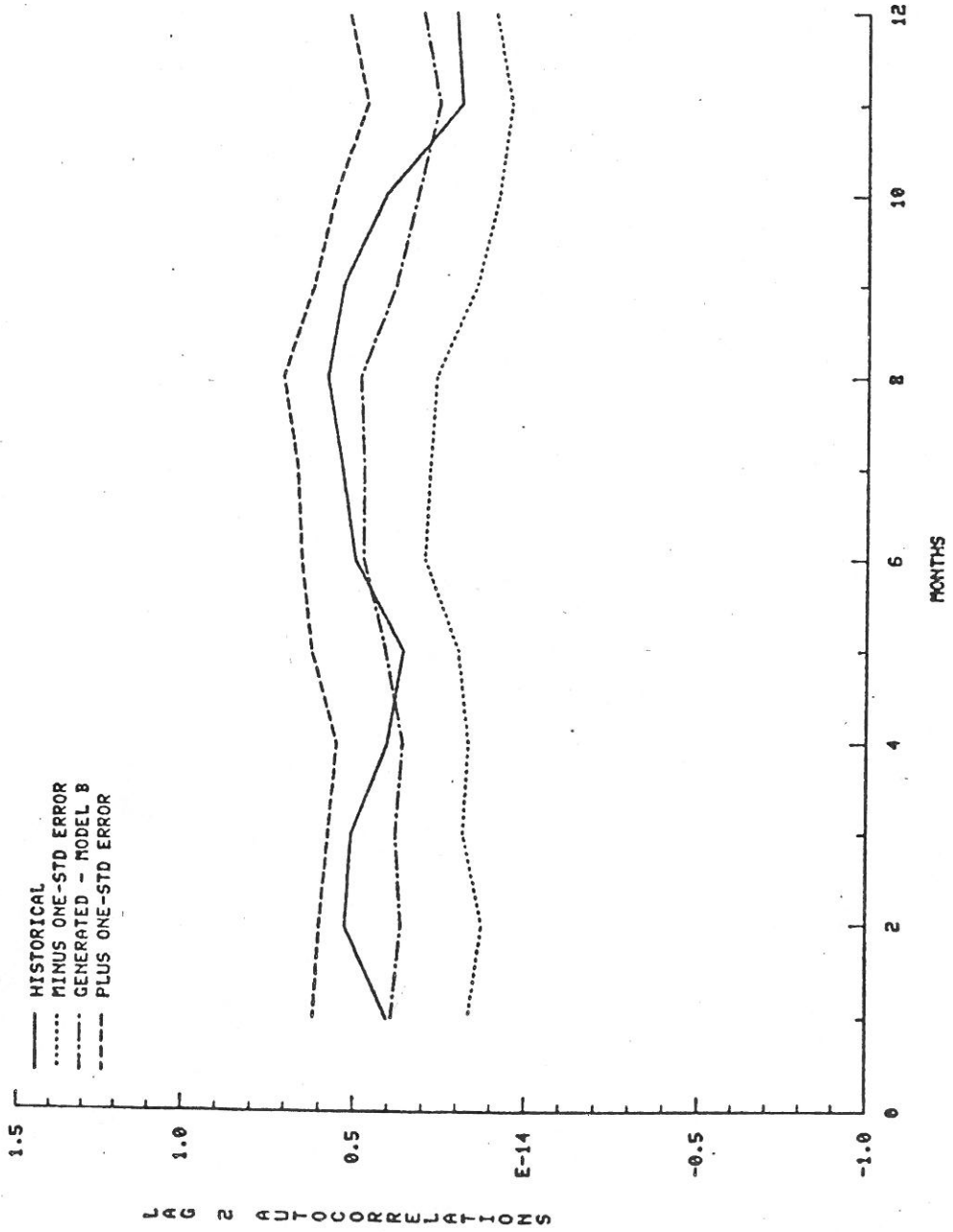


Figure 1.9.E.51. RANCHO - LAG 2 AUTOCORRELATIONS (LUH DOMAIN)

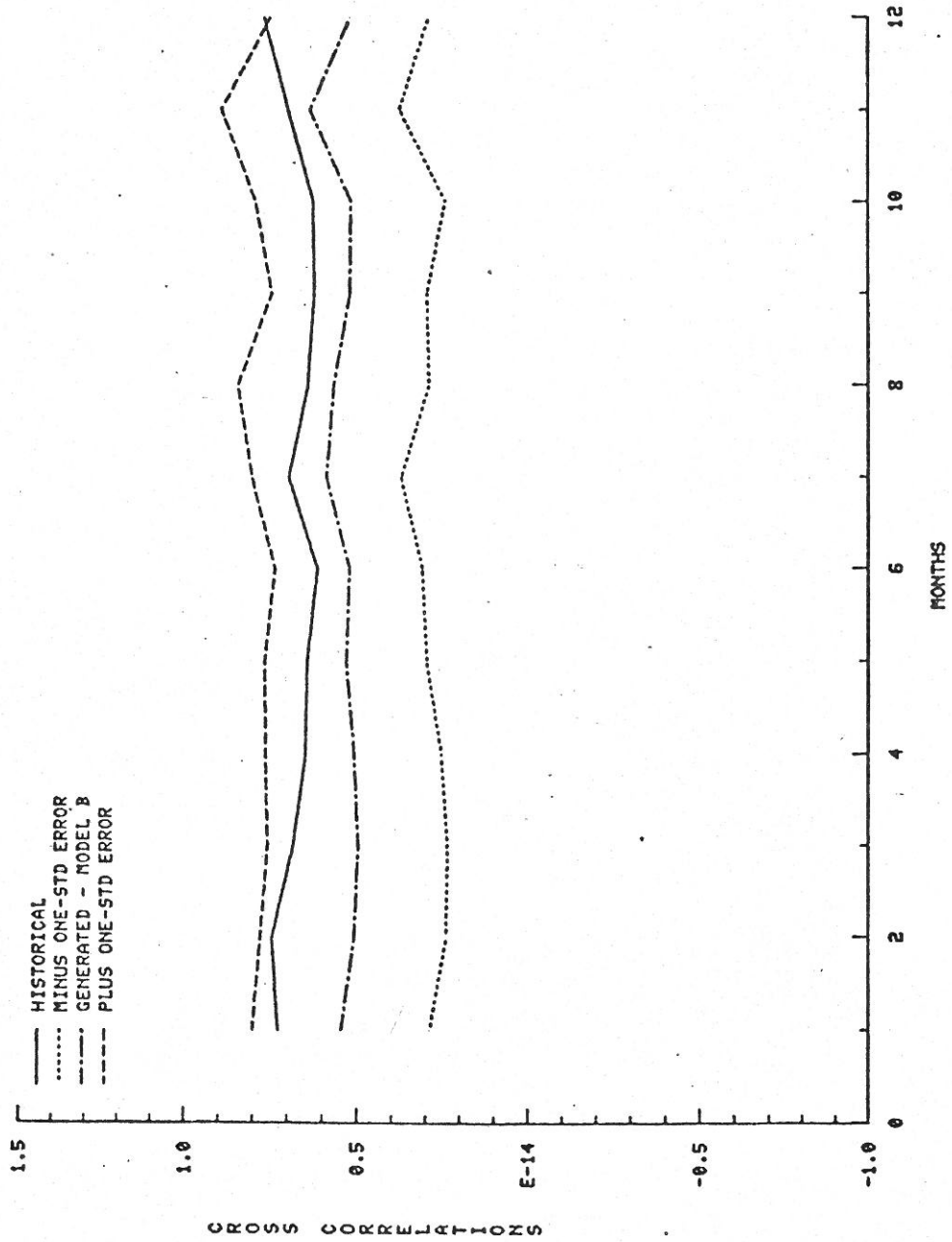


Figure 1.9.E.52. PALODE AND PASODE - CROSS CORRELATIONS (LWH DOMAIN)

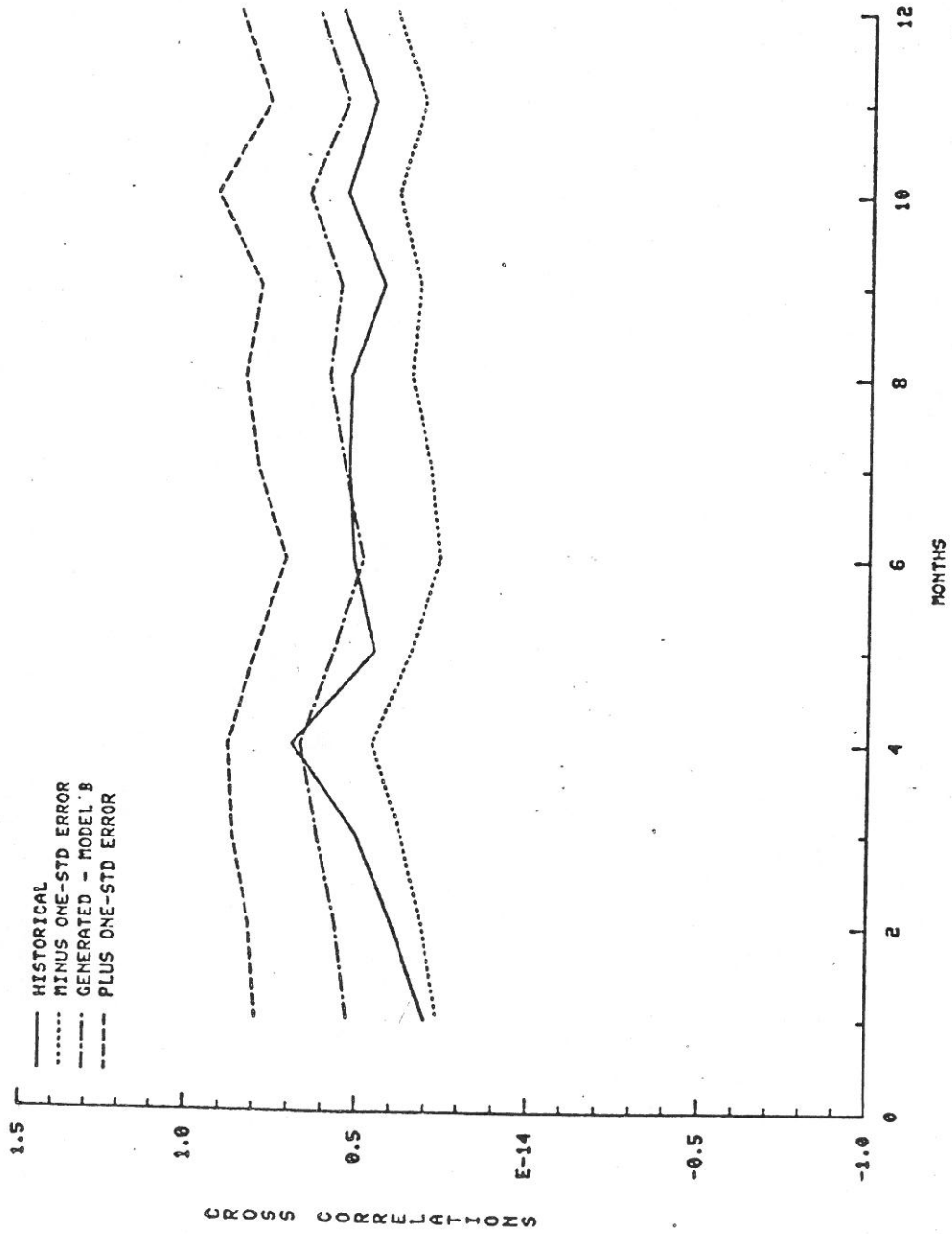


Figure 1.9.E.53. PALODE AND RANCHO - CROSS CORRELATIONS (LWH DOMAIN)

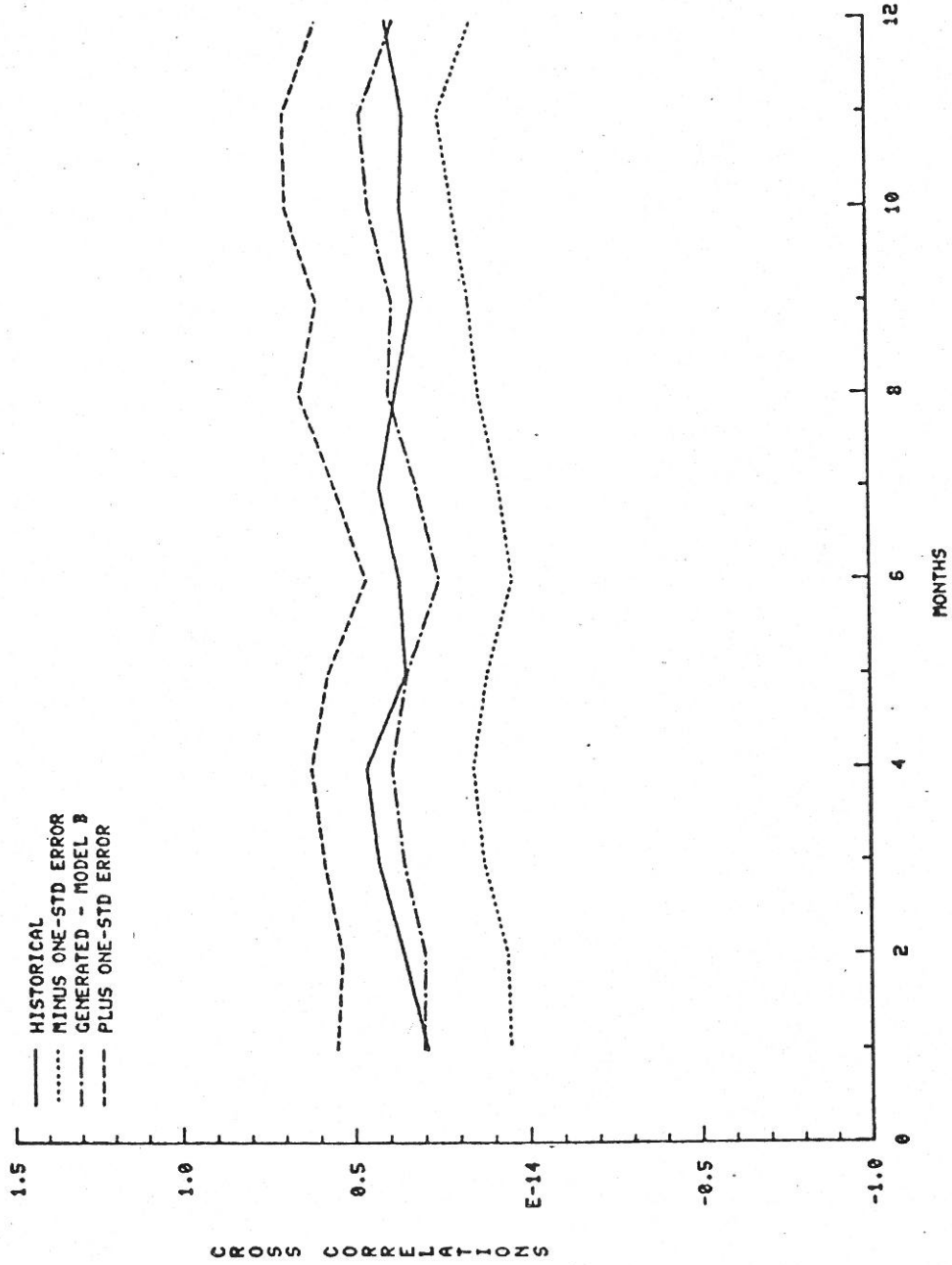


Figure 1.9.E.54. PASODE AND RANCHO - CROSS CORRELATIONS (LWH DOMAIN)

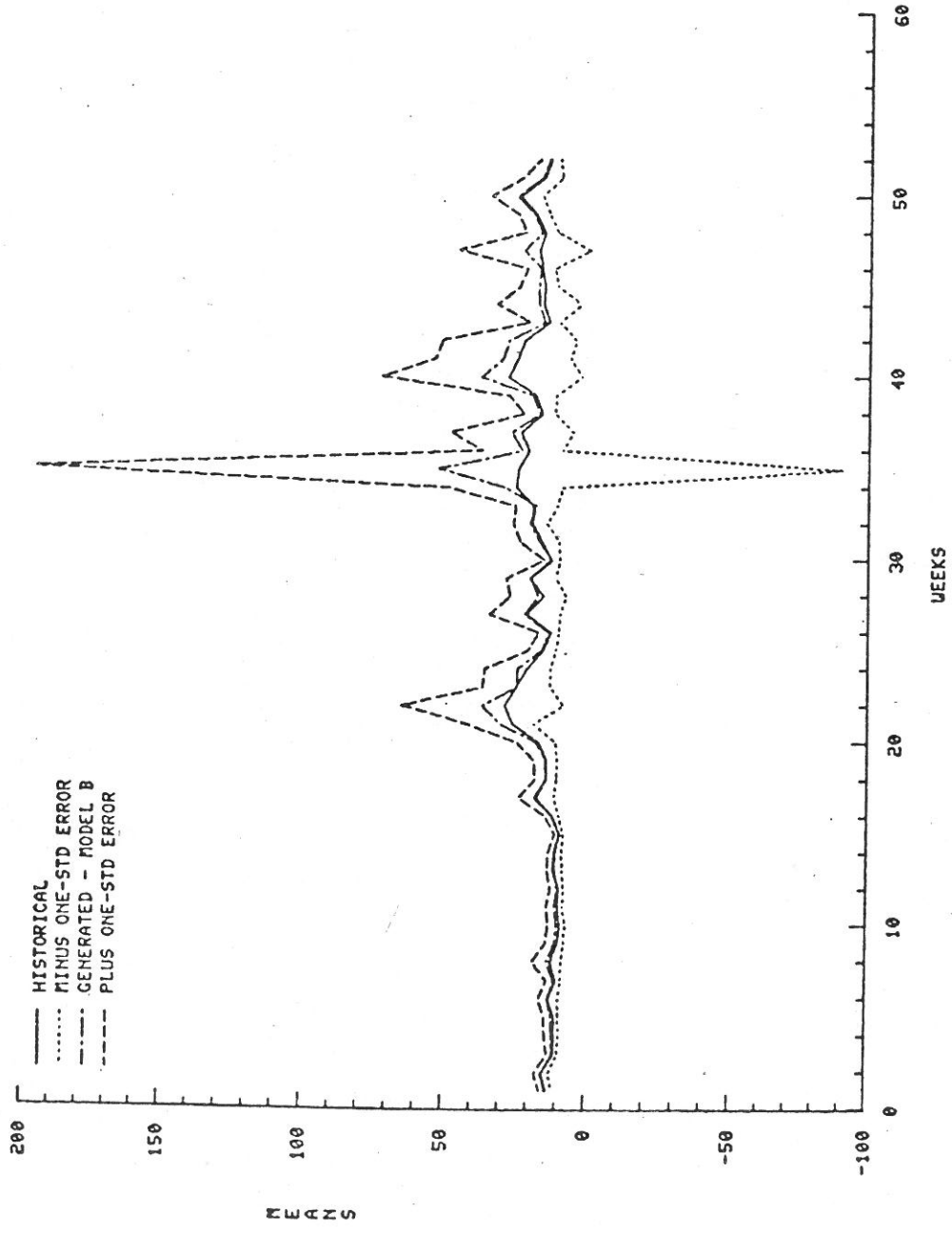


Figure 1.9.E.55. PALODE - MEANS (ORIGINAL DOMAIN)

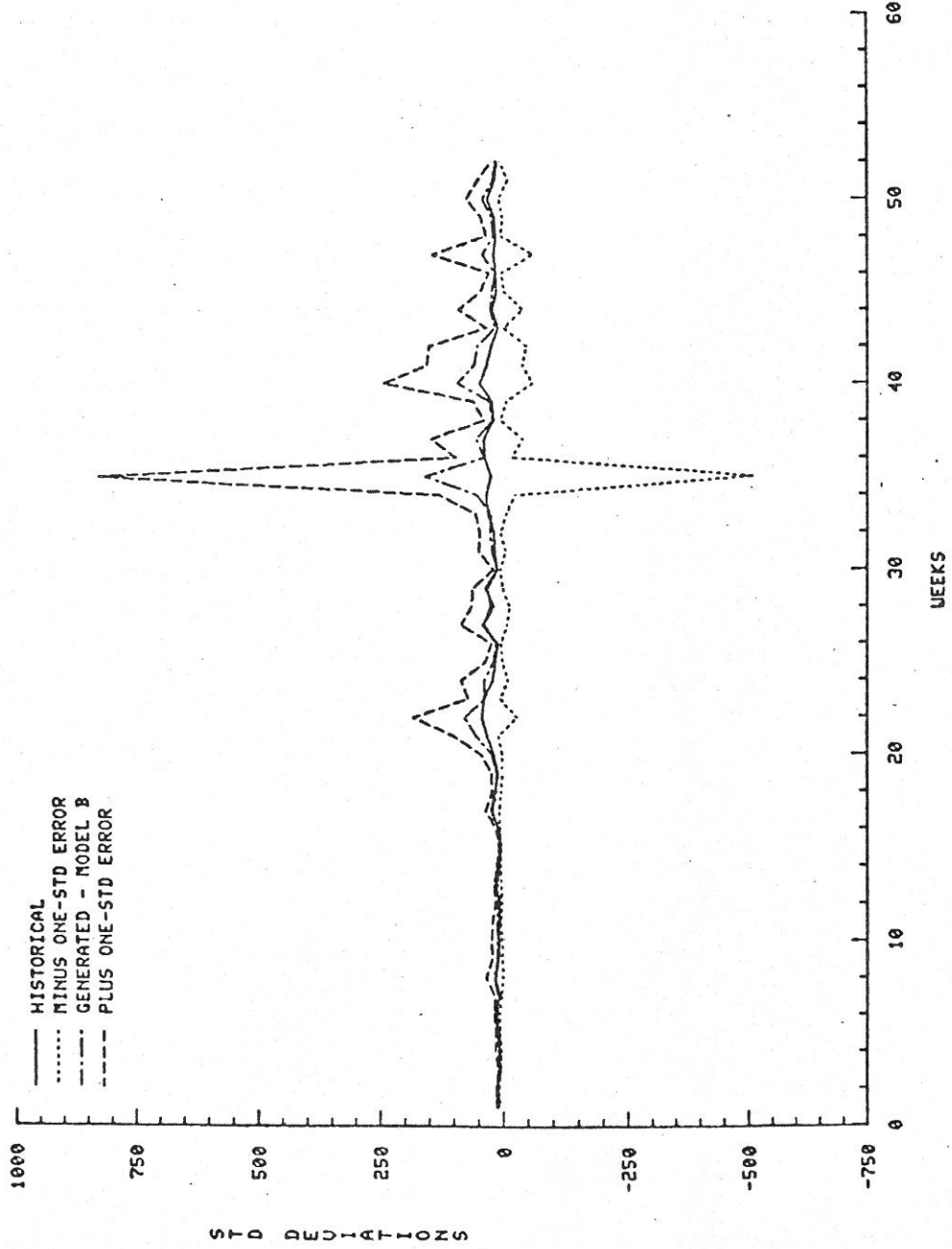


Figure 1.9.E.56. PALODE - STD DEVIATIONS (ORIGINAL DOMAIN)

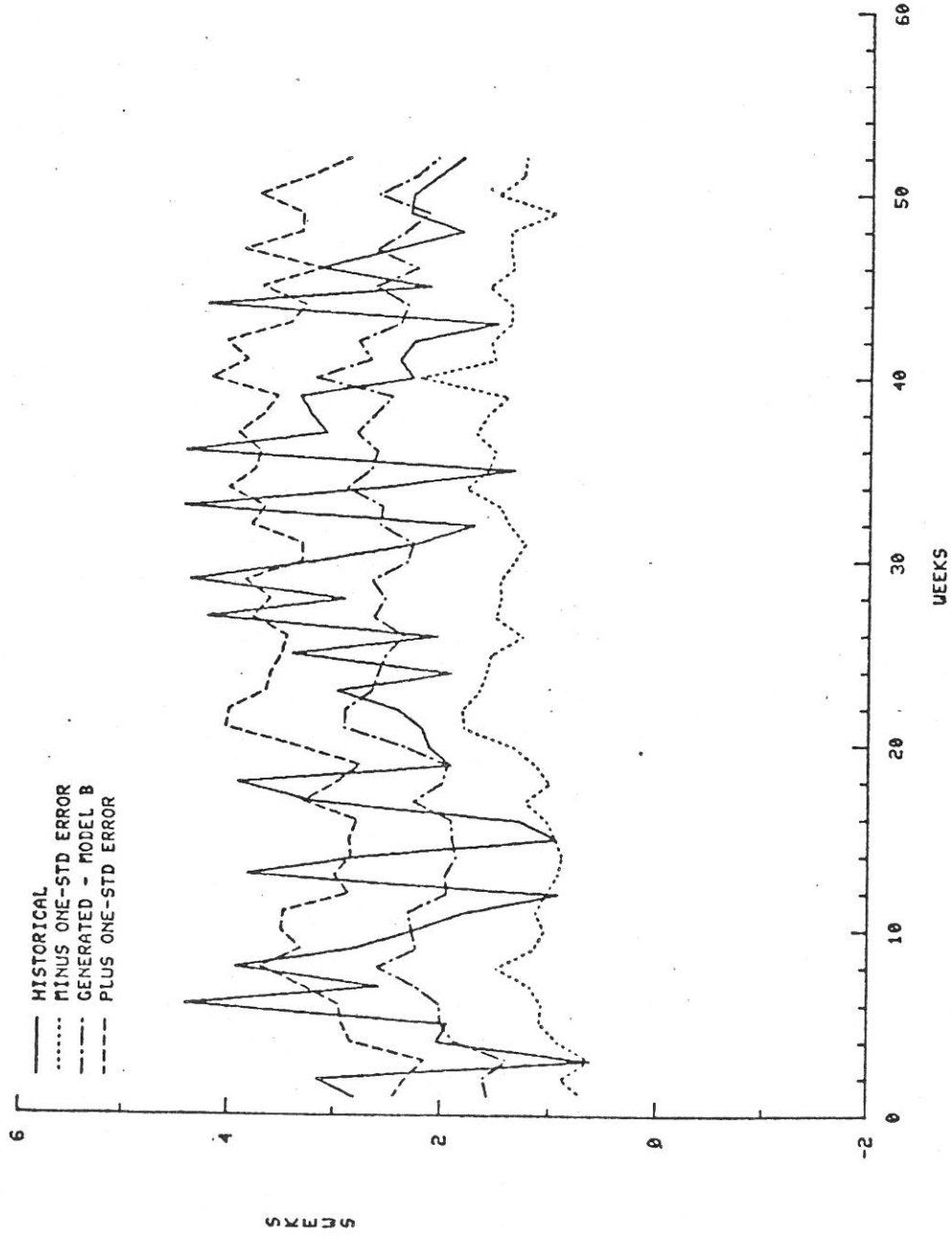


Figure 1.9.E.57. PALODE - SKEWS (ORIGINAL DOMAIN)

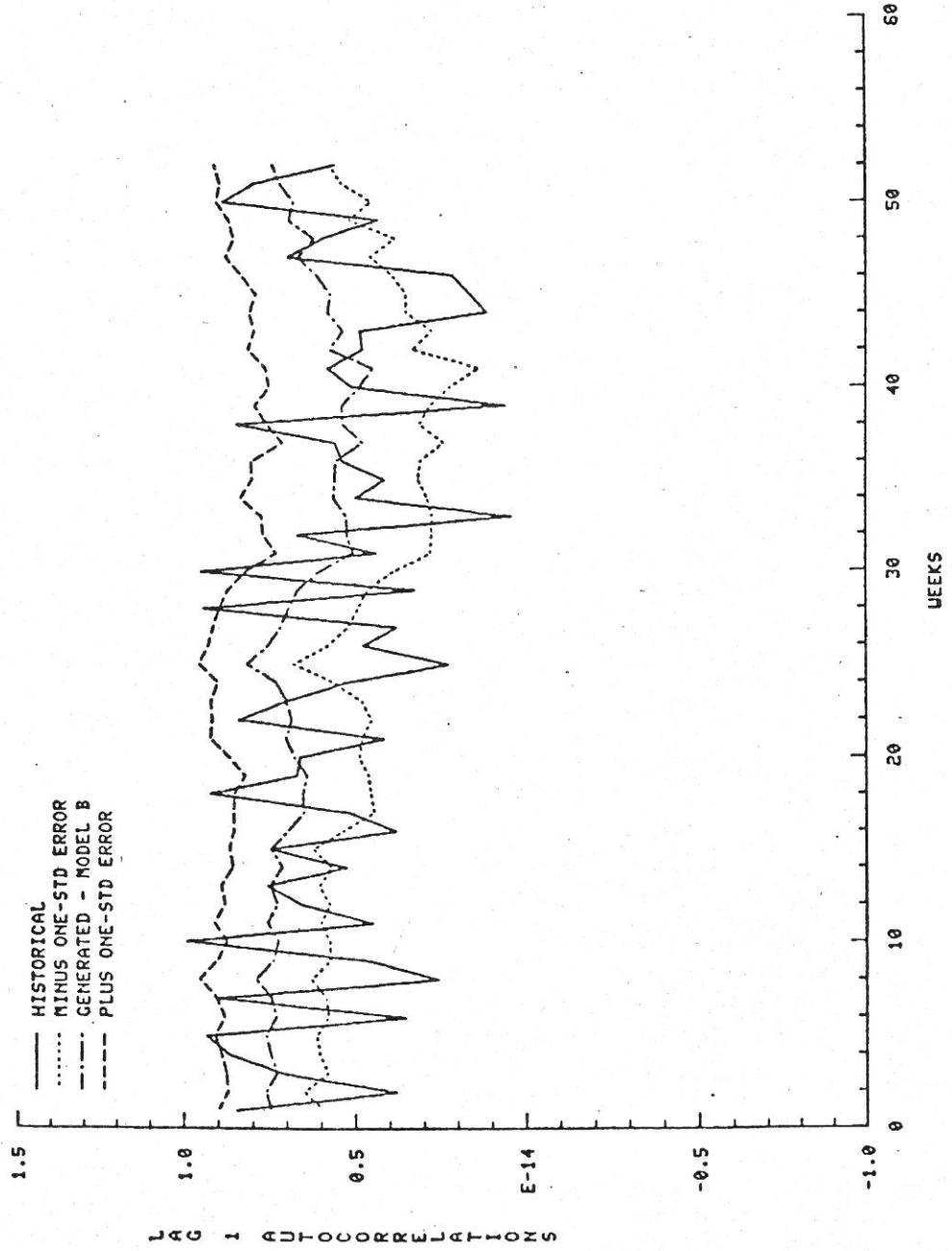


Figure 1.9.E.58. PALODE - LAG 1 AUTOCORRELATIONS (ORIGINAL DOMAIN)

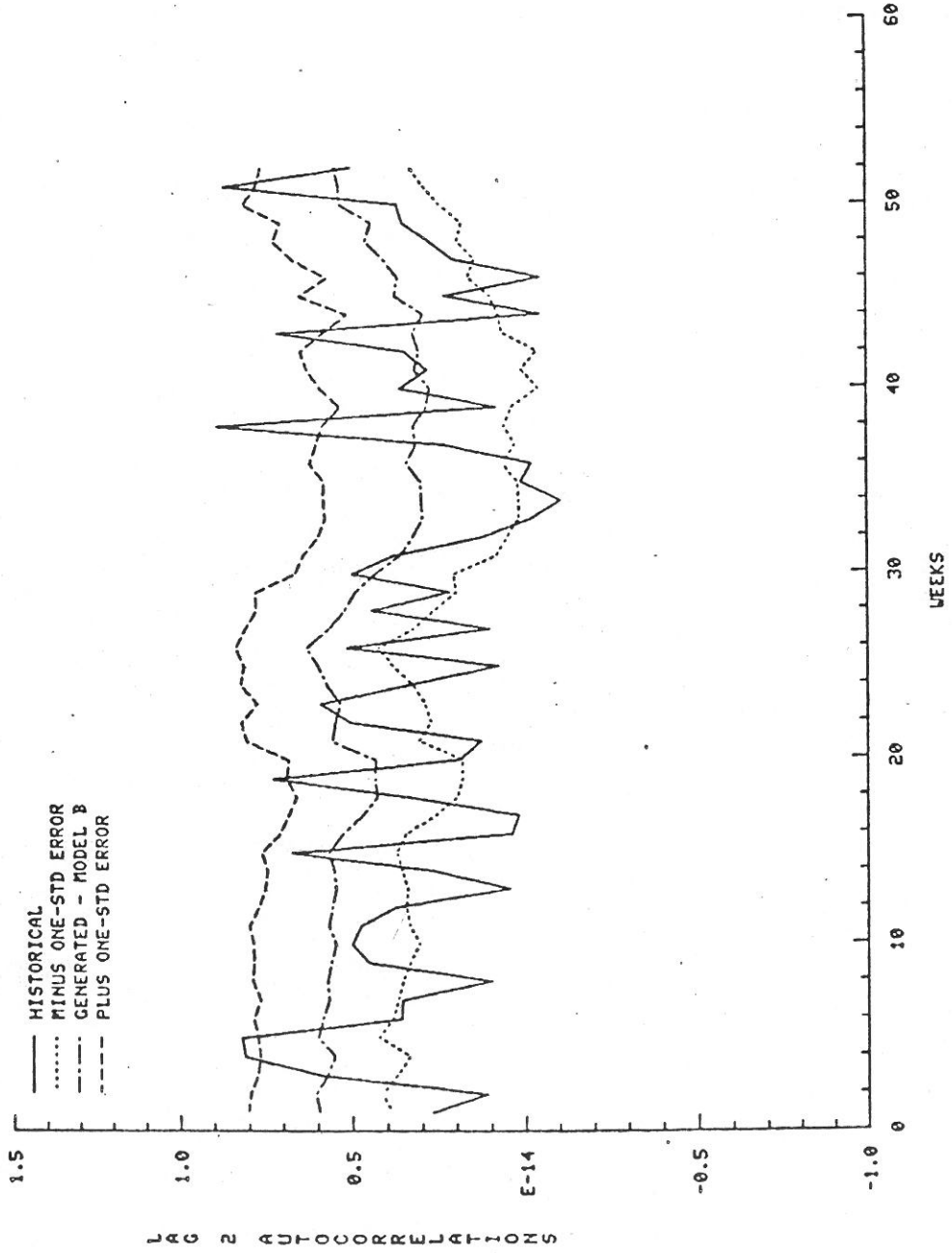


Figure 1.9.E.59. PALODE - LAG 2 AUTOCORRELATIONS (ORIGINAL DOMAIN)

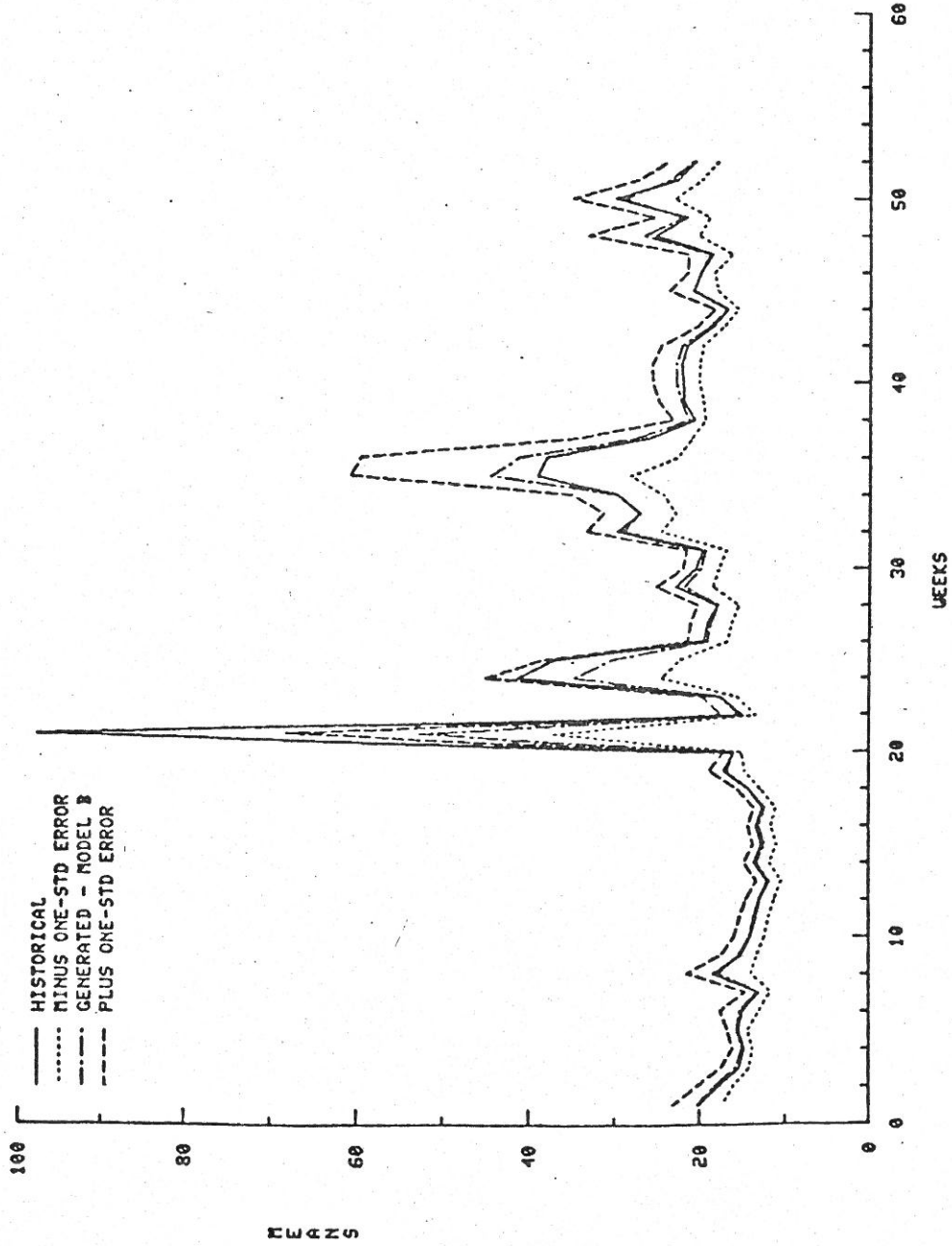


Figure 1.9.E.60. PASODE - MEANS (ORIGINAL DOMAIN)

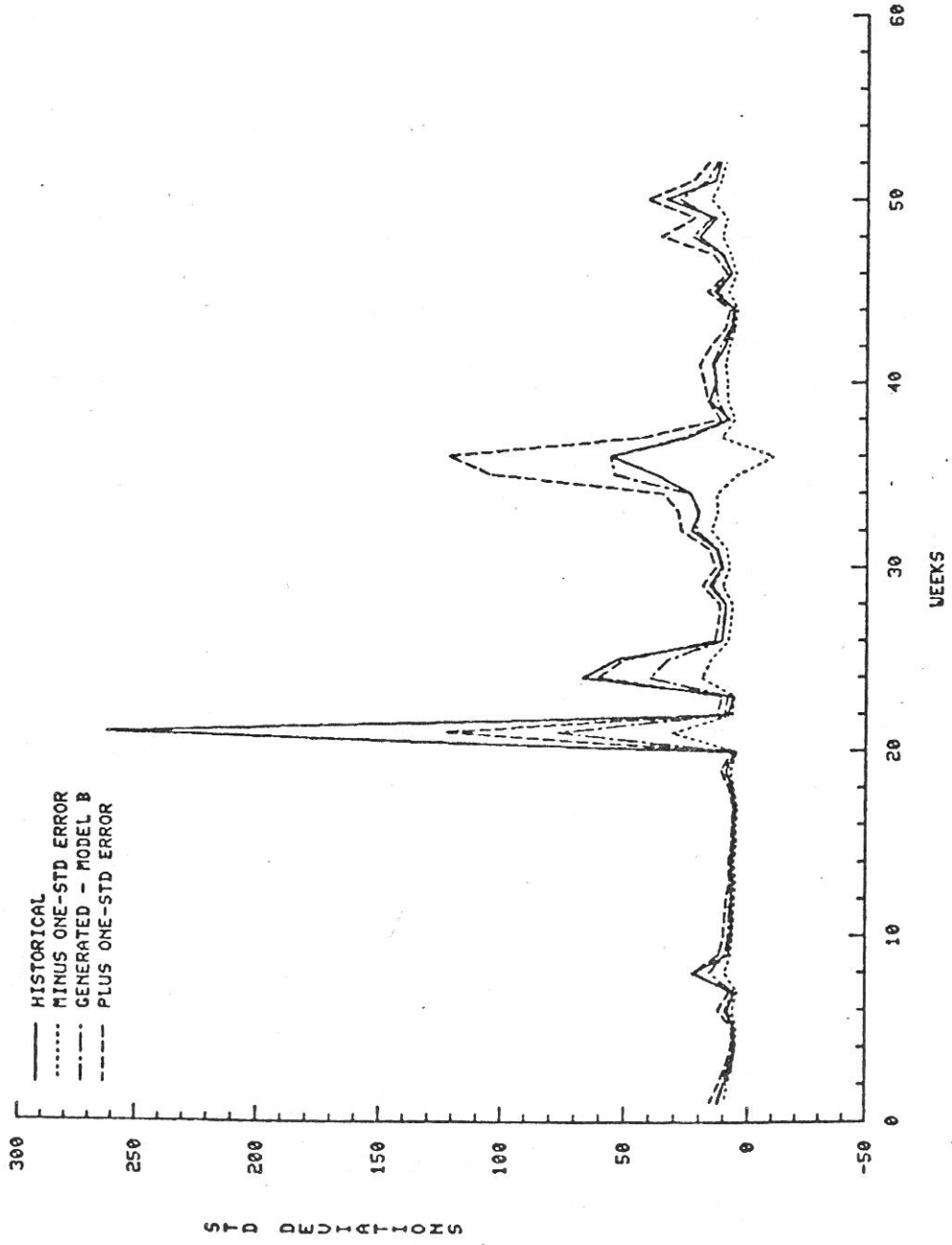


Figure 1.9.E.61. PASODE - STD DEVIATIONS (ORIGINAL DOMAIN)

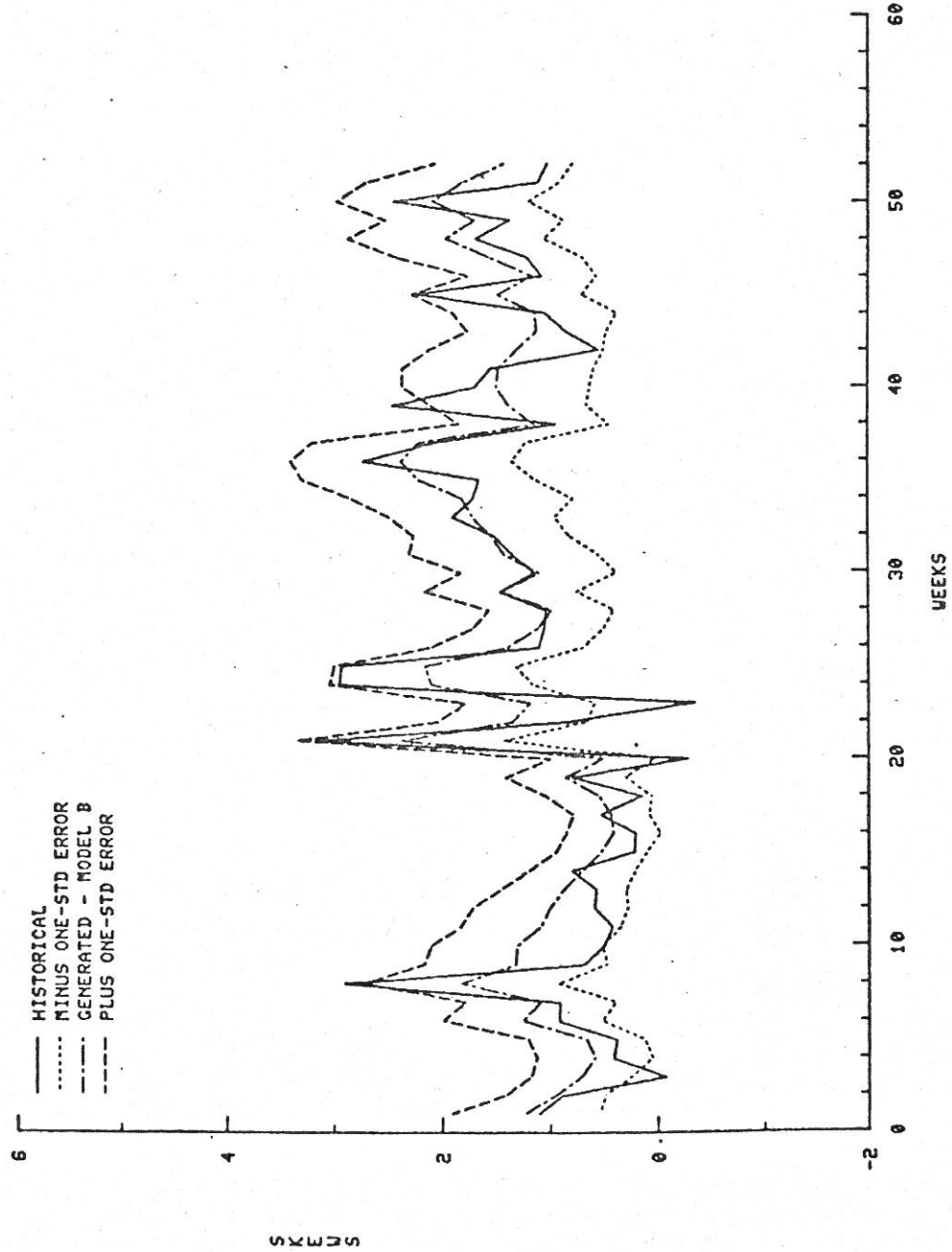


Figure 1.9.E.62. PASODE - SKEWS (ORIGINAL DOMAIN)

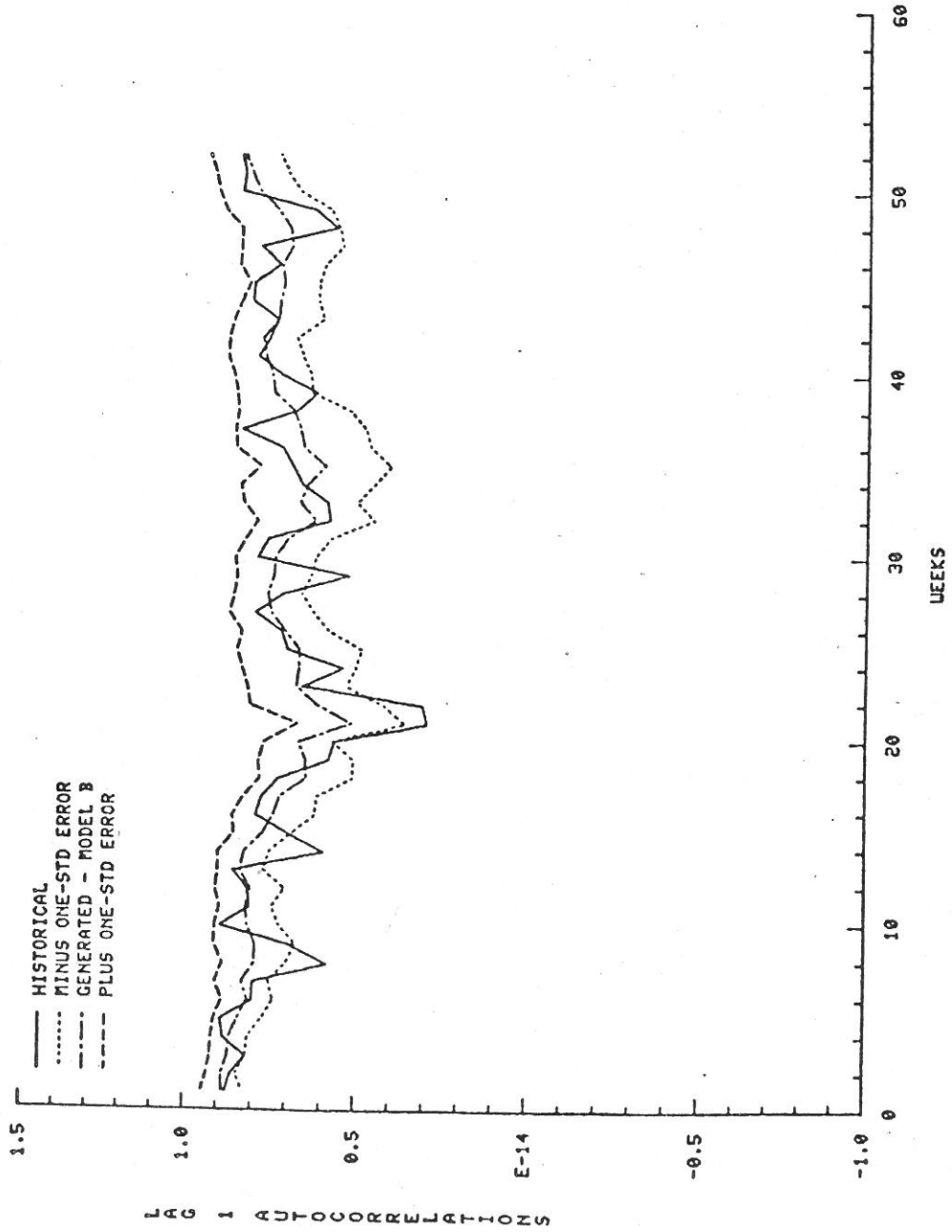


Figure 1.9.E.63. PASODE - LAG 1 AUTOCORRELATIONS (OR

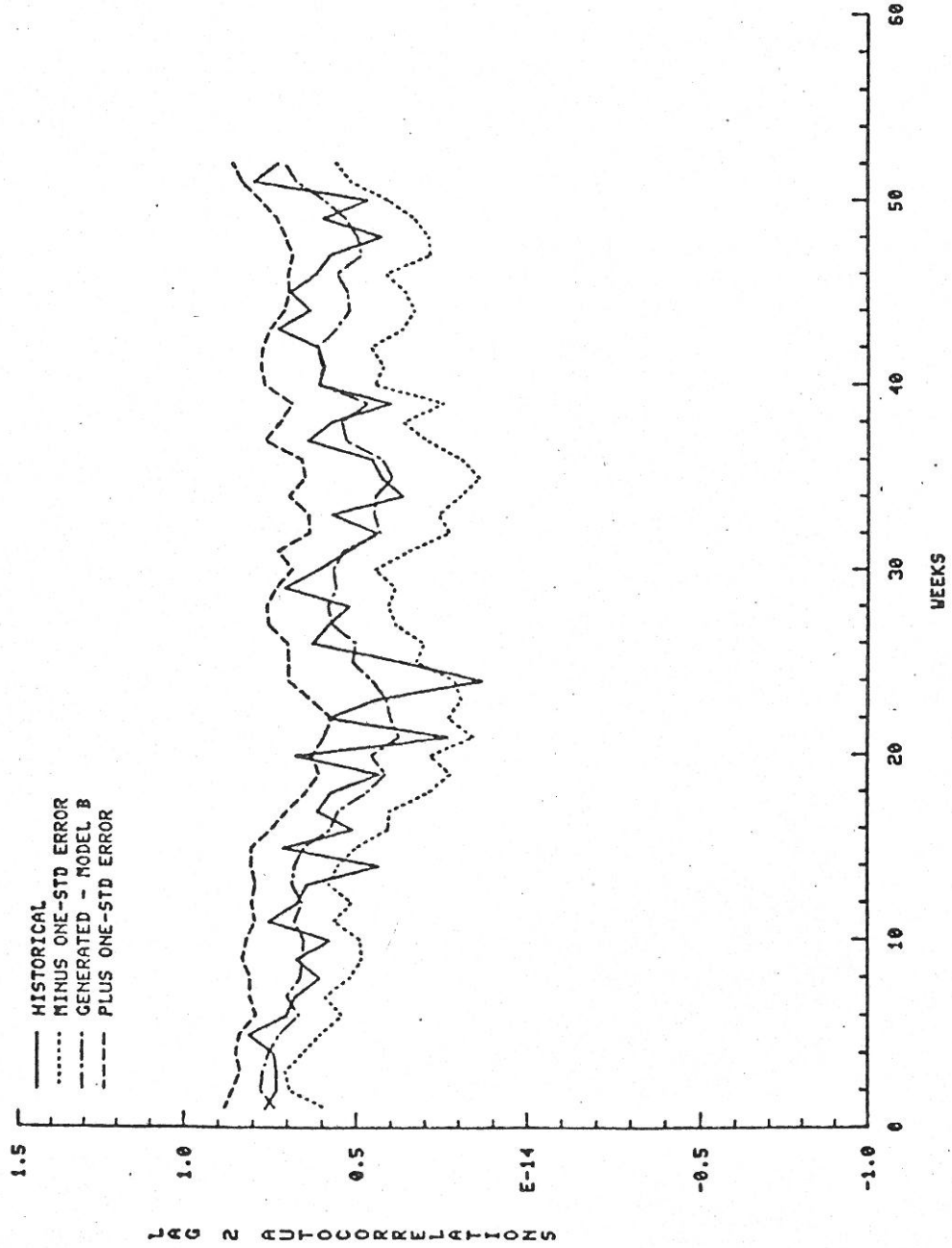


Figure 1.9.E.64. PASODE - LAG 2 AUTOCORRELATIONS (ORIGINAL DOMAIN)

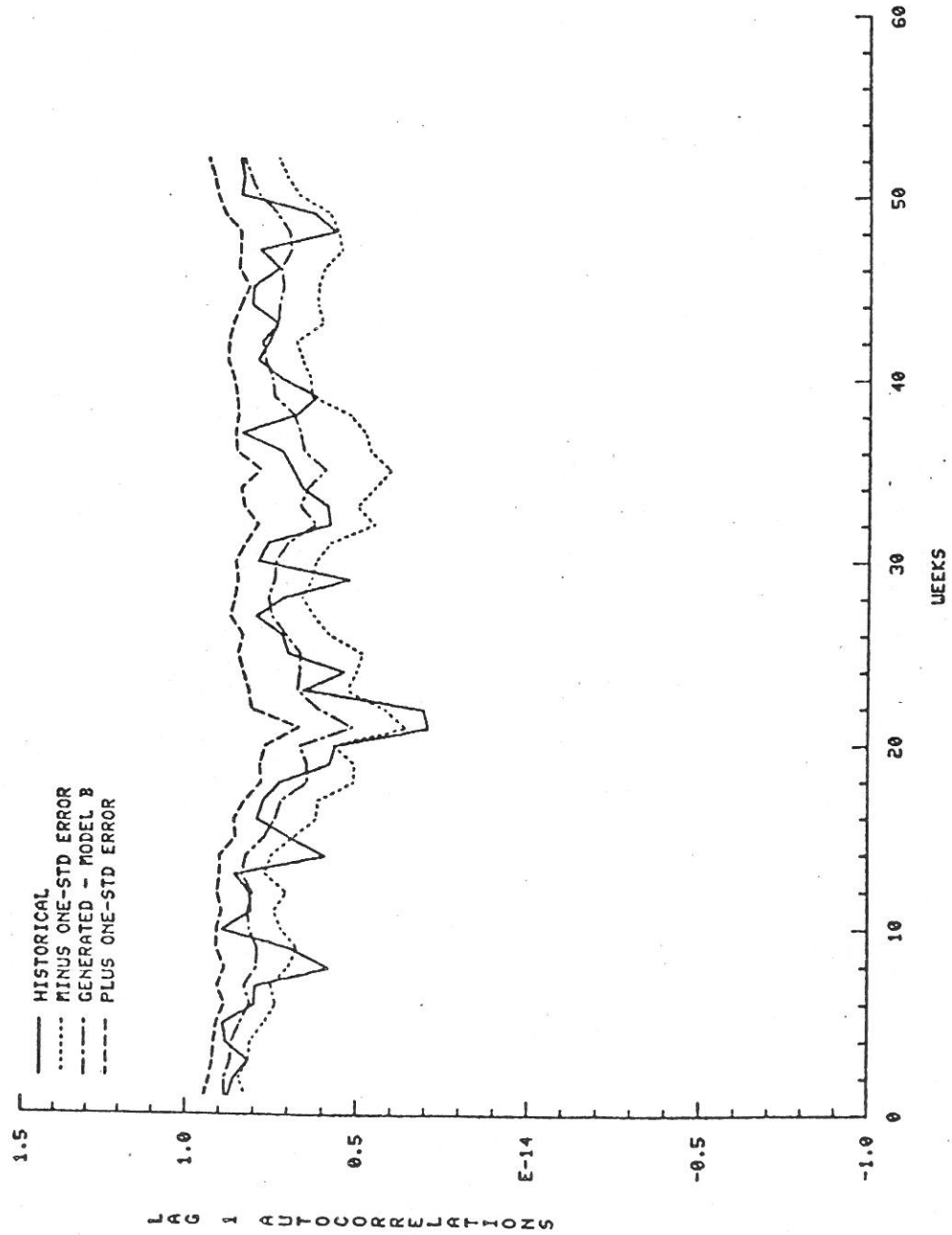


Figure 1.9.E.63. PASODE - LAG 1 AUTOCORRELATIONS (OR

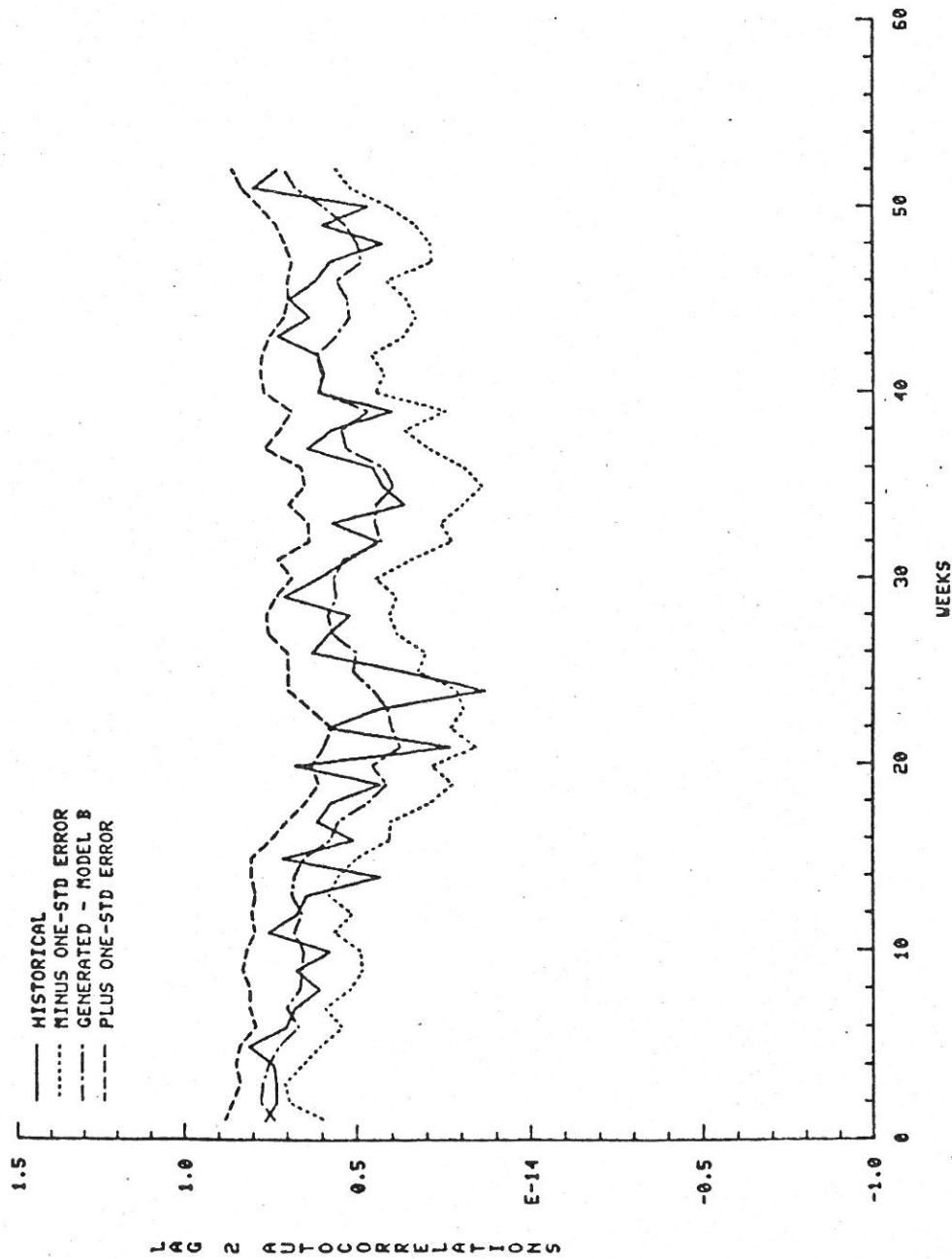


Figure 1.9.E.64. PASODE - LAG 2 AUTOCORRELATIONS (ORIGINAL DOMAIN)

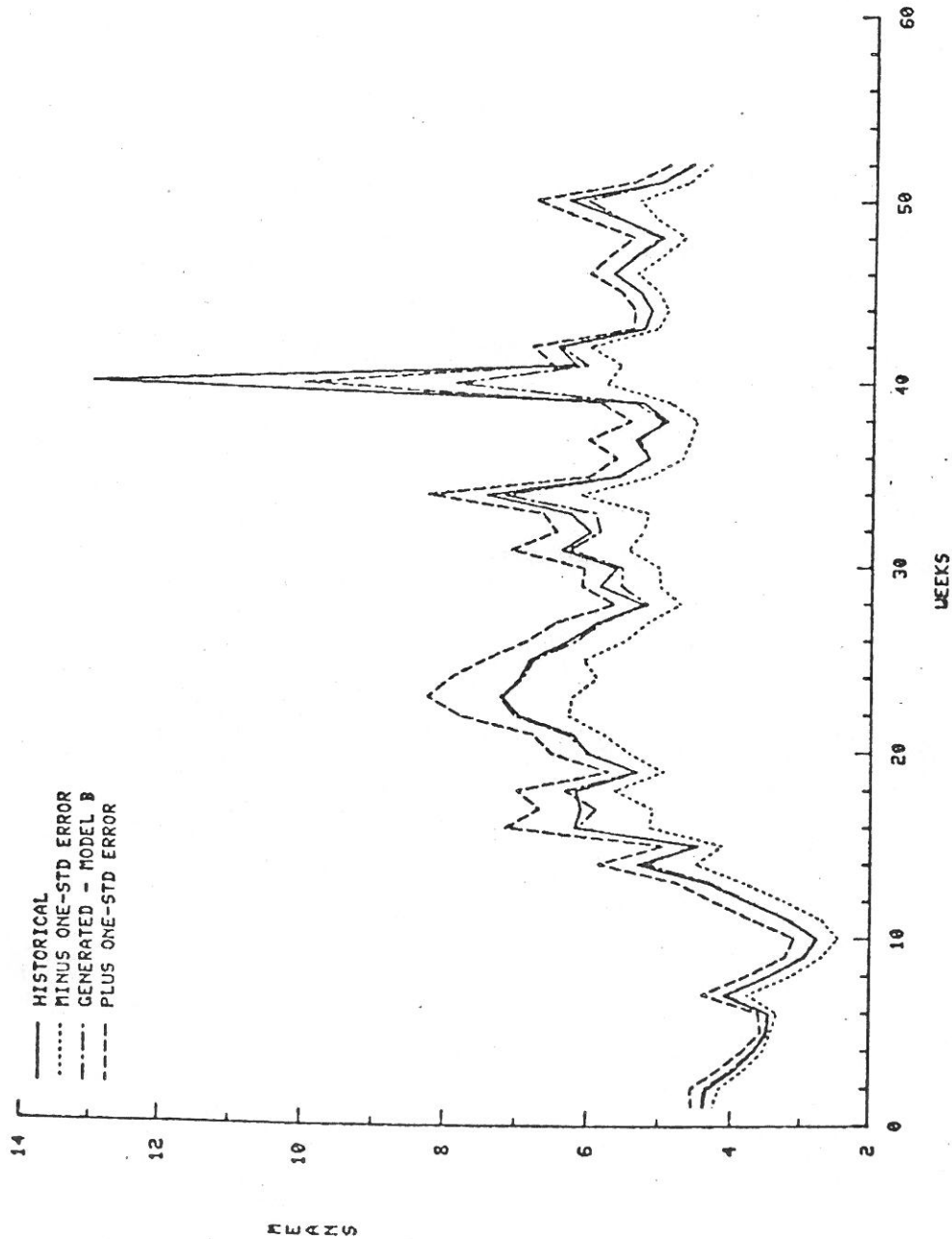


Figure 1.9.E.65. RANCHO - MEANS (ORIGINAL DOMAIN)

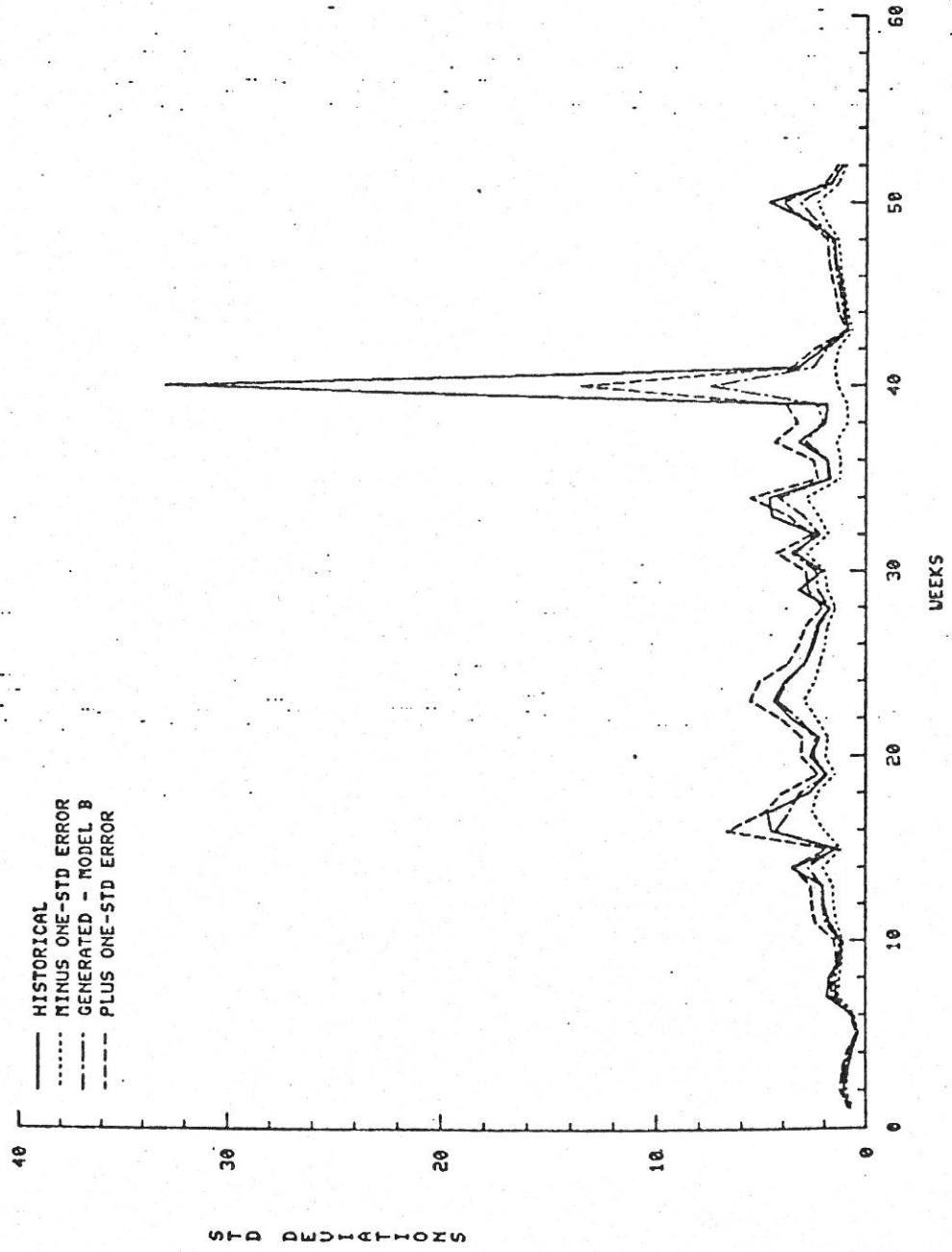


Figure 1.9.E.66. RANCHO - STD DEVIATIONS (ORIGINAL DOMAIN)

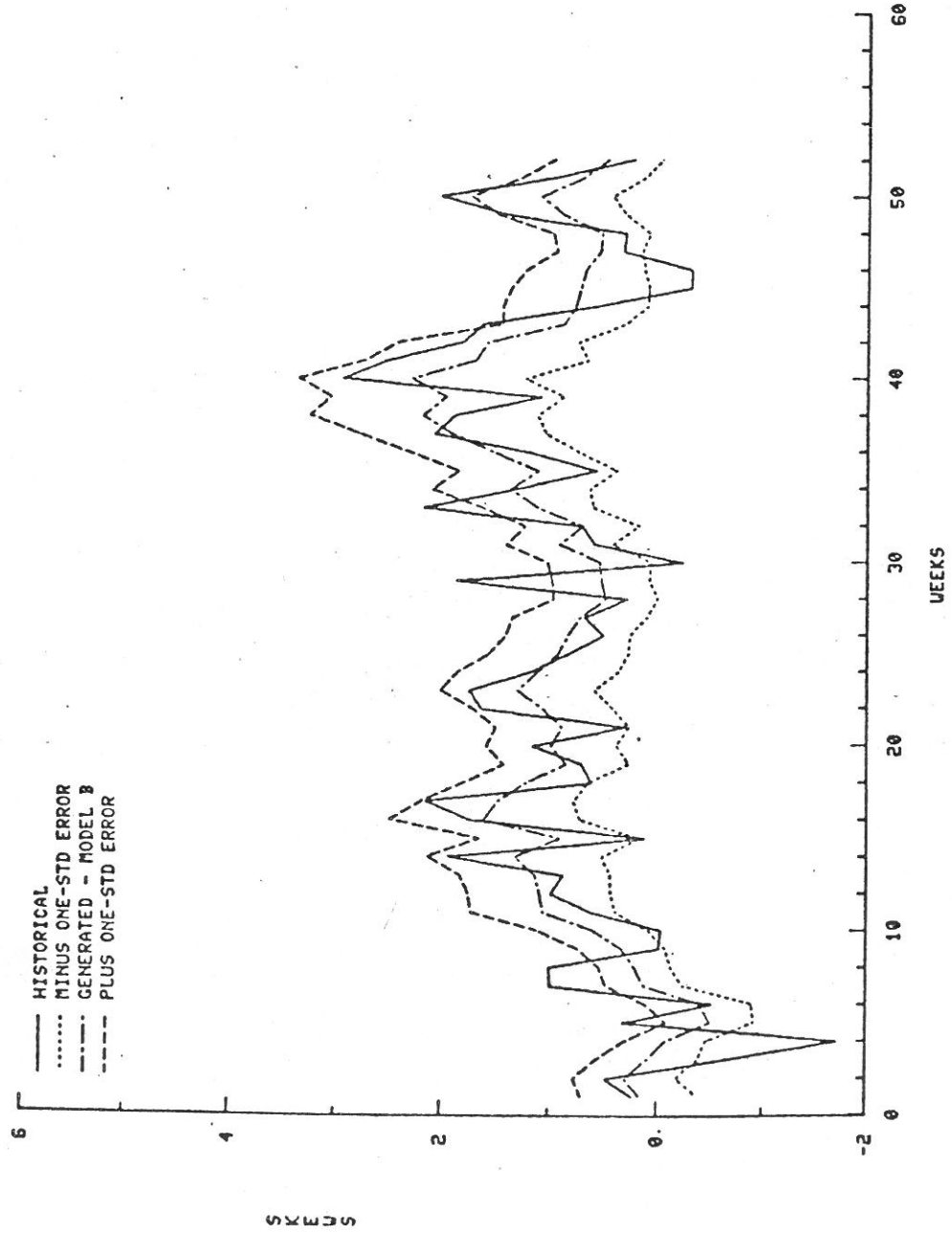


Figure 1.9.E.67. RANCHO - SKEWS (ORIGINAL DOMAIN)

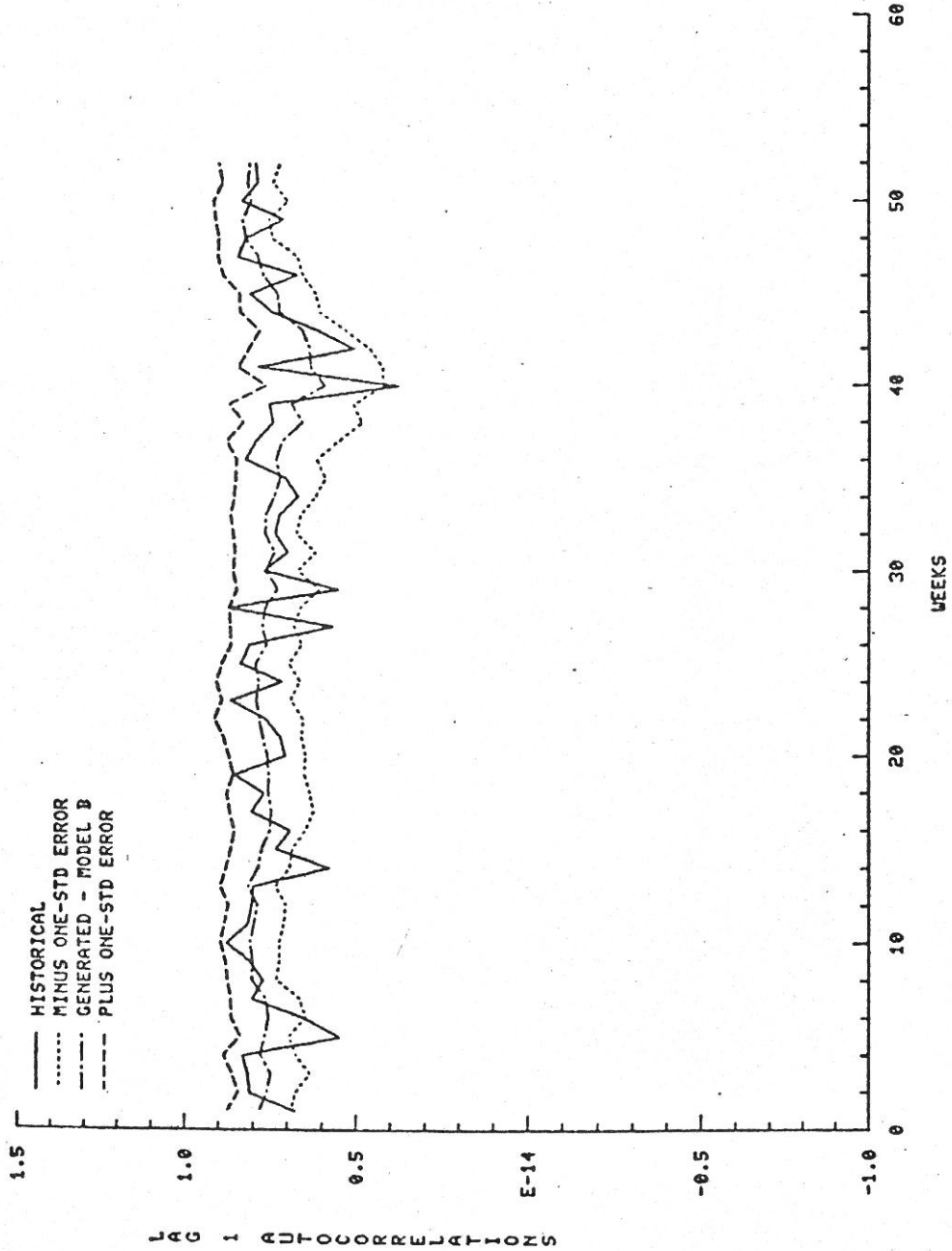


Figure 1.9.E.68. RANCHO - LAG 1 AUTOCORRELATIONS (ORIGINAL DOMAIN)

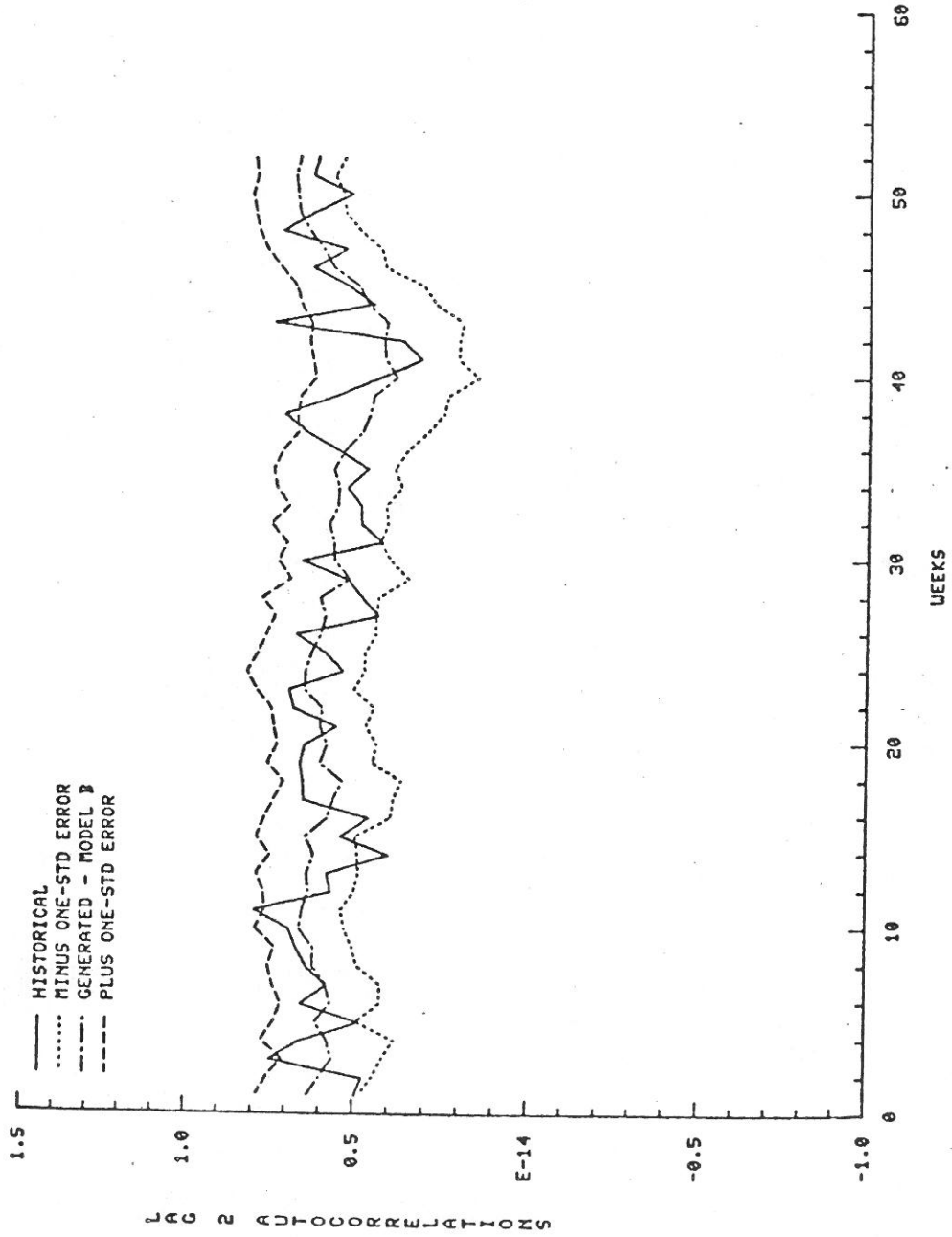


Figure 1.9.E.69. RANCHO - LAG 2 AUTOCORRELATIONS (ORIGINAL DOMAIN)

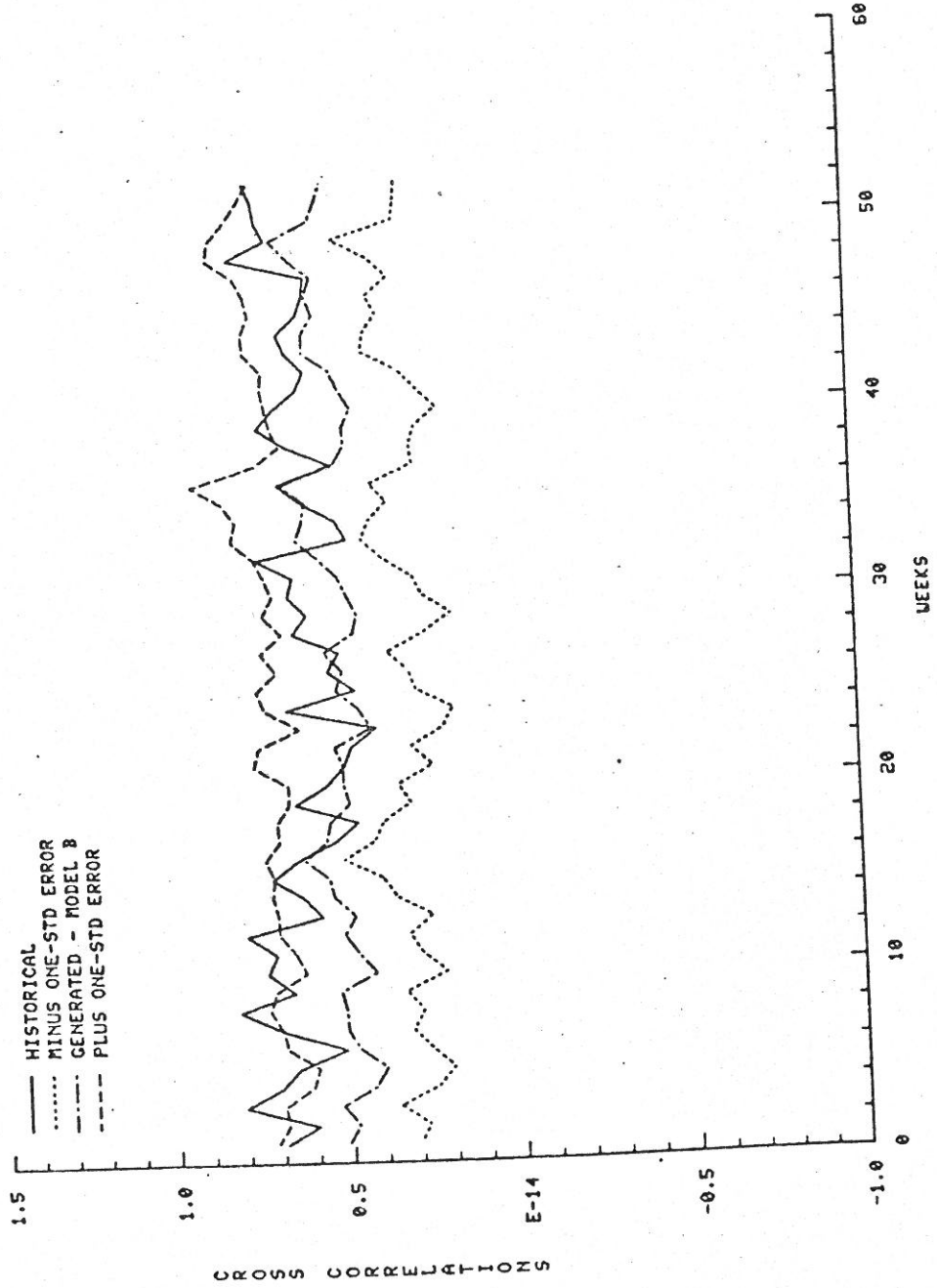


Figure 1.9.E.70. PALODE AND PASODE - CROSS CORRELATIONS (ORIGINAL DOMAIN)

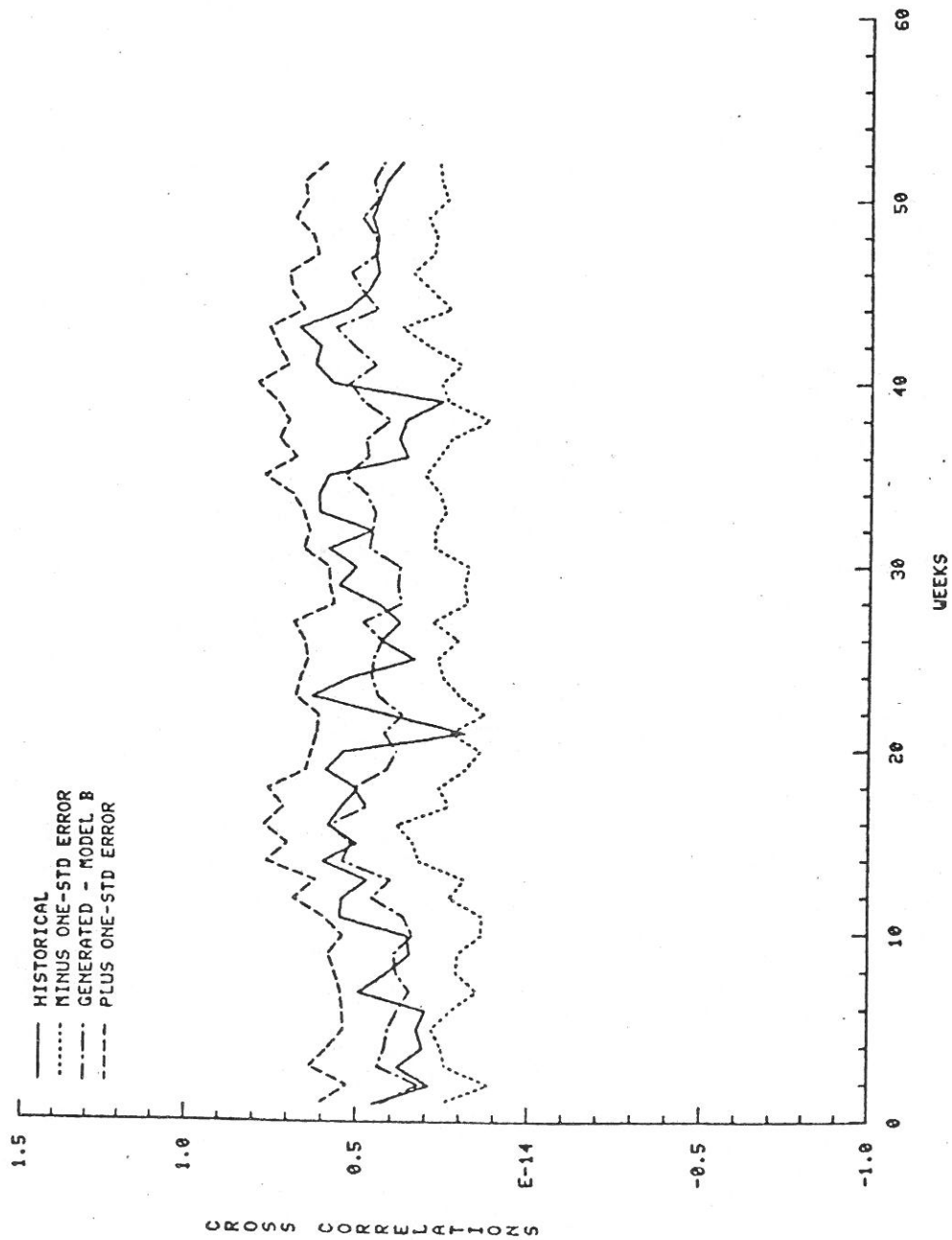


Figure 1.9.E.71. PALODE AND RANCHO - CROSS CORRELATIONS (ORIGINAL DOMAIN)

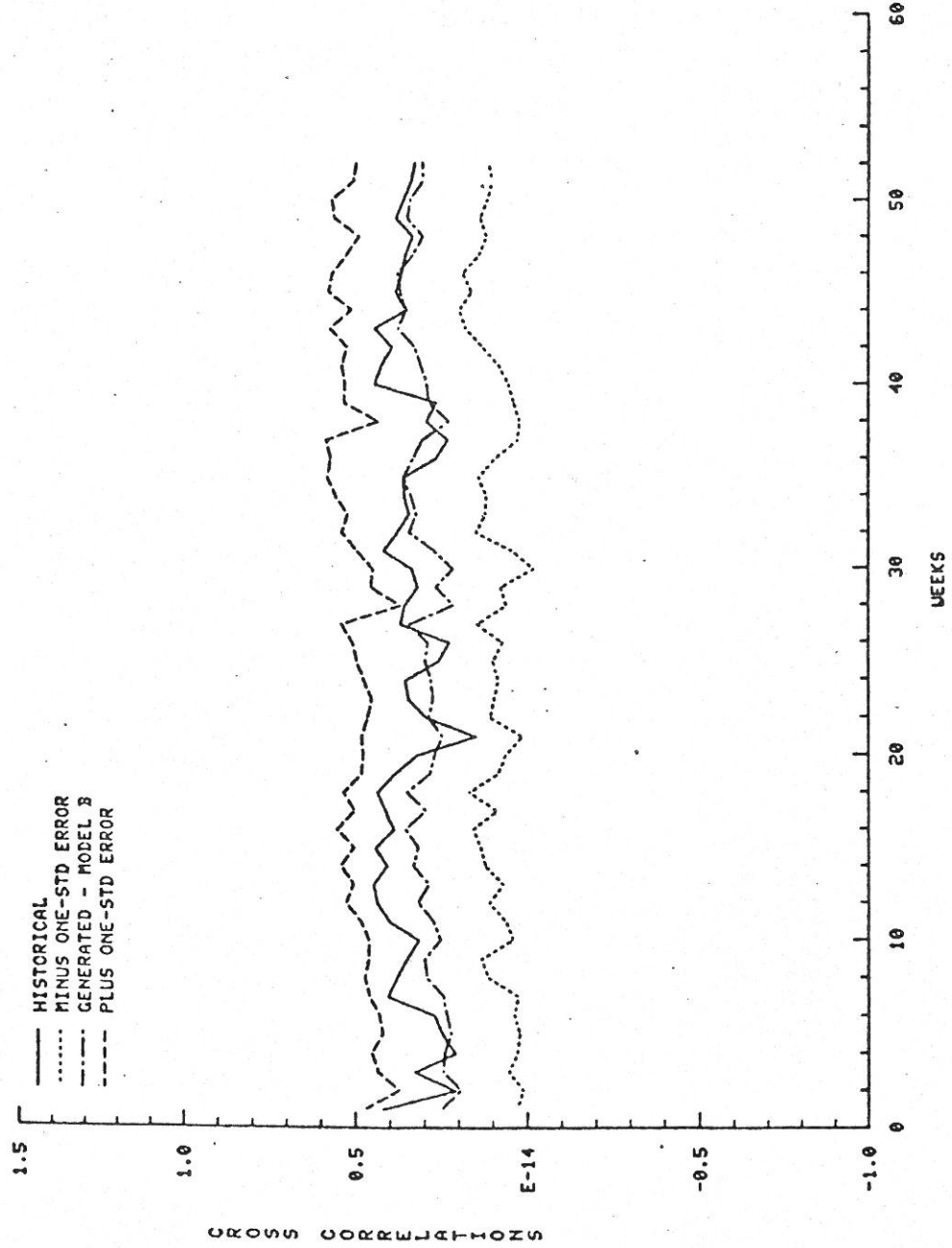


Figure 1.9.E.72. PASODE AND RANCHO - CROSS CORRELATIONS (ORIGINAL DOMAIN)

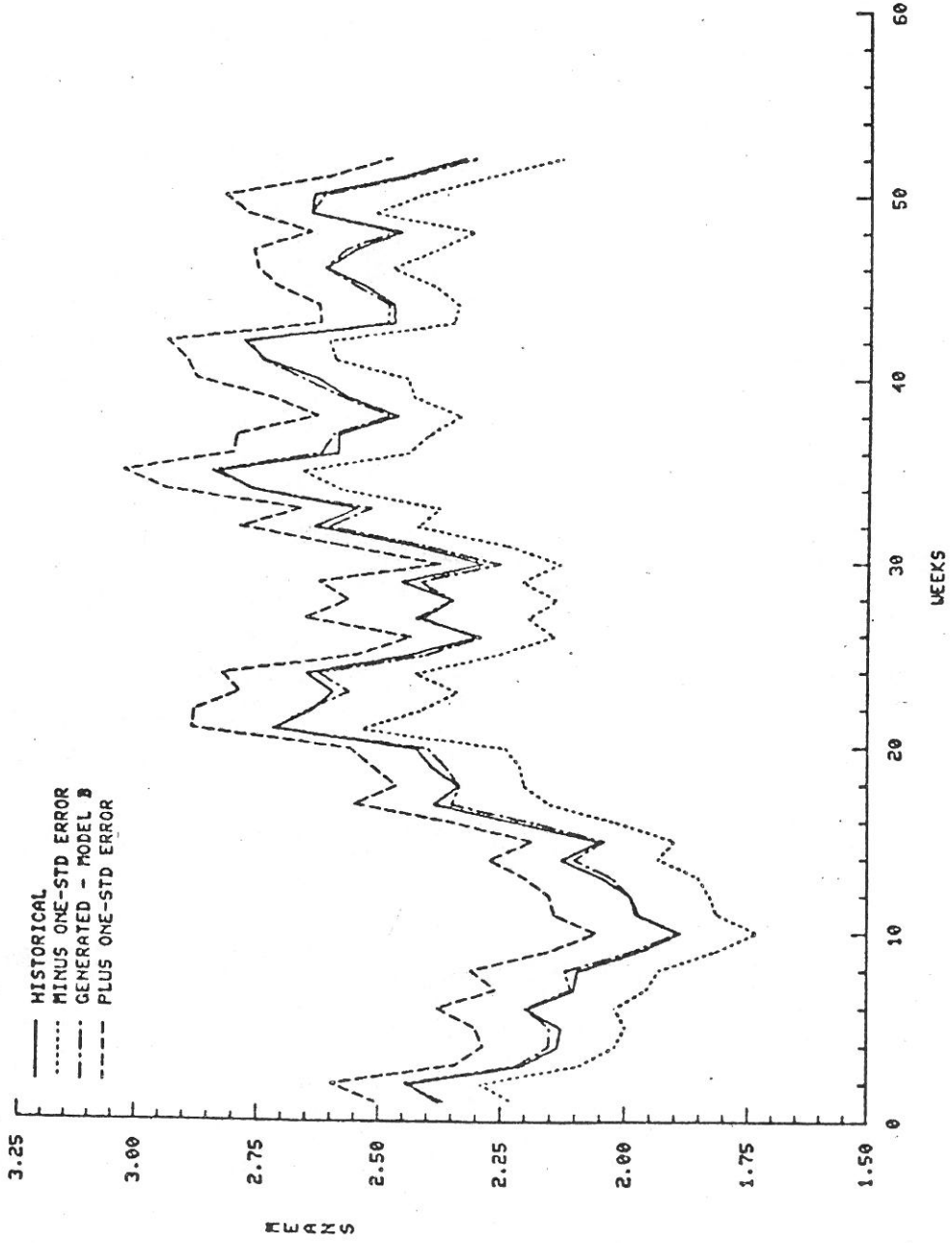


Figure 1.9.E.73. PALODE - MEANS (LOG DOMAIN)

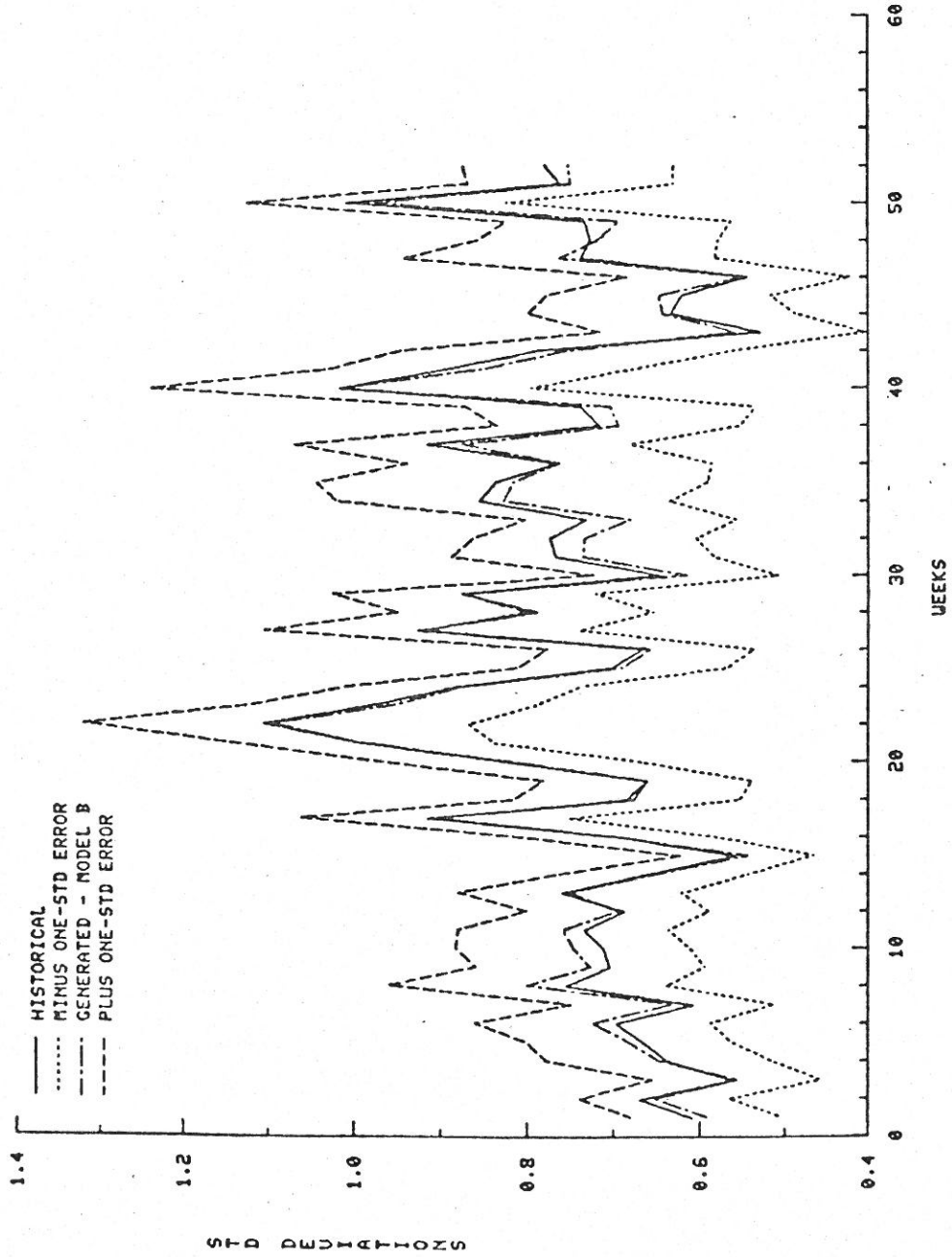


Figure 1.9.E.74. PALODE - STD DEVIATIONS (LOG DOMAIN)

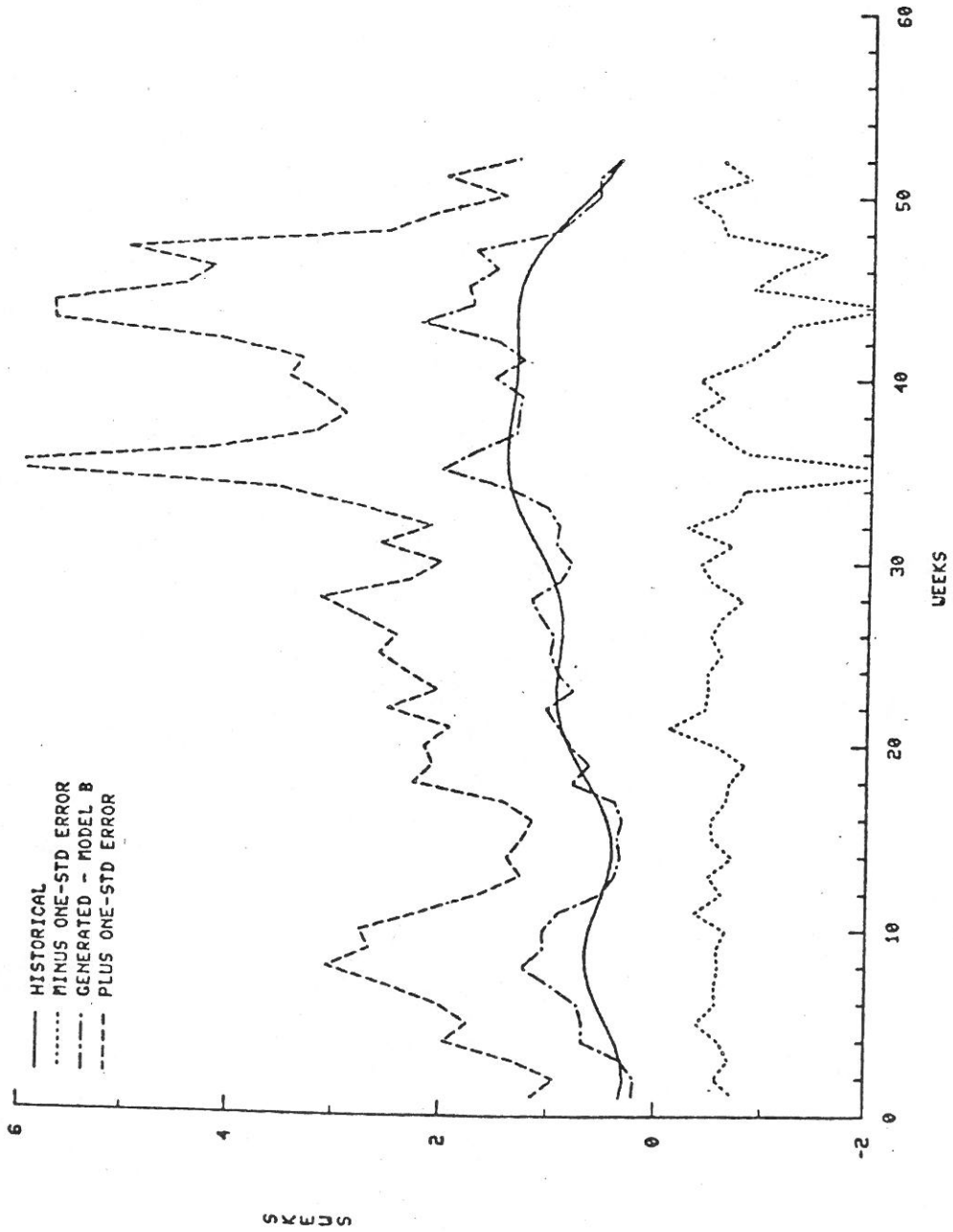


Figure 1.9.E.75. PALODE - SKEWS (LOG DOMAIN)

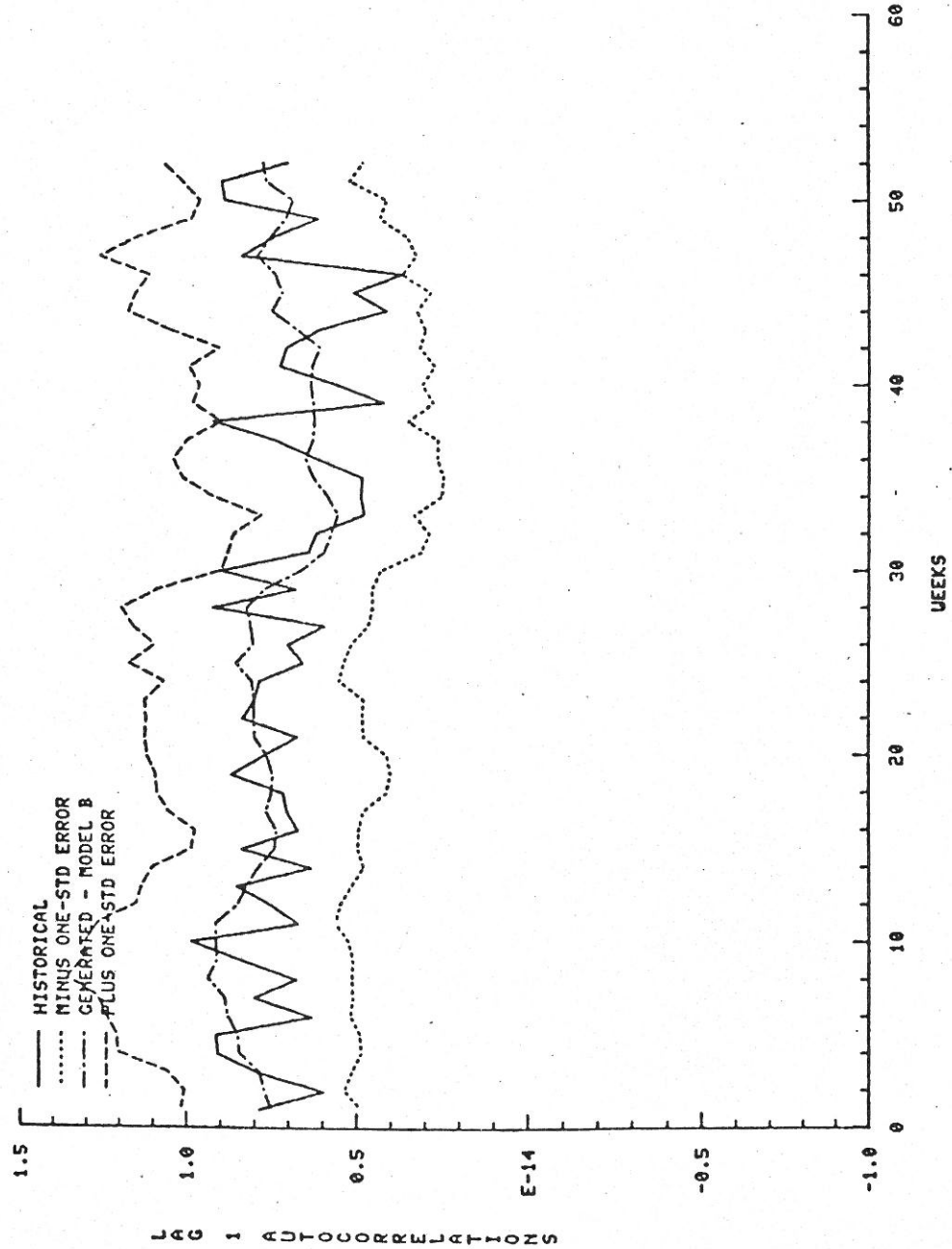


Figure 1.9.E.76. PALODE - LAG 1 AUTOCORRELATIONS (LOG DOMAIN)

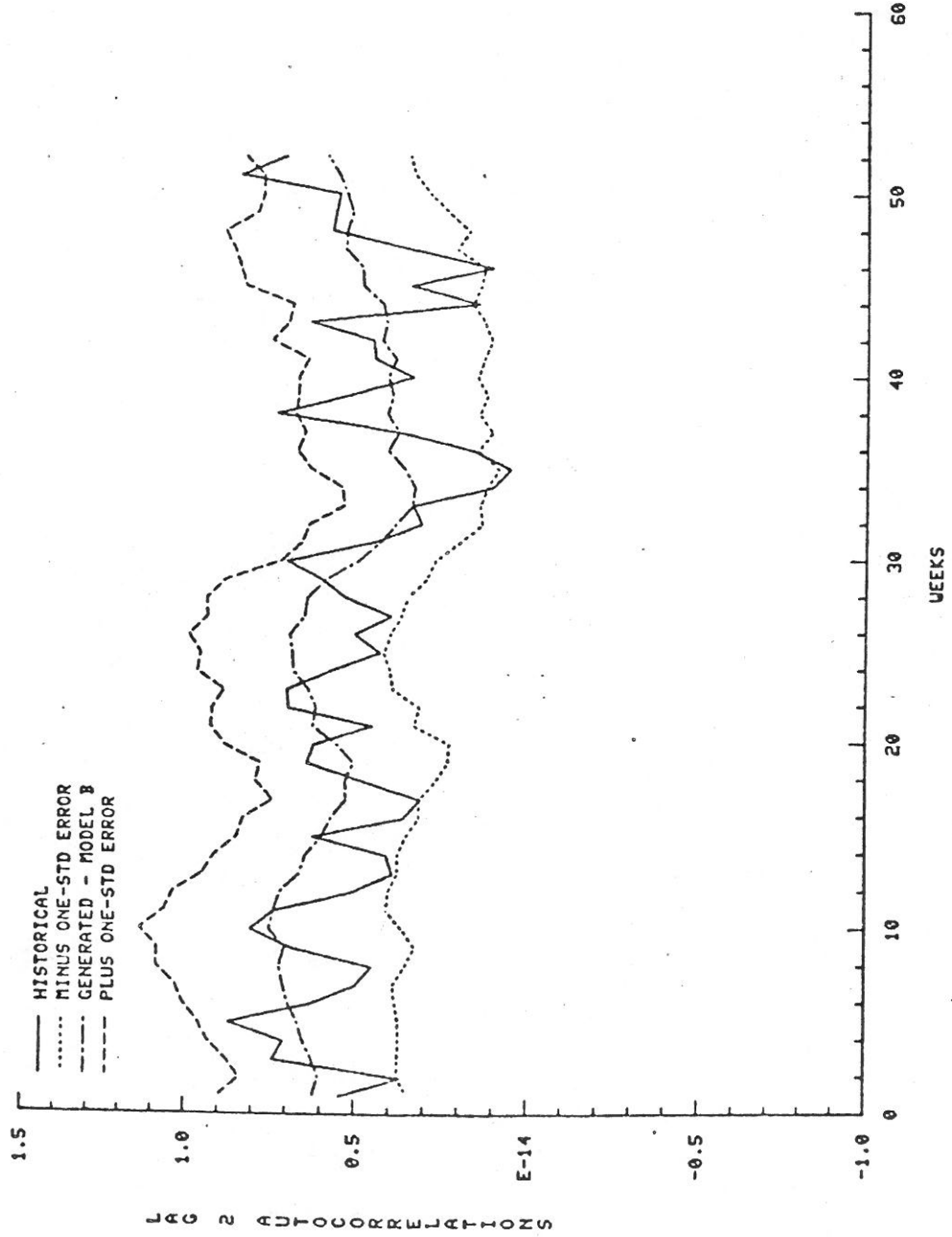


Figure 1.9.E.77. PALODE - LAG 2 AUTOCORRELATIONS (LOG DOMAIN)

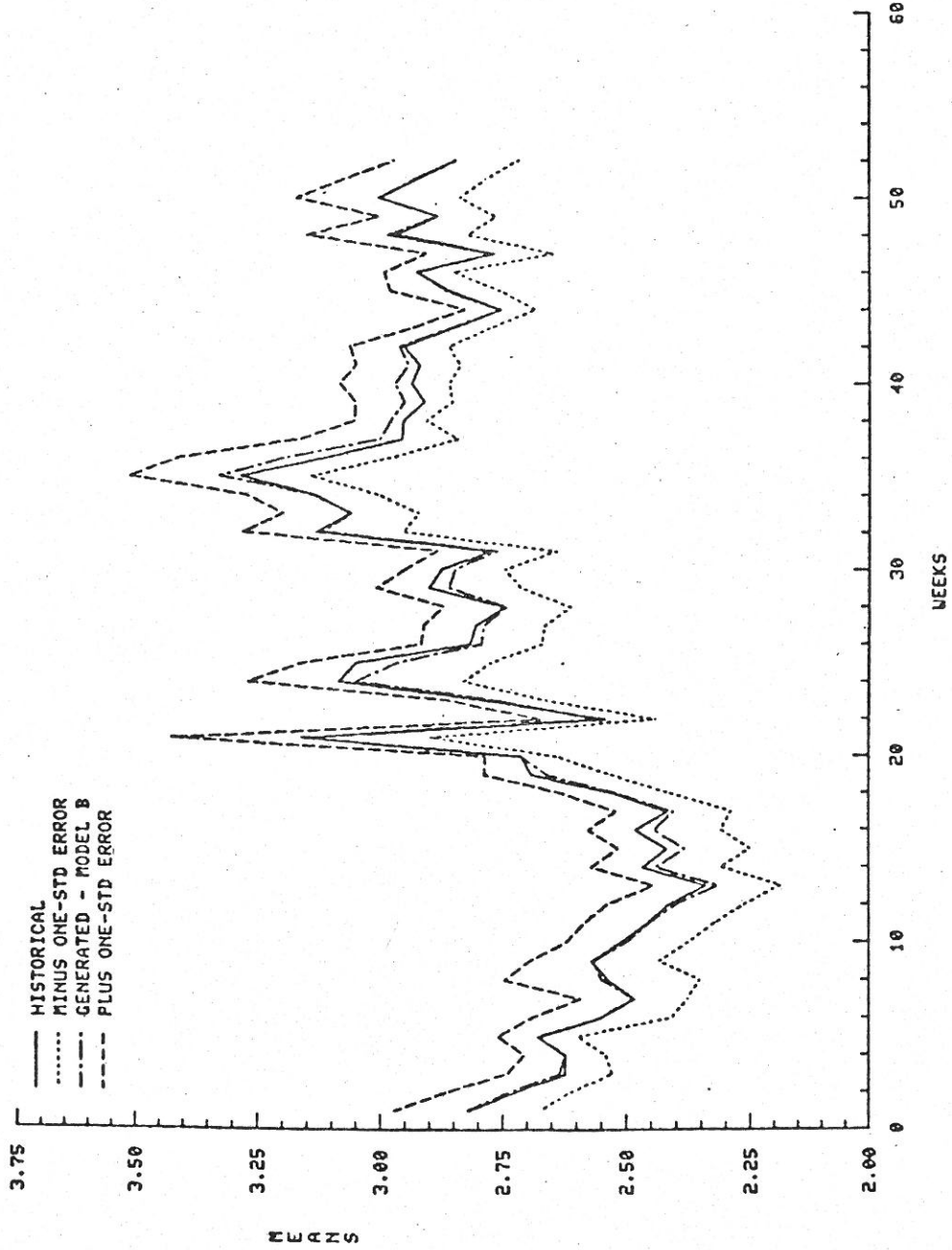


Figure 1.9.E.78. PASODE - MEANS (LOG DOMAIN)

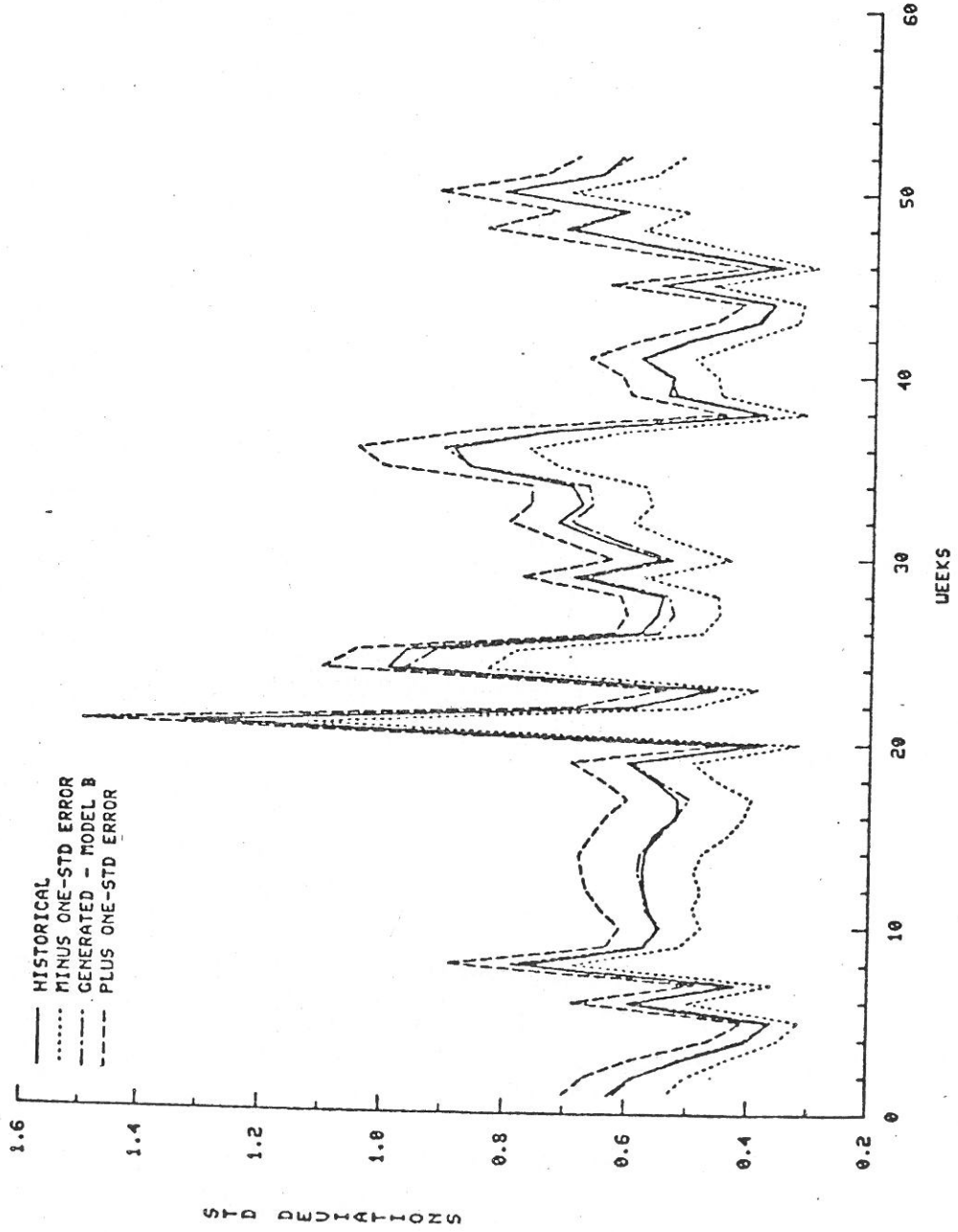


Figure 1.9.E.79. PASODE - STD DEVIATIONS (LOG DOMAIN)

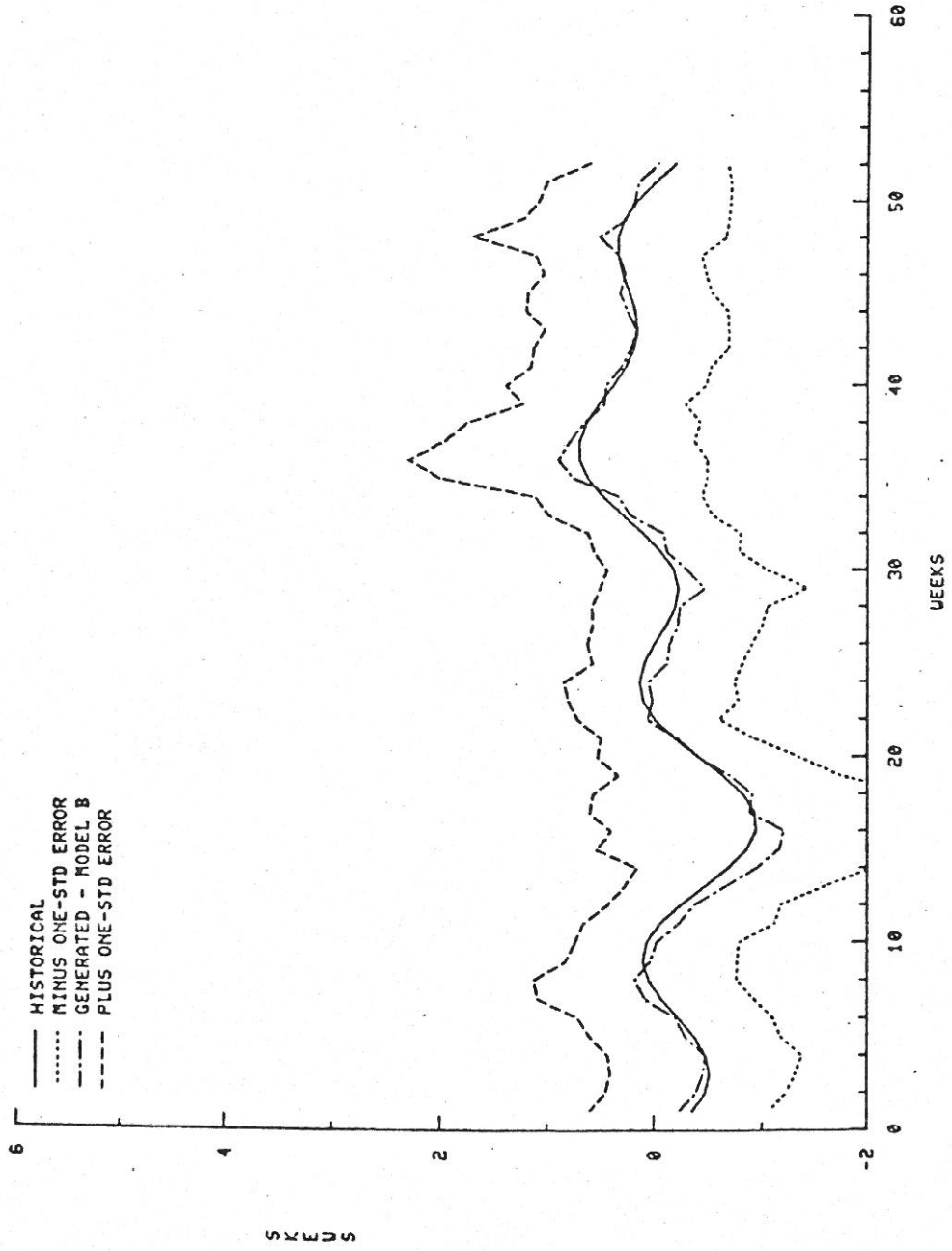


Figure 1.9.E.80. PASODE - SKEWS (LOG DOMAIN)

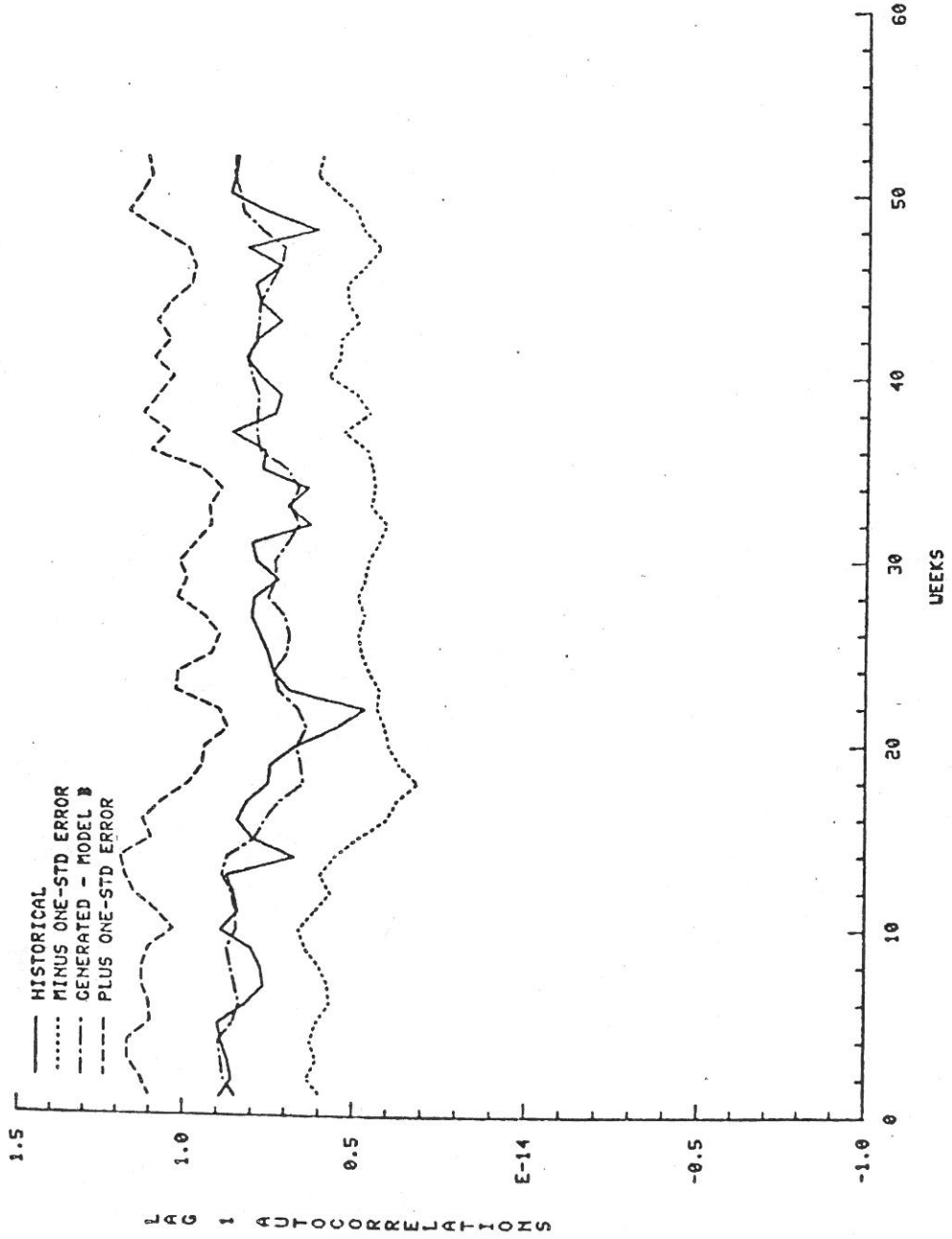


Figure 1.9.E.81. PASODE - LAG 1 AUTOCORRELATIONS (LOG DOMAIN)

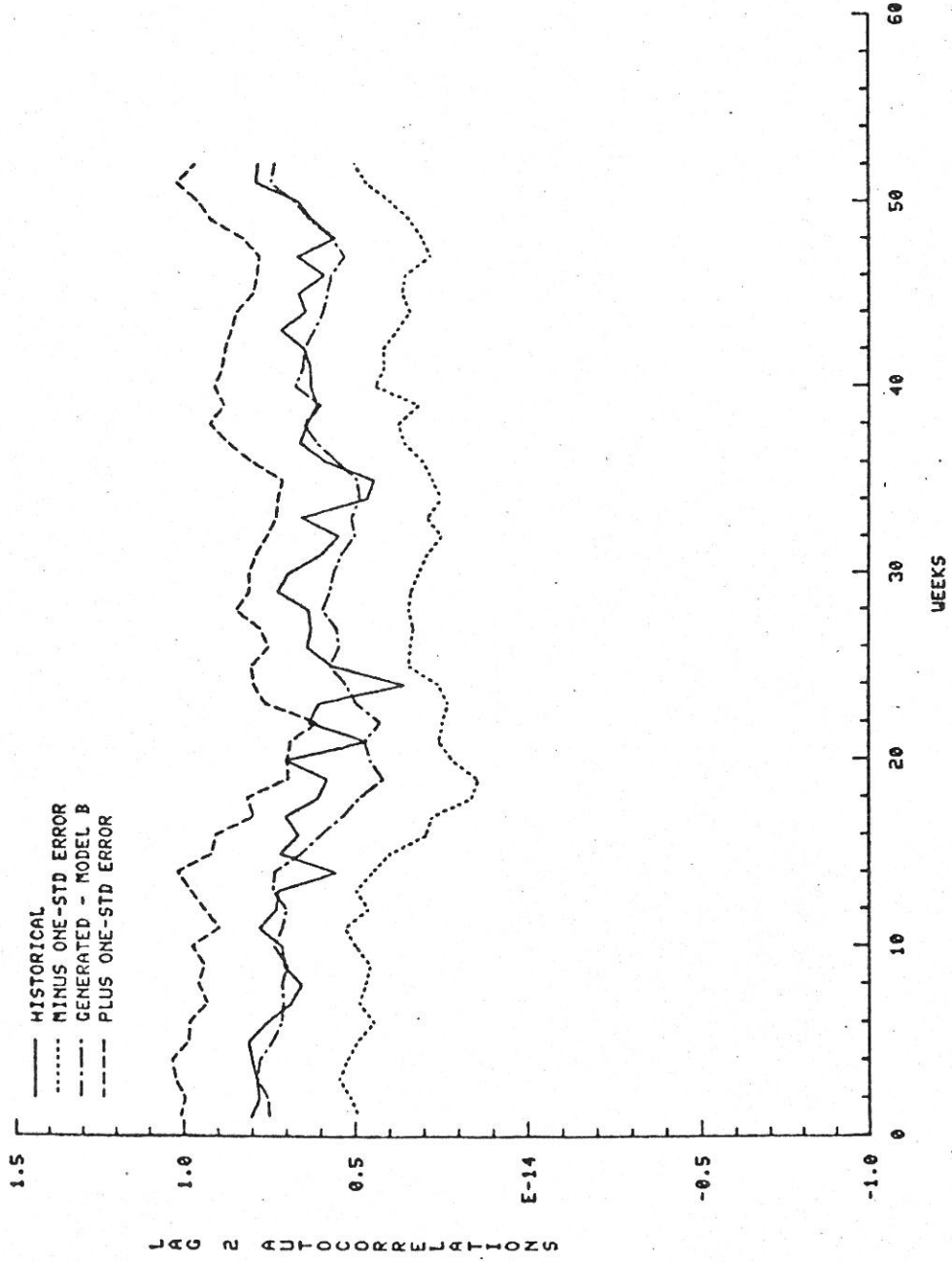


Figure 1.9.E.82. PASODE - LAG 2 AUTOCORRELATIONS (LOG DOMAIN)

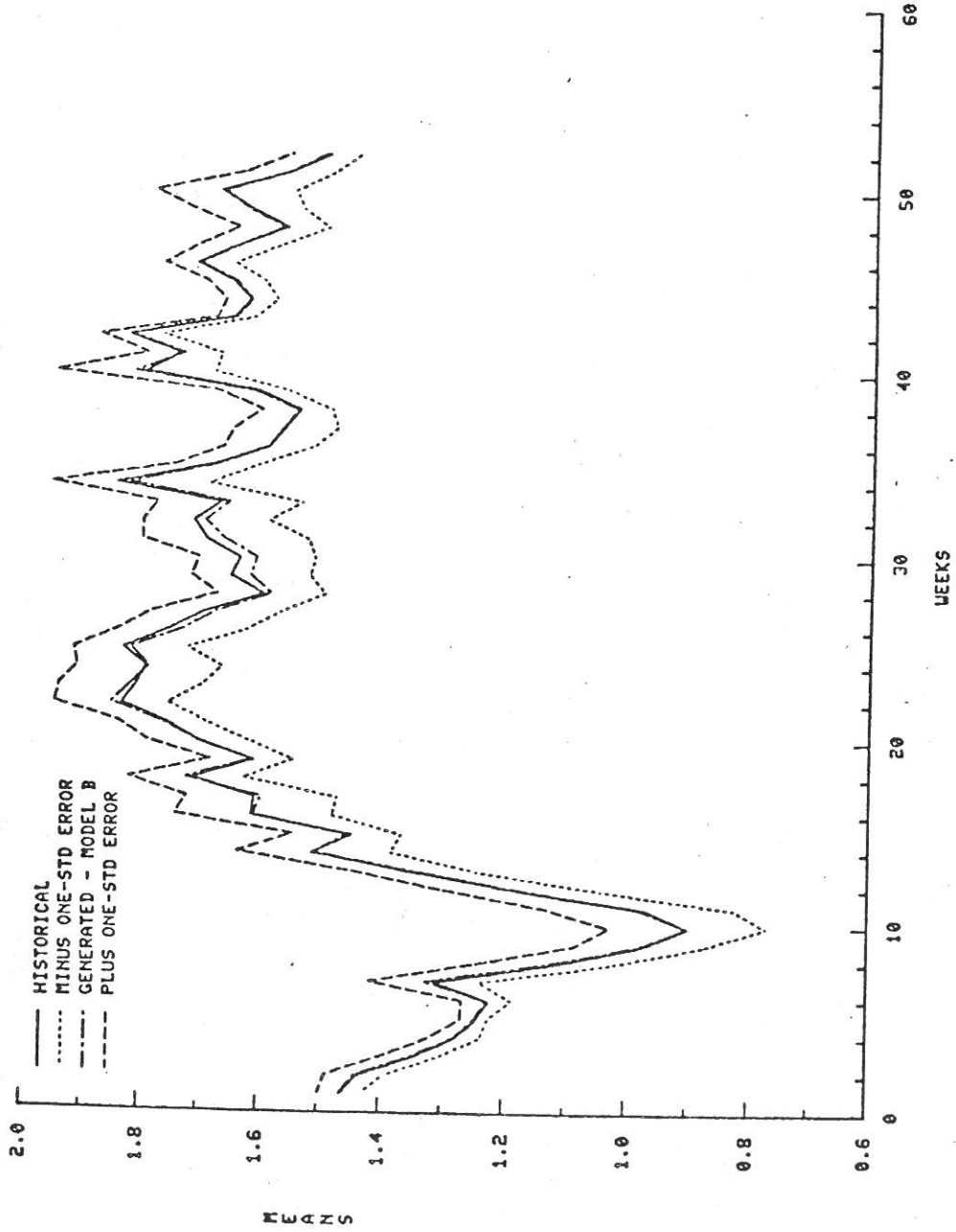


Figure 1.9.E.83. RANCHO - MEANS (LOG DOMAIN)

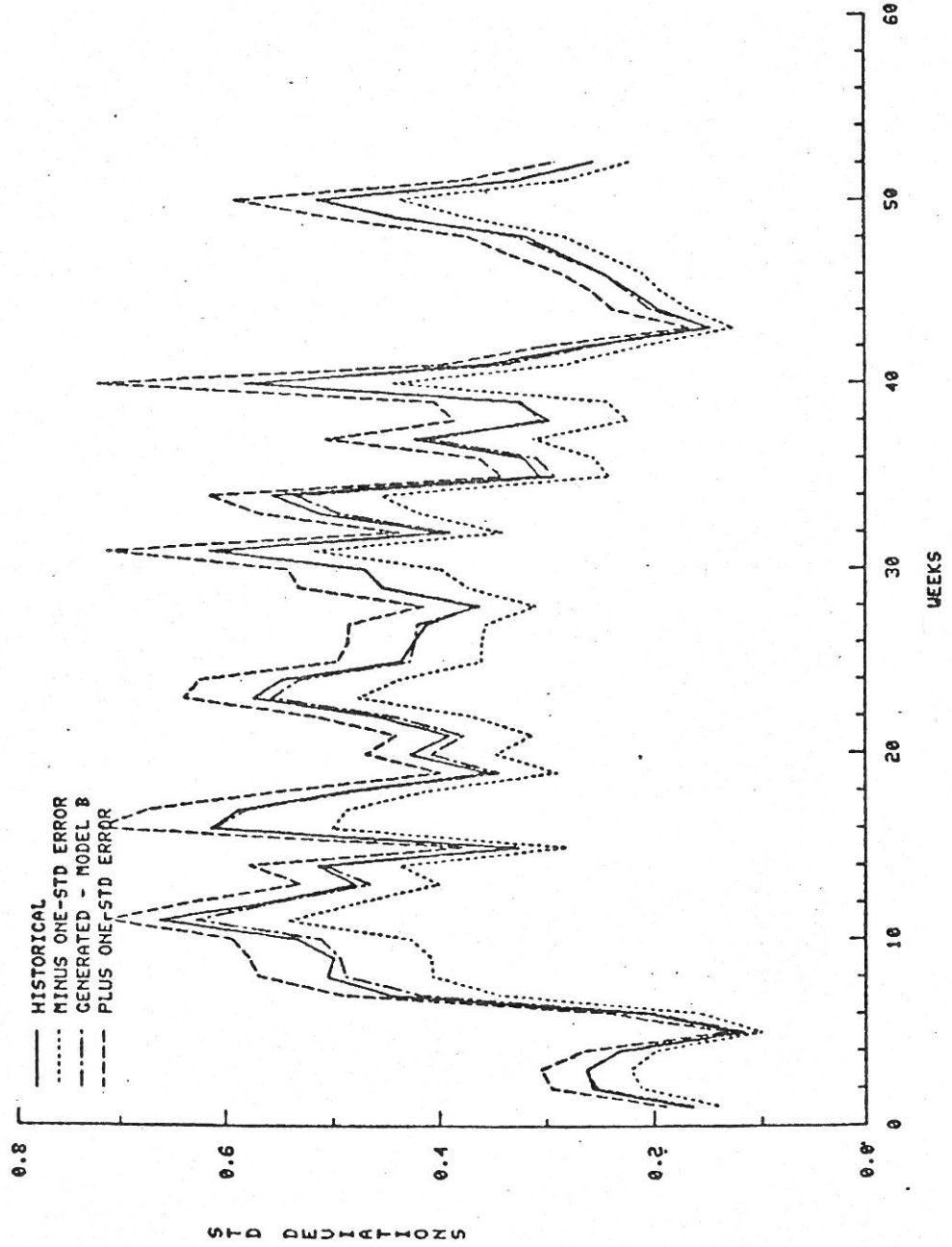


Figure 1.9.E.84. RANCHO - STD DEVIATIONS (LOG DOMAIN)

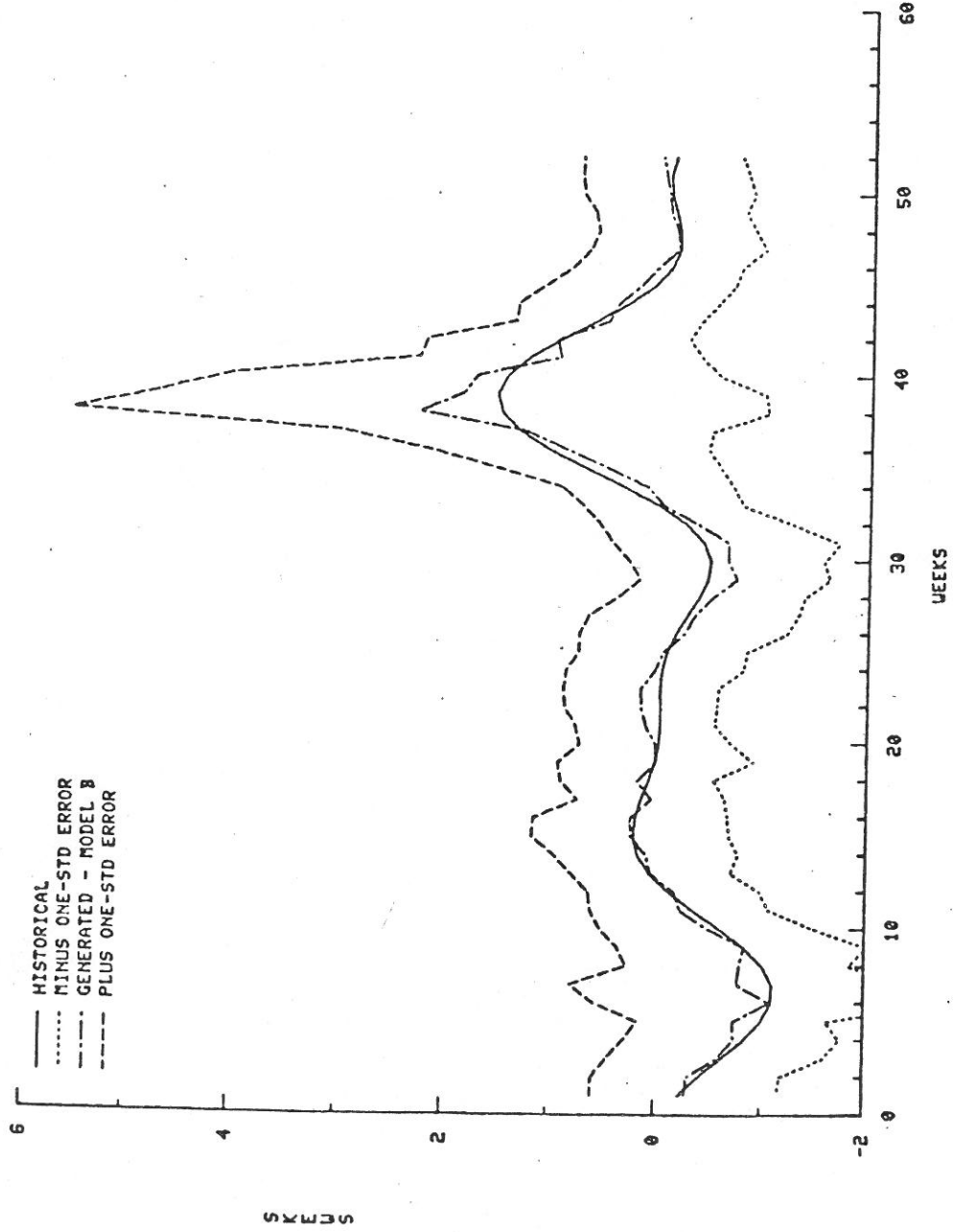


Figure 1.9.E.85. RANCHO - SKEWS (LOG DOMAIN)

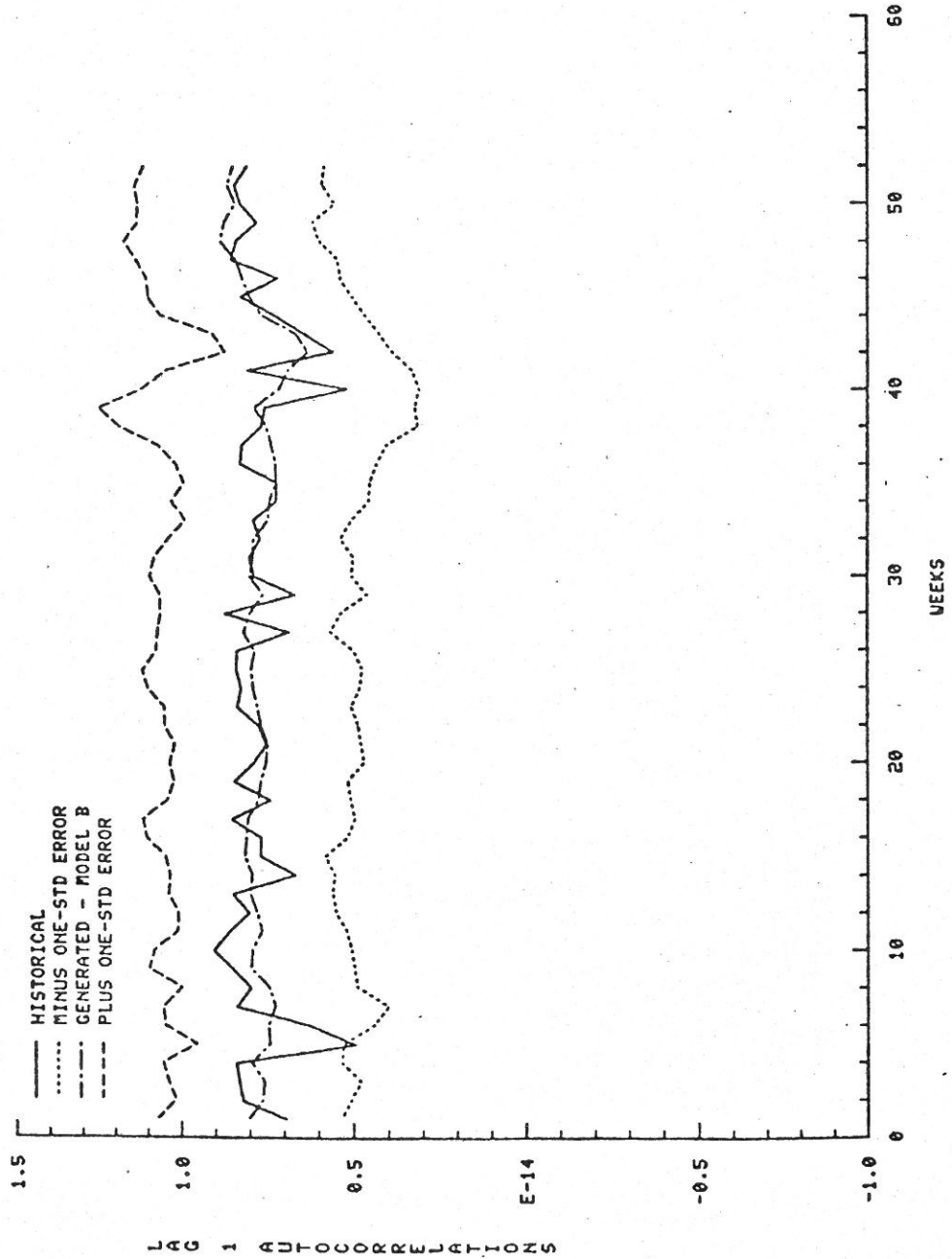


Figure 1.9.E.86. RANCHO - LAG 1 AUTOCORRELATIONS (LOG DOMAIN)

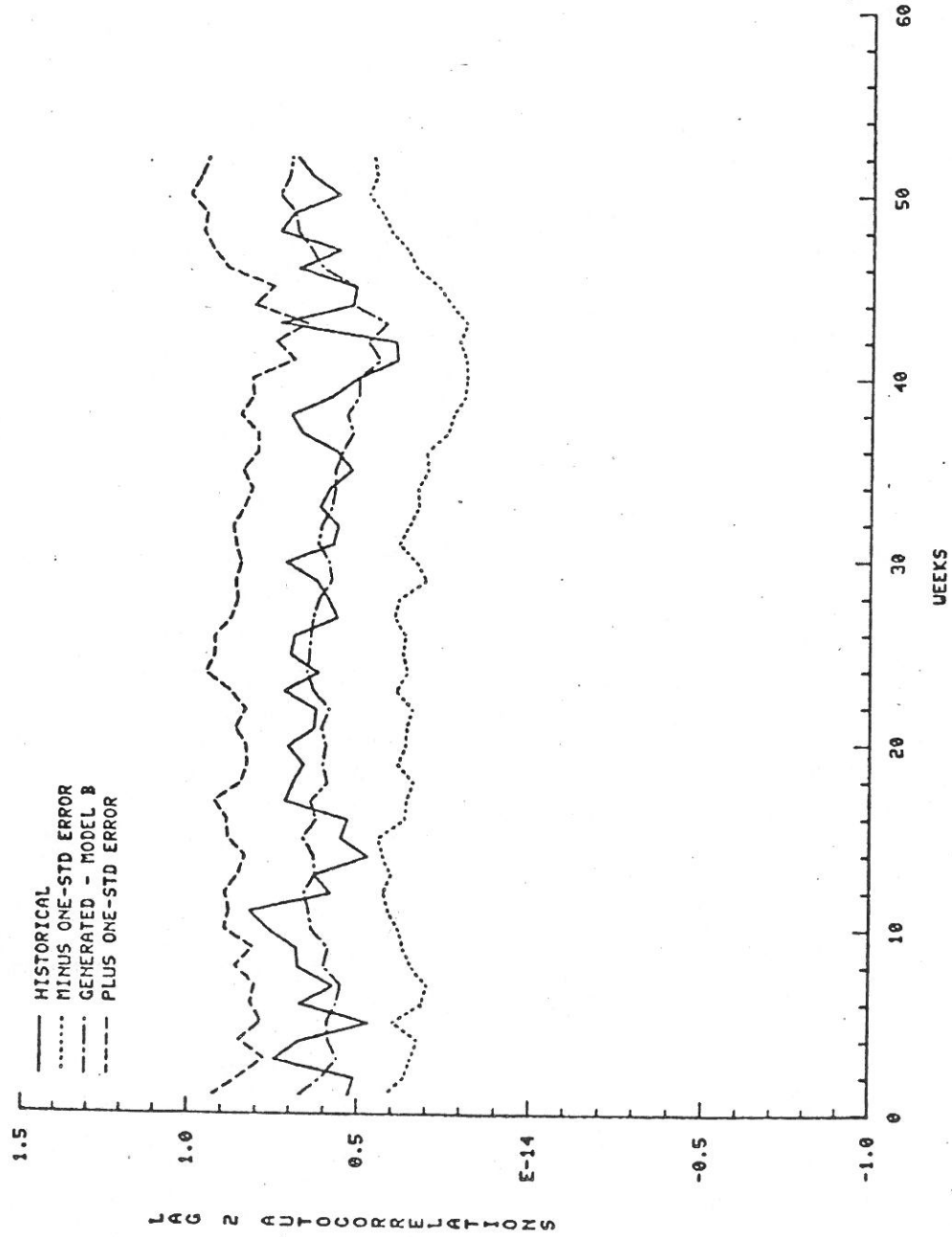


Figure 1.9.E.87. RANCHO - LAG 2 AUTOCORRELATIONS (LOG DOMAIN)

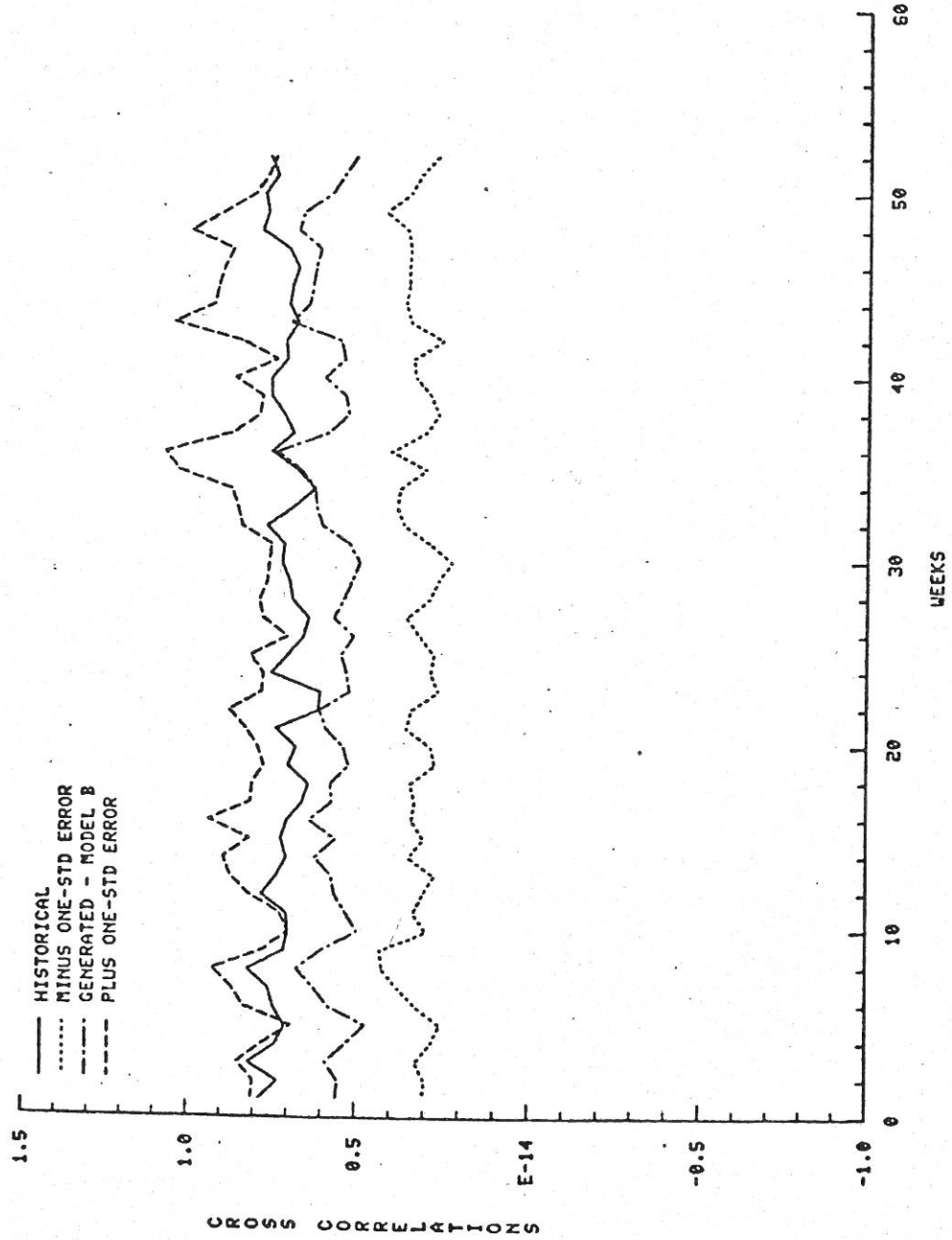


Figure 1.9.E.88. PALODE AND PASODE - CROSS CORRELATIONS (LOG DOMAIN)

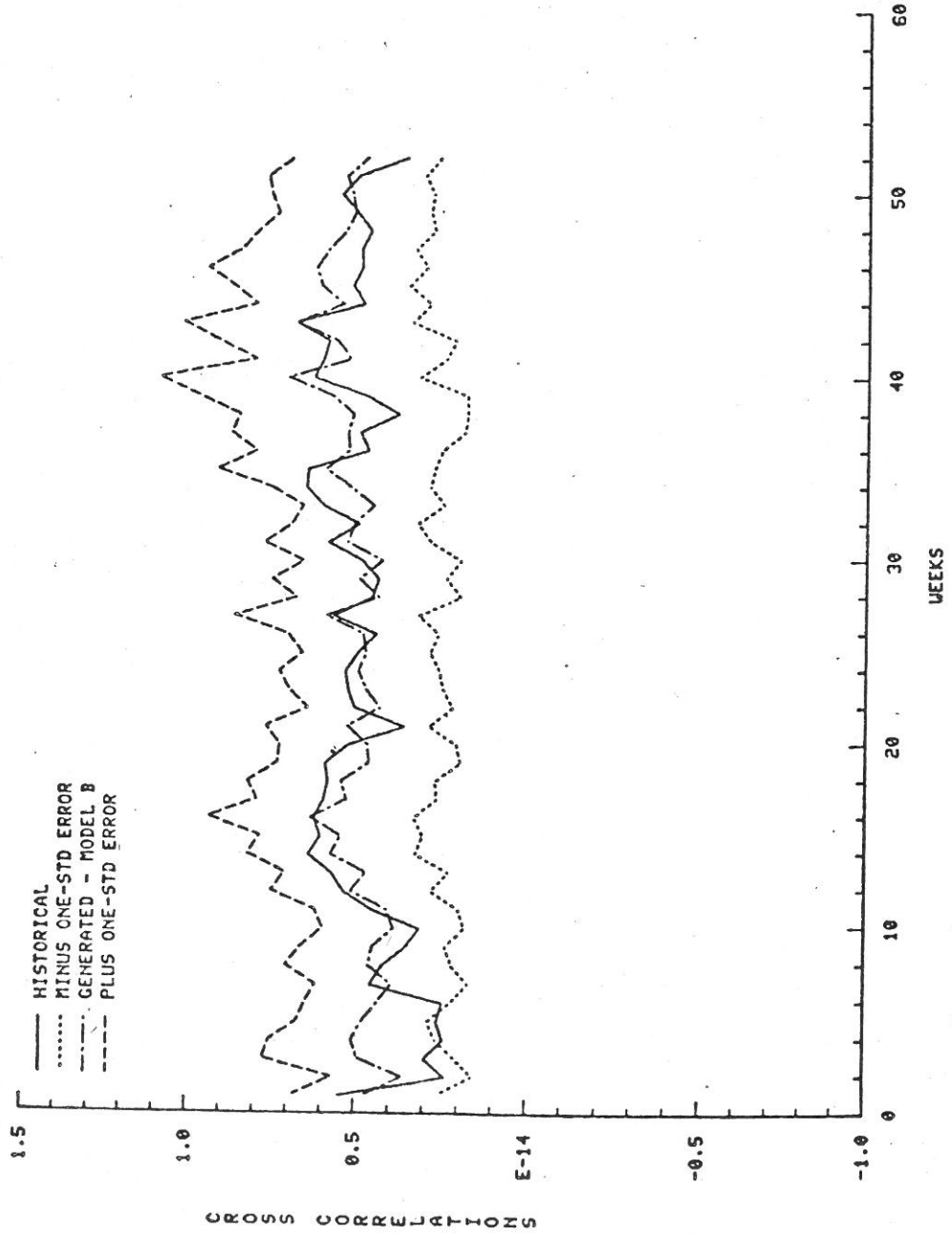


Figure 1.9.E.89. PASODE AND PASODE - CROSS CORRELATIONS (LOG DOMAIN)

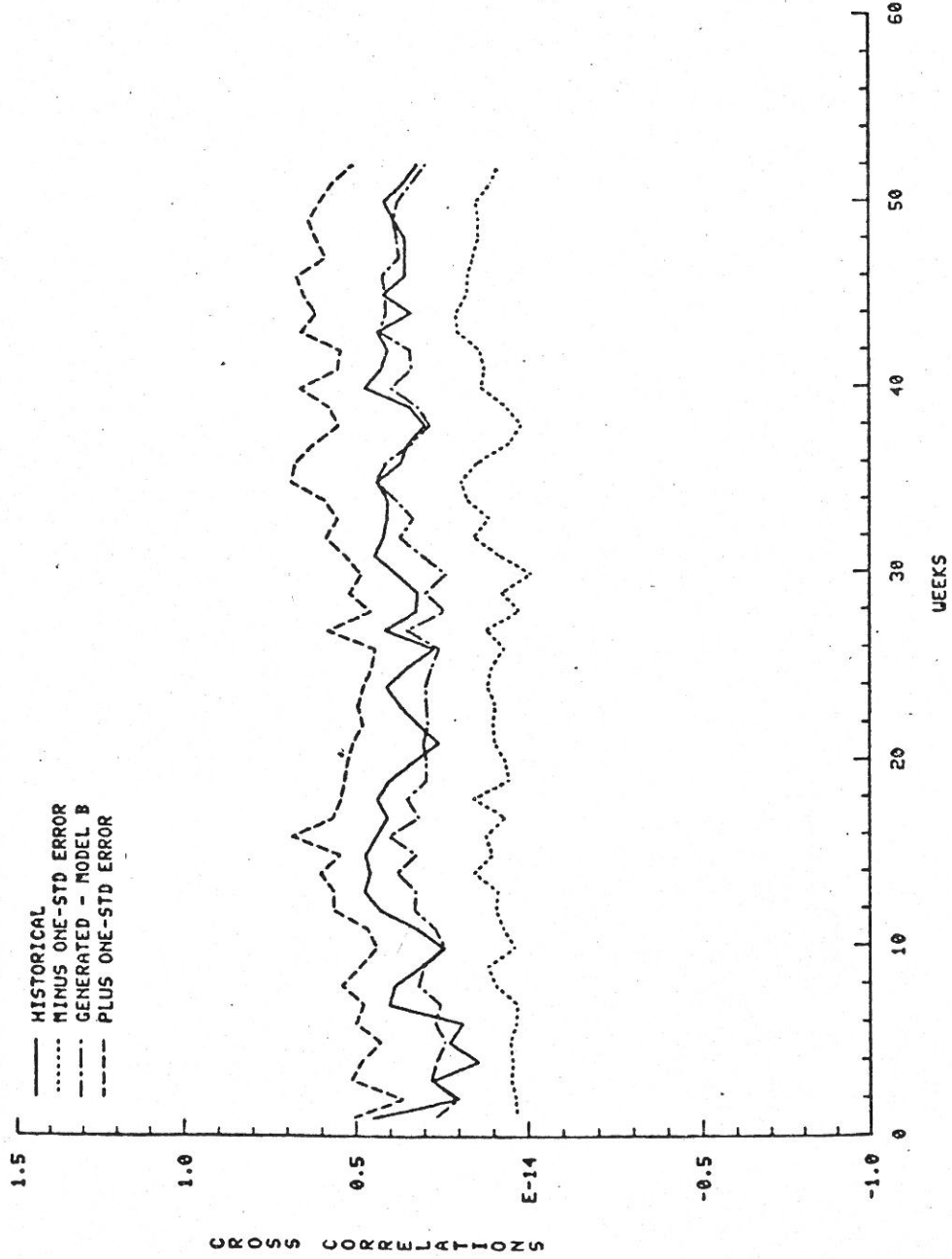


Figure 1.9.E.90. PASODE AND RANCHO - CROSS CORRELATIONS (LOG DOMAIN)

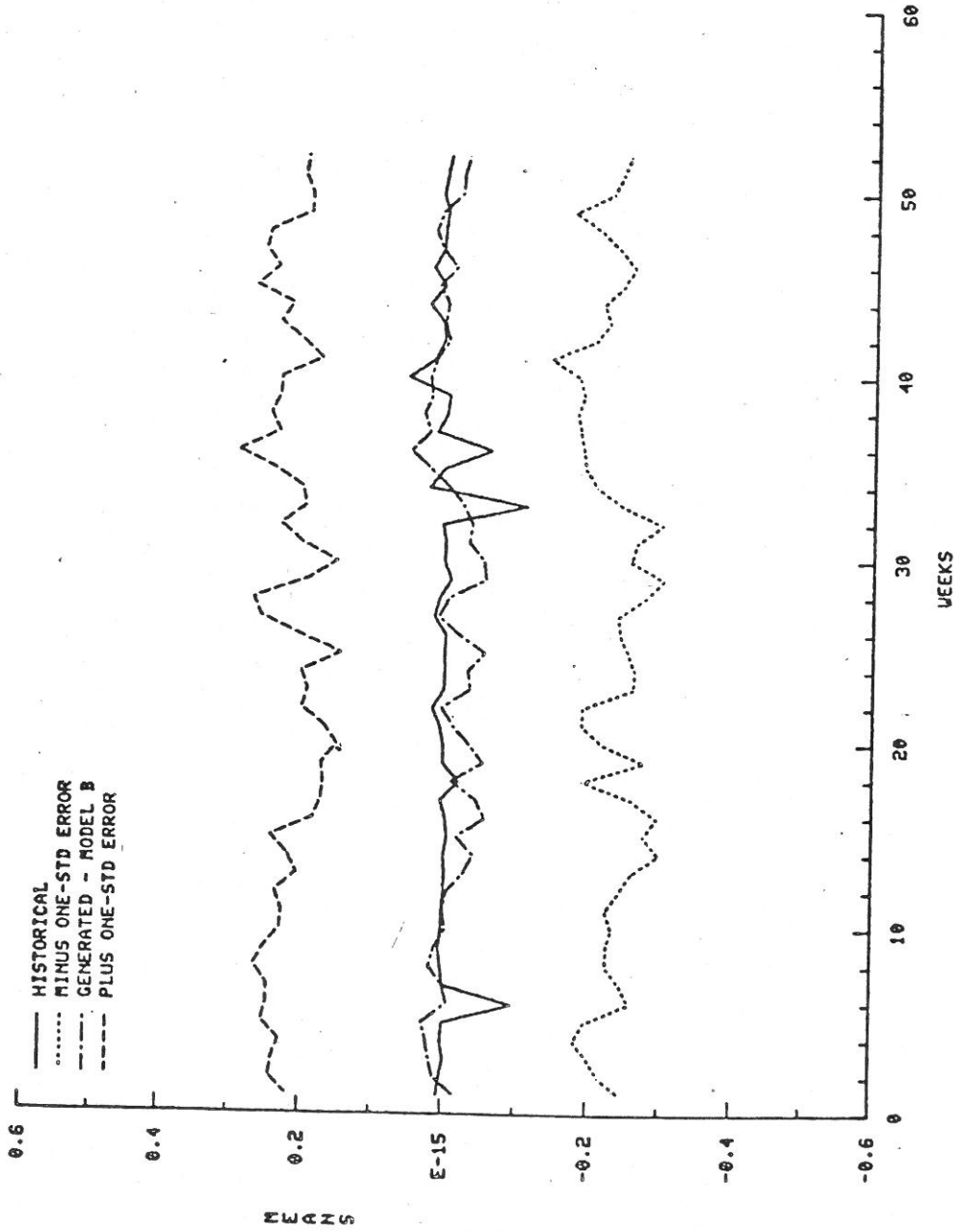


Figure 1.9.E.91. PALODE - MEANS (LWH DOMAIN)

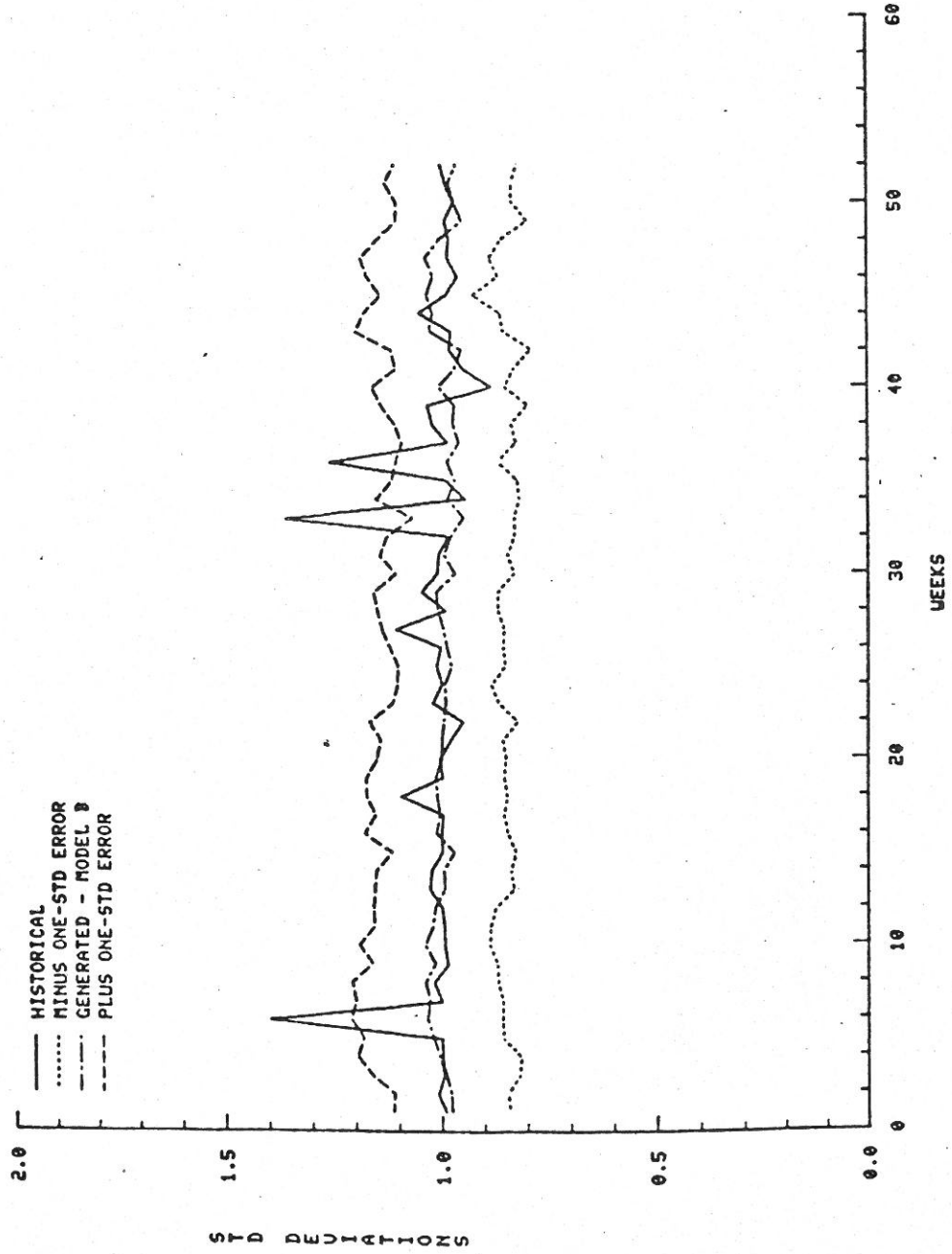


Figure 1.9.E.92. PALODE - STD DEVIATIONS (LWH DOMAIN)

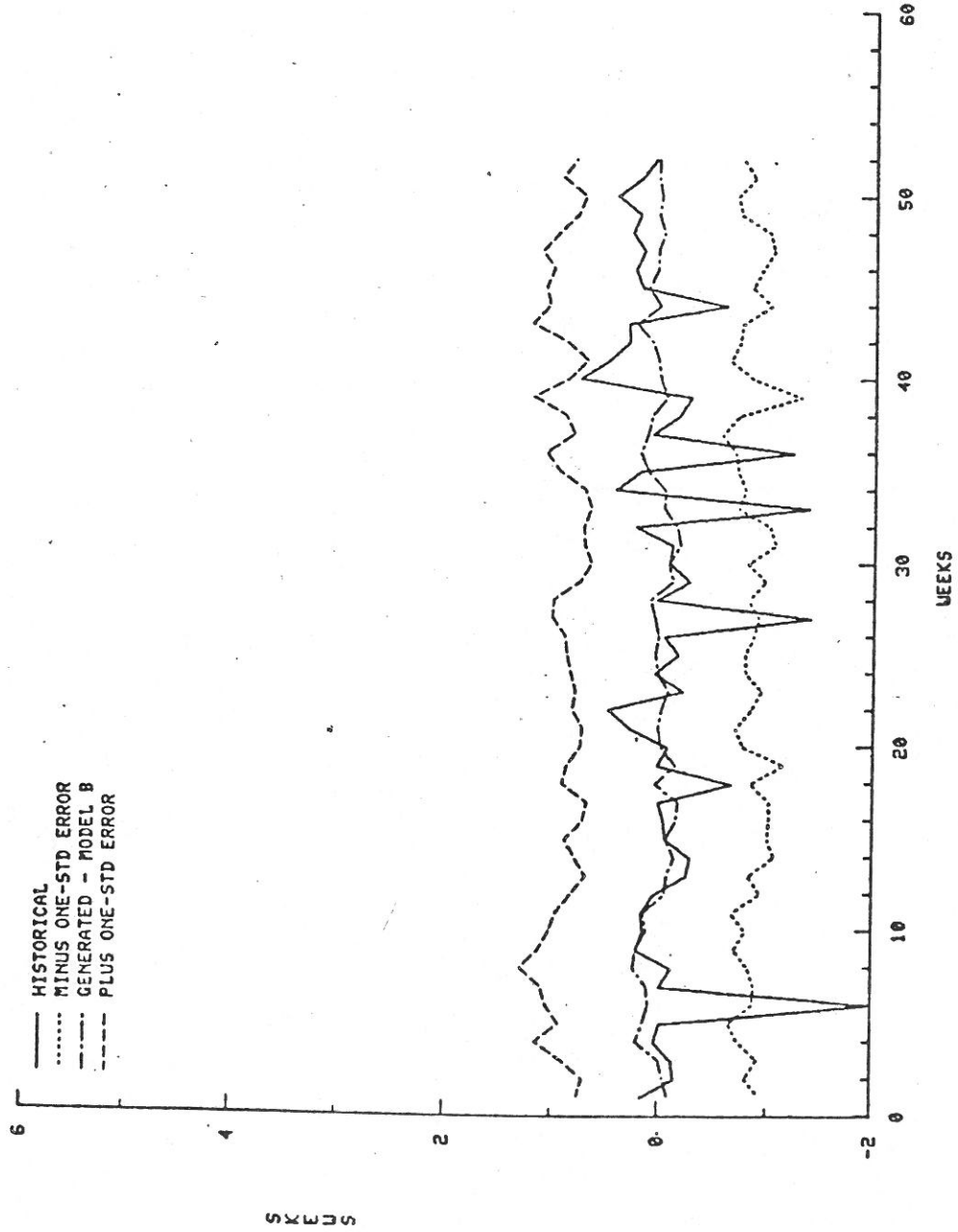


Figure 1.9.E.93. PALODE - SKEWS (LWH DOMAIN)

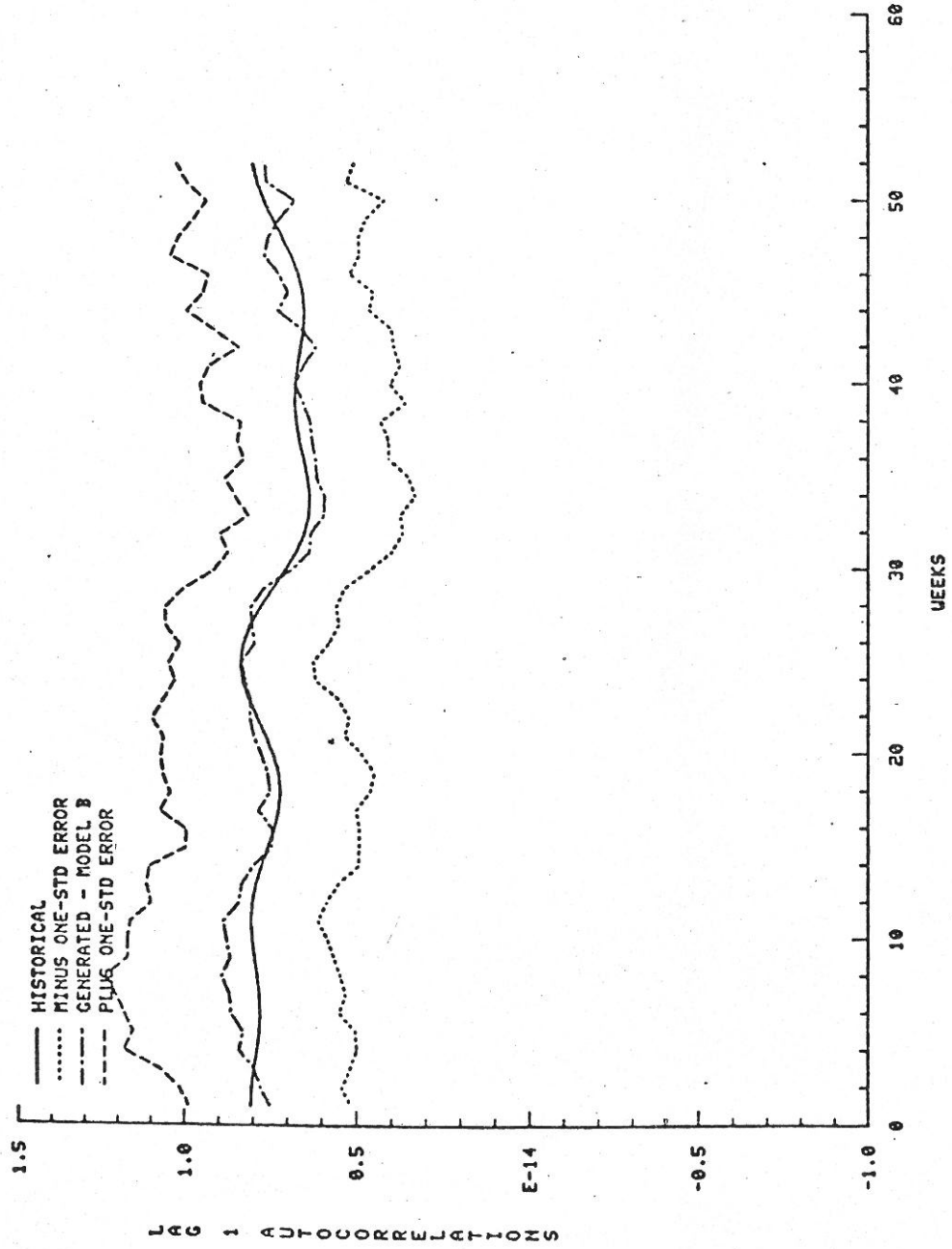


Figure 1.9.E.94. PALODE - LAG 1 AUTOCORRELATIONS (LWH DOMAIN)

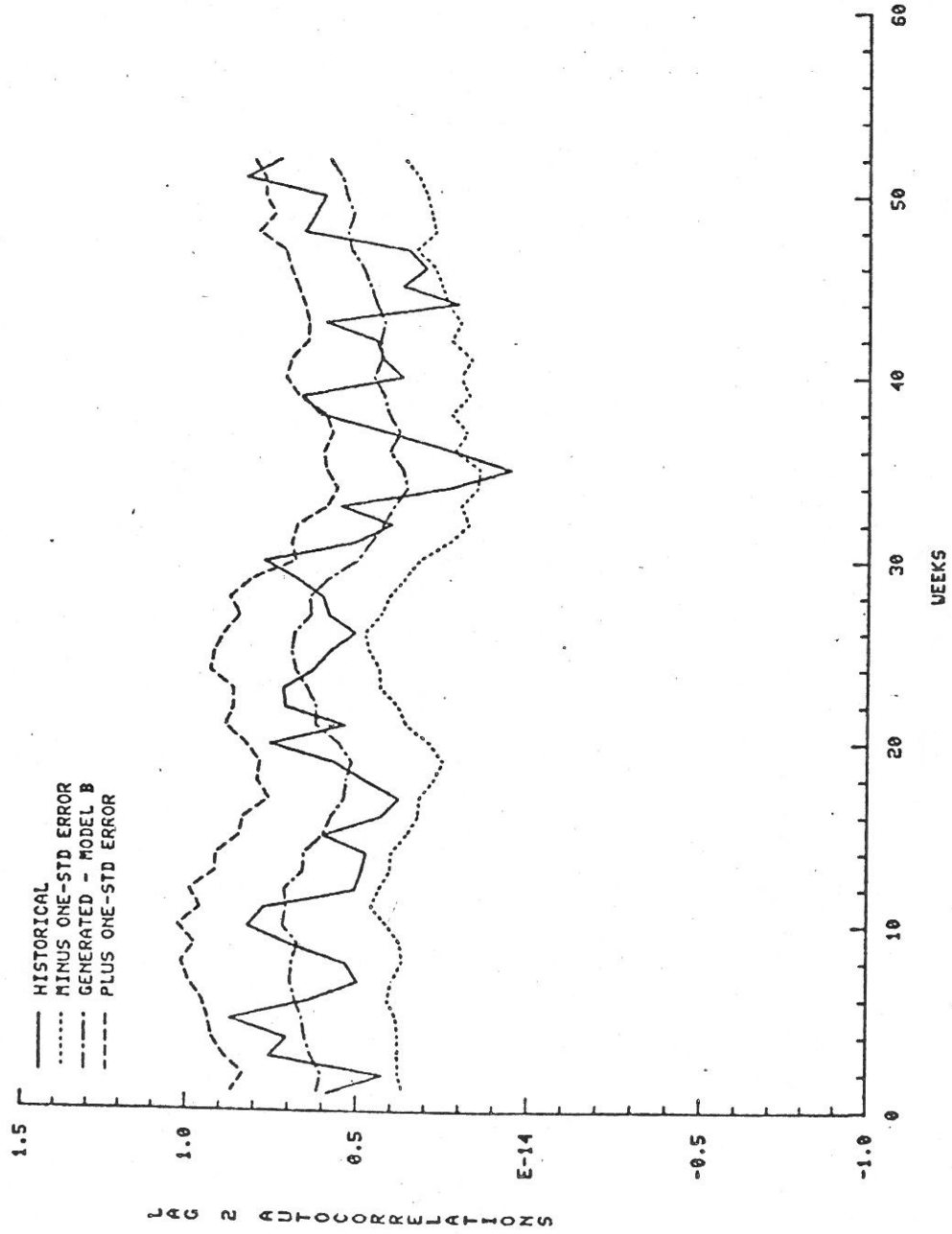


Figure 1.9.E.95. PALODE - LAG 2 AUTOCORRELATIONS (LWH DOMAIN)

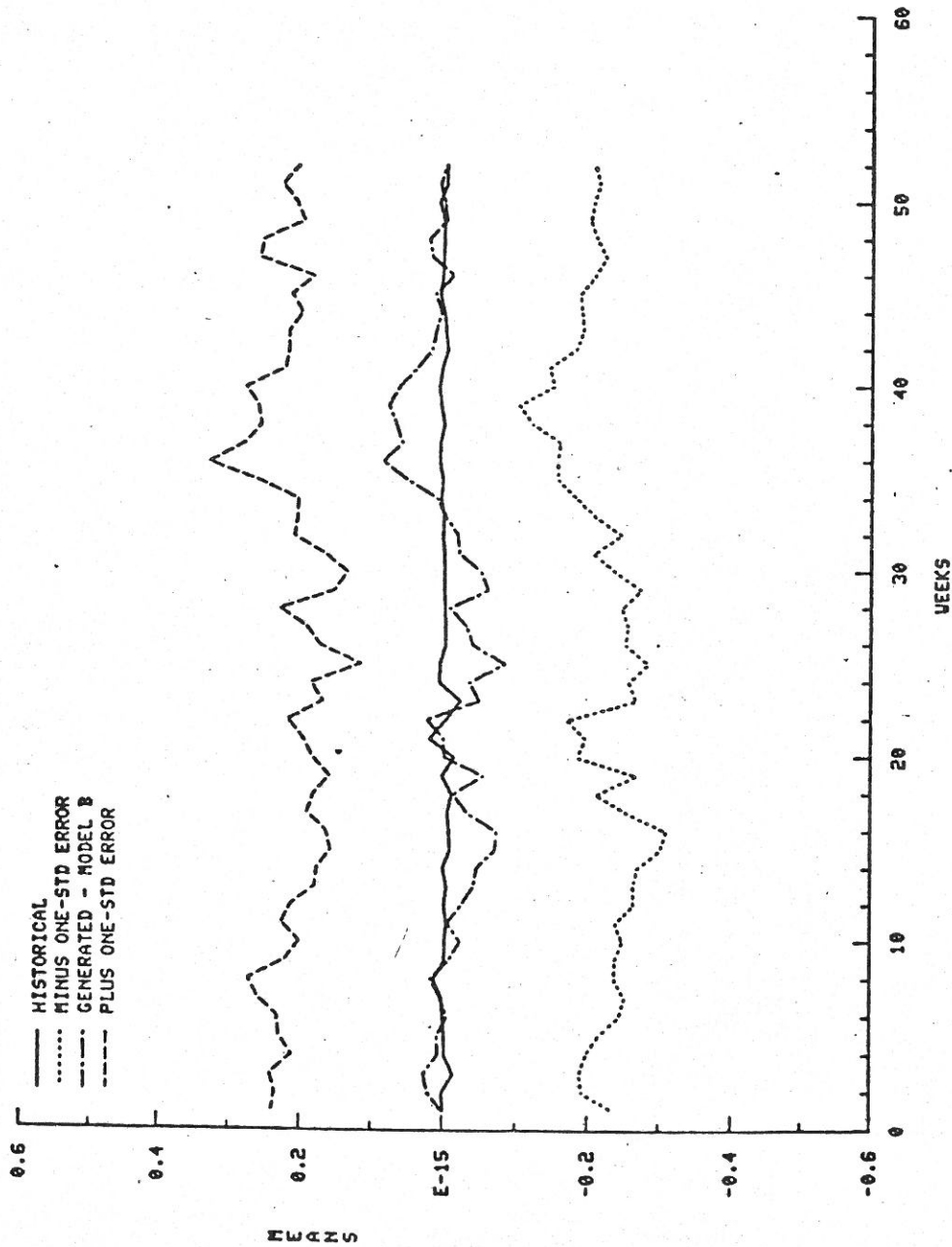


Figure 1.9.E.96. PASODE - MEANS (LWH DOMAIN)

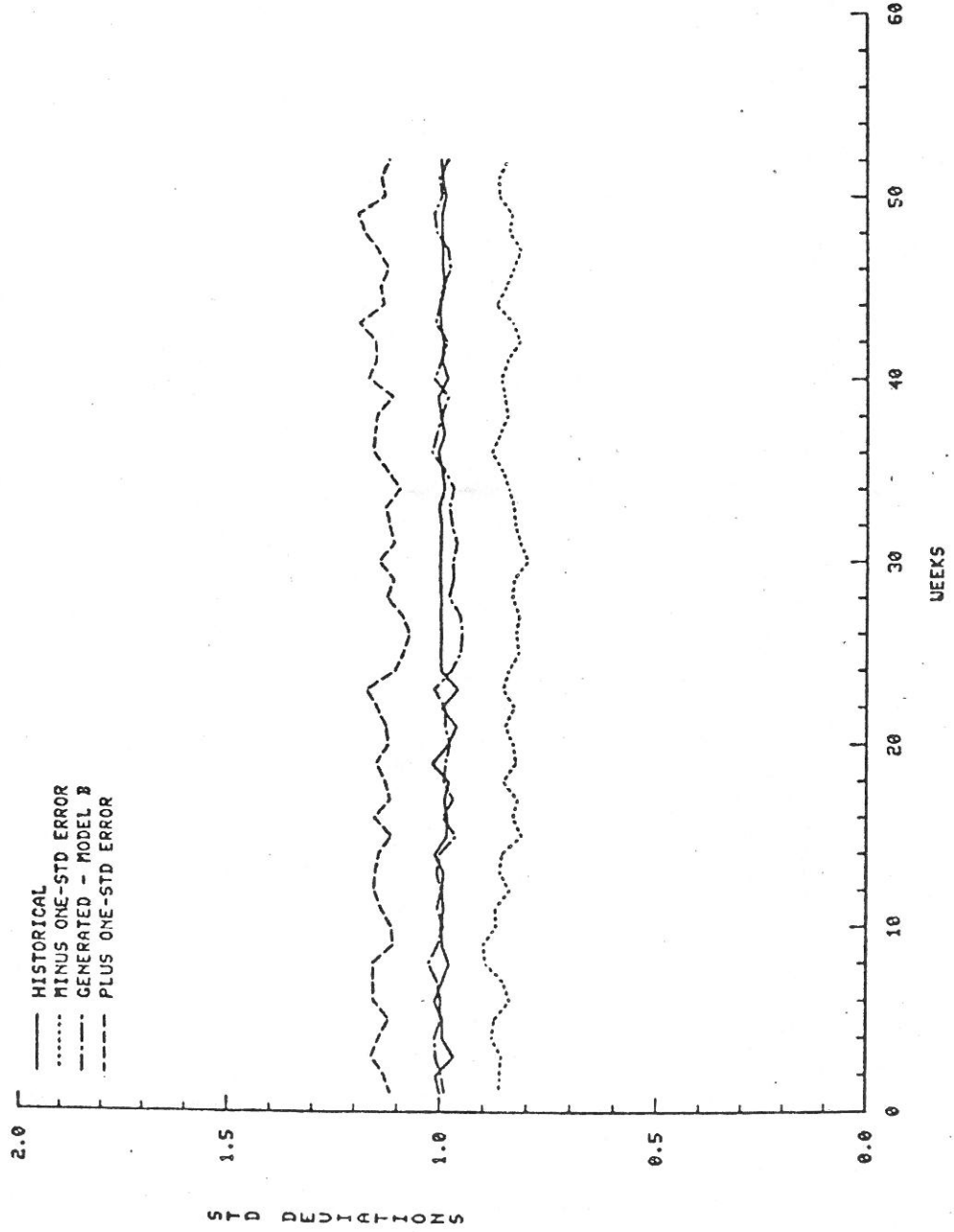


Figure 1.9.E.97. PASODE - STD DEVIATIONS (LUH DOMAIN)

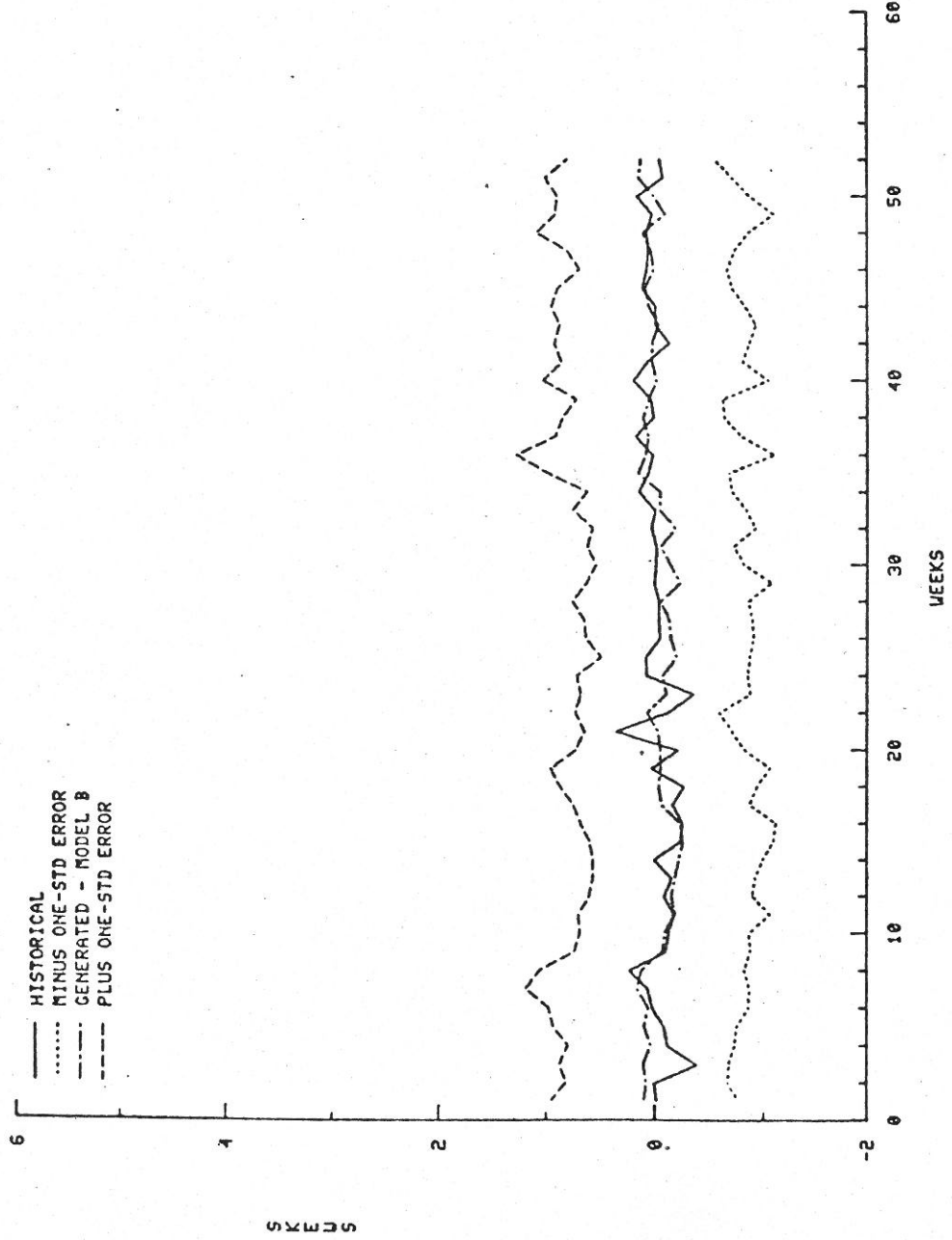


Figure 1.9.E.98. PASODE - SKEWS (LUH DOMAIN)

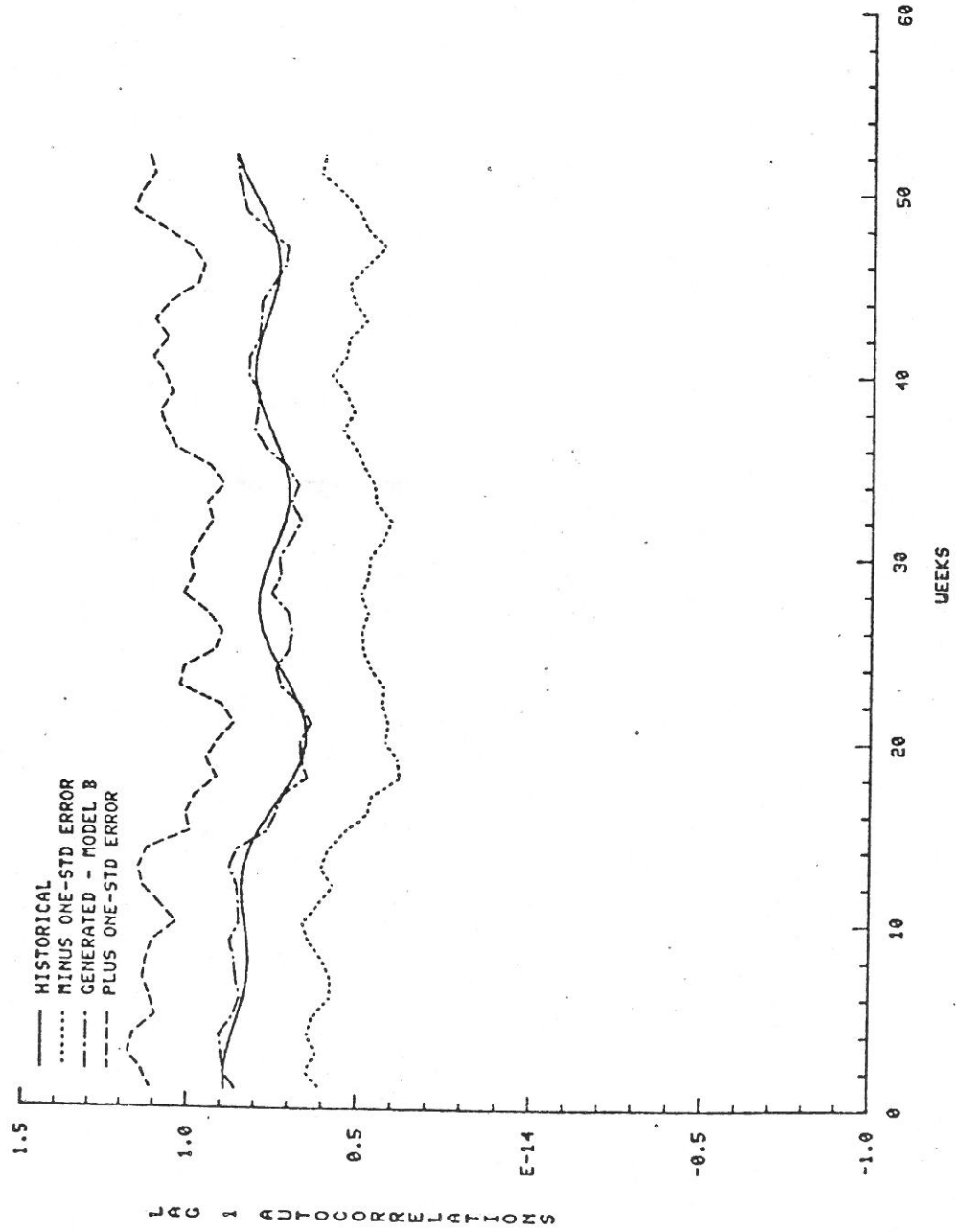


Figure 1.9.E.99. PASODE - LAG 1 AUTOCORRELATIONS (LWH DOMAIN)

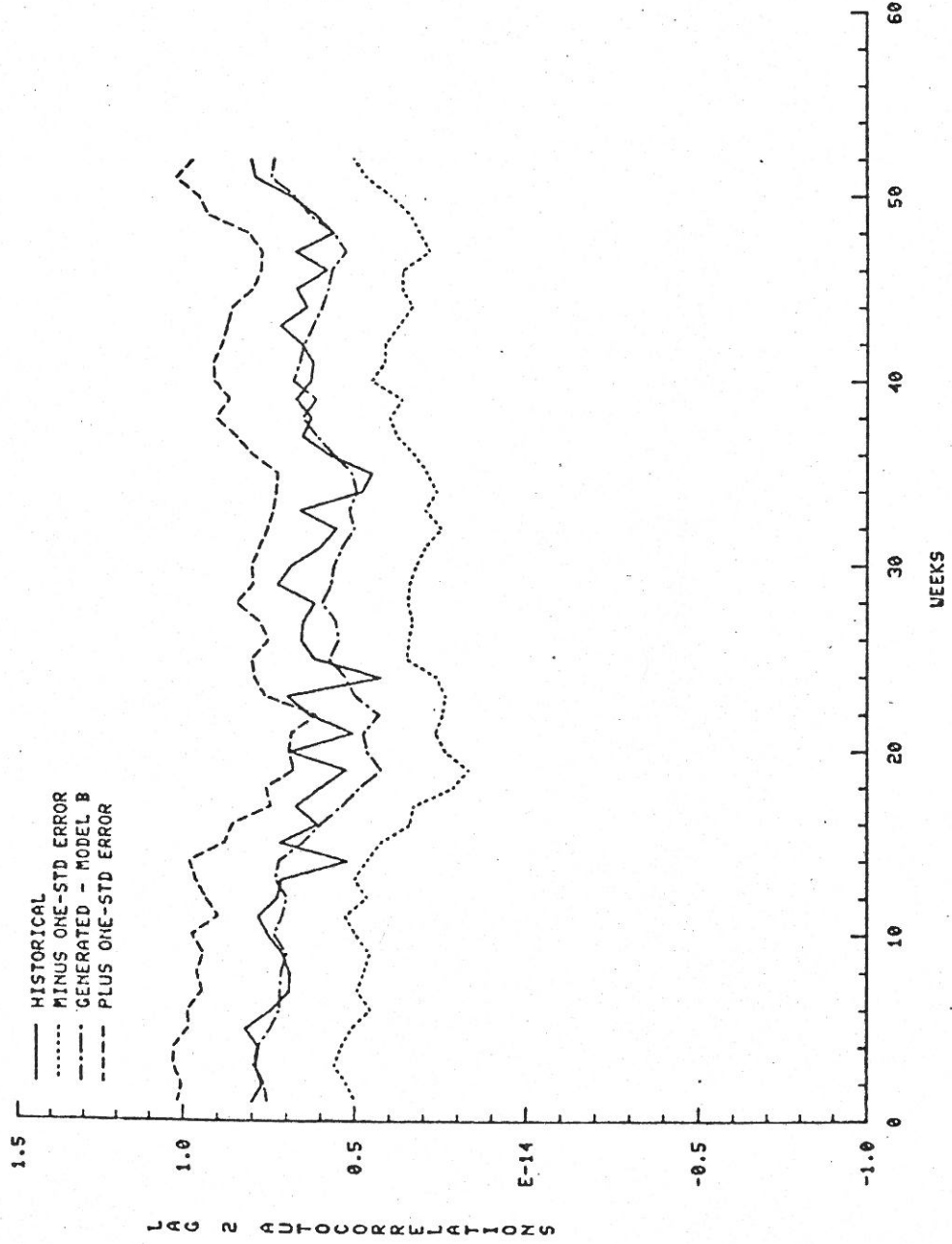


Figure 1.9.E.100. PASODE - LAG 2 AUTOCORRELATIONS (LUH DOMAIN)

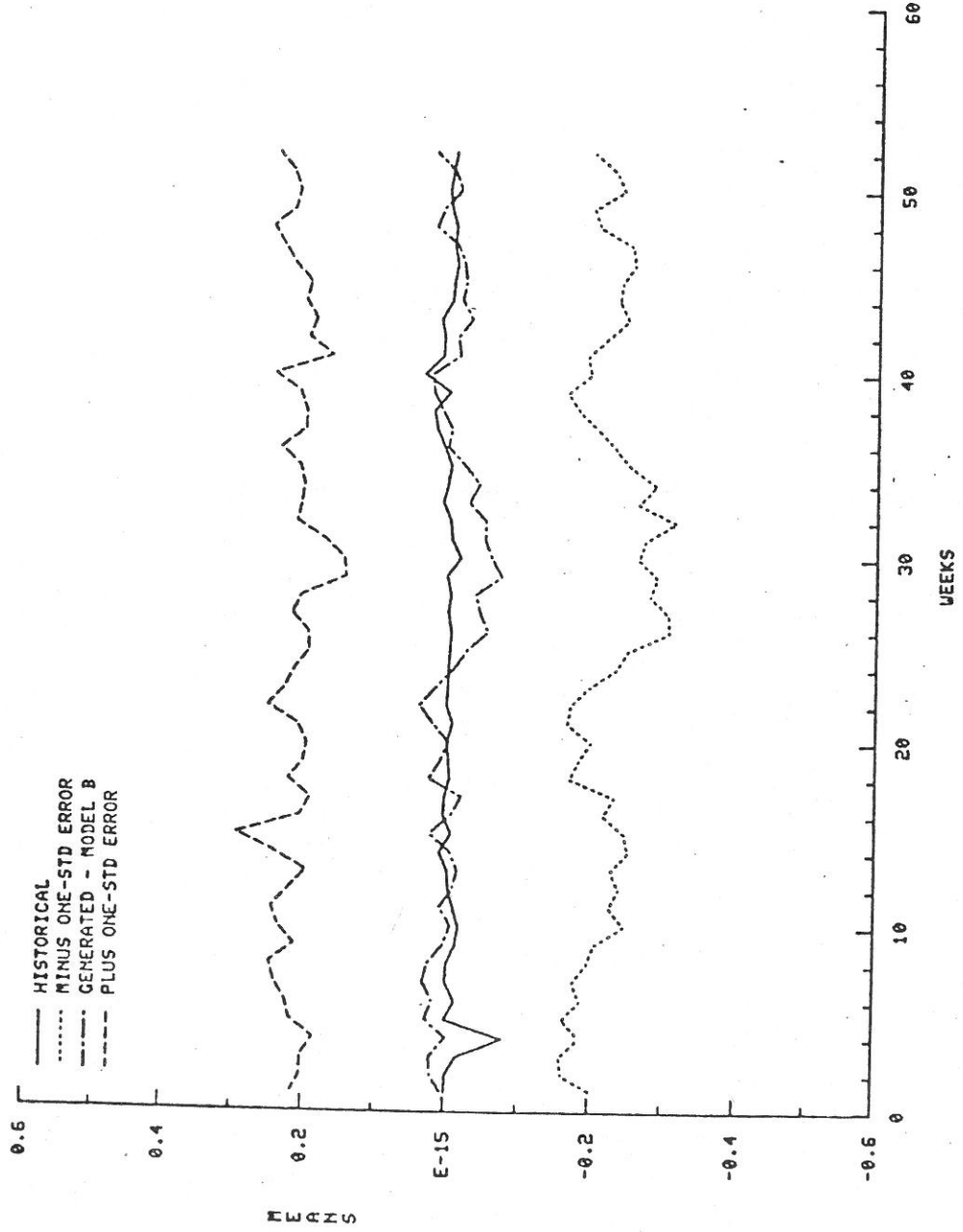


Figure 1.9.E.101. RANCHO - MEANS (LUH DOMAIN)

UNSATISFIED EXT P - 011666
EXCHANGE PACKAGE

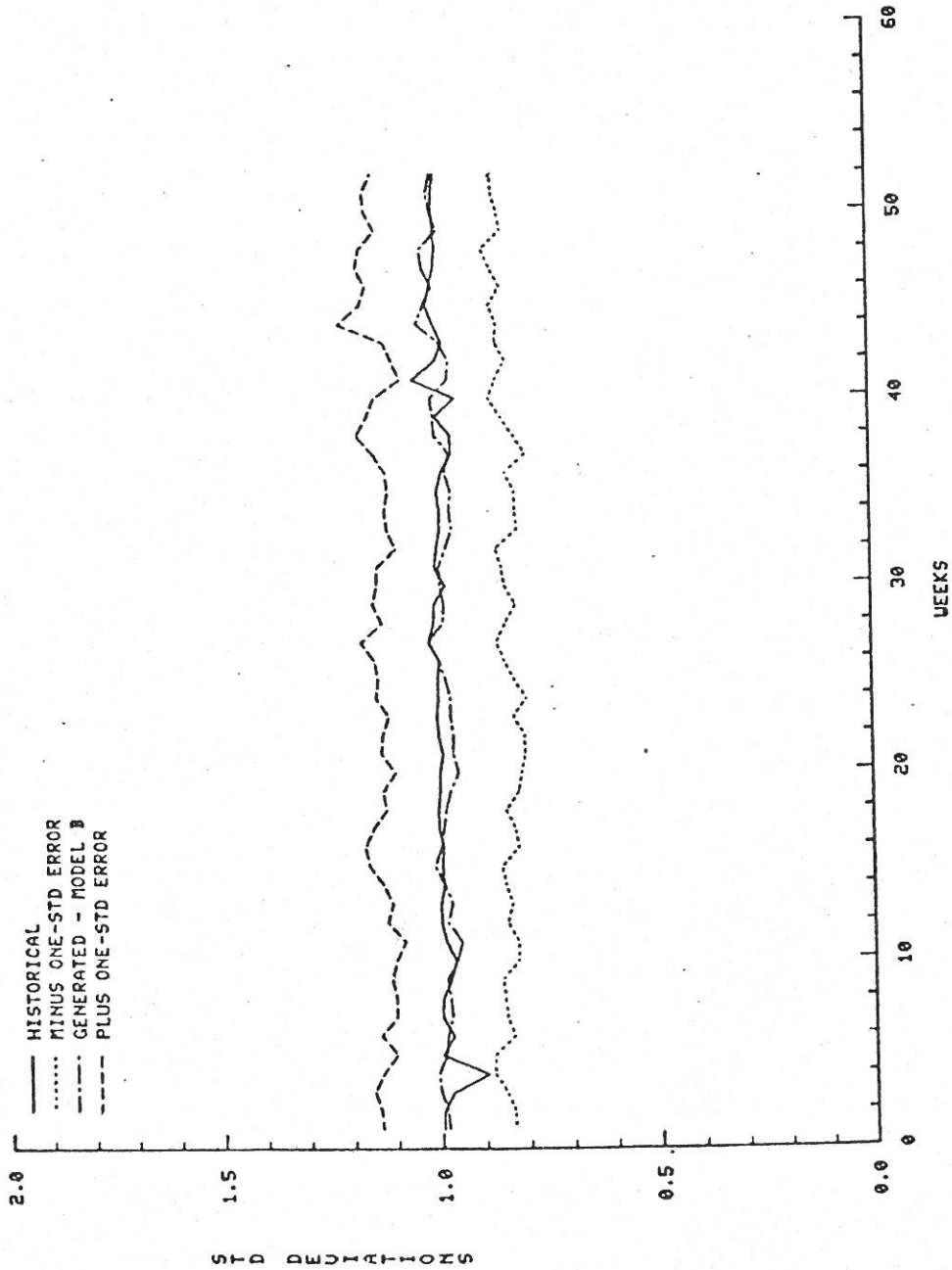


Figure 1.9.E.102. RANCHO - STD DEVIATIONS (LUH DOMAIN)

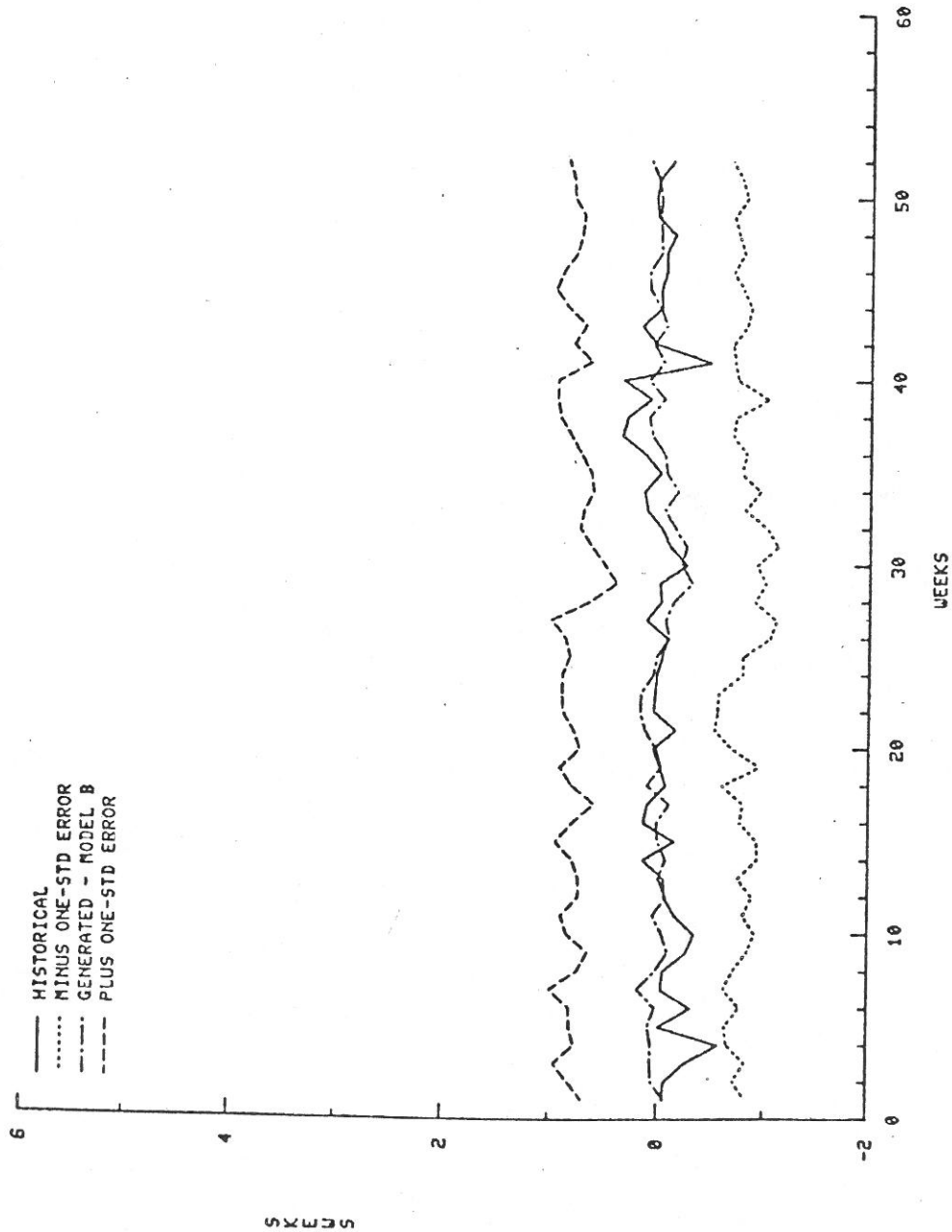


Figure 1.9.E.103. RANCHO - SKEWS (LWH DOMAIN)

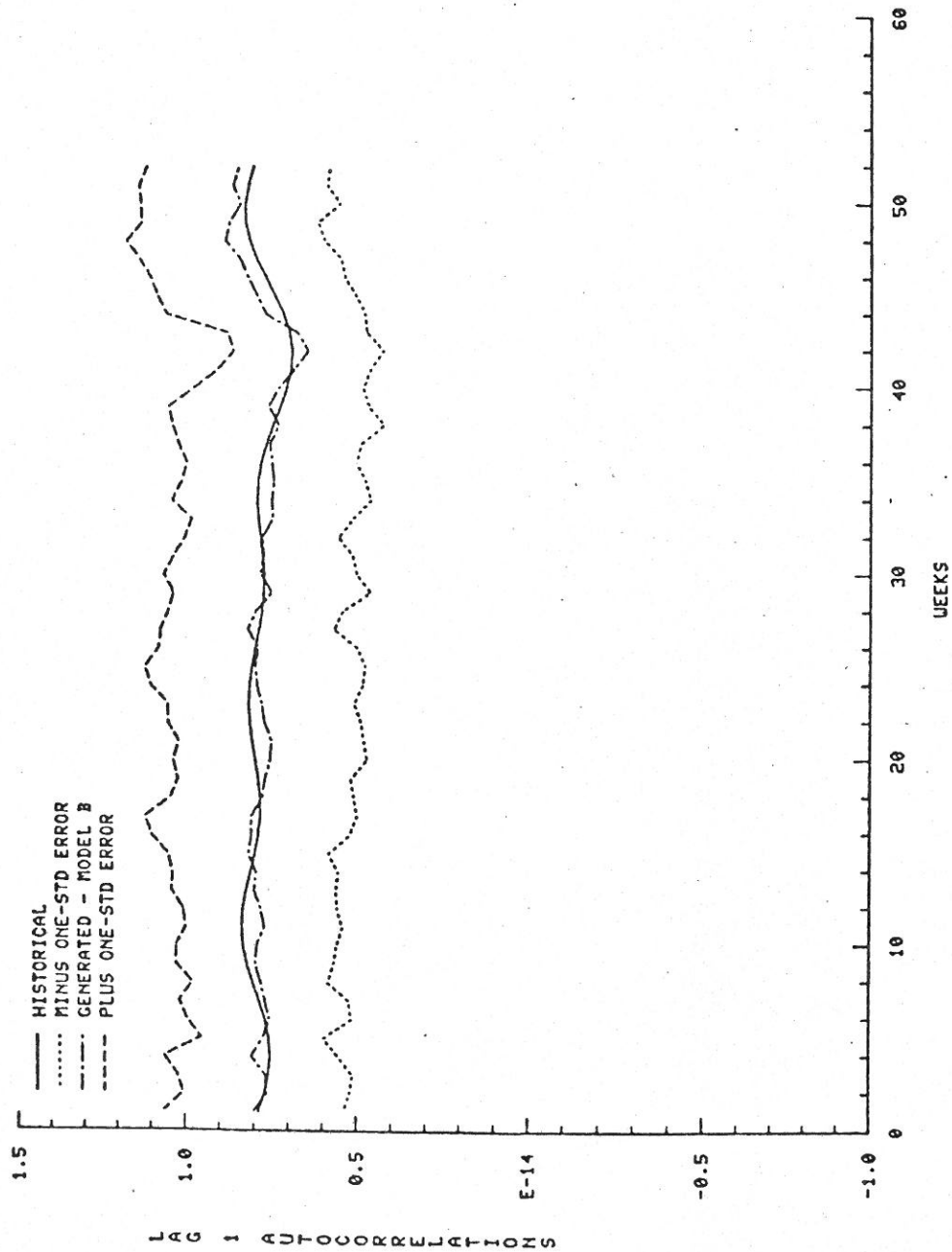


Figure I.9.E.104. RANCHO - LAG 1 AUTOCORRELATIONS (LWH DOMAIN)

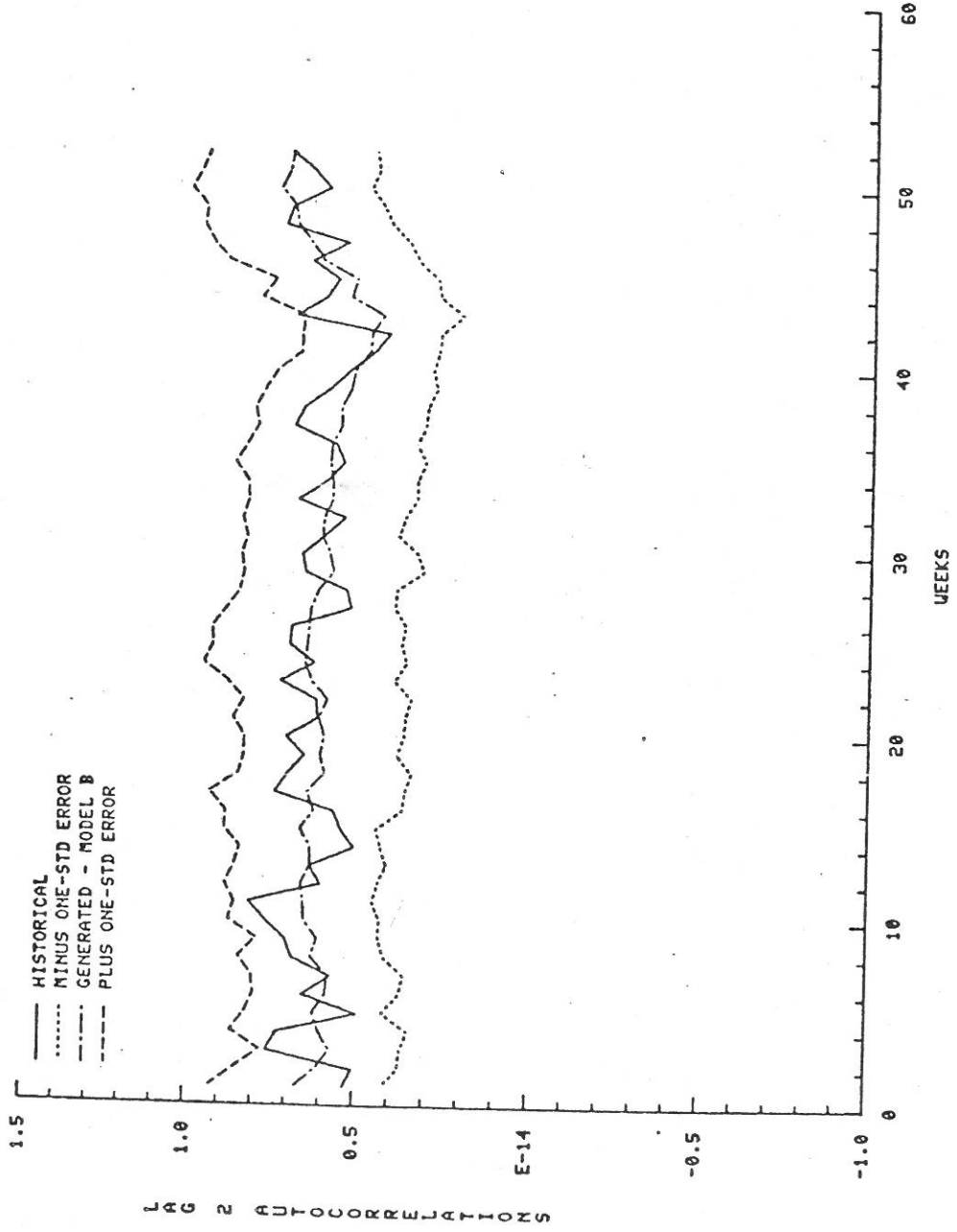


Figure 1.9.E.105. RANCHO - LAG 2 AUTOCORRELATIONS (LWH DOMAIN)

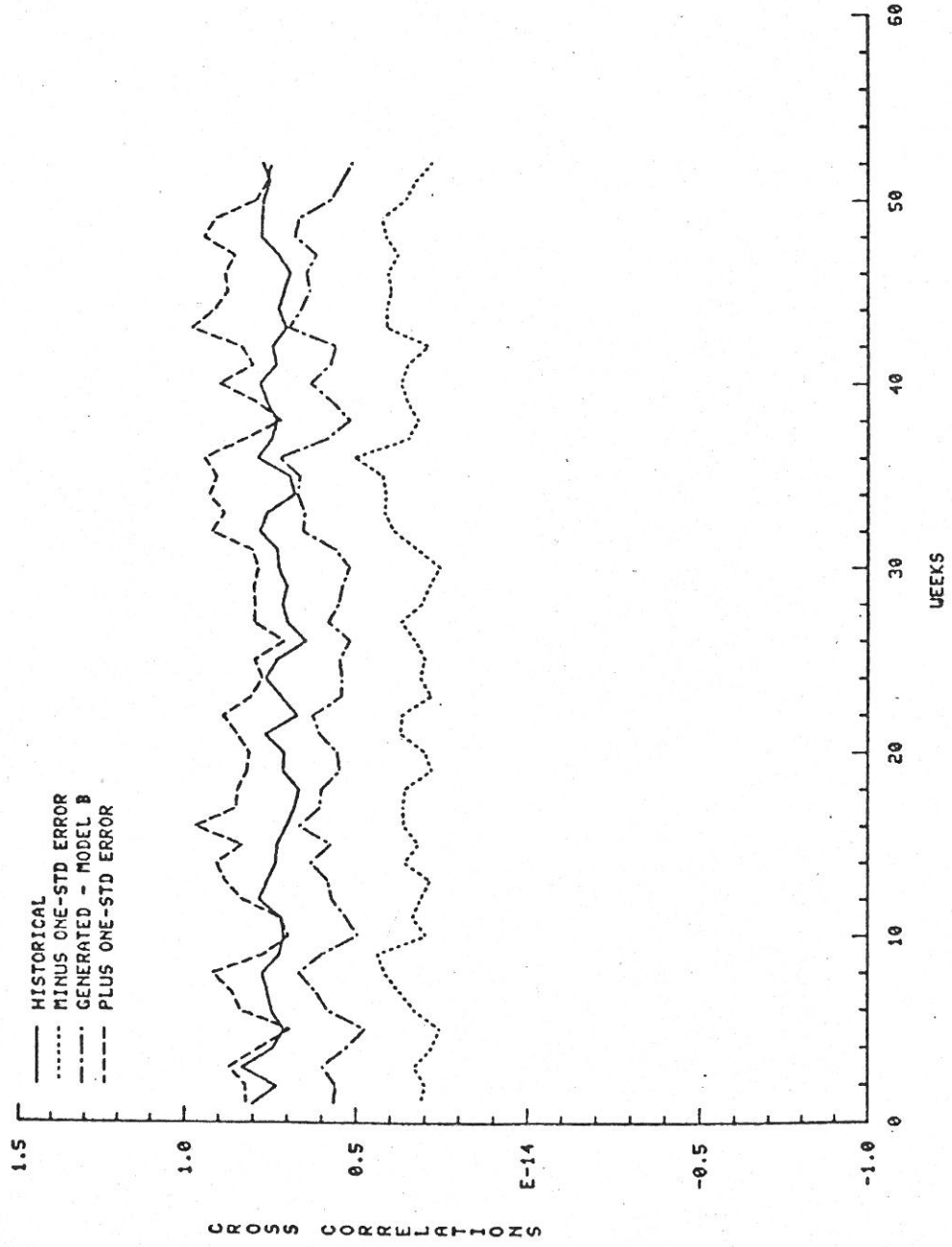


Figure 1.9.E.106. PALODE AND PASODE - CROSS CORRELATIONS (LJH DOMAIN)

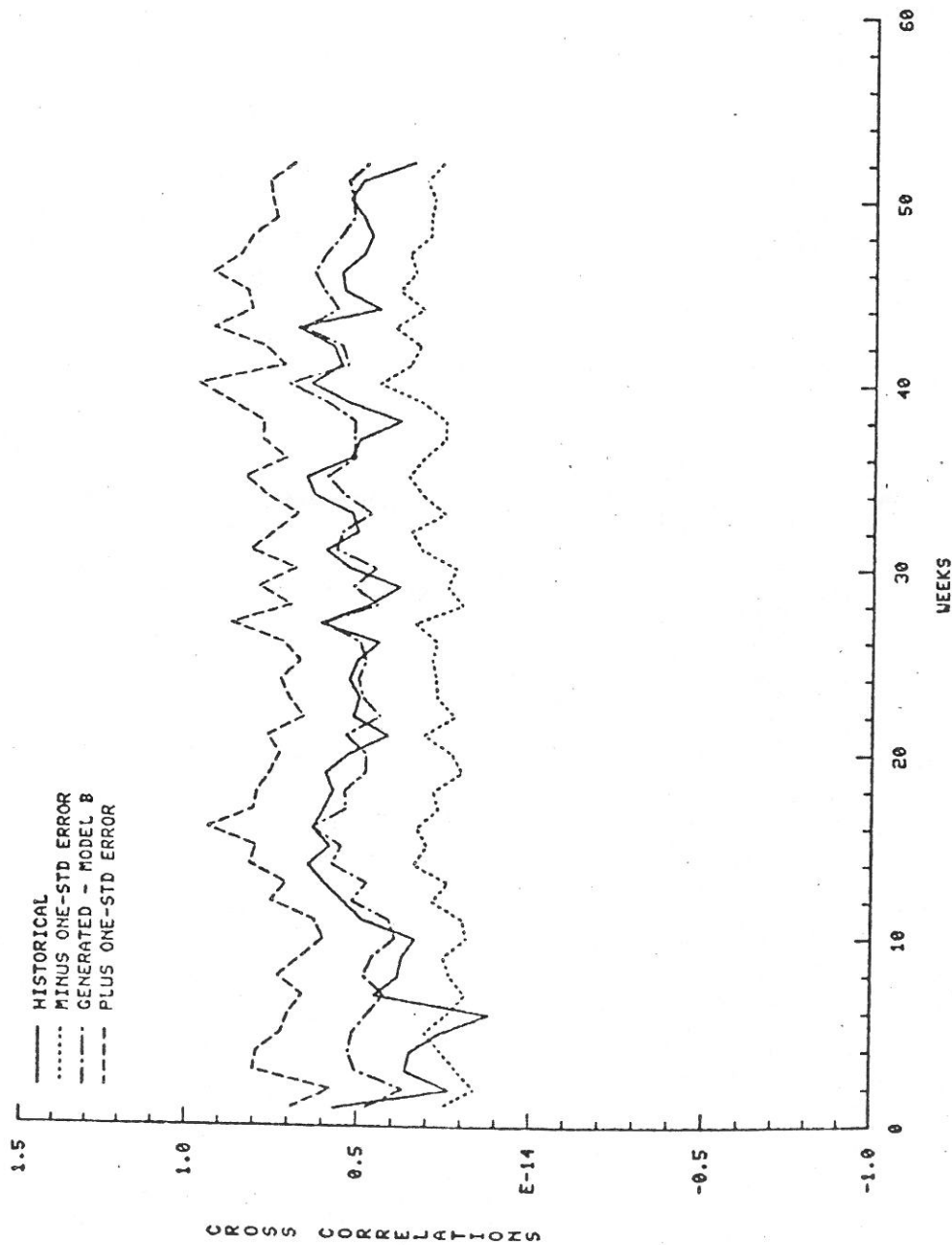


Figure 1.9.E.107. PALODE AND RANCHO - CROSS CORRELATIONS (LUH DOMAIN)

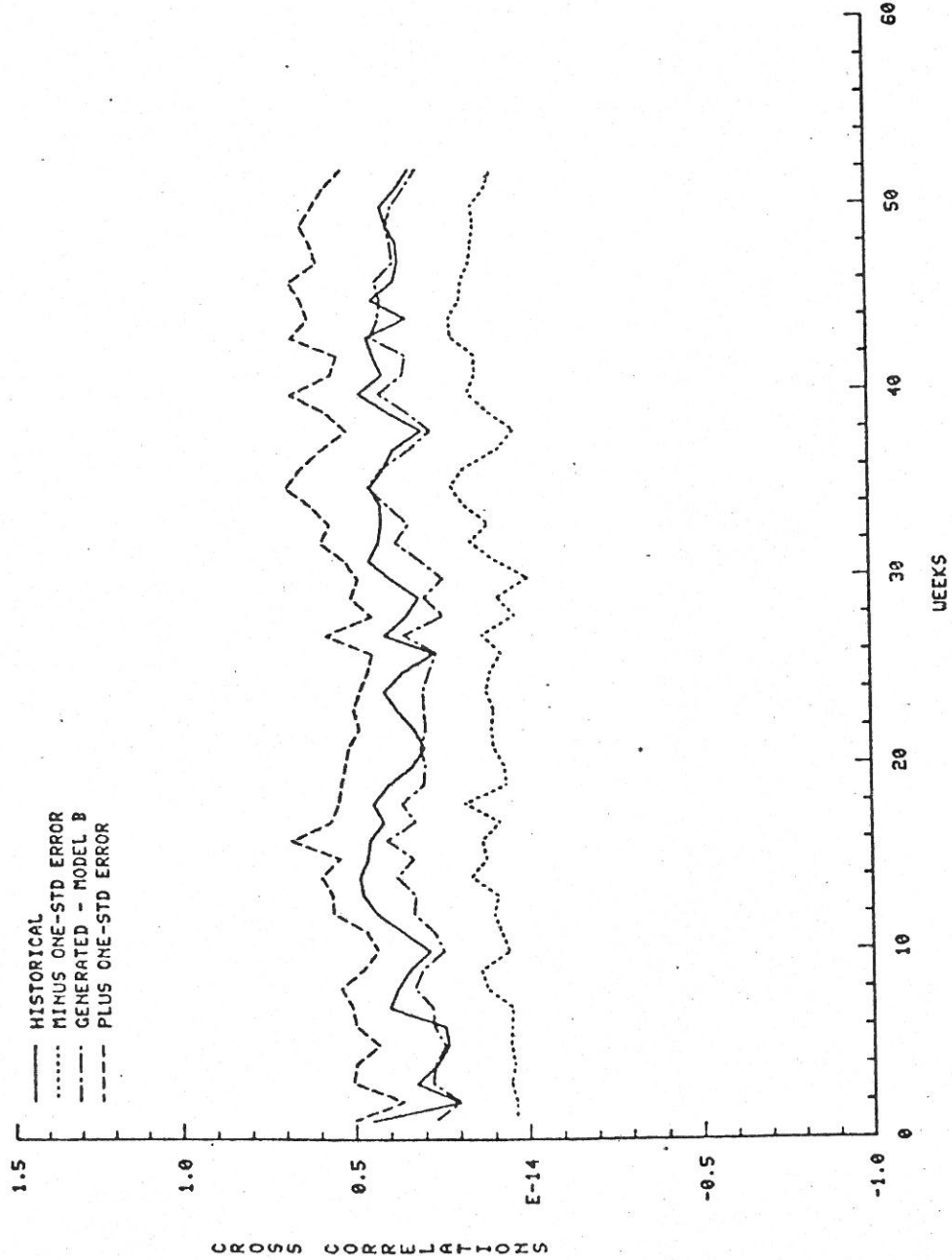


Figure 1.9.E.108. PASODE AND RANCHO - CROSS CORRELATIONS (LUH DOMAIN)

APPENDIX 1.9.F

HISTORICAL AND GENERATED STATISTICS OF MONTHLY TURBINE OPERATING
HOURS TIME SERIES OF VALDESIA RESERVOIR

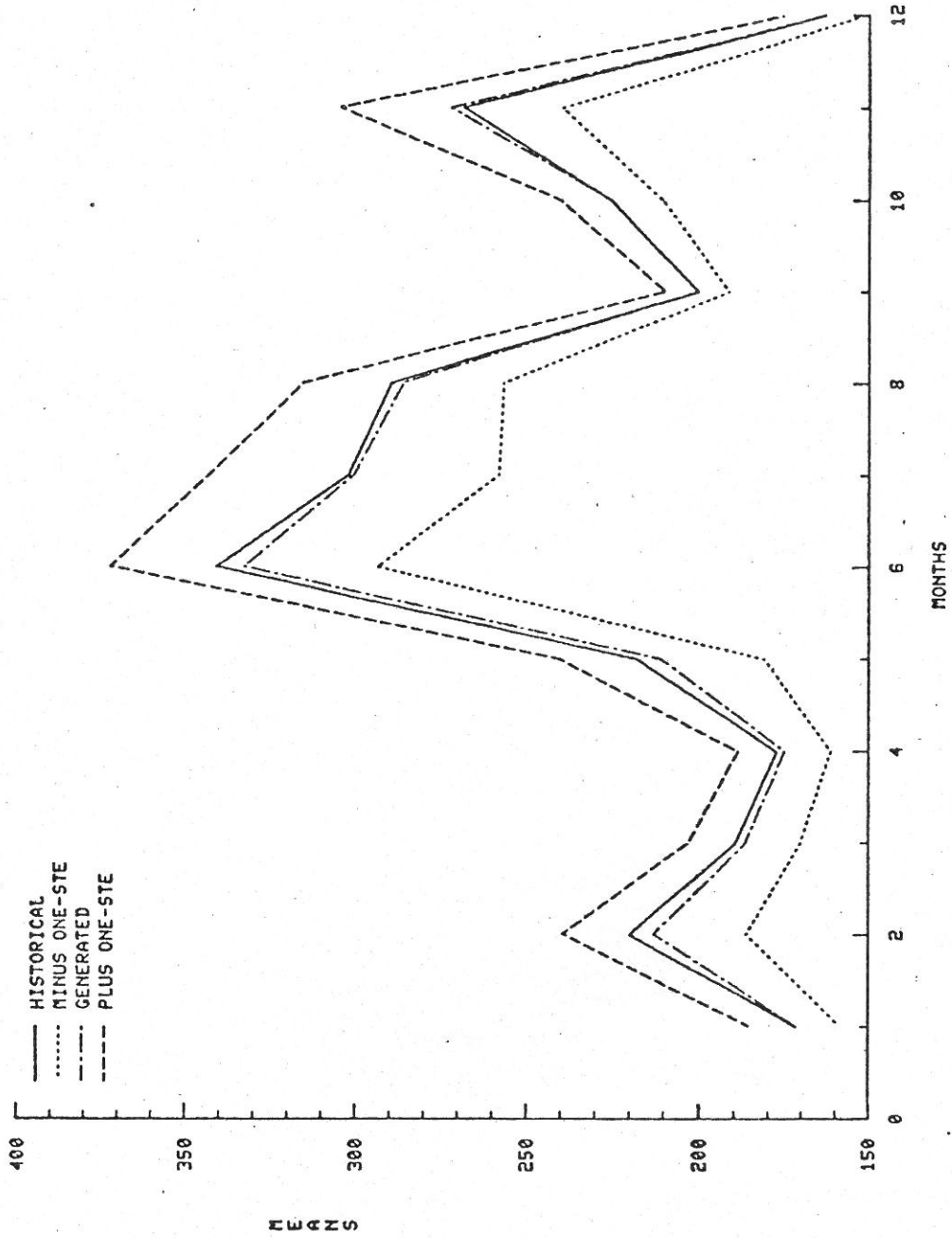


Figure 1.9.F.1. Historical and generated monthly means of turbine operating hours at Valdesia reservoir.

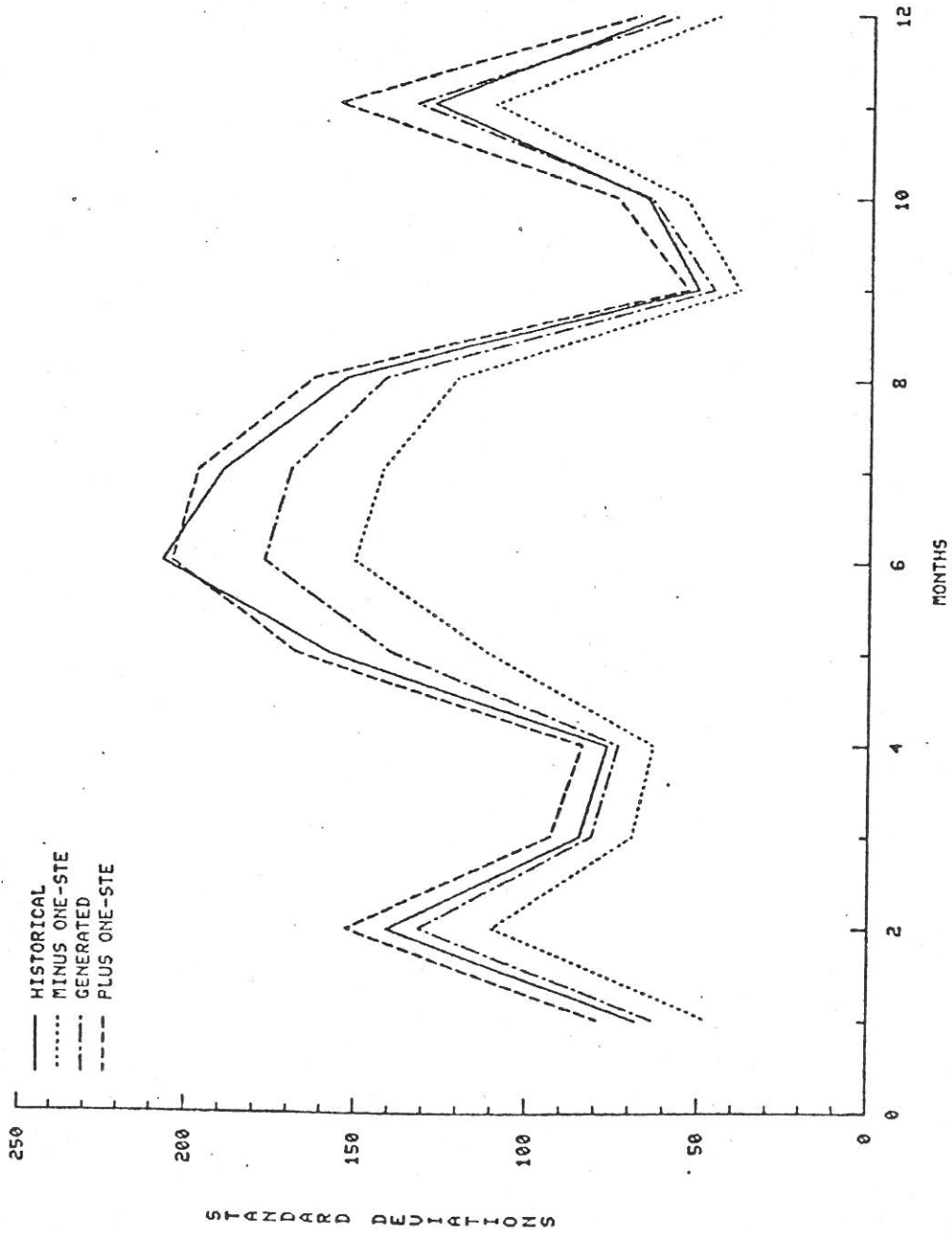


Figure 1.9.F.2. Historical and generated monthly standard deviations of turbine operating hours at Valdesia reservoir.

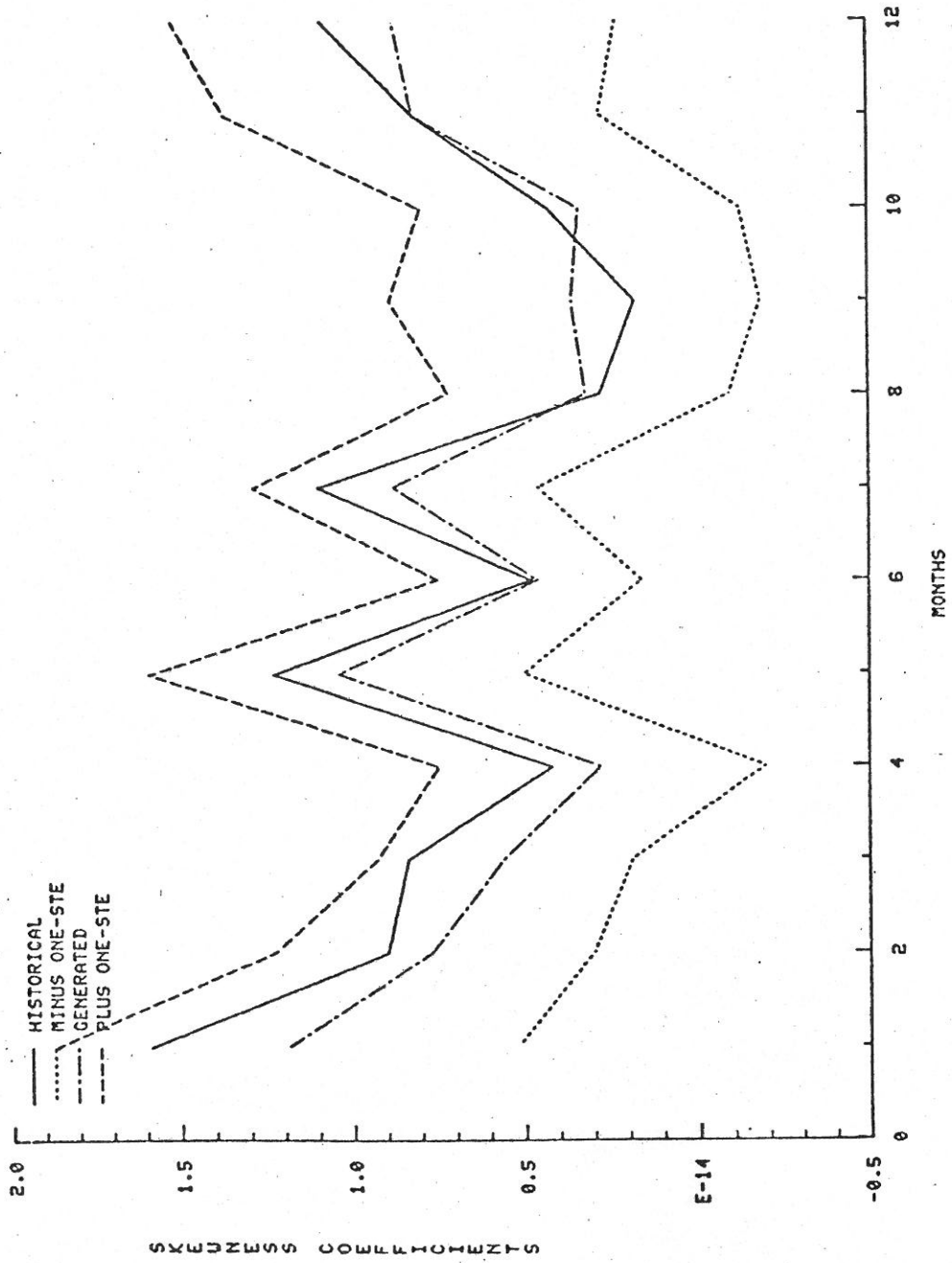


Figure 1.9.F.3. Historical and generated monthly skewness coefficients of turbine operating hours at Valdesia reservoir.

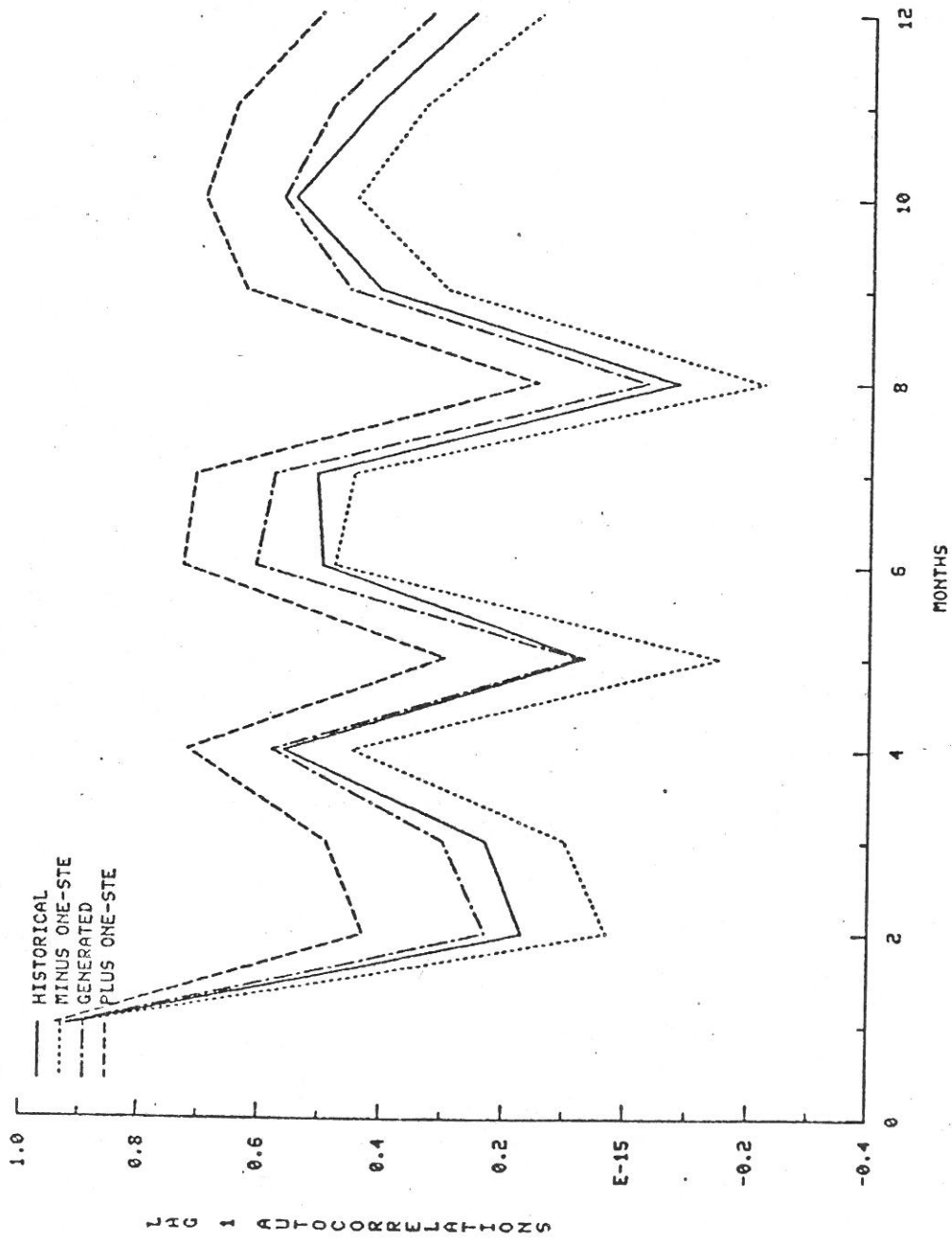


Figure 1.9.F.4. Historical and generated monthly lag-1 autocorrelation coefficients of turbine operating hours at Valdesia reservoir.

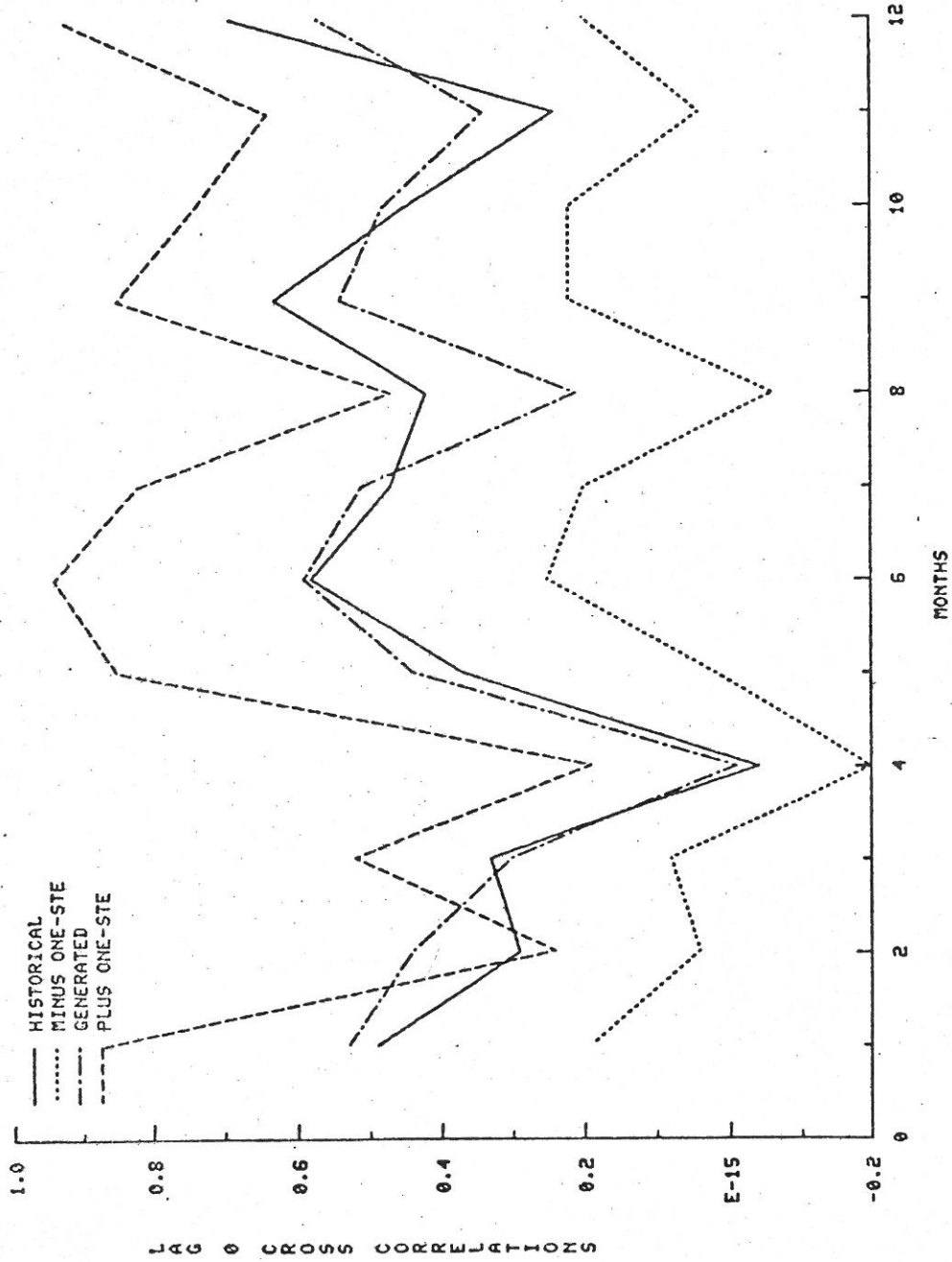


Figure 1.9.F.5. Historical and generated monthly lag-0 cross-correlation coefficients between turbine operating hours and streamflows.